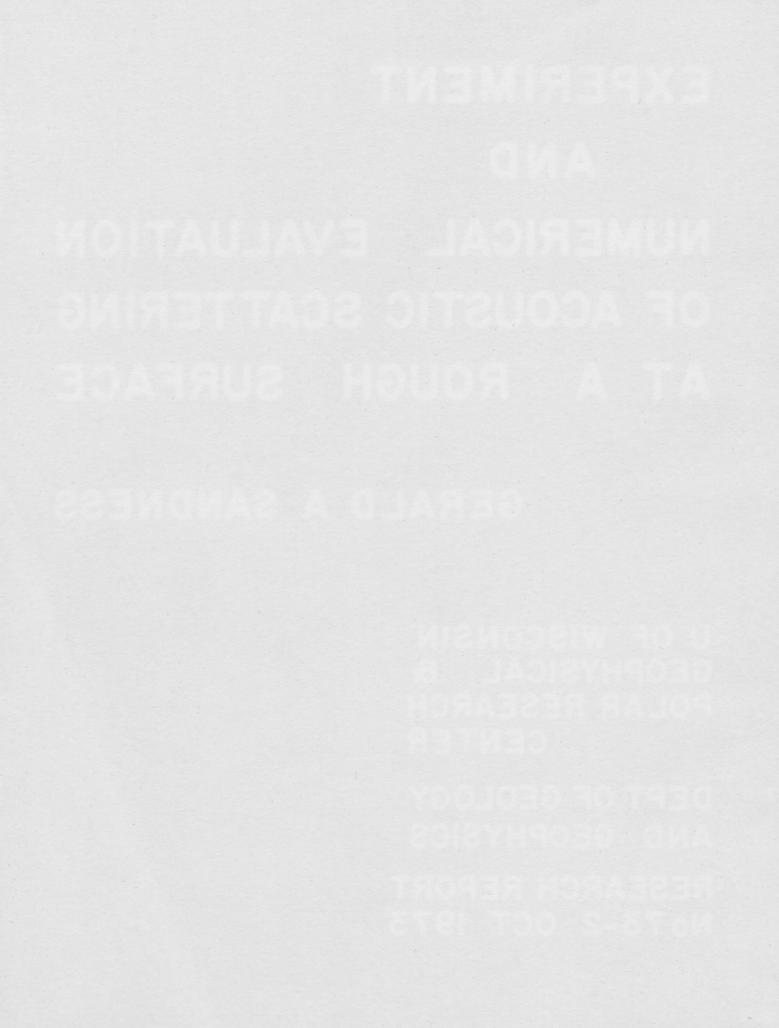
EXPERIMENT AND NUMERICAL EVALUATION OF ACOUSTIC SCATTERING AT A ROUGH SURFACE

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by

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Abstract

The Helmholtz integral is often used as the basis for theories on scattering of acoustic signals from rough surfaces, but several approximations must be made to obtain analytical solutions. Many of the mathematical difficulties in analytical computations can be avoided by numerically evaluating the Helmholtz integral. A numerical technique is described which utilizes measured values of the acoustic pressure of the incident beam, including the side lobes, thereby removing the uncertainty which normally results from approximating the form of the incident beam. A comparison of numerical, analytical, and experimental results is made for the case of a vertically incident beam and a moderately rough surface. The numerical computations are shown to yield values of the scattered pressure which agree well with experimental data for both rough and smooth surfaces.

These initial results indicate that the pressure reflected or scattered from a smooth, or slightly rough surface is sensitive to the phase of the side lobes of the incident beam. The rms pressure scattered from a very rough surface is apparently less sensitive to the form of the approximation used to specify the incident pressure.

The geometrical slope factor, $f(\theta_1, \theta_2)$, is often used in the literature to account for the effect of finite surface slopes. Its validity is verified numerically for moderate scattering angles and a sinusoidal surface, and the limitations on its validity are discussed. A further study of the effect of surface slopes, and a study of the general boundary value problem, including the effects of shadowing and the limitations of the Kirchhoff approximation, are suggested as useful extensions of this work.

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I. Introduction

A. Background.

In the broadest sense, this report deals with the scattering of acoustic and electromagnetic waves from rough surfaces. In this sense the topic involves numerous and diverse areas of research. Without attempting to provide an exhaustive list of these areas, it may be useful to mention a few in order to illustrate the point, and to provide a background for this work.

The first major efforts in studying the effects of rough surfaces on propagating acoustic and electromagnetic waves were in connection with the reverberation or clutter which occurs in sonar and radar returns.¹⁻³ The problem of reverberation is a major one in target detection. A specific example of this kind of a problem is the detection and classification of fish in shallow water or at shallow depths in deep water. Similarly, interference and distortion are produced in communications signals which may impinge on and be affected by the surface of the sea, the sea floor, the surface of the earth, layers in the atmosphere, and so on. In problems such as these, effort is directed mainly toward calculating, eliminating, or compensating for the corruption of information-carrying signals which impinge on a scattering surface. This is the so-called 'direct' problem.

In the 'inverse' problem, the emphasis is on using scattering phenomena to obtain information about the scattering surface. An example of this approach is the attempt to determine by acoustic means the roughness and composition of the sea floor and, ideally, some of

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the properties of the sub-bottom layers. These are the objectives that have been emphasized in this laboratory. There is considerable interest also in scattering from the surface of the sea, with a view toward understanding the growth and propagation of ocean waves and the interaction of the wind and the surface of the water.⁴ Another example is found in the field of meteorology. Both acoustic and radar beams are scattered from layers in the atmosphere in order to study its structure and dynamics.⁵ And finally, radar signals back-scattered from the moon and the planets are analyzed for the purpose of determining some of the statistical roughness parameters of the surfaces of those bodies.⁶

The extent to which either the direct or the inverse problem can be solved depends largely on whether a tractable and complete theoretical description of the scattering process can be derived. One of the more formidable aspects of the theoretical problem is obtaining solutions that are applicable to the case of very rough surfaces (large surface slopes). Another is the inclusion of realistic transmitted beams.

In the following paragraphs a very brief historical view of the development of acoustic theory is given. For a more extensive review of scattering see References 7-9.

The first theoretical treatment of acoustic scattering was apparently by Rayleigh¹⁰ in 1895. He considered the case of a plane wave normally incident on a corrugated surface. He assumed that the scattered field can be expressed as a superposition of plane waves propagating away from the surface. After some controversy in the litera-

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ture, this assumption has been shown to be invalid at points close to the scattering surface.¹¹ More recently, Uretsky¹² has formulated an 'exact' solution for a sinusoidal surface, but his solution is limited to the case of an incident plane wave.

Several papers were published in the early 1950's which laid the foundation for most of the recent work in scattering theory. Several of these papers were published in Russian by Brekhovskikh¹³ and Isakovich.¹⁴ Of special importance in western literature was a paper published in 1953 by Eckart.¹⁵ He adapted the Helmholtz integral^{16,17} - a classical method of physical optics -- to the problem of acoustic scattering from a random surface. He introduced the statistics of the surface into the theory in a simple manner, and obtained the first and second moments of the scattered signal. His basic procedure was shown by LaCasce and Tamarkin¹⁸ to be useful also for corrugated surfaces, and it has subsequently formed the basis for much of the theoretical work found in the literature.

The basic validity of the Helmholtz integral approach has been well established. Its primary virtue is perhaps the fact that it is amenable to approximation techniques which lead to closed-form results. At the same time, its primary shortcoming is that mathematical difficulties <u>necessitate</u> numerous approximations. The degree of success that is ultimately realized in solving the direct and inverse problems with this approach will depend largely on the validity of those approximations.

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B. Approximations and difficulties in acoustic scattering theory.

For the purposes of this discussion, six factors can be enumerated as sources of difficulty in the application of the Helmholtz integral technique to the scattering of acoustic waves from rough surfaces. (The same, or similar, factors arise in the scattering of electromagnetic waves, but the specific concern of this report is with the acoustic case.) Some of these factors are not explicitly labled as approximations, but in each of these cases it will be clear that there is a direct correspondance with some approximation in the theory. The six are:

- 1) The Kirchhoff approximation.
- 2) The incident pressure.
- 3) Phase approximations.
- 4) The zero-slope approximation.
- 5) Description of the surface.
- 6) Shadowing.

These are all further discussed below, but first, it should be pointed out that they are not all independent. For example, 'The Kirchhoff approximation' and 'Shadowing' could be logically combined under the heading 'Boundary values'. Also 'The incident pressure' and 'Phase approximations' could be combined under some more general title. Nevertheless, the present arrangement is convenient, and is also consistent with current treatments of these topics in the literature.

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The Kirchhoff approximation -- Although the Helmholtz integral is formally exact, a fundamental limitation on its power is imposed by the need to know and to specify boundary values; specifically, the acoustic pressure of the scattered signal on the surface, and its derivative with respect to the local surface normal. In the Kirchhoff approximation, these quantities are replaced at each point by the values that they would have if the incident wave were reflected from a plane tangent to the surface at that point. ^{15, 16} Eckart¹⁵ was among the first to apply this approximation to acoustic scattering -- its use was a basic step in his formulation. This approximation is considered to be valid if the minimum radius of curvature of the surface is much larger than the acoustic wavelength.^{6,19,20} This condition is satisfied in the present case and the Kirchhoff approximation is used in both the numerical and analytical computations.

The incident pressure -- The standard procedure in the literature is to represent the incident pressure as a product of an amplitude term and a phase term. The amplitude term includes a directivity function for the incident beam, and for spherical waves, it includes a geometrical spreading factor. The directivity function usually approximates the main lobe of the incident beam by some simple analytical expression such as a Gaussian -- side lobes are ignored. Often, the amplitude of the incident wave is assumed to be constant, or uniform, within some finite ensonified area. Similarly, the phase distribution in the incident beam is approximated by simple expressions which correspond either to plane waves or to spherical waves. Two

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comments can be made regarding the use of these approximations. First, if plane waves or a uniform amplitude is assumed, or if the side lobes are ignored, one would expect the resulting theory to apply mainly to cases in which the scattering surface is of finite size and is completely ensonified by the main lobe of the beam. There are several examples of laboratory experiments using artificial surfaces in which this arrangement has been used. ¹⁸, ²¹, ²² This is probably not a useful procedure to pursue in applying theory to field measurements. Second, for a directional source, the amplitude and phase of the beam are complicated functions of position in the region close to the source. A spherical wave (point-source) approximation is therefore not valid except at large distances from the source. Its validity must be verified for a given application, particularly in laboratory experiments where it is often desirable to use large, highly directional sources.

Recently, in this laboratory, acoustic scattering from a windblown water surface was studied as a function of the width of the acoustic beam.²³ An analytical theory, derived from the formulation of Eckart¹⁵ was used to calculate the mean square pressure of the signal scattered from the rough surface as well as the pressure reflected from the smooth surface. The Fresnel phase approximation was used in the theory in the manner of Gulin²⁴ and Horton and Melton.²⁵ The surface was randomly rough and the spatial correlation functions of the surface are shown in Figure 1b. Computations using these correlation functions gave the solid line shown in Figure 1a. The experimental data are represented by the dashed curve, and the shaded area shows the scatter in the data. These curves and the results

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Figure 1. (a) Normalized mean square pressure versus beamwidth for a wind-blown water surface. p is the scattered acoustic pressure, P_0 is the pressure reflected from the smooth surface, and * means complex conjugate. The beamwidth is the angle between the halfpower points of the incident beam. The scattering angles were $\theta_1 = \theta_2 = 0^0$ (see Figure 2, p.15) and the incident beams were symmetrical about the z axis. The rms roughness of the surface was 0.14 cm, the acoustic frequency was 200 kHz, and the data were measured at depths of 50 and 100 cm. (b) Spatial correlation functions, $\Psi(\xi)$ and $\Psi(\eta)$, in the downwind and crosswind directions, respectively. ξ and η are the corresponding correlation parameters. The correlation functions are defined as follows:

$$\Psi(\xi) = \langle \zeta(\mathbf{x}, \mathbf{y}) \ \zeta(\mathbf{x} + \xi, \mathbf{y}) \rangle_{\mathbf{x}} / \langle \zeta^{2}(\mathbf{x}, \mathbf{y}) \rangle,$$

$$\Psi(\eta) = \langle \zeta(\mathbf{x}, \mathbf{y}) \ \zeta(\mathbf{x}, \mathbf{y} + \eta) \rangle_{\mathbf{y}} / \langle \zeta^{2}(\mathbf{x}, \mathbf{y}) \rangle,$$

where $\zeta(x, y)$ is the surface displacement, and $\langle \rangle_x$ denotes an ensemble average taken over the x coordinate; and similarly for y.

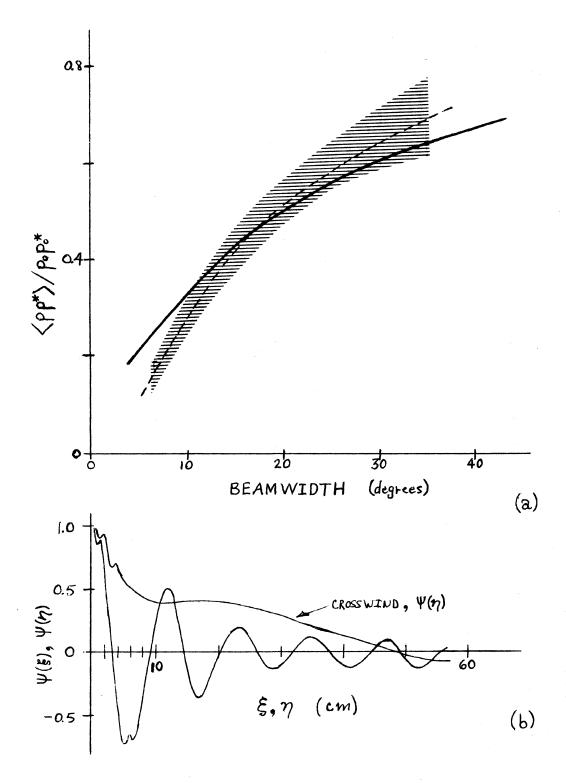


Figure 1. Sound scattered at a randomly rough windblown surface. Vertical incidence.

> (a) Normalized mean square pressure versus beamwidth. Hashed - experiment; solid - theory.

(b) Spatial correlation functions.

 $\sigma = 0.14 \text{ cm}$ f = 200 kHz

of subsequent measurements show a persistent and apparently systematic discrepancy between theory and experiment which worsens as the beamwidth is reduced. The determination of the source of this discrepancy was part of the motivation for doing a numerical integration of the Helmholtz integral. It appeared to be largely attributable to the use of an approximation for the incident pressure. But it was not clear how the errors were apportioned between the calculated pressures for the smooth and rough surfaces. This matter is studied numerically in the present work. Measured values of the incident pressure are inserted directly into the computation in order to avoid the problem of specifying its form mathematically. The normal derivative of the incident pressure is obtained from simple approximations (described in Section II. B.). The accuracy of this procedure is checked by comparing computed and measured pressures of signals reflected from the smooth water surface.

<u>Phase approximations</u> -- A well-known problem in physical optics is the approximation of the phases of the incident and scattered waves in ways which lead to tractable integrals. More specifically, the problem involves approximating the path lengths of the incident and scattered rays. Thus, there is at least an operational distinction between this process and the specification of the phase distribution in the incident beam. Discussions of the applicability and effects of the Fraunhofer and Fresnel approximations in acoustic scattering have been given by Melton and Horton, ²⁶ Horton and Melton, ²⁵ and by MacDonald and Spindel.²⁷ In this work, the Fresnel approximation is used in the derivation of the analytical theory (Section III). The primary emphasis is on phase approximations made in the specification of the incident pressure.

The zero-slope approximation -- The applicability of most scattering theories is limited to surfaces of moderate roughness.⁸ In Eckart's original formulation, and in most of the subsequent variations on that theory, this limitation is due, in part, to the approximation that the surface slope is zero everywhere. A technique for including surface slopes in this kind of a theory has been described for the case of the Fraunhofer approximation by Beckmann and Spizzichino ⁶ and by Tolstoy and Clay.²⁸ They used an approximate integration by parts to evaluate the terms involving the surface slope. The result is a simple geometrical 'slope factor' which multiplies the remaining integral. It is often used for both periodic and random surfaces, but the conditions of its validity are not clear from its derivation. A slope correction curve obtained by using this factor is compared in Section V. B. with a curve obtained from numerical computations. Other theoretical formulations which take account of surface slopes are found in References 29-32.

Description of the surface -- In the Eckart theory, ¹⁵ it is necessary to obtain the bivariate probability density function (PDF) of the surface displacement. If the surface has a Gaussian PDF, the rms scattered pressure can be expressed in terms of a spatial correlation function. In the literature it has been argued that exponential

or linear correlation functions are improper due to the non-zero values of the first derivatives at the origin. 33, 34 In practice, however, it is often observed that these functions yield results that closely approximate measured values. The important point seems to be simply that the scattered pressure may depend strongly on the correlation function near its origin. This is demonstrated in Reference 23 (see Figure 1) where a numerical integration utilizing measured values of the correlation function yielded good results, whereas a quadratic approximation failed badly. Clay has shown, using the same data, that a linear approximation yields equally good results. Also, Clay, Medwin, and Wright have shown that the coherent pressure (the ensemble average, $\langle p \rangle$) is strongly affected by the form of the PDF of a rough water surface. These problems are perhaps not well suited to a numerical analysis. On the other hand, the exact shape of the scattering surface (to the degree that it can be measured) can be utilized in the numerical calculation and its effect can be studied. This is done to some extent in the present work, where the sinusoidal surface used in the analytical theory is an approximation of the real surface.

<u>Shadowing</u> -- At large angles of incidence, portions of a rough surface are not ensonified by the direct transmitted beam. Geometrical shadowing factors have been derived to account for this effect, ³⁷⁻⁴¹ but examples of their use are not numerous. A recent example is a paper by Gardner, ⁴² who considered the effect of shadowing on backscattering from the sea surface at incident angles greater than 80°. He found that shadowing can account for effects previously ascribed to scattering from bubbles or biological scatterers. Clay, <u>et al</u>³⁶ have used Wagner's³⁹ shadowing factor in a study of coherent scattering from a wind-blown water surface. They comment that it seems to work although, in this instance, it is not clear why. It is possible that their result can be explained by the diffraction of sound into the geometrical shadows, and by multiple scattering at the surface. These phenomena have apparently not yet been dealt with and no attempt is made to do so here. A study of shadowing is also beyond the scope of the present work, and shadowing does not occur for the surface and scattering angles used here.

C. Objectives.

The effects of none of the factors just discussed have yet been fully evaluated; and in fact, substantial difficulties are encountered in attempts to do so.

Scattering theories are often tested in the laboratory, where the scattering parameters can be controlled and measured more easily and more accurately than is possible in the field. Yet, even under those circumstances, it is difficult to satisfy assumptions made in the derivation of the theory and to isolate, either analytically or experimentally, the effects of any one of the factors that are sources of difficulty. For this reason it is worthwhile to consider a numerical evaluation of the Helmholtz integral. Most of the mathematical difficulties encountered in an analytical solution do not exist in this case. Although attention must be paid to the errors inherent in numerical computations (truncation and roundoff errors), these

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errors are shown in a later section to be negligible in the computations reported here.

There are two main limitations on the usefulness of the numerical technique. First, the displacement and slope of the scattering surface must be known at each sample point. This suggests that an artificially produced surface be used. Also, restricting the roughness of the surface to a function of only one variable greatly simplifies the problem without altering the basic physics of the scattering pro-Bourianoff and Horton 43 implemented this approach by computer cess. simulation of the sea surface in a numerical study of the statistics of acoustic backscattering. Second, because the integrand is an oscillatory function of surface position, the required number of computations (sample points) may be prohibitively large. That number is roughly proportional to the number of ensonified Fresnel zones on the surface, and can be minimized for a given geometry and signal wavelength by using highly directional transmitted beams. In laboratory experiments those parameters often have values such that numerical computations are feasible. Comparisons of analytical, numerical, and experimental results are thus possible. Because a linear scaling law applies to acoustic experiments, conclusions concerning the effects of the factors enumerated earlier can be expected to be valid for scattering in the sea as well as for scattering in the laboratory.

In this work, acoustic scattering from a periodic surface is studied in order to obtain initial results and to evaluate the practicality and usefulness of extending the work. Measurements of the pressure of the scattered acoustic signals were made in a laboratory

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water tank, with the transmitter and receiver restricted to the x-z (y = 0) plane. Periodic, long-crested waves were produced on the water surface by a mechanical wave generator. The geometry of the experiment is shown in Figure 2. A cross-section of the water waves very closely approximated a sinusoid, so the numerical and experimental results are compared with values obtained from an analytical expression derived for a sinusoidal surface. The numerical computations are limited to the case of vertical incidence $(\theta_1 = 0^\circ)$, but measurements of the backscattered signal were also made and are compared with the analytical results. The emphasis in this work is on the effects of phase approximations, particularly in the incident pressure, and on the effect of neglecting surface slopes.

The chart shown in Figure 3 shows the computations that have been made. The approximations and conditions used in each computation are shown in blocks and are arranged vertically in four columns. Each of the first three columns on the left corresponds to a distinct representation of the incident pressure. The fourth column corresponds to the analytical solution.

Very briefly, the three representations of the incident pressure were chosen for the following reasons: 1) In order to evaluate the numerical procedure, it is important to avoid as many approximations as possible. A set of directly measured values of both phase and amplitude is the best available representation of the incident pressure. 2) The most reasonable analytical representation of the phase of the incident pressure is that of a point

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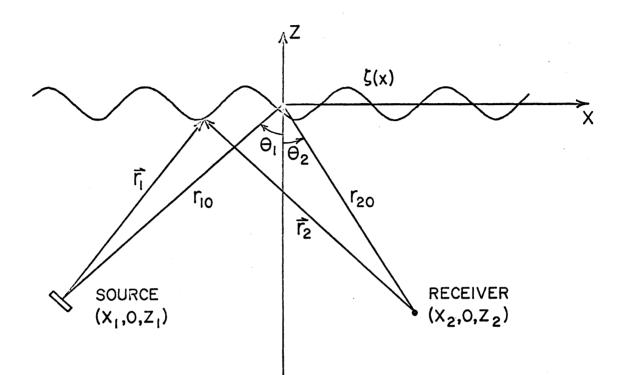


Figure 2. The scattering geometry, shown, for simplicity, in two dimensions. The transmitter and receiver are restricted to the x-z plane, but the surface corrugations effectively extend to $\pm \infty$ in the y direction. The waves propagate in the x direction and are assumed to be periodic.

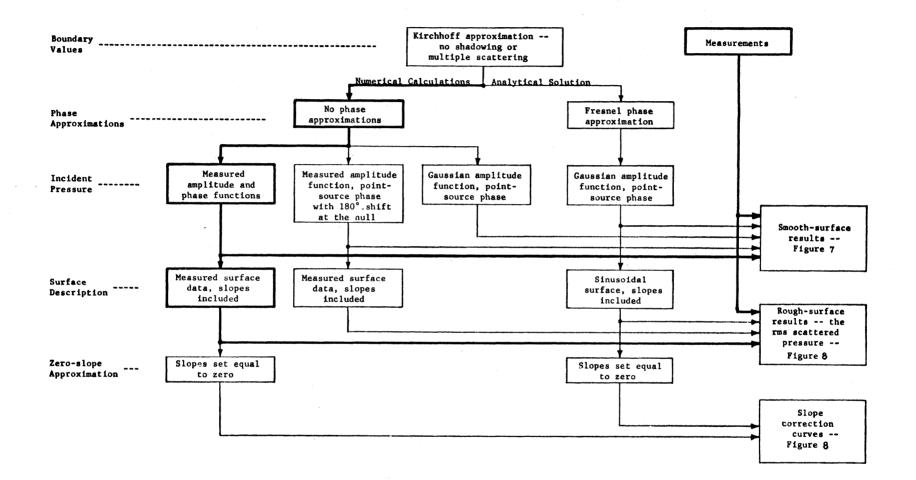


Figure 3. A chart outlining the computations and showing the approximations and conditions used. The relationship of the measured data and the organization of the results are also shown.

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source. But in order to include the side lobes of the beam, it is necessary to add 180° phase shifts at points corresponding to the amplitude nulls between the lobes of the beam. This kind of a function deviates slightly from the measured values (see Figure 5, p. 24). The effect of this deviation is evaluated for both the smooth and rough surfaces. 3) A Gaussian amplitude function and a point-source phase term are often used to specify the incident pressure. The effect of neglecting the side lobes is of particular interest here.

The blocks on the far right represent the graphs shown in Section V. The numerical computation using the measured values of both the amplitude and phase of the incident pressure is of primary interest. For the purpose of reference, this will be referred to as the 'exact' computation, and the parts of the chart which pertain to it are drawn with heavy lines. The results of all of the computations are compared with measured data -- also shown by heavy lines. Further explanations of the computations, and of the notations within the blocks, are given in following sections of the text. Results of analytical computations and measurements for the case of backscattering are also shown in Section V but are not included in the chart.

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II. Numerical evaluation of the Helmholtz integral.

A. Theory

For the case of scattering of an acoustic wave from a pressurerelease surface, the monochromatic Helmholtz integral, in the Kirchhoff approximation, is 6,15,16

$$P = -(4\pi)^{-1} \int_{S} n^{\bullet} \nabla(p'G) dS$$
(1)

where p is the scattered acoustic pressure at a point receiver,

p' is the incident pressure at the scattering surface,

$$G = \exp(ikr_2)/r_2 , \qquad (2)$$

 r_2 is the length of a vector from the observation point,

 $(x_2, 0, z_2)$, to the scattering point,

 ∇ is the gradient operator,

 \underline{n} is a unit vector normal to the surface at the scattering point, S is the surface,

the time factor exp(-iwt) is ignored.

The geometry is shown in Figure 2.

Let the displacement of the water surface be $\zeta(x,y) = \zeta(x)$, and let its x derivative be $\zeta_x(x)$. Then the normal unit vector and the differential surface area are

$$n_{\tilde{z}} = (\zeta_{x\tilde{z}} - z) / (1 + \zeta_{x}^{2})^{\frac{1}{2}}$$
(3)

and
$$dS = (1 + \zeta_x^2)^{\frac{1}{2}} dsdy$$
, (4)

where x and z are unit vectors.

Substituting into Equation 1 and performing the indicated operations yields

$$p = (4\pi)^{-1} \iint_{-\infty}^{\infty} G(p'U + V) dxdy , \qquad (5)$$

where

$$U = r_2^{-1}(ik - r_2^{-1})[\zeta - z_2 - \zeta_x(x-x_2)]$$
, (6)

$$v = \partial p' / \partial z - \zeta_x \partial p' / \partial x , \qquad (7)$$

k is the acoustic wavenumber = $2\pi/\lambda$.

The scattered pressure is computed for many equally-spaced positions of the surface. (For the computations reported here, twenty-three positions were used.) Then the rms scattered pressure is obtained from λ_{c}

$$p_{\rm rms} = (\lambda_{\rm s}^{-1} \int_{0}^{s} p \, p \star dx)^{\frac{1}{2}}$$
, (8)

where λ_s is the wavelength of the surface. Simpson's rule is used to evaluate the integrals in both Equation 5 and Equation 8. A discussion and listing of a computer program to perform this computation are presented in Appendix A (program SCATTER).

The zero-slope approximation consists in setting $\zeta_x(x) = 0$ (i.e., $\underline{n} = -\underline{z}$). The error resulting from the use of this approximation can be easily evaluated numerically. The computation of p_{rms} is performed twice; once using values of $\zeta_x(x)$ computed from measured values of $\zeta(x)$, and again, with $\zeta_x = 0$. The error is the difference of the two results. In the remainder of this report, it will be referred to in terms of a 'slope correction' to be applied to values of the scattered pressure calculated in the zero-slope approximation.

Finally, the pressure reflected from a smooth surface is obtained by setting $\zeta = \zeta_{\rm v} = 0$.

A monochromatic incident signal has been assumed in this derivation. This assumption is permissible in a pulsed-mode experiment if the pulse (or ping) length is sufficiently long that the surface is completely ensonified for a period greater than the duration of the receiver gate aperture. In this experiment, a pulse length of 400 μ sec was used. This pulse length was long enough to satisfy this criterion and yet so short that the wave surface could be regarded as stationary for the period of one measurement. (Frequency shifted components of the scattered signal are not of interest here.)

B. Computational procedure.

Two sets of input data are read into the computer: first, the surface displacement, $\zeta(x)$, and its x derivative, $\zeta_x(x)$; and second, the incident pressure and its x and z derivatives on the mean water surface.

Surface data -- To obtain a representative set of data for use in the computation, the surface displacement was measured in the center of the ensonified area and recorded on analog tape for a period of 5 minutes. That record was then digitized and analyzed to find an 'overall' rms value (the rms displacement for the entire 5 minute record). An 'average' block of 600 data (values of ζ), corresponding to a wave train approximately 60 cm long (1.5 sec., ~ 7 λ s), was then

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selected using the criterion that its rms value be within 0.5% of the overall rms displacement.

The assumption was made that the quantity measured, $\zeta(t, x = 0)$, is insignificantly different from the desired quantity, $\zeta(t = 0, x)$, which is much more difficult to measure. This is justifiable because dispersion and attenuation do not appreciably alter the waveform as it propagates across the ensonified area.

The wave spectrum is sharply peaked at ~ 4.5 Hz and the waveform closely approximates a sine wave with a slight flattening of the troughs and sharpening of the peaks. This is the waveform that one would expect for gravity waves with a very narrow frequency spectrum. In Figure 4, surface displacement probability density functions obtained from the 'average' block and from the entire 5 minute record are compared with the PDF of a cosine wave. The computer program used to compute the PDF's of the surface data is presented in Appendix C.

Each digitized value of the surface displacement corresponds to a scattering point in the numerical integration. The sampling rate was chosen so that the truncation error in a numerical integration of the scattering function in the x direction would be acceptably small. Using the usual estimate of error for Simpson's rule, ⁴⁴ with point spacings $\Delta x = 0.0998$ cm and $\Delta y = 0.2$ cm, the upper bound on truncation errors in the integration is about 1%. The computation is normally performed with single precision arithmetic -- which corresponds to nine decimal digits in the computer used, a Univac 1108. Test computations in double precision arithmetic (eighteen decimal

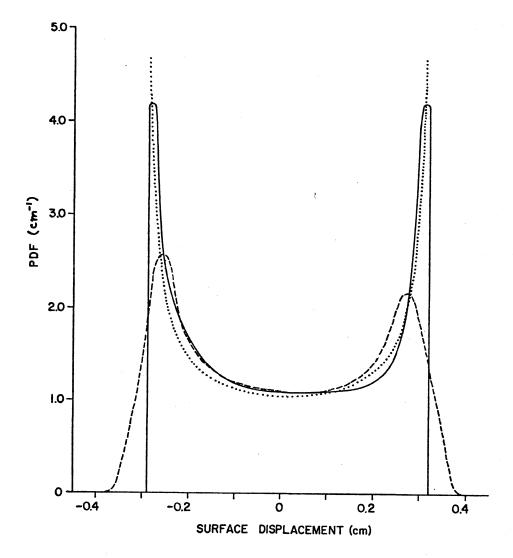


Figure 4. Normalized probability density functions (PDF's) for surface displacement. The solid curve is derived from the 600 work 'average' block; the dashed curve, from the entire 5 minute record; and the dotted curve, from a cosine function having an amplitude of 0.305 cm. The cosine PDF has been shifted on the abscissa to facilitate a comparison with the other curves. digits) have shown that roundoff errors arise only in the fourth significant digit of the computed scattered pressure.

A value of ζ_x was computed for each of the 600 ζ 's in the 'average' block. The surface was sampled 88 times per surface wavelength, so the simplest approximation formula was adequate. That is,

$$\zeta_{xi} \simeq \frac{\zeta_{i+1} - \zeta_{i-1}}{2\Delta x} . \tag{9}$$

For further details of the analysis of the surface data, and of the computation of the x derivatives, see Appendix A, program TRANSLATE.

Incident pressure data -- It was mentioned earlier that for this calculation the transmitted beam is required to be both vertically incident and symmetrical about the z axis. In addition, the depth of the transducer was fixed at 50.1 cm. Consequently, when the water surface is smooth, the incident pressure distribution at the surface is a function only of ρ , where ρ is in the surface and is the radial distance from the beam axis. The incident pressure was measured by placing the transducer at a depth of 100.0 cm and traversing the field 50.1 cm above it with a small hydrophone. Measured values of the amplitude and phase of the pressure are shown in Figure 5.

The measured pressure can be written as

$$p'(\rho, z) = |p'| \exp(i\varphi)$$
 (10)

At each point on the radius the partial derivatives with respect to ρ and z are

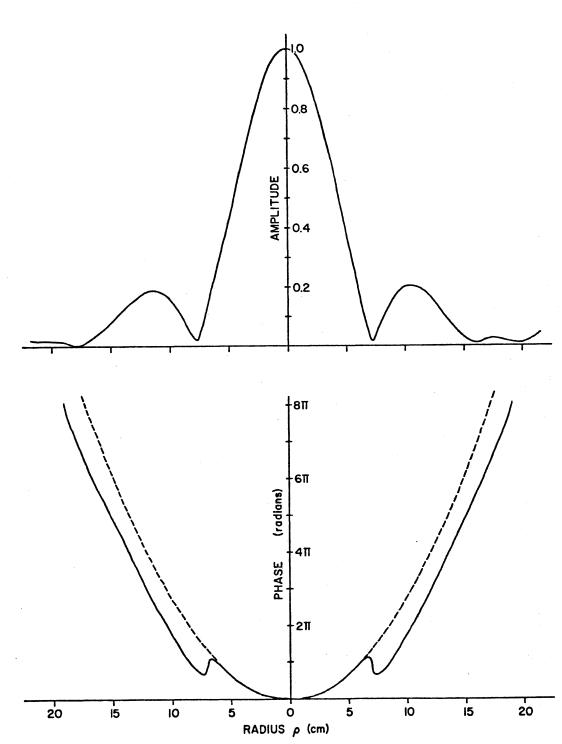


Figure 5. Normalized values of the amplitude and phase of the incident pressure measured in a plane parallel to the surface of the transducer at a distance of 50.1 cm. The dashed line represents the phase of a point source.

$$\frac{\partial \mathbf{p'}}{\partial \rho} = \left[\left(\frac{\partial |\mathbf{p'}|}{\partial \rho} \right)^2 + |\mathbf{p'}|^2 \left(\frac{\partial \mathbf{p}}{\partial \rho} \right)^2 \right]^{\frac{1}{2}} \exp[\mathbf{i}(\varphi + \tan^{-1}w)], \quad (11)$$

$$\frac{\partial \mathbf{p'}}{\partial \mathbf{z}} \simeq \frac{-\mathbf{kr}_{10} |\mathbf{p'}|}{\mathbf{r}_{1}} \exp \left[\mathbf{i}(\varphi + \pi/2)\right] , \qquad (12)$$

where
$$w = |p'| \left(\frac{\partial |p'|}{\partial \rho}\right)^{-1} \frac{\partial p}{\partial \rho}$$
 (13)

At each scattering point the x derivative of p' is obtained from the identity $\partial p'/\partial x = (\partial p'/\partial p)(\partial p/\partial x)$, where $\partial p/\partial x = x/\rho$. The partials with respect to ρ of |p'| and ϕ are obtained from approximation formulas equivalent to Equation 9 (see Appendix A, program PGEN2). The data, shown in Figure 5, were measured at increments of $\Delta \rho = 0.25$ cm. Test computations have shown that this is adequate even for points near the null in the amplitude profile; the scattered and reflected pressures are insensitive to the values of the partials at the null.

At points on the rough surface, the incident pressure and its derivatives are obtained by perturbing the phases of the nearest smooth-surface values. In the neighborhood of each point, the phase shifts are approximately equal to $k \triangle r_1$. This was checked by numerically computing the incident pressure and its derivatives by a method similar to that used by Zemanek⁴⁵ to compute the pressure in the nearfield of a piston source. On this basis, phase errors resulting from the perturbation procedure are estimated to be typically an order of magnitude smaller than the uncertainty (± 0.2 radian) in the measurement of φ . Amplitude corrections were ignored because they are approximately proportional to $\Delta r_1/r_1$, and are an order of magnitude

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less than the uncertainty $(\pm 5\%)$ in the measured values of |p'|.

It is clearly necessary to limit a numerical integration to a finite surface area. For this computation, the area of integration was chosen to include the main lobe and the first side lobe of the incident beam; the maximum radius was set at 17.0 cm. Errors resulting from this choice are likely to be most easily seen in a comparison of computed and measured pressures reflected from a smooth surface. Such a comparison is made in Section V. A. and the error is shown to be small.

III. Analytical solution - sinusoidal surface.

Three approximations are made here in addition to the Kirchhoff approximation already used in the derivation of Equation 5. First, the Fresnel approximation is used in computing the phases of the incident and scattered rays. Second, the geometrical spreading factor, r_2^{-1} , is set equal to r_{20}^{-1} and is removed from the integral. Third, the incident pressure is approximated by the product of a Gaussian directivity function and a point source phase function as follows:

$$p' = \exp\left[-\left(\frac{x \cos\theta_1}{L_x}\right)^2 - \left(\frac{y}{L_y}\right)^2\right] \exp(ikr_1) .$$
 (14)

 L_x and L_y are the linear beamwidths in the x and y directions, measured, as before, in a plane parallel to the face of the transducer and separated from it by a distance r_{10} . As in the numerical calculation, the time factor, exp(iwt), is not explicitly included.

In the Fresnel approximation the total path length of a scattered ray is

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$$r_1 + r_2 = r_{10} + r_{20} + A_1 x^2 + A_2 x + A_3 y^2 + A_4 \zeta$$
 (15)

$$A_{1} \equiv \frac{\cos^{2}\theta_{1}}{2r_{10}} + \frac{\cos^{2}\theta_{2}}{2r_{20}} , \qquad (16)$$

where

$$A_2 \equiv \sin \theta_1 - \sin \theta_2 , \qquad (17)$$

$$A_3 = \frac{1}{2r_{10}} + \frac{1}{2r_{20}} , \qquad (18)$$

$$A_4 \equiv \cos \theta_1 + \cos \theta_2 \tag{19}$$

Substitution of Equations 14 and 15 into Equation 5 yields

$$p = \frac{\exp[ik(r_{10} + r_{20})]}{4\pi r_{20}} \iint_{-\infty}^{\infty} [ikA_4 + (2a_x^2 - ikA_2)\zeta_x]$$

$$\exp[-(a_x^2 x^2 - ikA_2 x + a_y^2 y^2 - ikA_4 \zeta)]dxdy$$
, (20)

where

$$a_{x} \equiv A_{x} \exp(-i\varphi_{x}/2) , \qquad (21)$$

$$a_{y} \equiv A_{y} \exp(-i\varphi_{y}/2) , \qquad (22)$$

$$A_{x} \equiv (L_{x}^{-4} \cos^{4}\theta_{1} + k^{2}A_{1}^{2})^{\frac{1}{4}} , \qquad (23)$$

$$A_{y} \equiv (L_{y}^{-4} + k^{2}A_{3}^{2})^{4}$$
, (24)

$$\varphi_{\mathbf{x}} \equiv \tan^{-1}(\mathbf{k}A_{1}L_{\mathbf{x}}^{2}/\cos^{2}\theta_{1}) , \qquad (25)$$

$$\varphi_{y} \equiv \tan^{-1}(kA_{3}L_{y}^{2}) \qquad (26)$$

It is convenient at this point to also define the quantity

$$v^{2} \equiv 4A_{x}^{2} / (k^{2} \cos \varphi_{x})$$
 (27)

The scattered pressure can be written as the sum

$$\mathbf{p} = \mathbf{p}_{a} + \delta \mathbf{p} \quad , \tag{28}$$

where p_a is the approximate scattered pressure corresponding to the term ikA_4 in the square brackets of Equation 20, and δp is the slope correction obtained from the remaining terms in the brackets.

The scattering surface is

$$\zeta_{\rm m}({\rm x}) = \zeta_{\rm o} \cos K ({\rm x} - {\rm x}_{\rm m}) , \qquad (29)$$

where ζ_0 is the amplitude, and $K = 2\pi/\lambda_s$. The term x_m is provided as a means of translating the surface, and the subscript m is used below to refer to the mth position of the surface. Alternatively, a continuously propagating surface can be obtained by replacing Kx_m by $\omega_c t$, where ω_c is the angular frequency of the surface.

We now use the Bessel function expansion $\frac{46}{3}$,

exp[ia cos K(x-x_m)] =
$$\sum_{n=-\infty}^{\infty} J_n(a) \exp[inK(x-x_m)]$$
, (30)

where

$$a \equiv kA_4 \zeta_0 \qquad (31)$$

Substitution of Equations 29 and 30 into Equation 20 yields the scattered pressure in the small slopes approximation,

$$p_{am} = \frac{kA_4 \exp(i\phi_0)}{4\pi r_{20}} \sum_{n=-\infty}^{\infty} J_n(a) \exp(-inKx_m) \iint_{-\infty}^{\infty} \exp(-ia_x^2 x^2)$$

$$-i(kA_{2} + nK)x + a_{y}^{2}y^{2}] dxdy , \qquad (32)$$

and the slope correction term,

$$\delta p_{\rm m} = \frac{\exp i(\phi_{\rm o} - \pi/2)}{4\pi r_{20}} \sum_{\rm n=-\infty}^{\infty} J_{\rm n}(a) \exp(-inKx_{\rm m}) \iint_{-\infty}^{\infty} (2a_{\rm x}^2 x - ikA_2) \zeta_{\rm x}$$

exp -
$$[a_x^2 x^2 - i(kA_2 + nK)x + a_y^2 y^2] dxdy$$
 (33)

where
$$\varphi_0 \equiv k(r_{10} + r_{20}) + \pi/2$$
 . (34)

Integration of Equations 32 and 33 yields

$$p_{m} = p_{am} + \delta p_{m}$$

$$= \frac{kA_{4}}{4r_{20}A_{x}A_{y}} \sum_{n=-\infty}^{\infty} (1 + \epsilon_{n})J_{n}(a)\exp[-\alpha^{2}(n)] \exp[i\phi_{m}(n)], \quad (35)$$

where

$$\alpha^{2}(n) \equiv (A_{2} + nK/k)^{2}/v^{2}$$
, (36)

$$\varphi_{\rm m}({\rm n}) \equiv \varphi_{\rm o} + (\varphi_{\rm x} + \varphi_{\rm y})/2 - \alpha^2({\rm n})\tan\varphi_{\rm x} - {\rm n}K_{\rm x} + {\rm n}\pi/2$$
 (37)

The term ϵ_n is the fractional slope correction associated with the amplitude of the nth order diffraction lobe. Its form is

$$\epsilon_n = \left(\frac{nK}{kA_4}\right)^2 \qquad . \tag{38}$$

The rms pressure, obtained from Equations 8 and 35, is

$$P_{\rm rms} = \frac{kA_4}{4r_{20}A_xA_y} \left\{ \sum_{n=-\infty}^{\infty} (1+\epsilon_n)^2 J_n^2(a) \exp\left[-2\alpha^2(n)\right] \right\}^{\frac{1}{2}} .$$
 (39)

Computer programs for evaluating Equations 35 and 39 are presented in Appendix B.

As a function of receiver angle, p_{rms} has the form of a diffraction pattern. The directions of the maxima are given by the familiar diffraction grating formula,

$$\sin \theta_2 = \sin \theta_1 \pm nK/k \quad ; \tag{40}$$

and the amplitudes of the diffraction peaks are proportional to the Bessel functions, $J_n(a)$. The zeroth order terms in Equations 35 and 39 are the specular components of the pressure and the mean square pressure, respectively, and do not include slope corrections. The angular width of the diffraction peaks is dependent on the surface wavelength and, via the quantity v, on the beam width, the geometry and the acoustic wavelength. Much of the width is due to the curvature of the acoustic waves. This factor is accounted for in the Fresnel approximation by the term $(kA_1)^2$ in v (see Equations 23 and 27). Further, v is roughly proportional to the angular beamwidth ($\sim L_x/r_{10}$). Analagous to the case of a coarse diffraction grating in optics, broad diffraction peaks are obtained if the diameter of the ensonified area is not much larger than the surface wavelength. Further, for small values of the ratio K/k, the orders overlap and may not be resolved in a measurement of the rms pressure.

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For directions corresponding to diffraction maxima, ϵ_n can be rewritten in the geometrical form

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$$\epsilon(\theta_1, \theta_2) = \left(\frac{\sin \theta_2 - \sin \theta_1}{\cos \theta_2 + \cos \theta_1}\right)^2 \qquad (41)$$

Tolstoy and Clay^{28} define the term $f(\theta_1, \theta_2)$. This is the geometrical slope factor mentioned in the Introduction. Comparing their expression for the scattered pressure with Equation 35, we find that

$$\mathsf{E}(\theta_1, \theta_2) = \mathsf{A}_4[1 + \epsilon(\theta_1, \theta_2)]/2 \qquad (42)$$

Beckman and Spizzichino¹⁸ define a similar term which includes an additional factor of sec θ_1 . The differences are due to definitions and are of no consequence.

In those previous derivations, no distinction was made between periodic and random surfaces, although a geometrical form is perhaps not what one would expect for the slope factor in the latter case. However, this result is consistent with the concept that the pressure scattered in a specific direction from a rough surface is the resultant pressure diffracted from those wavelength components of the surface roughness which satisfy Equation 40.^{1, 47} The value of the fractional slope correction is the same for all of those components since they are all characterized by a constant value of the term nK. But a question of validity arises for the following reason. For the sinusoidal surface the slope correction at a given angle is a resultant value determined by the contributions of overlapping diffraction peaks. This is accounted for by the inclusion of the quantity $1 + \epsilon_n$ within the summations in Equations 35 and 39; and only for very narrow beamwidths is it justifiable to remove that term from the summations.

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In the references referred to above, the term $f(\theta_1, \theta_2)$ is used (for a random surface) as a coefficient external to the scattering integral. Apparently, implicit in this useage is the assumption that the beamwidth is small and that the diffraction peaks associated with each component of the surface spectrum are very narrow. For the beamwidths commonly used, it is not clear that this assumption can be made.

IV. Experiment

The measurements were made in a water tank 1.80 m in diameter and 1.50 m deep. Waves were produced at one end of the tank by vertically oscillating a horizontal, partially submerged, wooden rod 5.0 cm in diameter and 125 cm in length. The driving mechanism was a Slo-Syn stepping motor connected to the rod by a rack and pinion. A mechanical switch connected to the motor shaft was used to reverse the motor. The stepping rate of the motor was constant in the periods between reversals.

Measurement of the wave heights was by resistance probes, a standard method described in the literature.^{23, 48} To calibrate the wave probes, a micrometer head was mounted above the water surface. It was connected electrically in series with a battery, a resistor, and the water in the tank. Touching the water with a probe wire at the end of the micrometer completed the circuit and produced a voltage drop across the resistor. This was observed on an oscilloscope. Wave crests just touching the end of the wire produced a string of pulses on the oscilloscope. The wave height was obtained directly from the difference in micrometer readings for the rough and smooth surfaces. Small variations in the shape and amplitude of the waves occurred because it was not possible to totally eliminate reflections from the tank walls, or to completely stabilize the wave generation process. The mean amplitude of the waves, measured at different points within the ensonified area, did not vary more than the uncertainty in the measurement of the mean value, approximately \pm 5%.

The source transducer was a 5.1 cm diameter Clevite G7 ceramic disk with a thickness mode resonance at 205.6 kHz. The edge and bottom face of the disk were covered with 0.5 cm of styrofoam to eliminate transmissions in unwanted directions. For the measurements at vertical incidence the source transducer was mounted on a horizontal arm and positioned in the center of the tank at a depth of 50.1 cm.

The receiving hydrophones were (approximately) square pieces of ceramic about 1.5 mm on a side. These were cut from a thin disk having a thickness mode resonance at about 1200 kHz. A hydrophone was mounted on a horizontal arm suspended from pivots on the y axis at the water surface. The hydrophone was thus free to move along an arc in the x-z plane (see Figure 2). For these measurements the radius, r_{20} , of the arc was fixed at 49.2 cm.

For backscattering measurements, both the source transducer and a hydrophone were mounted on the pivoted arm, with the hydrophone offset 6° in the x-z plane.

A small piece of styrofoam was attached to the hydrophone on the side nearest the source. This effectively shielded the hydrophone from spurious signals scattered from the source, yet had no measureable effect on signals received directly from the water surface.

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The electronics are straightforward and are described in detail elsewhere.⁴⁹ Figure 6 is a block diagram which shows the basic features of the apparatus for the pulse generation and for the measurement of the mean square signal. Briefly, the signal from the hydrophone is amplified, squared, gated, and integrated. The gate is set to pass only a part of the signal scattered from the water surface, and is controlled by a digital time-delay generator. The output of the integrator is read on a digital voltmeter after a specified number of pulses have been transmitted.

Throughout the experiment, a system calibration was maintained by repeated measurements, at a fixed angle, of the pressure reflected from the smooth water surface. A disadvantage of this procedure is that in order to compare the measurements with theoretical results, a theoretical calculation must be made for the reflected pressure as well as for the scattered pressure.

V. Results

A. Smooth surface

The reflected pressure as a function of receiver angle was calculated numerically using three different representations of the incident pressure. The results are shown in Figure 7.

The solid curve is the 'exact' result obtained using the measured values of both the amplitude and phase of the incident pressure. The accuracy of this calculation is confirmed by the measured values of the reflected amplitude which have been fit by least squares to the calculated curve. These measurements were limited by the size of

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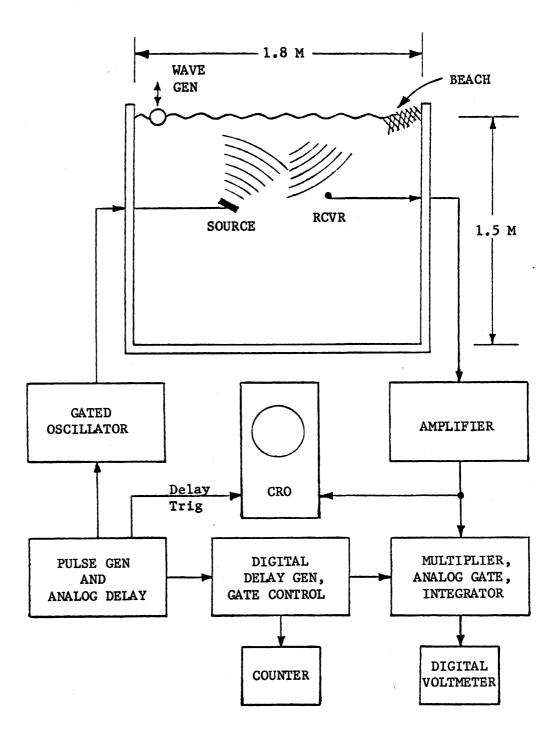


Figure 6. Apparatus for laboratory scattering measurements.

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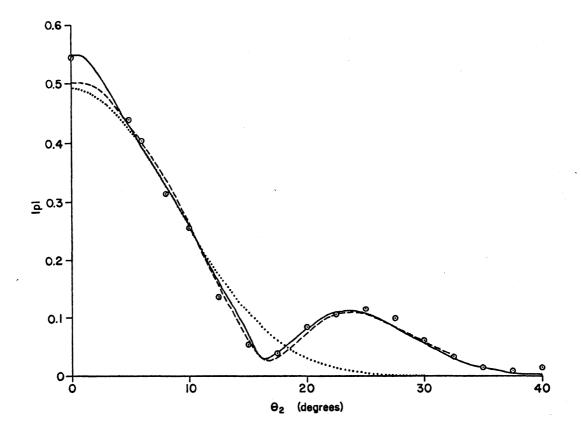


Figure 7. The amplitude of the pressure reflected from the smooth surface. The computed curves were obtained by numerical integration using the measured values of the incident pressure (solid curve), the measured amplitude and a point-source phase with a 180° shift (dashed curve), and a Gaussian amplitude function with a pointsource phase (dotted curve). The circles are measured values. the source transducer to receiver angles greater than about 6°. Therefore, measurements were also made of the direct transmitted pressure using a geometrical configuration consistent with the concept of an image source. These data are indistinguishable from the reflected data, as expected, and the data shown at 0° and 5° were taken from these measurements.

In the second calculation (dashed curve), values of the measured incident amplitude were used, but the phase was approximated by the function kr_1 with a 180° phase shift added at the radius corresponding to the null in the amplitude profile. It is apparent in Figure 5 that approximating the phase in this way is accurate to within 0.5 radian out to a radius of 15 cm, where the amplitude is 95% down from the peak value. This phase error produces an error of approximately 8.5% in the computed amplitude at $\theta_2 = 0^\circ$. The phase errors in this approximation are not too different from the errors in the Fresnel approximation, where, at a radius of 15 cm, the phase error for $\theta_2 = 0^\circ$ is about 0.9 radian.

In the third calculation (dotted curve), the incident pressure was approximated by Equation 14, a Gaussian amplitude function and a point-source phase function. The width of the Gaussian was obtained by fitting it to the half-amplitude points of the measured amplitude profile of Figure 5. The effect of this approximation is to ignore the side lobes of the incident beam, and the resulting error in the reflected amplitude at $\theta_2 = 0^\circ$ is about 10%. One might expect larger errors if fewer Fresnel zones were ensonified by the incident beam. The analytical solution, Equation 35, was also evaluated for this case.

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The essential difference between this calculation and the numerical integration is that the latter does not involve the Fresnel approximation. The difference in the results is very small, however, because the Fresnel phase errors are less than 0.2 radian within the area ensonified by the significant part of the Gaussian beam.

B. Rough Surface

The effective roughness of a scattering surface is often specified by the quantity 50

$$g = k^2 \sigma^2 (\cos \theta_1 + \cos \theta_2)^2 , \qquad (43)$$

where σ is the rms displacement of an ensemble of surface points about their mean. A surface is normally regarded as 'rough' if g > 1. For these measurements, $\sigma = 0.21$ cm, k = 8.73 cm⁻¹, and θ_1 , $\theta_2 < 45^\circ$; so the range of g is approximately 6.8 - 13.6.

Scattering at vertical incidence -- Calculated and measured values of the rms scattered pressure are shown in Figure 8. Calculated values of the slope correction are also shown. These curves were computed using the same representations of the incident pressure that were used in computing the curves of Figure 7. However, here, the dashed curves are the analytical results, obtained from Equation 39; and the dotted curve is the numerical result obtained with the measured incident amplitude and the modified (180° shifts) point-source phase function. The solid curves again refer to the 'exact' numerical computation with the qualification that the solid slope correction curve is the difference between the 'exact' result and the equivalent result with the slopes set to zero.

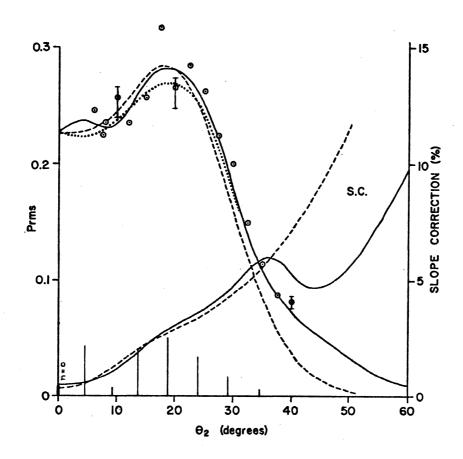


Figure 8. The rms scattered pressure and the slope correction (the curves labled s.c.) at vertical incidence. The solid and dotted curves were computed numerically; the dashed curves are from the analytical solution, Equation 39; and the circles are measured values. Calculated relative amplitudes of the diffraction peaks are shown at the base of the graph.

The agreement of the pressure curves at small values of θ_2 is somewhat better than that of the corresponding curves for the smooth surface. The approximations of the incident beam apparently become increasingly valid as the roughness of the surface is increased. This may be due to a relative decrease in the coherence of the point scattemers illuminated by the side lobes. In a more detailed view, we note that the small peak evident in the numerical curve at $\theta_2 \simeq 4^\circ$ is not resolved in the other two curves. It is clear from Equation 40 that this peak is a diffraction maximum of order n = 1. Failure to resolve it is a consequence of approximating the incident pressure. It was pointed out in Section III that the non-planar nature of the incident acoustic waves has a strong broadening effect on the diffraction peaks of the scattered signal. Comparison of the solid and dotted curves shows that small phase errors, particularly in the representation of the side lobes, can have a significant effect on the computed widths of the diffraction peaks. In the examples shown here, the effect of phase approximations is to reduce the resolution of the individual diffraction peaks, but an erroneous enhancement of the resolution is also possible. For example, it is easy to show that in a Fraunhofer or plane wave approximation the peaks are dramatically narrowed.

Each of the data in Figure 8 is an average of 20,000 - 30,000returns accumulated over a period of 2-3 minutes. The scatter in the data is due to longer-period fluctuations in the amplitude and shape of the water waves, and to possible effects of air bubbles near the surface of the water and on the transducers. The vertical bars on the data points at 10° , 20° , and 40° show the scatter resulting from

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five measurements at each of those angles. It was desirable to forgo longer averaging times in order to minimize problems which tended to arise in long-term measurements -- namely, problems with bubbles and with instability in the rather unreliable wave generator. Nevertheless, on the average, the measured data deviate from the computed values (solid curve) by less than 4%.

The data were multiplied by the same constant used to fit the smooth-surface data to the 'exact' curve in Figure 7. All three curves fit the data reasonably well when $\theta_2 < 20^\circ$. For larger values of θ_2 , the data increasingly diverge from the analytical curve, whereas a good fit is maintained with the numerical curves. This is partially attributed to the effect of the side lobe of the incident beam. In the geometry of this measurement, a portion of the side lobe is forward-scattered in the direction of the receiver with an angle of incidence of approximately 12°. With the aid of Equations 31 and 40 and a table of Bessel functions, it is easy to show that the diffracted orders associated with the side lobe have their maximum amplitudes in the neighborhood of θ_2 = 40°, with orders of lesser amplitude occurring at both smaller and larger values of θ_2 . Small adjustments in the parameters used to calculate the analytical curve can improve its fit with the data in part of the angular range, but not without degrading the fit elsewhere. The amplitude used for the surface roughness was $\zeta_0 = 0.305$ cm, and was taken from Figure 4.

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<u>Backscattering</u> -- The numerical integration was restricted to the case of a vertically incident beam. Therefore, only analytical curves and experimental data are shown in Figure 9. The data have been matched to the numerical curve (computed for vertical incidence) at the point $\theta_1 = 0^\circ$.

If is evident upon comparison of Figures 8 and 9 that the same kinds of comments that were made with regard to the scattering at vertical incidence are also appropriate in this case. First, the agreement between the computed curve and the data is reasonably good, at least at small values of θ_1 . Second, the discrepancy between the computed curve and the datum at $\theta_1 = 2.5^\circ$ suggests again that the analytical solution yields a lesser degree of resolution than that which is physically and numerically realizable. The accuracy of that datum is supported by other data that are not shown here. Finally, side lobe contributions are clearly evident at values of θ_1 greater than 20° .

<u>Slope correction</u> -- Slope correction curves, given in terms of percentage contribution, are plotted in Figures 8 and 9. In Figure 8 it is possible to compare the analytical result with that obtained by numerical integration. The agreement is excellent at receiver angles less than 36°. At larger angles the curves diverge with the numerical curve clearly showing the reduced slope dependence appropriate to the forward-scattered side lobe.

This follows from Equations 38 and 39 where we note that each diffraction order has a slope correction term associated with it. Further, the correction terms increase rapidly with increasing order.

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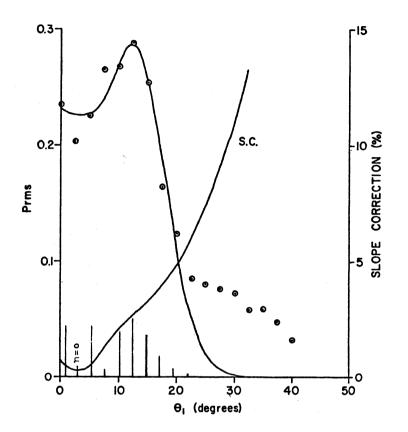


Figure 9. The rms scattered pressure and the slope correction for backscattering. The curves are from the analytical solution, Equation 39, and the circles are measured values. Calculated relative amplitudes of the diffraction peaks are shown at the base of the graph. The prominent side lobe contributions apparent in Figures 8 and 9 consist of low-order components, but occur at the same angles as higherorder main lobe components. Thus, in this angular range, the percentage slope comection will reflect the dominance of the lower-order side lobe components. This is in agreement with an intuitive argument; namely that in forward scattering, the average surface slope tends to zero and results in very small values of the slope correction. Maximum values of the slope correction occur in backscattering where the average surface slope is increasingly biased as the angle of incidence is increased.

According to Equation 41, the fractional slope correction in backscattering increases rapidly at angles of incidence greater than those of the present measurement. At large angles, it may be an important factor in backscattering from the sea surface and sea floor. For example, at $\theta_1 = -\theta_2 = 60^\circ$, a moderate angle, the slope correction is $\epsilon = \tan^2 \theta_1 = 3$. It would be useful to extend the numerical part of this work to backscattering from random surfaces and to a study of the slope correction at large angles of incidence.

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VI. Conclusion

A comparison bf numerical, analytical, and experimental results has been made for the case of a vertically incident beam and a moderately rough surface. The numerical computations have been shown to yield values of the scattered pressure which agree well with experimental data. Simple approximation formulas, applied to values of the incident pressure on the mean scattering surface are apparently satisfactory (within the limits of the Kirchhoff approximation) for computing the scattered pressure and its normal derivative on the rough surface. The use of measured values of both the phase and amplitude of the incident pressure makes possible the accurate calculation of the scattered acoustic pressure for both smooth and rough surfaces. This approach appears to be an effective means for mininizing some of the difficulties inherent in studies of acoustic scattering; particularly, difficulties associated with the specification of the incident pressure and with phase approximations in both the incident and scattered waves. Side lobes of the incident beam are included in the numerical calculations and are not a source of uncertainty in a comparison between theoretical and experimental results.

The present computations indicate that the pressure reflected or scattered from a smooth, or slightly rough surface is sensitive to the phase of the side lobes of the incident beam. Phase approximations result in errors in the calculated pressure that are difficult to avoid in analytical scattering theories. The rms pressure scattered from a very rough surface is apparently less sensitive to

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the form of the approximation used to specify the incident pressure. This is a useful conclusion in terms of both analytical computations and further applications of the numerical technique, and its limitations should be determined.

The validity of the geometrical slope factor, $f(\theta_1, \theta_2)$ (Equation 42), has been verified numerically for moderate scattering angles and a sinusoidal surface. But there are at least two difficulties associated with its use. First, the side lobes of the incident beam are not included in the derivations leading to $f(\theta_1, \theta_2)$ -- so it is not valid at scattering angles where the side lobes contribute significantly to the scattered pressure. Second, its validity is uncertain for the case of a random surface because the effect of the widths of the overlapping diffraction peaks is not accounted for. (This effect is accounted for in the case of a sinusoidal surface.) It is important to extend the study of the effect of surface slopes to backscattering and to non-sinusoidal surfaces.

Other extensions of this work are feasible. A study of the general boundary value problem, which includes the effects of shadowing and the limitations of the Kirchhoff approximation, would probably be most useful. For that purpose, surfaces of greater roughness should be used and the computation should be extended to large angles of incidence. In that case, it may be useful to utilize the pointsource phase approximation. Errors associated with its use are not large in the present case, and they may be reduced by moderately increasing (in comparison with the present value) the distance of the source from the center of the scattering surface. The use of

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Appendix

Computer Programs

The purpose of this appendix is to document the computer programs used in the work described in the text. For the most part, these programs have been written for a specific purpose so are not expected to be useful in a general sense. However, they may be helpful in an extension of the present work or in making similar kinds of computations. A brief description of each program is given in terms of the purpose of the program and the method used. Definitions of terms and details of the programs are given mainly by numerous comment cards within the programs. The programs were written in Fortram V for a Univac 1108 computer.

A. Programs used in the numerical integration of the Helmholtz integral.

The numerical integration is performed by program SCATTER. In addition, two auxilliary programs provide input data: program TRANSLATE provides values of the surface displacement, J, and its x derivative, J_x ; and program PGEN2 provides values of the incident acoustic pressure, p', and its partial derivatives on the scattering surface. The relationships of these auxilliary programs to the measured input data and to program SCATTER are shown in Figure A1. Modifications of the computational procedure will be necessary in order to apply the method of numerical analysis to other geometrical configurations. But it is hoped that the techniques used here will

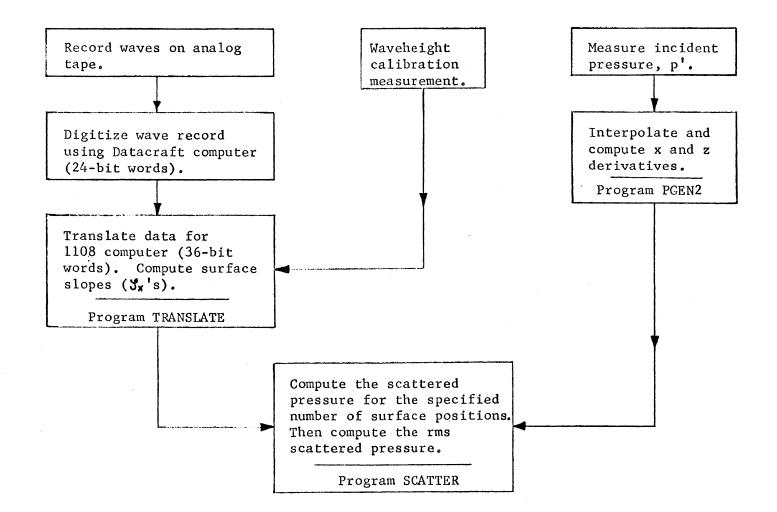


Figure A1. Data, procedures, and programs used in the numerical evaluation of the Helmholtz integral.

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be both useful and instructive in extensions of this work.

<u>Program SCATTER</u> -- The real and imaginary parts of Equation 5 in the text are computed separately, but concurrently. Using Equations 6 and 7, the real part of the scattered pressure is

$$p_{re} = \frac{1}{2\pi} \int_{-X_{max}} \int_{0} \left[T_1 \cos(kr_2) - T_2 \sin(kr_2) \right] r_2^{-1} dy dx \quad ; \quad (A1)$$

and the imaginary part is

$$p_{im} = \frac{1}{2\pi} \int_{-X_{max}} \int_{0} \left[T_1 \sin(kr_2) - T_2 \cos(kr_2) \right] r_2^{-1} dy dx \quad ; \qquad (A2)$$

where
$$T_1 = \left(\frac{kp'_{im}}{r_2} + \frac{p'_{re}}{r_2^2}\right) \left[(x - x_2) \mathbf{y}_x - (\mathbf{y} - z_2)\right] - \left[\mathbf{y}_x \left(\frac{\partial p'}{\partial x}\right)_{re} - \left(\frac{\partial p'}{\partial z}\right)_{re}\right],$$
 (A3)

$$T_{2} = \left(\frac{kp're}{r_{2}} - \frac{p'im}{r_{2}^{2}}\right) \left[(x-x_{2})''_{x} - ('j-z_{2}) \right] + \left[\psi_{x} \left(\frac{\partial p'}{\partial x} \right)_{im} - \left(\frac{\partial p'}{\partial z} \right)_{im} \right]$$
(A4)

The limits of integration are determined by the (arbitrarily defined) dimensions of the ensonified area on the surface. The length of the ensonified area in the x direction is $2 \chi_{max}$. Because of the symmetry which exists when both the source and the receiver

are in the x-z (y=0) plane, the integration on y is performed only over the half-plane y > 0.

The notation used for position coordinates in SCATTER differs slightly from that used in the text and in Equations A1-A4; The source coordinates are $(0, 0, z_0)$; the receiver coordinates are (x, 0, z); and the surface coordinates are (x', y', J). Also, the distance from the scattering point to the receiver is R, rather than r_2 .

The computational procedure described here was chosen because it is direct and because of the symmetry in the scattering geometry. Many variations are possible, and some are probably necessary in order to make similar calculations for other geometrical configurations. Throughout the present program, parallel computations are made for terms corresponding to positive and negative values of the surface coordinate, x'. This may not be a generally useful procedure. Equations A1-A4 can be simplified along with the entire computational procedure if the phase of the incident pressure can be approximated by the function kr1. It was decided to read the input data from punched cards exclusively because the number of input data required in the present computations was not large (402 and 602 cards were supplied by PGEN2 and TRANSLATE, respectively). Data access from tape, drum, or disk files is more efficient than cards for other scattering geometries and for random surfaces.

HAIN 1. C----THIS-IS-PROGRAM-SCATTER. 2 . C IT COMPUTES THE SCATTERED ACOUSTIC PRESSURE AND THE RMS SCATTERED 3. PRESSURE BY NUMERICALLY INTEGRATING THE HELMHOLTZ INTEGRAL. THE 4. C COMPUTATION IS LIMITED TO THE SPECIAL CASE OF SYMMETRY ABOUT THE X 5. AXIS .-- ALSO -- THE OBSERVATION -POINT-IS RESTRICTED-TO -THE-X+2 PLANE+ ð i C C THE PRESSURE DISTRIBUTION ON THE MEAN SURFACE IS REQUIRED TO BE 7. C-SYMMETRICAL-ABOUT-THE-Z-AXIS -- VALUES-OF-THE-INCIDENT-PRESSURE AND-ITS-8. GRADIENT ARE SPECIFIED AT POINTS ALONG A RADIUS, RHO, AND ARE SUPPLIED ON CARDS BY PROGRAM PGEN2. A DIMENSION STATEMENT LIMITS THE 40+ c 11. C NUMBER OF CARDS TO 500. THE SCATTERING SURFACE-IS ASSUMED TO BE PERIODIC. 127 c--13. VALUES OF THE SURFACE DISPLACEMENT AND ITS X DERIVATIVE ARE C 140 SPECIFIED AT POINTS ON THE X-AXIS, AND ARE SUPPLIED ON CARDS BY Ĉ 15. PROGRAM TRANSLATE. С THE SCATTERING POINTS ARE ARRANGED IN ROWS PARALLEL TO THE Y-AXIS 16. c THE COMPUTATIONAL PROCEDURE IS TO FIND THE CONTRIBUTION OF POINTS 17. C ON A ROW, THEN THE ROW, THEN ALL ROWS, INTEGRATION IS BY MEANS OF 18. C C SIMPSON'S RULE. 19. 20-RESULTS ARE COMPUTED FOR SEVERAL OBSERVATION ANGLES. THETA. -FOR EACH c 21. ANGLE. THE SCATTERED SIGNAL IS CALCULATED FOR MANY POSITIONS OF THE SURFACE. 22. C THE RMS SCATTERED PRESSURE IS OBTAINED AFTER THE SCATTERED PRESSURE 23. HAS BEEN-CALCULATED FOR THE SPECIFIED NUMBER OF SURFACE POSITIONS 24--25. C THE SQUARED PRESSURE MAGNITUDES ARE AVERAGED BY ANOTHER APPLICATION 24f-OF SIMPSON+S-RULE-27. C THE ARGUMENTS OF SIN AND COS ARE APPROXIMATED IN THE INNER DO LOOPS 28. TO GAIN SPEED. ARRAYS SSIN AND CCOS ARE COMPUTED AND STORED. ALUES 29. c ARE AT INCREMENTS OF 1. DEGREE. 30. 31. C ALL NAMES PREFIXED BY A 'W' REFER TO VALUES FOR NEGATIVE XP. *P --- YP --- AND-- ZETA -ARE--THE-SURFACE COURD INATES. -32+ 33. THE CARTESIAN COURDINATES OF THE OBSERVATION POINT ARE X,0...Z. c TRE AND TIM ARE THE REAL AND IMAGINARY CONTRIBUTIONS TO THE 34. r 35. C 36. SCATTERING INTEGRAL PROM A SINGLE POINT, XP, YP, ZETA 37. RE AND RIM ARE THE REAL AND IMAGINARY CONTRIBUTIONS OF ONE ROW (AFTER С 38. INTEGRATING WRT YPT PRE AND PIH ARE THE REAL AND IMAGINARY COMPONENTS OF THE SCATTERED c 39. PRESSURE FOR ONE POSITION OF THE SURFACE. 40. С 41. C 42. 43. PARAMETER MW MUST' BE .GE. THE NUMBER OF SURFACE DATA CARDS. c 44. PARAMETER MAXI MUST BE GE. THE NUMBER OF XP+S USED IN THE C CALCULATION. 45. C 96. PARAMÈTER MAXJ MUST BE «GE« THE NUMBER OF SCATTERING POINTS-USED-IN C 47. C THE Y DIRECTION. PARAMETER MNT IS THE MAXIMUM NUMBER OF OBSERVATION ANGLES. PARAMETER MS IS THE MAXIMUM NUMBER OF SURFACE POSITIONS. 48. C 40. C THESE PARAMETERS HUST BE THE SAME IN ALL SUBROUTINES. 50. C 51. c 52. PARAMETER MW=600 PARAMETER MAXI=250, MAXJ=100 53. 54. PARAMETER MNT=14+MS=25 55. c

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56.	c
57	Common/integi/nyP; []; NXP; DXP; DYP; TWoP1
58.	COMMON/INTEG2/TRE(MNT,MAXJ),TIM(MNT,MAXJ),QTRE(MNT,MAX,I)
59.	COMMON/INTEG3/QTIMIMNT,MAXJJJREIMNT,MAXIJ,RIM(MNT,MAXI), QREIMNT
60.	IMAXI) • QRIM(MNT • MAXI)
61.	CGMKON/TERHS/PREIMNTIMSI PIMIMNTIMSI
62. 63.	COMMON/PERTB/SSIN(360),CCOS(360)
64.	COMMON NTHETA,NS DIMENSION XP(M#);ZETA(MW),ZETAX(MW)
65, -	DIMENSION TERMI (MNT), OTERMI (MNT)
66.	DIMENSION ZETHZ(MNT, MW), XPMX(MNT), QXPMX(MNT)
67.	DIMENSION ZEHZINNIJANJARAJANIJ
68.	DIMENSION XP2(MW), YP2(MAXJ), THET(MNT)
69.	DIMENSION RHO(500), PMAG(500) PPHAZ(500), DXMAG(500), DXPHAZ(500)
70.	DIMENSION DZMAG(SOD), DZPHAZ(SOD)
71	DIMENSION PP(HS)
72.	DATA TWOP1/6+2831853/
73.	C
74.	C
75.	C NUMB IS THE NUMBER OF POSITIONS OF THE SCATTERING SURFACE FOR WHICH
76.	C THE HTK INTEGRAL IS EVALUATED. IT MUST BE ODD.
77. 78.	C KSHIFT IS THE NUMBER OF INCREMENTS (DXP) THAT THE SURFACE IS SHIFTED.
	C SET KSHIFT AND NUMB SO THAT KSHIFT. (NUMB-1) DXP 15 AS CLOSE AS
79. 80.	C POSSIBLE TO ALAGTH. C SET IROUGH # D IF THE CALCULATION IS TO BE MADE FOR A SMOOTH SUBFACE.
81.	C SET IROUGH # O IF THE CALCULATION IS TO BE MADE FOR A SMOOTH SURFACE. C THE DATA CARDS FROM TRANSLATE ARE READ EVEN IF IROUGH # D.
82.	C APMAX IS THE MAXIMUM VALUE OF XP AND YP ACTUALLY USED IN THE
83	C CALCULATION.
84.	C DYP IS THE SPACING OF SCATTERING POINTS IN THE Y DIRECTION.
85.	C THETA IS THE OBSERVATION ANGLE WAT VERTICAL
86.	C NTHETA IS THE NUMBER OF THETAS. IT MUST BE LEE MNT.
87.	C RADIUS IS THE DISTANCE FROM THE CENTER OF THE COORDINATE SYSTEM TO
68.	C THE OBSERVATION POINT.
89	C NW IS THE NUMBER OF WAVE DATA READ IN.
90. 91	C DXP IS THE SPACING OF THE WAVE DATA.
92.	C WENGTH IS THE WAYELENGTH OF THE WATER WAVES
93	C ZETA IS THE VERTICAL DISPLACEMENT OF THE SCATTERING SURFACE.
94	C SEE PROGRAM PGEN2 FOR DEFINITIONS OF RHO, PMAG, PPHAZ, DZMAG, ETC.
95	C - C - C - C - C - C - C - C - C - C -
96.	C
97	C INPUT CARDS+ +++++++++++++++++++++++++++++++++++
98.	c c c c c c c c c c c c c c c c c c c
99.	READ 1000, NUMB, KSHIFT, IROUGH
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01	READ-1020, XPHAXIDYP
02.	1020 FORMAT(2F10.0)
03.	READ 1030, NTHETA, RADIUS
04.	READ 1025, (THET(NT), NT=1,NTHETA)
05.	1025 FORMAT(5F10.0)
56.	¢
	C DATA FROM PROGRAM THANSLATE
08.	
10.	READ 1030, NW.DXP.WLNGTH
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12.	READ 1040, KARD,ZETA(1),ZETAX(1)

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13. C 13. A REAL K 13. A REAL K 14. A REAL K 14. A REAL K 14. A REAL K 15. A REAL K 16. A REAL K 16. A REAL K 17. A REA		
135. RETWOINE 137. C 138. PRINT 1070, APMAX. 139. 1070 FORMATIVITHE RADIUS OF THE SCATTERING AREA IS NOMINALLY XPMAX =', 140. PRINT 1080, NADIUS 141. PRINT 1080, NADIUS 142. 108. FORMATIV'T THE OBSERVATION-POINTS LIE ON AN ARC OF RAD' 35+, F4+2, 1 143. 1080 FORMATIV'T THE OBSERVATION POINTS LIE ON AN ARC OF RAD' 35+, F4+2, 1 144. PRINT 1080, NH, DXP, WLNGTH 145. 1080 FORMATIV'T THE NUMBER OF SURFACE DATA READ IN IS NW =', 15, '. THEI 145. 1080 FORMATIV'T THE NUMBER OF SURFACE DATA READ IN IS NW =', 15, '. THEI 146. PRINT 1070, NH, DXP, WLNGTH 147. 2106TH -', F4+3, '. CH.'' 148. PRINT 100 FORMATIV'T THE FIRST AND LAST SURFACE DATA CARDS ARE'/7X, 'CARD ',4 150. 11/2CTMATIC'/ THE FIRST AND LAST SURFACE DATA CARDS ARE'/7X, 'CARD ',4 150. PRINT 1100, (1, 2ETA(1), 2ETA(1), 2ETAX(1)-TIET, NH, FINCT 151. 1100 FORMATIC'/' THE ACOUSTIC WAVELENGTH IS NL -', F6+4, ' CM+'/' THE TRAN 152. PRINT 1100, NHOODAND 153. 1120 FORMATIC'/' THE ACOUSTIC WAVELENGTH IS NL -', F6+4, ' CM+'/' THE TRAN 154. TIBO FORMATIC'/' THE ACOUSTIC WAVELENGTH IS NL -', F6+4, ' CM+'/' THE TRAN 155. 1120 FORMATIC'/' THE ACOUSTIC WAVELENGTH IS NL -', F6+4, ' CM+'/' THE TRAN 156. TIBO FORMATIC'/' THE ACOUSTIC WAVELENGTH IS NL -', F6+4, ' CM+'/' THE TRAN 156. TIBO FORMATIC'/' THE FIRST AND LAST PRESURE DATA CARDS ARE'/7X, 'KARO ', 167. THE SPACING IS DANO -', F7+5', 'CM') 168. TIBO FORMATIC'/' THE FIRST AND LAST PRESURE DATA CARDS ARE', 7X, 'KARO ', 167. THE SPACING IS DANO -', F7+5', 'CM') 164. TIBO FORMATIC'/' THE FIRST AND LAST PRESURE DATA CARDS ARE', 7X, 'KARO ', 165. PRINT TIBO', 'THE FIRST AND LAST PRESURE DATA CARDS ARE', 7X, 'KARO ', 167. THE SPACING IS DANO -', F7+5', 'CM') 166. C ', KA, 'OCPAAZ'/) 167. THE SIGN OF ZO IS CHANGED BECAUSE THE COORDINATE SYSTEM NOW BEING 168. C THE SIGN OF ZO IS CHANGED BECAUSE THE COORDINATE SYSTEM NOW BEING 169. C THE SIGN OF ZO IS CHANGED BECAUSE THE COORDINATE SYSTEM NOW BEING		* • • • • • • • • • • • • • • • • • • •
137. C 138. PRINT 1070, APMAX. 139. 1070 FORMAT(1)THE RADIUS OF THE SCATTERING AREA IS NOMINALLY XPMAX =', 140. PRINT 1080, RADIUS 141. PRINT 1080, RADIUS 142. 1080 FORMAT(//) THE OBSERVATION POINTS LIE ON AN ARC OF RAD')SLIFAL2. 143. 1080 FORMAT(/) THE NOVER WINGTH 144. PRINT 1090, NN,02P, WINGTH 145. 1090 FORMAT(/) THE NOVER OF SURFACE DATA READ IN IS NW =',15,'. THEI 146. PRINT 1090, NN,02P, WINGTH 145. 1090 FORMAT(/) THE NOVER OF SURFACE DATA READ IN IS NW =',15,'. THEI 146. PRINT 1090 147. 21.07TH = ',76.3.' CH.') 147. 21.07TH = ',76.3.' CH.') 148. PRINT 1100 149. 1100 FORMAT(/) THE FIRST AND LAST SURFACE DATA CARDS ARE'/7X.'CARD ',4 150. 1110 FORMAT(IN ,110.2E12-5) 151. 1116 CHART(IN' THE ACOUSTIC WAVELENGTH IS NU =',F6.4.' CM.'/' THE TRAN 152. PRINT 1100, THE FIRST AND LAST PRESSURE DATA CARDS ARE'/7X.'THE TRAN 153. 1100 FORMAT(IN' THE ACOUSTIC WAVELENGTH IS NU =',F6.4.' CM.'/' THE TRAN 154. 1100 FORMAT(IN' THE ACOUSTIC WAVELENGTH IS NU =',F6.4.' CM.'/' THE TRAN 155. 1100 FORMAT(IN' THE ACOUSTIC WAVELENGTH IS NU =',F6.4.' CM.'/' THE TRAN 156. 1100 FORMAT(IN' THE ACOUSTIC WAVELENGTH IS NU =',F6.4.' CM.'/' THE TRAN 157. PRINT 1100, NNDORHO 158. 1100 FORMAT(/' THE ACOUSTIC WAVELENGTH IS NU =',F6.4.' CM.'/' THE TRAN 159. PRINT 1100, NNDORHO 160. PRINT CHE'NORHO '',F7.'', 'CM' BLOW THE HEAN WATER SUMFACE'' 160. PRINT 1100 161. 1100 FORMAT(/' THE FIRST AND LAST PRESSURE DATA CARDS ARE'/7X.'KARD ', 162. 100 FORMAT(/' THE FIRST AND LAST PRESSURE DATA CARDS ARE'/7X.'KARD ', 163. 1100 FORMAT(/' THE FIRST AND LAST PRESSURE DATA CARDS ARE'/7X.'KARD ', 164. 1100 CHARMOSTIC/' THE FIRST AND LAST PRESSURE DATA CARDS ARE'/7X.'KARD ', 165. PRINT 1100. (1,RRO(1),PMAG(1),PMAG(1),DAHAG(1),DAHAG(1),DAHAG(1),DAHAG(1), 165. PRINT 1100. (1,RRO(1),PMAG(1),PMAG(1),DAHAG(1),DAHAG(1),DAHAG(1), 165. C THE SIGN OF ZO IS CHANGED BECAUSE THE COORDINATE SYSTEM NOW BEING 167. C THE SIGN OF ZO IS CHANGED BECAUSE THE COORDINATE SYSTEM NOW BEING	135. REAL K	
<pre>13e</pre>		
140.	138. PRINT-10701-XPMAX	HP EFATTEDING ADEA IS NOMINALLY YOMAY OF.
-42. 1080-FORMATI//-THE -085ERVATION-POINTS L+E-ON-AN-ARC OF RAD'JS+F6+2.1 143. (CH.*) 144. PRINT 1070, NW DXP, WUNGTH 145. 1090 FORMATI//* THE WUNBER OF SURFACE DATA READ IN IS NW **,15,*. THEI 146. 14 SPACING-15 DAP **,F6+4+1_CM+J/* THE *AVELENGTH OF THE *AVES IS * 147. 2140TH **,F6+3,*. CM.*) 148.		The sector is the is the second s
143. 1(M+1) 143. PRINT 1070, NW,DXP,WLNGTH 145. 1000 FORMATL//* THE NUMBER OF SURFACE DATA READ IN IS NW =*,15,*. THEI 146. IS SPACING IS DAP =*,F6.9.4. CM-1/J. THE WAVELENGTH OF THE WAVES IS # 147. 2LNGTH =*,F6.3,* CM.*] 148. PRINT 1100 149. PRINT 1100 149. ID FORMATL//* THE FIRST AND LAST SURFACE DATA CARDS ARE*/7X,*CARD *,4 150. IX,*ZETA *,6x,*ZETAX-*/) 151. INCCMN-1 152. PRINT 1110, (1,ZETA(I),ZETA(I), THE TRAN 152. PRINT 1110, (1,ZETA(I),ZETA(I), THE TRAN 153. IIIO FORMATL//* THE KOUSTIC WAVELENGTH IS WL =*,F6.4.4.* CM.*/* THE TRAN 154. PRINT 1110, NM-0RHO 155. IIZO FORMATL//* THE KOUSTIC WAVELENGTH IS WL =*,F6.4.4.* CM.*/* THE TRAN 156. SIJCO FORMATL//* THE WUMBER OF PRESSURE DATA READ IN IS NRHO =*,15,*. T 157. PRINT 110, NM-0RHO 158. TISO FORMATL//* THE FIRST AND LAST PRESSURE DATA CARDS ARE*/7X,*KARD *, 160. PRINT 1140 161. II40 FORMATL//* THE FIRST AND LAST PRESSURE DATA CARDS ARE*/7X,*KARD *, 162. PRINT 1140 163. ARC************************************		0N-POINTS LIE ON AN ARC OF RADIUS
195. 1090 FORMATI//* THE NUMBER OF SURFACE DATA READ IN IS NW **,15,*. THEI 145. 24.05TH **,FA:3,* CH.*! THE WAVELENGTH OF THE WAVELENGTH OF THE WAVELENGTH OF THE WAVES IS # 197. 21.05TH **,FA:3,* CH.*! PRINT 1100 197. 1100 FORMATI//* THE FIRST AND LAST SURFACE DATA CARDS ARE*/7X,*CARD *.4 190. 1x*'ZETA ** 6x*ZETAX**// 151. 1100 FORMATI(/*) THE FIRST AND LAST SURFACE DATA CARDS ARE*/7X,*CARD *.4 152. PRINT 1110, (1,ZETA*) 153. 1110 FORMATI(/*) THE FIRST AND LAST SURFACE DATA CARDS ARE*/7X,*CARD *.4 154. PRINT 1110, (1,ZETA*) 155. 1120 FORMATI(/*) THE ACOUSTIC WAVELENGTH IS WL **,F6.4,** CH.*/* THE TRAN 156. 1120 FORMATI(/*) THE ACOUSTIC WAVELENGTH IS WL **,F6.4,** CH.*/* THE TRAN 156. 1130 FORMATI(/*) THE FIRST AND LAST PRESSURE DATA CARDS ARE*/7X,***********************************	143• ICH+*)	
144	145. 1090 FCRMAT(//* THE NUMBER OF	SURFACE DATA READ IN IS NW .15, . THEI
146.	146	
150. 11, *2ETA *, 6x, *2ETAX-*/) 151. 114(E*N#-1 152. PRINT 1110, (1,2ETA(1),2ETAX(1), 1ETAX(1), 1ETAX(1), 1 153. 1110 FURNAT(1H ,110,2E12.5) 154. PRINT 1130, NRH0,02HO 155. 1120 FURNAT(//* THE ACOUSTIC WAVELENGTH IS WL =*,F6.4.4.* CM.*/* THE TRAN 156. ISDUCEK 19*,F6.2,* CM BELUW THE THEAN WATER SURFACE++) 157. PRINT 1130, NRH0,02HO 158. TI30 FURNAT(/* THE NUMBER OF PRESSURE D_ATA READ IN 15 NRHO =*,15,* T 159. IHEIR SPACING IS DRHO =*,F7.5.* CM.*) 160. PRINT 1140 161. 1140 FURNAT(/* THE FIRST AND LAST PRESSURE DATA CARDS ARE*/7X.*KARD *, 162. St.*RHO *77.*PMAG *,***PPMAZ *,***DXHAG *,****DXPMAZ*,**** 163. 2AG *.***PPMAZ *,****PPMAZ *,****DXHAG *,****DXPMAZ*,**** 164. 1140 FURNAT(1), THAG (1),PPMAZ(1),DAHAG(1),OXPHAZ(1),DZHAG(1) 164. 1102PHAZIT), 1=1.NRH0TINC; 165. PRINT 1150, (I,RHO(1),PMAG(1),PPMAZ(1),DXHAG(1),OXPHAZ(1),DZHAG(1) 166. 1102PHAZIT), 1=1.NRH0TINC; 167. C THE SIGN OF ZO IS CHANGED BECAUSE THE COORDINATE SYSTEM NOW BEING	148PRINT-1100	
151. INC=NA=1 152PRINT 110. (1,ZETA(1),ZETA(X(1), -1=1,N#,11NG) 153. IIIO FORMAT(1/) IIO.2E12.5) 154		
153. 1110 FORMAT(1H :110:2E12:5) 154. PRINT :120:-mty20 155. 120 FORMAT(7)' THE ACOUSTIC WAVELENGTH IS WL =*,F6.4.* CM.*/* THE TRAN 156. ISDUCER 15*;F6.2.* CM BELUW-THE-MEAN-WATER-SURFACE++) 157. PRINT 1130, NRH0.DRH0 158. 1130 FORMAT(7)' THE FORMAER-OF PRESSURE DATA READ IN 15 NRH0 =*,15.*. 159. 1HEIR SPACING IS DRH0 =*,F7.5.* CM.*) 160. PRINT 1140 161. 1140 FORMAT(7)' THE FIRST AND LAST PRESSURE DATA CARDS ARE*/7X.*KARD ', 162. 164:*RH0-*;7X:*PRHAG *;6X:*PPHAZ *;6X;*DXHAG *;6X:*DXPHAZ *;6X:*DZPHAZ*;6X:*DZPHAZ 163. 2AG ':6X:*DZPHAZ'/) 164. 1110CWTRH0-1 165. PRINT 1150, (I,RH0(I),PMAG(I),PPHAZ(I),DXHAG(I),DXPHAZ(I),DZHAG(I) 166. 1150 FORMAT(7); 10:SEL3:6:F12:5:E13:6:F12:5:E13:6:F12:5] 168. C 167. 1150 FORMAT(7); 20 IS CHANGED BECAUSE THE COORDINATE SYSTEM NOW BEING	151. IINC=NN-1	
<pre>155. 1120 FORMAT(//* THE ACOUSTIC WAVELENGTH IS WL =*,F6.4.* CM.*/* THE TRAN 156</pre>		&TAX+++
156.		WAVELENCTH IS WI MI FARMAN CHAILS THE TRAN
158. 1130 FOPMAT(//*-THE-NUMBER-OF-PRESSURE D _A TA-READ IN IS-NRHO **,15,**-T 159. IHEIR SPACING IS DRHO **,67,*5,* CM**) 160. PRINT 1140 161. 1140 FORMAT(//* THE FIRST AND LAST PRESSURE DATA CARDS ARE*/7%,*KARD *, 162. 167.*RHO *77X;*PMAG *,6X;*PPHAZ *,6X;*DXHAG *,6X;*DXHAZ*,6X;*DZHAZ*,6X; 163. 2AG ',6X;*DZHAZ'/) 164. IINC*NRHO=1 165. PRINT 1150, (I,RHO(I),PMAG(I),PPHAZ(I),DAMAG(I),DXPHAZ(I),DZMAG(I) 166. 102PHAZ(I), I=1,NRHOTIINC; 167. 1150 FORMAT(IH ;110;F10;5:E13)*6;F12:5:E13;*6;F12:5:E13;*6;F12:5] 168. C 169. C THE SIGN OF ZO IS CHANGED BECAUSE THE COORDINATE SYSTEM NOW BEING		
159. IHEIR SPACING IS DRH0 =*, F7.5,* CM.*) 160. PRINT 1140 161. 1140 FORMAT(//* THE FIRST AND LAST PRESSURE DATA CARDS ARE*/7X,*KARD *, 162. 164: *RN0-*T7x;*PMAG *, 64x;*PPHAZ *,64x;*DXHAG *;64x**DXPHAZ*;64x;*DZHAG* 163. 2AG *,64x;*DZPHAZ*/) 164. 1.10C=MRH0=1 165. PRINT 1150, (1,RH0(1),PMAG(1),PPHAZ(1),DXHAG(1),DXPHAZ(1),DZHAG(1) 165. PRINT 1150, (1,RH0(1),PMAG(1),PPHAZ(1),DXHAG(1),DXPHAZ(1),DZHAG(1) 166. 1:DZPHAZT(); 1=1,NRH0;11NC; 167. 1150 FORHAT(1H +:110:F10:5:E13:6:F12:5:E13:6:F12:5:E13:6:F12:5) 168. C 169.	157. PRINT 1130, NRHO, DRHO 158	-PRESSURE DETA-READ IN-IS-NRH0
161. 1140 FORMAT(//* THE FIRST AND LAST PRESSURE DATA CARDS ARE*/7X,*KARD *, 162. 164. 164. 177.5*PMAG *,62.5*PPMAZ *,62.5*DXMAG *,62.5*DXPMAZ*;62.5*DZMAG 163. 2AG *,62.5*DZPMAZ*/) 164. 11NC=NRHO=1 165. PRINT 1150. (1,RHO(1),PMAG(1),PPMAZ(1),DXMAG(1),DXPMAZ(1),DZMAG(1) 166. 10DZPMAZT1), 1=1,NRHOTINC; 167. 1150 FORMAT(1H +,110.5).E13*6*F12*5*E13*6*F12*5*E13*6*F12*5 168. C 167. C THE SIGN OF ZO IS CHANGED BECAUSE THE COORDINATE SYSTEM NOW BEING	159. IHEIR SPACING IS DRHO =", F	
162:	161. 1140 FORMAT(//" THE FIRST AND	
164. IINC=NRH0=1 165. PRINT II50. (I,RH0(I),PMAG(I),DPHAZ(I),DXMAG(I),DXPHAZ(I),DZMAG(I) 165. I;DZPHAZII), I=1,NRH0;IINC; 167. II50 FORMAT(IH ;II0;FI0:5:EI3:6:FI2:5:EI3:6:FI2:5:EI3:6:FI2:5) 168. C 167. C THE SIGN OF ZO IS CHANGED BECAUSE THE COORDINATE SYSTEM NOW BEING	162	▼*₽₽₭₳₴─*ৢ6%ৢ* ₽X₦₳₲──*₮₳₭₮*₽Х₽₦₳₴*ৢ6₭ _₮ *₽₴₦────
167. 1150 FORMAT(1H +110+F10+5+E13+6+F12+5+E13+6+F12+5+E13+6+F12+5) 167. 1150 FORMAT(1H +110+F10+5+E13+6+F12+5+E13+6+F12+5+E13+6+F12+5) 168. C 168. C 169. C THE SIGN OF ZO IS CHANGED BECAUSE THE COORDINATE SYSTEM NOW BEING	169. IINC=NRHO=1	
167. 1150 FORHAT(1H +110+F10+5+E13+6+F12+5+E13+6+F12+5+E13+6+F12+5) 168. C 169. C THE SIGN OF ZO IS CHANGED BECAUSE THE COORDINATE SYSTEM NOW BEING 169. C THE SIGN OF ZO IS CHANGED BECAUSE THE COORDINATE SYSTEM NOW BEING	165. PRINT 1150, (1,RHO(1),PM/	AG(1),PPHAZ(1),DXMAG(1),DXPHAZ(1),DZMAG(1) C)
169. C THE SIGN OF ZO IS CHANGED BECAUSE THE COORDINATE SYSTEM NOW BEING	167. 1150 FORMAT(1H +110+F10+5+E13)	
		ECAUSE THE COORDINATE SYSTEM NOW BEING
		•

	A HEED IS DISEEPENT FROM THAT HEED IN REPUS.
•	C USED IS DIFFERENT FROM THAT USED IN PGEN2.
•	20=-20
	PRINT 1160, DXP, DYP
,	1140 FORMAT (// THE SPACING OF THE SCATTERING POINTS IS DAP # , FA. 4,
	IM IN THE X DIRECTION, AND DYP =+,F6.4, CM IN THE Y DIRECTION.*)
•	
•	C CONSTRUCT 360-ELEMENT ARRAYS FOR THE SIN AND COS APPROXIMATIONS.
•	C
•	ANGL#=DANGL/2.
:	DU 120 J=1,360
	ANGL#ANGL+DANGL
•	SSIN(J)=SIN(ANGL)
	[20 CCOS(J)=COS(ANGL)
	c
•	C MXP AND KSHIFT ARE PARAMETERS USED IN MANIPULATING THE SURFACE DATAS
•	C NXP IS THE NUMBER OF XP*S IN XPMAX.
•	
•	NXP=XPMAX/DXP+1.5
• 2 •	1F(NXP+EQ+(NXP/2)+2) NXP=NXP+1 MXP=NXP
 	na = 0xr hxP0#nxp
	XPMAX=(MXP+1)+DXP
•	140 PRINT 1190, MXP.NUMB.KSHIFT
• ****	IIO FORMAT(//*-MXP-=*+I4+*-SVRFACE-POIN TS-ARE INITIALLY-ASSIGNED TO TH IIE *XP DATA SET+*/* THE H=K INTEGRAL WILL BE EVALUATED FOR NUMB =*;
). /	2131+ SURFACE POSITIONS++/+ FOR EACH SHIPT OF THE SURFACE - HXP WILL
	3 BE INCREASED BY KSHIFT #1,13,1.1)
	C. COMPUTE AND STORE THE X COORDINATES OF THE SCATTERING POINTS,
•	C ALSO COMPUTE, STORE, AND PRINT THE NY'S, WHERE NY(1) IS THE NUMBER OF
•	C SCATTERING POINTS IN THE ITH RON. THEY MUST ALL BE ODD.
•	c
, ,	00 150 J=1 (MAXJ
	150 YP2(J)=((J-1)+DYP)++2
,	DO 155 1=1;MXP XP(1)=(1=1)+DXP
	xP(1)=(]=(]=0xP
	HY=5QRT(XPMAX++2=XP2(1))/DYP+1+5
	1F(HY - EQ. 1) HY=3
•	IF (MY + EQ+ (MY/2)+2) MY=MY+1
,	IF (XP2(1)+((NY-1)+0YP)++2-+6T+-RH0(NRH0)++2) 60-T+-500
,	155 NY(1)=MY
	PRINT 1198
	1198 FORMAT(//* THIS LIST IS THE ARRAY Ny(1).*/)
•	PRINT 1199; INT(1); I=1;HXP)
•	1199 FORMAT(1H ,2016)
	C COMPUTE THE COOPDINATES OF THE OPSERVATION POINTS, AND THE ZETWINS.
	C COMPUTE THE COORDINATES OF THE OBSERVATION POINTS, AND THE ZETHZ S.
	C ZEINZ NEANS "ZEIA MINUS Z"
	156 DO 158-NT#1.NTHETA
	IDO VU IDU NICIA

•	X(NT)=RADIUS+SIN(THETA)
	Z (NT)=-RADIUS+COS (THETA)
	DO 158 I=1.NW
•	158 ZETMZ(NT,1)=ZETA(1)=Z(NT)
	$X_{P}(1) = 00001$
	C A LOOP ON SURFACE POSITION BEGINS HERE.
•	C THE SURFACE IS SHIFTED BY-RETURNING-TO-STATEMENT-160. THE-INGREMENT-
•	C KSHIFT IS ADDED TO MXP, AND THEN THE INDICES OF ZETA AND ZETAX ARE
•	C CHANGED BY THAT AMOUNT.
•	c .
•	
•	GO TO 170
÷	160 NS=NS+1
•	MxP=MXPO+(NS-1)+KSHIFT
• • • •	C ESTABLISH THE SURFACE COORDINATE XP+ THAT IS, BEGIN A LOOP ON THE
•	
•	C INDEPENDENT OF THE COORDINATE YP.
:	C II IS THE INDEX OF A RON-
:	C NNXP IS THE +XP SURFACE DATA INDEX FOR THAT ROW+
	C- NNXQ IS THE CORKESPONDING -XP INDEX-
•	
	NYP=NY(II)
• •	NNXP=MXP+1=11
,	NIXQ=HXP-1+11
	XPMX(NT)=XP(11)=X(NT)
•	TERHI(NT)=ZETAX(NNXP)+XPHX(NT)=ZETM7(NT,NNXP)
•	
•	C
• •	C JJ IS THE INDEX OF A POINT IN THE ROW.
:	00 280 JJ-1,NYP
•	e
	C THE RADIAL DISTANCE FROM THE ORIGIN TO THE PROJECTION OF THE
,	C SPECIFIED AT A SET OF RADII. RHO(1).
·	CPINDTHERHO-CLOSEST TO THE CURRENT SCATTERING POINT+
	C LET 1 BE ITS INDEX.
	C
•	RH2=XP2([1])+YP2(JJ)
,	RH=SQRT(RH2)+++00001
	R1=SQRT(RH2+(ZETA(NNXP)-ZO)++2)
,	QR1=SQRT1RH2+12ETA1NHXQ1=201++2;
	I=RH/DRHO+1.5
	PGENR=SQRT (RH01(17)=2+20=+21
•	DF=K+(R1-PGENR)
• ``	QDF=K+(QR1=PGENR)
•	
•	C OBTAIN THE INCIDENT PRESSURE AT THE SCATTERING POINTS BY PENTURBING
•	C THE PHASES OF THE SMOOTH-SURFACE VALUES SUPPLIED BY PGEN2. C Subroutine-Prtb-Perform s This function.
•	
	C C C C C C C C C C C C C C C C C C C

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•	P=PMAG(1)
	P2=D2MAG(1)
•	P X = U X M A G (1)
•	Fz=DZPHAZ(1)
	FX=DXPHAZ(I)
•	CALL PRTD(P,F,DF,PKEP,PIMP)
	CALL PRTBIP, F. QDF, QPREP, 4PINP) CALL PRTBIPZ, FZ, DF, DZREP, 0ZIMP)
	CALL PRTB (PZ, FZ, 4DF, 9DZREP, 9DZIHP)
•	CALL PRTB(PX,FX,DF,DXREP,DXIMP)
•	CALL PRYBIPX,FX,4DF,4DXREP,40X1HP) C
	C TERM IS THE SINE OF THE ANGLE BETWEEN AN AND THE Y ANIS. IT IS
•	C NEEDED TO COMPUTE THE PARTIAL OF P WRT X.
,	
	TERH=XP(11)/RH
•	DXHEP=TERM+DXREP
•	
• •	DXINP=TERM+DXIMP QDXIMP=-TERM+QDXIMP
•	c
•	C COMPUTE INTEGRAND TERMS WHICH DON'T INVOLVE THETA.
	C TERMY#ZE1AX(NNXP)#DXREP=DZREP
•	GTERM4=ZETAX (NHXQ) +QDXREP-QDZREP
	TERM5=ZETAX(NNXP)+DXIMP+DZIMP QTERM5=ZETAX(NNXQ)+QDXIMP+QDZIMP
•	
•	C LOOP ON THE VALUES OF THETA (OBSERVATION POINTS).
•	DO 280 NT=1,NTHETA
	C R2 MEANS R-SQUARED.
•	R2=XPMX(NT)++2+YP2(JJ)+ZETMZ(NT,NNXP)++2
•	₩2=QxPmxtntt=+YP2{ JJ}+ZETHZINTyNnxQ}++2 R=Suktir2]
•	QR=SQHT(QR2)
•	TERH2=PREP/R2+K+PIHP/R
•	─────────────────────────────────────
	QTEXH3=K+QPRCP/QR-QPIHP/VK2
•	TI=TERM1(NT)+TERM2-TERM4
•	QT1=QTERM1(NT)=QTERM2=QT ERM4 T2=TERM1(NT)=TERM3+TERM5
	QT2=QTERM1(NT) • QTERM3+QTERM5
•	c
	C FIND THE PHASES OF THE SCATTERED RAYS. C REMOVE INTEGER MULTIPLES OF TWOPI FROM KOR, THEN APPROXIMATE SIN AND
	C COS BY SSIN AND CCOS.
•	C THE FUNCTION AINT REMOVES THE FRACTIONAL PART OF G.
	G=R/XL
	40+64/WL
•	NG=(G-AINT(G))+360.+1.
•	NDG#(QG=#INY(QG))#360++1+ S=SSIN(NG)

•

· · · ·

41.	QS=SSIN(NQG)
12	C=CCO\$(NG)
13.	QC=CCOS(NQG)
44	
45.	C THE FOLLOWING TERMS ARE VALUES OF THE INTEGRAND AT THE SCATTERING
46	C POINTS.
47.	c
48	TRE(NT,JJ)=(T1+C+T2+S)/R
49.	$QTRE(NT,JJ) = (QT1 \circ QC + QT2 \circ QS) / QR$
50	TIM(NT,JJ)=(TI+S=T2+C)/R
51.	280 QTIM(NT, JJ)=(QT1*QS-QT2*QC)/QR
52 53.	C INTEGRATE WET VO. THAT IS FIND THE CONTRIBUTION OF THE DOWE AT
	C INTEGRATE WRT YP. THAT IS . FIND THE CONTRIBUTION OF THE ROWS AT
54	C
55.	C
56	C 290 CALL- INTGY2
58	CINTEGRATE-WRT-XP+THAT-IS-INTEGRATE-NVER-ALL-THE-ROWS-TO-OBTAIN-THE
59.	C SCATTERED PRESSURE FOR THE CURRENT SURFACE POSITION.
60.	
61.	CALL INTGX2
62.	
63.	C ALL TERMS HAVE BEEN COMPUTED FOR THIS SURFACE POSITION.
64.	C EITHER SHIFT THE SUNFACE AND REPEAT THE CALCULATION, OR STOP
65.	C C
66.	IF(NS .EQ. NUMB) 60-TO-400-
67.	GO TO 160
68	
69.	C FOR EACH RECEIVER ANGLE FIND THE MAGNITUDE OF THE SCATTERED PRESSURE
70	
71.	C THE P2'S AND TAKE THE SQUARE ROOT TO GET PRMS.
72	
73.	400 D0 460 NT=1.NTHETA
74	
75+	// 1200 FORMAT('!=************************************
77.	PRINT 1290
78	
79.	18x, *PF**/)
80	D0-440-NS=1-NUMB
81.	A #PRE(NTINS)
82.	B=P1M(NT,NS)
83.	P2=A+A+B+B
84	PHAGN=SQRT (P2)
85.	PP(NS)=P2
86	CALL ANGLE (A) BIPHASE
87.	440 PRINT 1300, NS.A.B.PMAGN.PHASE.P2
89.	1F(NUMB +LT+ 3) GO TO 460
90. 91.	C FIND PHMS AND PRINT IT.
72.	
93.	H=1./(NUMB-1)
74.	CALL INTEGLIPPINUMBIHIPMS - PRMS = SQRT (PMS)
95.	TRINI 1310) FRN3
96.	1310 FORMAT(82%)F14491+ # PRN5+1
97.	460 CONTINUE

----398. STOP c-----399. 400. 401+ STOP 402. 510 PRINT 1330 1330 FORMAT(///' SURFACE DATA CARDS NOT IN ORDER. QUIT.") STOP 403. .404. 405. 406. 468. STOP -407. END 410. С NO DIAGNOSTICS. END OF COMPILATION: , ï

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1.	c	
2.		SUBROUTINE INTG2
3.	ç	
4		IS SUBROUTINE USES SIMPSON'S RULE TO INTEGRATE THE SCATTERING
5.	C TE	RMS CORRESPONDING TO ROWS (ENTRY INT _g y2) and strips (entry intgx2).
	C	
7.		PARAMETER MNT=14+MS=25 PARAMETER MAXI=250+MAXJ=+00
8		COMMON/INTEGI/NYP,IIINXPIDXPIDYPITWOPI
9. 10		-COMMON/INTEGI/TRE(MNT,MAXJ),TIM(MNT,MAXJ),GTRE(MNT,MAXJ)
11.		CUMMON/INTEG3/GTIM(MNT, MAXJ), RE(MNT, MAXI), RIM(MNT, MAXI), GRE(MNT,
12.		1MAX1 + GRIM (MNT + MAX1
13.		COMMON/TERHS/PRE(MNT,MS),PIN(MNT,MS)
14		- COMMON-NTHETATNS
15.	c	
10		- ENTRY-INTGY2
17.		IMAX=NYP=3
18.		- #=DYP/3,
19.	C	
20		-Do 200 NT=1 NTHETA
21.		FI=TRE(NT,1)+TRE(NT,NYP)
22.		-G1=T1M(NT;);+T1M(NT;NTP)
23.		4F1=QTRE(NT,1)+QTRE(NT,NYP)
24.		-4G1=4T1M(NT,1)+4T1M (NT,NYP) F2=TRE(NT,NYP=1)
26		-62=T1H(NT\$NYP=1)
27,		OF2=GTRE(NT,NYP-1)
28		- QG2=QTIN(NT,NYP=1)
29.		F3=0.
30	· · · · · · · · · · · · · · · · · · ·	-63=0.
31.		9F3=0.
32		- QG3=0+
33.	C ·	
34.		-D0 100 1=2:1MAX;2
35.		F2=F2+TRE(NT,1)
36		- G2=G2+T1M(NT,1)
37.		4F2=QF2+QTRE(NT,1) QG2=QG2+QT1M(NT,1)
38		F3=F3+TRE(NT+1+1)
39.		- 63=63+11H(NTTI+1)
40. 41.	•	QF3#QF3+QTRE(NT,1+1)
42		- 463=063+071M(NT, 1+1)
43.		RE(NT,11)=w+(F1+4,+F2+2++F3)
44		R1H(NT,11)=#+(61+4++62+2++63)
45.		GRE(NT, 11)=#+(GF1+4.+GF2+2.+GF3)
461	200	- ORIMINT; 111 ##+1961+4++992+2+4493}
47.		RETURN
48.	с — с	
49.		ENTRY INTGX2
50.		1MAX=NXP=3
51.		N=DXP/3.
52.		
53.		DO 400 NT=1.NTHETA
540		$F_{I}=RE(NT, I \rightarrow RE(NT, NXP)$ GI=RIM(NT, I)+RIM(NT, NXP)
55.		V4-n4/11/14/14/14/14/14/14/14/14/14/14/14/14

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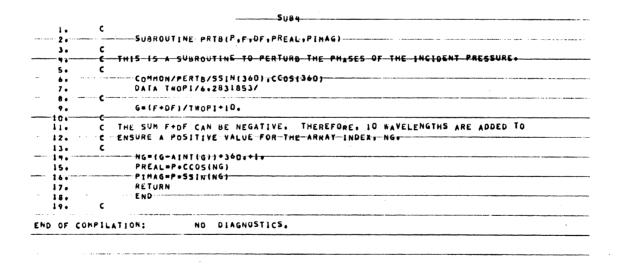
•

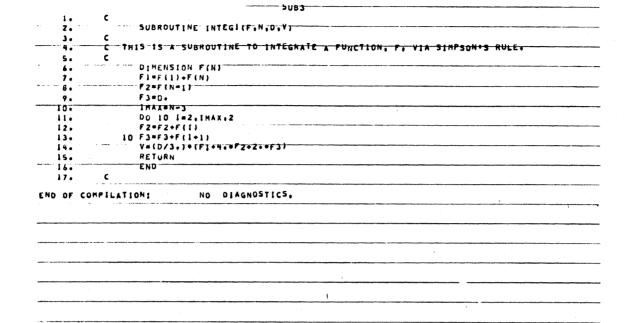
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	QF1=QRE(NT+1)+Q1					
57.	QG1#QRIMINT,I)+	RIMINTONAPI			·····	
	F2=RE(NT;NXP=1)					
9. 0.	G2=RIM(NT;NXP=1 GF2=GRE(NT;NXP=					
2.	F3=0.	••				
3.	G3=0.					
4.	QF 3=0.					
5.	963≠0•					
6. (-		1		
7.	DO 300 1=2,1MAX F2=F2+RE(NT,1)					
9.	G2=G2+R1M(NT,I)					
0.	QF2=QF2+QRE(NT+)				
1.	QGZ=QGZ+QRIM(NT	1,				
2.	F3=F3+RE(NT,1+1)					
3.	G3=G3+RIMINT,I+		· · · · · · · · · · · · · · · · · · ·			
4.	QF3=GF3+GHE(NT)					
5.	"300" QG3=4G3+QRIMINT PPRE=F1+4.+F2+2					
6. 7	PPIM=G1+4.+G2+2					
78. (
9		TEMENTS, THE Q	PRESSURE COMPON	ENTS ARE ADDI	ED TO	
30. (
	THE +XP TERMS.					
32.	PRE(NT,NS)=(PPR)					
32.	PRE(NT+NS)=(PPR) 400 PIM(NT+NS)=(PPI)					
32 • 3 3 •	PRE(NT,NS)=(PPR) 400 PIM(NT,NS)=(PPI) RETURN					
31. (32.) 33. (33.) 34. (35.) 35. (34.)	PRE(NT,NS)=(PPR) 400 PIM(NT,NS)=(PPI) Return End					
32 • 33 • 34 • 35 •	PRE(NT,NS)=(PPR 400-P1M(NT;NS)=(PPt) RETURN END					
2 . 3 . 4 . 5 .	PRE(NT,NS)=(PPR 400-P1M(NT;NS)=(PPt) RETURN END	*QG1+4; * €62+2•				
2. 3. 4. 5.	PRE(NT,NS)=(PPR 400-P1M(NT;NS)=(PPt) RETURN END	*QG1+4; * €62+2•				
2 . 3 . 4 . 5 .	PRE(NT,NS)=(PPR 400-P1M(NT;NS)=(PPt) RETURN END	*QG1+4; * €62+2•				
2 . 3 . 4 . 5 .	PRE(NT,NS)=(PPR 400-P1M(NT;NS)=(PPt) RETURN END	*QG1+4; * €62+2•				
2 . 3 . 4 . 5 .	PRE(NT,NS)=(PPR 400-P1M(NT;NS)=(PPt) RETURN END	*QG1+4; * €62+2•				
2 . 3 . 4 . 5 .	PRE(NT,NS)=(PPR 400-P1M(NT;NS)=(PPt) RETURN END	*QG1+4; * €62+2•				
2 . 3 . 4 . 5 .	PRE(NT,NS)=(PPR 400-P1M(NT;NS)=(PPt) RETURN END	*QG1+4; * €62+2•				
2 . 3 . 4 . 5 .	PRE(NT,NS)=(PPR 400-P1M(NT;NS)=(PPt) RETURN END	*QG1+4; * €62+2•				
2 . 3 . 4 . 5 .	PRE(NT,NS)=(PPR 400-P1M(NT;NS)=(PPt) RETURN END	*QG1+4; * €62+2•				
2. 3. 4. 5.	PRE(NT,NS)=(PPR 400-P1M(NT;NS)=(PPt) RETURN END	*QG1+4; * €62+2•				
2. 3. 4. 5.	PRE(NT,NS)=(PPR 400-P1M(NT;NS)=(PPt) RETURN END	*QG1+4; * €62+2•				
2. 3. 4. 5.	PRE(NT,NS)=(PPR 400-P1M(NT;NS)=(PPt) RETURN END	*QG1+4; * €62+2•				
2. 3. 4. 5.	PRE(NT,NS)=(PPR 400-P1M(NT;NS)=(PPt) RETURN END	*QG1+4; * €62+2•				
2. 3. 4. 5.	PRE(NT,NS)=(PPR 400-P1M(NT;NS)=(PPt) RETURN END	*QG1+4; * €62+2•				
2 . 3 . 4 . 5 .	PRE(NT,NS)=(PPR 400-P1M(NT;NS)=(PPt) RETURN END	*QG1+4; * €62+2•				
2 . 3 . 4 . 5 .	PRE(NT,NS)=(PPR 400-P1M(NT;NS)=(PPt) RETURN END	*QG1+4; * €62+2•				
2 . 3 . 4 . 5 .	PRE(NT,NS)=(PPR 400-P1M(NT;NS)=(PPt) RETURN END	*QG1+4; * €62+2•	*Q63+*(#/T#OPI)			
2 . 3 . 4 . 5 .	PRE(NT,NS)=(PPR 400-P1M(NT;NS)=(PPt) RETURN END	*QG1+4; * €62+2•	*Q63+*(#/T#OPI)			
2. 3. 4. 5.	PRE(NT,NS)=(PPR 400-P1M(NT;NS)=(PPt) RETURN END	*QG1+4; * €62+2•	*Q63+*(#/T#OPI)			
32 . 3 . 4 . 3 5 .	PRE(NT,NS)=(PPR 400-P1M(NT;NS)=(PPt) RETURN END	*QG1+4; * €62+2•	*Q63+*(#/T#OPI)			

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· · ·	<u>د</u>	SUBROUT	NE ANGLE	Y						
3.	c		TE ANGECTA							
		15-15-A-5	UBROUTINE-	TO-COMPUT	E-THE-PHA	SE ANGLE	OF A COMPL	EX NUMBER	•	
· 6. ···	C	P1=1-144	59265							
7.).GT.1.E-1		BS(Y).GT.	1.E-10) G	0 70 5			
	····C···IF	BOTH X A	ND -Y-ARE-2	ERO A WI	LL-BE-SET					
9. 		IF (ABS(X) . LT . 1 . E . 1	0 .AND. A	BS(Y).LT.	1.E-10) (O TO 130			
11+			}•LT•I•E=1 }•LT•1•E=1					· · · · · · · · · · · · · · · · · · ·		
12.		IF (ABS(Y)+LT+1+E=1	D-ANDA X	+LT+0+}-G	0-10-120-				
13.		IF [A 6 5 [Y • A = A T A N (Y)+LT+1+E-1	O +AND+ X	•GT+0+) G	0 TO 130				
15.	2	1F(A) 10								
16		-1F1x}-20								
17.	20	A=A+P1 G0 T0 20	^							
19.	30	A=A+2P								
- 20		GO TO 20					<i></i>			
21.		1F(X) 20 A=135+P1								_
23.		GO TO 20								
24.	110	A#+5+P1-								
- 26		GO TO 20								
27.		GO TO 20	0							
28,		RETURN							-	
30.		END		• • • • • • • • • • • • • • • • • • •	1		· · · · ·			
31.										
ID OF CO	C MPILATIO)N ;	ND DIÂG	NOSTICS.						
ID OF CO)N ;	ND DIAG	N05TIC5.	· · · · · · · · · · · · · · · · · · ·					
ID OF CO)N;	ND DIAG	NOSTICS.				· · · · · · · · · · · · · · · · · · ·		
)N :	ND DIAG	NOSTICS.				· · · · · · · · · · · · · · · · · · ·		
)N :	ND DIAG	NOSTICS.				· · · · · · · · · · · · · · · · · · ·		
4D OF CO)N :	ND DIAG	NOSTICS.						
)N :	ND DIAG	NOSTICS.						
)N :	ND DIAG	NOSTICS.						
)N :	ND DIAG	NOSTICS.						
)N :	ND DIAG	NOSTICS.						
)N :	ND DIAG	NOSTICS.						
)N :	ND DIAG	NOSTICS.						
)N :	ND DIAG	NOSTICS.						





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GAGT			
MAP 017K=05/26=20:52	· · · · · · · · · · · · · · · · · · ·		
	· · · · · · · · · · · · · · · · · · ·		
	,		
ADDRESS LIMITS DOLODO (014447 040000 1544	101	
STARTING ADDRESS 012735 -			
PORDS DECIMAL 5728	18ANK 37170 DBAI	• K	· · · ·
			·
			×
SEGHENT MAIN	001000 0144	447	04000n 154401
NTABS/FORIO	1 001000-001610	2	040000 040044
UWERR3/NAGFORFUND2 NFFT1s/FOR10	1 001611 002667	2	040505 040736
VFFTOS/FORIO	1 002670 003256		
ERUS		•	
SORTS/NAGFORFUNIO	-1-003257-003326	- 2	-041001 041012
SINCOS\$/NAGFORFUNID	1 003327 003474	2	041013 041042
FOR1052/FORIO	-1-003475-006167		- 041043 043547
NISTHS/FORIO	1 006170 006267	2	043550 043554
ATA45/HAGFORFUNIO	1-006270 006530	2	-043555 043614
FORIOSI/FORIO	1 006531 011730	2	043615 044703
TERHS-ICONMON-BLOCKI			
INTEG3 (COMMON BLOCK)			046200 104247
INTEG2 (COMMON BLOCK)		4	-104250-114417
INTEGI (COMMON BLOCK)			114420 114425
BLANKSCOMMON{COMMON-BLOG	«)		
PERTB (COMMON BLOCK)	· · · · · · · · ·	_	114430 115747
5094	-101+731-011772		
		2	PERTB
5083	-1-011773 012072	0	-115765-116012
5082	1 012073 012330	0	116013 116030
SUB1	-10123 31-012734 3 INTEG1		
		<u> </u>	
	-SINTEG3	6	TERMS
		o	116110 154401
	-+		
MA SH	1 012735 014447	-	
#A I N	3 INTEGI	2	BLANKSCOHMON
MAIN	3 INTEGI	2	TERMS

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SAMPLE OUTPUT

THE RADIUS OF THE SCATTERING AREA IS NOMINALLY XPMAX = 17.0 CM.
THE OBSERVATION POINTS LIE ON AN ARC OF RADIUS 49.20 CH.
THE NUMBER OF SURFACE DATA READ IN IS NW = 600+ THEIR SPACING IS DXP = +09/7 CM+ The #Avelength of the waves is #lngth == 8;800-CH+.
THE FIRST AND LAST SURFACE DATA CARDS ARE
i +21983+00 =+14639+00 600 ++28540+00 =+33782=01
THE TRANSDUCER IS 50.10 CM BELOW THE MEAN WATER SURFACE.
THE NUMBER OF PRESSURE DATA READ IN-IS-NRHO
THE FIRST AND LAST PRESSURE DATA CARDS ARE KAND RHO PHAG PPHAZ DXHAG DXPHAZ DZMAG DZPHAZ
i00000i00000+013+78667000000+000000872907+015+35746 401 20+00000 +170000−01 2+78814 +550442=01 4+28620 +137818+00 4+35893
MXP = 175 SURFACE POINTS ARE INITIALLY ASSIGNED TO THE +XP DATA SET. THE H=K INTEGRAL WILL BE EVALUATED FOR-NUMB = -23-SURFACE-PASITIONS. FOR EACH SHIFT OF THE SURFACE, MXP WILL BE INCREASED BY KSHIFT = 4.
#7 87 <
$\begin{array}{cccccccccccccccccccccccccccccccccccc$
51 51 51 49 49 49 47 47 47 45 45 43 43 43 41 41 39 39 37 37 35 33 33 31 29 29 27 25 23 21 19 17 15 11 3

THETA =	LS.	•					
	NS	PRE	PIM	PMAGN	PHASE	PP+	
	1	.075735444	•074120705	.105970451	.774623	.011229736	
	2	.105919360	.052027809		. 456585		
	3	159945803	•088086168	·182597460	.503400	.033341832	
	5	.368657842	+309901230	.481609151	.699023	.231947375	
		-322320325-					
	7	.168334100	.062726230	.198505197	.321503	.039404313	
		.109896071				012185808	
	ĕ	+112101615	+028317080	+115622789	.247426	.013368629	
			+000476201	.109485161	.004349	011987001	
	• -	055204276	-+066643352	086538132	5.404181	.007488848	
	11			084706120		.007175127	
	12	120596207	-+048163686	.129858328	3.521544	.016863185	
			-+079950394	-201359764	-3.549896		
	17	202438163	-+190323496	.277856514	3.896156	.077204242	
	15		-+ 277034733				
	17	167170521	-+270591632	+318065736	4.158987	.101165812	
				-257782191	3.986587		
	18	-+145663049	+114758985	+185438260	3.808874	.034387348	
	19				• • • • •		
	20-	+008301147	+128079873	128348598	4.777111	.016473363	
	21			072503055			
	22,		•047663454	080938539	.629678	+006551047	
	23	1003119411	- 0 17 003 40 4			.127339	- PRMS

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			NUMB = 23
		USER:	NTHETA = 7
- ITEM	AHOUNT-	COST DOLLARS	
	14144323		
1/0 REQUESTS	396	\$0.28	
		\$0.13	
CORE USAGE	11+515	\$3.80	
CARDS IN			
PAGES PRINTED	16	\$0,33	
JOB CHARGE	tttt	s0,12	
· · · · · · · · · · · · · · · · · · ·			
THE ABOVE DOLLAR ANOUNTS-ARE	APPROXIMATE -	AND-ARE-BASED-ON-RATES	FOR CONVENIENCE RUNS

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<u>Program TRANSLATE</u> -- The essential function of this program is to read a set of digitized values of the surface displacement from magnetic tape, compute approximate surface slopes for a selected subset of the data, and punch cards containing the displacements and slopes for that subset. In the work reported in the text, the surface data were digitized and recorded on tape with the aid of a small Datacraft computer. This machine uses 24-bit words, whereas a Univac 1108 computer, for which this program was written, is a 36-bit machine. Thus, much of the program is concerned with simply reading the tape and translating the data into the proper format. Subroutine RREC is the primary program unit for performing this task.

The digitized data on the tape are positive binary integers with 10 significant bits in the present case. Numbers proportional to the displacement about the mean surface level are obtained by subtracting the number $2^{10}/2$ from each datum. A proportionality factor is then used to convert these numbers to surface displacements in centimeters (cards 234-254).

The 'selected subset' of data referred to above is an N-word block of data which is considered to be representative of the entire data set; that is, it is an 'average' block. It is selected using the criterion that its rms displacement be very close (normally within .5%) to the rms displacement of the entire data set (cards 176-194).

The surface slope, \mathcal{J}_{x} , (the x derivative of the displacement,

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 \mathbf{Y}) is computed for each value of \mathbf{Y} by means of Equation 9 in the text. Subroutine XDERIV performs this function.

Definitions of terms, including the quantities required on input cards, are given by comment cards within the program.

Subroutine PLDATA was supplied by B. Sternberg of this laboratory.

	TARC
0001	WASG, TH 177 PASSWORD ., BC9, REEL ID
00002	GFOR,ASI MAIN
66663	C
00004	C THIS IS PROGRAM TRANSLATE. IT READS A MAGNETIC TAPE CONTAINING
00005	C BLOCKS OF NUMBERS RELATED TO THE DISPLACEMENT OF THE WATER SURFACE.
60006	C THE NUMBERS ARE CONVERTED TO WAVEHEIGHTS IN CENTIMETERS AND
60607	C X-DERIVATIVES ARE CALCULATED. THE DATA IN A SELECTED BLOCK ARE
00008	C PUNCHED ON CARDS FOR USE IN PROGRAM SCATTER.
66669	C THIS PROGRAM HAS BEEN WRITTEN FOR THE PARTICULAR CASE WHERE THE DATA
00010	C ARE RECORDED ON 9-TRACK TAPE, IN 24-BIT WORDS, AND IN N-WORD BLOCKS.
00011	C THE DATA , ORDS ARE CONVERTED FROM 24 BITS TO 36 BITS.
00012	C THE NEAN IS FOUND AND IS SUBTRACTED FROM EACH DATUM.
00013	C AN KMS VALUE IS COMPUTED FOR EACH BLOCK.
60014	C THE MEAN OF THE RHS VALUES IS CALCULATED, AND AN "AVERAGE" BLOCK
00015	C HAVING AN RMS VALUE CLOSE TO THE MEAN IS FOUND. ONLY THE DATA IN
00016	C THAT ELOCK ARE USED FURTHER.
00017	C THE NUMBERS IN THAT AVERAGE BLOCK ARE CONVERTED TO CM.
06018	C X DEFIVATIVES ARE CALCULATED.
60019	C CARDS ARE PUNCHED FOR USE BY PROGRAM SCATTER.
00020	
00021	C N IS THE NUMBER OF WORDS IN A DATA BLOCK.
60622	C FOR THIS PROGRAM, N MUST BE AN INTEGER MULTIPLE OF 3.
60623	C N CAN BE CHANGED BY CHANGING THE PARAMETER STATEMENTS.
60024	C NFILES IS THE NUMBER OF TAPE FILES TO BE SKIPPED.
00025	C HE IS A NUMBER , GE, THE NUMBER OF BLOCKS (OR RECORDS) IN ANY FILE TO
06026	C BE SKIPPED.
00027	C SRATE IS THE EFFECTIVE DIGITIZER SAMPLING RATE.
00028	C EFFECTIVE MEANS (DIGITIZER SAMPLING RATE)+(RECORDER RECORD SPEED/
00029	C RECORDER PLAYBACK SPEEDI.
00030	C SIGN IS SET .LT. D. IF THE DATA ARE INVERTED ON THE TAPE.
00031	C PKHT IS THE ASSUMED PEAK HEIGHT OF THE WATER WAVES.
00032	C SPEED IS THE SPEED OF THE WATER WAVES.
00033	C WLNGTH IS THE WAVELENGTH OF THE WAVES.
00034	C NB IS THE NUMBER OF N-WORD BLOCKS IN THE FILE.
00035	C - IB IS THE INDEX OF THE AVERAGE BLOCK. IT CAN HAVE ANY VALUE FROM
00036	C 1 TO NB.
00037	C EPS IS THE FRACTIONAL DIFFERENCE BETWEEN THE AVERAGE RMS AND THE RMS
00038	C OF THE AVERAGE BLOCK.
00030	C RMSHT IS THE RMS SURFACE DISPLACEMENT IN CH FOR THE DATA IN THE
00040	C AVERAGE BLOCK.
00041	C DXP IS THE HORIZONTAL SEPARATION IN CH OF THE SURFACE DATA.
00042	C ZETA 15 THE SURFACE DISPLACEMENT IN CH.
60643	C ZEYAX IS THE APPROXIMATE & DERIVATIVE OF ZETA.
00044	
66045	
00046	PARAMETER N=600
66647	COMMON/REC/IY(N)
00048	COMMON/DERIV/DXP,ZETA(N),ZETAX(N)
00048	DIMENSION LABEL (7), RMS(1000)
66650	DIMENSION TEMP(N), DUN(N)
00050	
00052	REAL MEAN(1000)
66653	NPAN
60054	C

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0055	C INPUT CARDS.
-	ç
Cu57	READ 10, NFILES,MB
0058	10 FORMAT(2110)
6059	REAC 20, SRATE
0060	READ 20, SIGN
0061	READ 20, PKHT, SPEED, WLNGTH
00062	20 FORMAT(3F10.0)
10063	
	- C
10065	č
00666	PRINT 30, NFILES
0067	30 FORMAT(1H1, "NFILES =", 12," TAPE FILES HAVE BEEN SKIPPED.")
66068	PRINT 40, SRATE SIGN
0069	40 FORMAT(1HO, THE EFFECTIVE SAMPLING RATE 15", F6.1. SAMPLES PER SEC
0070	10ND. // SIGN =', F4.0//)
0071	PRINT 50, PKHT
0072	TO EQUMATLY THE DEAY HT OF THE AVEDACE WAVE 154 PL 3 4 CM 43
0073	PRINT 60, SPEED
0074	60 FORMAT(' THE SPEED OF THE WAVES IS', F6.2, ' CH PER SECOND.")
10075	PRINT 70, WLNGTH
00076	70 FORMAT(' THE #AVELENGTH OF THE WATER WAVES IS', F6.3, ' CH. ')
	SKIP FILES THAT ARE NOT OF INTEREST. LABELS ARE FILES.
· · · · · · · · · · · · · · · · · · ·	
0060	IF(NFILES .NE. D) CALL SKIPF(NFILES,MB)
CONTRACTOR A DESCRIPTION	
	C READ AND PRINT THE TAPE LABEL.
	C INTPIN IS AN MACC 1108 LIBRARY SUBROUTINE TO READ A DATA BLOCK FROM
	C MAGNETIC TAPE. SEE THE REFERENCE MANUAL FOR THE 1108 COMPUTER.
	C IN THIS CASE THE LABEL IS STRICTLY NUMERIC AND IS RECORDED AS 24-BIT
0086	C WORDS. FLD MANIFULATES THE BITS TO READ THE WORDS IN A 36-BIT (1108)
	C FORMAT.
	c c c c c c c c c c c c c c c c c c c
06089	CALL IOTPIN(17,1,1Y,N,LNZ, \$80)
040	80 LABEL(1)=FLD(0,24,1Y(1))
6691	LABEL(2)=FLD(0,12,17(2))
0092	LABEL(3)=FLD(12,24,IY(2))
0093	LÁGEL(4)=FLD(0,24,IŸ(3))
0074	LABEL(S)=FLD(0,12,1Y(4))
06695	LABEL(6)=FLD(12,24,1Y(4))
0096	LABEL(7)=FLD(0,24,IY(5))
0097	PRINT 90,(LABEL(J), J=1,7)
0098	90 FORMAT(///' TAPE NUMBER', 16, 10X, 'DATA TAKEN ON ', 12, '/', 12, '/', 12/
0699	128X, DATA DIGITIZED ON ',12,'/',12,'/',12)
and a second sec	¢
	C AN EOF SHOULD FOLLOW THE LABEL.
	C FIND THE LABEL EGF.
	C SUBHOUTINE INTPSP SPACES LOGICAL UNIT 17 (HERE) PAST THE NEXT DATA
	C BLOCK. IT TRANSFERS CONTROL TO STATEMENT 100 IF AN EOF 1S FOUND.
0105	c
0106	CALL 10TPSP(17,5100)
,0107	PRINT 95
0108	95 FORMAT(//' +++++ LABEL EOF NOT FOUND. +++++)

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00109	<u> </u>
60110	BEGIN THE ANALYSIS OF THE DATA.
00111	C REAC THE DATA BLOCK BY BLOCK. FOR EACH BLOCK COMPUTE THE RMS AND THE
06112	C MEAN. ALSO COMPUTE THE AVERAGE RNS AND THE AVERAGE MEAN. FINALLY,
00113	C COMPUTE THE STANDARD DEVIATIONS OF THE RMS'S AND MEANS.
00114	C BECAUSE THE WAVES ARE APPROX PERIODIC, USE ONLY THAT NUMBER OF DATA
00115	C INS) WHICH CORRESPOND TO AN INTEGER NUMBER OF WAVELENGTHS.
00116	C SUBROUTINE REEC CONVERTS 24-BIT WORDS TO 36-BIT WORDS. CONTROL GOES
60117	C TO STATEFENT INC IF AN EOF IS FOUND, AND TO 160 IF A TAPE ERROR IS
00118	C FOUND.
00119	
00120	
00121	SPAL=WLNGTH/DXP
00122	
	WLPEL=AINT(T+DXP/WLNGTH)
00124	NS=SPHL+HLPBL+1.
00125	J=0
00126	
00127	SUM2=0,
00128	SUH3=0,
00129	SU(14#0,
00130	SUM5=0,
00131	100 SUK=0,
00132	SUNI=D,
0133	J=J+1
00134	CALL RREC(NW,\$180,\$160)
60135	
00136	00 120 I=1,NS
00137	SUM-SUM+FLOAT(IY(II)
00138	120 SUN1=SUM1+FLOAT(IY(I))++2
00139	F INS
60140	MEAN(J)=SUM/F
00141	RH5(J)=SQRT(SUH17F=MEAN(J)++2)
00142	SUN2=SUM2+MEAN(J)
60143	SUN3=SUN3+MEAN(J)++2
00144	SUN4=SUN4+RMS(J)
66145	SUNS=SUNS+RMS(J)++2
60146	60 10 166
66147	160 PRINT 170, J
06148	170 FORMAT(//* +++++ TAPE ERROR IN BLOCK NO, *, 14, * +++++//)
60149	MEAN(J)=D.
60150	RMS(J)=0.
66151	GO TO 100
06152	180 FK=K
60153	NB=J-1
00154	AVGMN=SUM2/FK
66155	SOMEAN=SQRT((SUM3=FK+AVGMN++2)7(FK-1+))
C0156	AVGRKS=SUN4/FK
66157	SDRMS=SQRT((SUM5-FK+AVQRMS+#2)/(FK=1+))
66158	PRINT 190, NS, WLPBL
66159	190 FORMAT(/// THE FIRST', 14, WORDS IN EACH BLOCK ARE USED TO COMPUT
00160	1E THE MEAN AND THE RMS. THESE WORDS CORRESPOND TO', F5.1, ' WAVELEN
00161	267H5,*)
00162	PRINT 200, K.AVGHN, SDMEAN

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60163	200 FORMAT(/// THE AVERAGE OF THE MEANS OF ,14, FRROR-FREE BLOCKS IS
CD164	1' FIG.5/' THE STANDARD DEVIATION OF THE HEANS IS', FIG.5)
66165	PRINT 210, AVGRHS, SDRMS
00166	210 FORMAT ('OTHE AVERAGE OF THE RMSES 15', F10.5/' THE STANDARD DEVIATI
66167	ION OF THE RMSES IS', FID.5)
00168	PRINT 220
00169	220 FORMAT(THIS IS ARRAY HEAN. //)
00170	PRINT 225, (HEAN(J), J=1,NB)
00171	225 FORHAT(1H ,10F12.5)
00172	PRINT 240 240 FORMAT(ITHIS IS ARRAY RMS. 7/)
00174	
00175	PRINT 225, (RMS(J), J=1,NB)
00176	C FIND A BLOCK CONTAINING DATA WITH AN RMS CLOSE TO AVGRMS.
00177	
60178	EPS0=.005
00179	EPS=EPS0
00180	260 J=0
00181	270 J=J+1
0182	IF (ABS (RNS (J) / AVGRHS-1.) .LT. EPS) GO TO 300
00163	1F(J-HB) 270,280,280
00164	280 EPS=EPS+EPS0
00185	IF(EP5-10.+EP50) 260,260,330
60186	300 PRCNT=100.+EPS
00187 00188	BRUNT SOU TO PURCENT
00189	PRINT 320, 16, RHS(18), PRCNT
60190	320 FORMAT(//' BLOCK NUMBER',14," HAS AN RHS VALUE OF',F7.2,", THIS I 15 #1THIN',F4.2," PERCENT OF THE MEAN RHS.")
00191	60 TO 340
00192	330 PRINT 335
00193	335 FORMAT(// +++++ CANY FIND AN RHS(J) CLOSE TO AVGRHS, QUIT. ++++)
00194	CALL EXIT
66195	ζ
60196	C REWIND THE TAPE AND RETURN TO THE BEGINNING OF THE WAVE DATA,
60147	2
00198	340 CALL IOTPRW(17)
06199	NFP1=NF1LES+1
00200 66201	CALL SKIPF(NFPI,MB)
00201	C GO TO BLOCK NUMBER IB, READ IT, AND COMPUTE THE MEAN AND RMS VALUES
00203	C AGAIN, AS A CHECK.
60264	
00205	360 IMAX-10-1
00206	DO 370 I=1, IMAX
66207	370 CALL 107PSP(17)
00208	CALL RREC(NW, \$480, \$440)
06209	00 360 I=1,N
00210	380 TEHP(I)=FLOAT(IY(I))-MEAN(18)
00211	SUM=0.
00212	SUN1=0,
00213	DO 390 INI,NS
0214 50215	SUM=SUM+FLOAT(IY(I)) 390 SUMI=SUMI+FLOAT(IY(I))++2
50215	JAN SOUTH-SUNT-FLOAT (1116-2) F=KS

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CU217	CKNEAN_SUM/F
66218	CKRMS=SQRT(SUM1/F-CKMEAN++2)
00219	PRINT 400, CKNEAN, CKRHS
C0220	400 FORMAT(//' CHECK THE MEAN AND RMS OF BLOCK IB."/10X,"MEAN =",F10.5
60221	1,10X, 'RHS = 'F10.57
00222	PRINT 420, NW
66223	420 FORMAT(77" CHECK IF NHEN. THE NUMBER OF DATA READ INTO ARRAY TEMP
00224	1 15',15,'.')
00225	60 10 500
60226	440 PRINT 445, IB
60227	445 FORMAT (7/ ***** TAPE ERROR IN BLOCK IB =*,14,*, TRY ANOTHER BLOC
00228	1K
0227	60 TO 270
00230	460 PRINT 490
60231	490 FORMAT(// ***** EDF FOUND BY RREC, QUIT: ******
00232	CALL EXIT
LC 233	c
00234	C FIND THE LARGEST POSITIVE AND REGATIVE VALUES IN THE AVERAGE BLOCK.
66235	C URSECH IS AN MACE LIBRARY SUBROUTINE TO FIND THE LARGEST OR SHALLEST
60236	C ELEMENT IN AN ARRAY.
UU237	2
60238	500 CALL UKSRCH(0,N,TEHP,LP,TMPMAX,0,DUM)
00239	CALL URSPCH(1,N,TEMP,LN,TMPMIN, 0, DUM)
00240	PRINT 520, LP,TMPMAX
00241	S20 FORMATT'ODATUM NUMBER, 14, IS THE LARGEST POSITIVE DATUM IN THE A
00242	IVERAGE BLOCK. IT HAS THE VALUE', F10.4)
00243	PRINT 530, LN, TMPHIN
00244	530 FORMAT('ODATUH NUMBER', 14, ' IS THE LARGEST NEGATIVE DATUH IN THE A
06245	VERAGE BLOCK. IT HAS THE VALUE FIG. 4)
66246	ζ
60247	C THE FACTOR OF WILL CONVERT THE NUMBERS IN ARRAY TEMP TO CH.
66248	c
60249	
66250	IF(SIGN .GT. D.) TMP=TMPMAX
00251	CF=PKHT/TMP
66252	RMSHT=ABS(CF)+RMS(IB)
CC 253	00 550 I=1,N
66254	550 ZETA(I)=CF+TEMP(I)
00255	
00256	PRINT SOO, CF
60257	580 FORMAT (' CTHE WAVEHT CONVERSION FACTOR IS CF = ',EID.4)
66258	PRINT 590, RMSHT
60259	590 FORPAT(" THE RMS WAVENT OF THE AVERAGE BLOCK IS RMSHT, F5.3, CH
00260	1.**)
00261	PRINT 600, DXP
00262	600 FORMATI' THE HORIZONTAL SEPARATION OF DATA ON THE SURFACE IS DXP =
00263	1',F6.4,' CM.')
00264	PRINT 620
00265	620 FORMAT ('ITHIS IS ARRAY ZETA, '77)
U0266	PRINT 630.(ZETA(I), I=1.N)
00267	630 FORMAT(1H ,10F12,5)
00268	C
66269	C COMPUTE X DERIVATIVES AND CREATE OUTPUT.
00270	Construction of the second

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06271	CALL XLERIV
0272	PRINT 650
0273	650 FORMAT(1H1,8X,"1 ",7X,"ZETA ",7X,"ZETAX")
.6274	PRINT 670, (1,ZETA(I),ZETAX(I), 1=1,50)
0275	670 FURMAT(1H ,110,2F12.5)
0276	CALL PLDATA(N, DUH, TEMP, 1)
16277	FUNCH 680,NN,UXP,WLNGTH
0278	680 FORMAT(110,2F10,5)
0279	PUNCH 700, (1,2ETA(1),2ETAX(1), 1=1,NN)
0281	700 FORMAT(110,2E12,5) END
0282	
0283	GFOR, SIA SUB1
0284	
0285	SUBROUTINE XDERIY
0286	C Sobreetine Aberry
0287	C THIS SUBROUTINE COMPUTES THE APPROXIMATE X DERIVATIVE OF FUNCTION
0288	C ZETA.
0289	L
6290	c
0291	PARAHETER N=600
0292	COMMON/DERIV/DXP,ZETA(N),ZETAX(N)
0293	C
0294	1MAX=N=1
0295	U=-2.+UXP
0296	ZETAX(1)=2.+(2ETA(2)=ZETA(1))/D
0297	ZETAX(N)=2.+(ZETA(N)-ZETA(N-1))/D
0296	DO 20 1=2,1MAX
0299	20 ZETAX(1)=(ZETA(1+1)-ZETA(1-1))/D
0300	RETURN
0301	END
0302	C 4FORISIA SUB2
0303	C C C C C C C C C C C C C C C C C C C
6305	SUBROUTINE RREC(NWORDS, 5, 5)
C 3 0 6	c
0307	C THIS IS A SUBROUTINE TO READ & RECORD OF TAPE FROM UNIT NO. 17 AND
6308	C CONVERT 9-TRACK, 24 BIT WORDS TO 36 BIT WORDS.
0309	۲
0310	PARAMETER N=600
6311	OIMENSION IX(N)
0312	COMMGN/REC/IY(N)
0313	¢
0314	M=2•N/3
0315	CALL 10TPIN(17,1,1X,H,LN2,520,530)
0310	C GO TO 20 IF AN EOF IS FOUND.
C318	C GO TO 20 IF AN EOF IS FOUND.
0319	C C C C C C C C C C C C C C C C C C C
0320	NVORDS=LNZ/3
6321	IMAX=NkORDS=2
6322	
6323	DO 1G 1=1,1MAX,3
0324	$J=2 \circ (1+2)/3-1$

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C325	1711)#FLD(0,24,1X(J))
0326	1Y(1+1)=FLD(24,12,1X(J))+N84+FLD(0,12,1X(J+1))
0327	10 1Y(1+2)=FLD(12;24;TXtJ+;1)
,0328	RETURN
00329	20 RETURN 2
00320	30 RETURN 3
10331	ENU
0332	C C
10333	WFCR, SIA SUB3
16334	c
0335	SUBROUTINE PLDATATLENT, A, B, NPLINE)
10336	c
10337	C THIS IS A SUBROUTINE TO PLOT (VIA PRINTER) THE DATA IN THE AVERAGE
0338	C BLOCK.
0339	C
10340	PARAMETER NC=18
16341	INTEGER M(NC)
6342	IMPLICIT INTEGER (A-Z)
0343	REAL A(2),8(2)
0344	C
0345	CALL URPRIMIDIOT
0346	PRINT 4000
0347	4000 FORMAT(25%, THE AVERAGE BLOCK. UNITS = DIGITIZER BITS'7
0348	-T6,'-1000',T16,'-800',T26,'-600',T36,'-400',T46,'-200',T58,'0'.
6349	-167, 200, 177, 400, 187, 600, 197, 800, 1107, 1000, 1113/
0350	-5X,10(************************************
0351	2
0352	C CLEAR PRINTER BUFFER.
0353	
0354	DO 21 1=1,NC
0355	
0356	DO 4 I=1,LENT,NPLINE
0357	
0358	CALL COL(U.D.1H.)
0359	DO 3 JEL, NPLINE
0360	
0361	IF(IJ.GT.LENT) GO TU 3
0362	CALL COL(B(1J), 1H+)
.0363	CALL COLIA(IJ), IH.)
0364	B=B([J) X=A([J])
0365	3 CONTINUE
0366	
0368	PRINT 4001, 1, M, 18, 14 4001 ECEMAT(18, 14, 18, 1744, 42, 2110)
12369	4001 FORMAT(1X,14,1X,17A6,A2,2110)
0370	
0371	C CLEAR PLOT BUFFER OF SYMBOLS.
0372	
0373	DO 31 K=1,NC 31 M(K)=6H
0374	4 CONTINUE
0375	CALL URPRIM(6,3)
0376	RETURN
C377	
0378	SUBROUTINE COL(VAL, SYM)

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CC379	REAL VAL	
06380	INTEGER SYM, COLV	
16231	COLV=(VAL+1049+1/20+	
66382	COLV=MAX(1,COLV)	
60383	COLV=HIN(104,COLV)	
66364	CALL CHSTORISYM, COLV, M)	
66385	RETURN	
66386	C	
00387	SUBROUTINE CHSTORICHAR, PO	S, ARRAY)
66388	DIMENSION ARRAY(2)	
66389	IMPLICIT INTEGER(A-Z)	
66390	WORD=(POS+5)/6	
66391	BIT=ABS(MUD((PO5-1)+6,36)	
60392	FLD(BIT,6,ARRAY(WORD)) =FL	D(O,6,CHAR)
60393	RETURN	
60394	END	
60395		
00 396	WFOR, SIA SUB4	
00397		
00378	SUBROUTINE SKIPF(NFILES,M	B J
00400	-	TARE DATA ETLES THAT ARE NOT OF
60401	C CURKENT INTEREST.	TAPE DATA FILES THAT ARE NOT OF
60402	C	· · · ·
66403	DO 60 1=1,NFILES	
00404	20 CALL IOTPSP(17,560)	
06405	DO 20 J=1,MB	
66406	40 FORMAT(///* NO EOF FOUND	BY IOTPSP. QUIT.")
60407	PRINT 40	
00408	CALL EXIT	
60409	60 CONTINUE	
60410	RETURN	
00411	END	
00412 (0413	<u>c</u>	
60414	BX9T	•
00415	1 250	
00416		ATA ARAF (TAIL. T
00417	-1.	DATA CAROS (INPUT)
00418	.32 39.1 8.80)	
00419	GFIN	ter en angelen en die en een een een de een de een de geelde een de een de eerste een de eerste een de eerste
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<u>Program PGEN2</u> -- This program was written specifically for a vertically incident beam with cylindrical symmetry about the beam axis. Its function has already been described. Very briefly, it reads input data, expands the number of data by linear interpolation, computes partial derivatives, and punches cards. The input data are values of the amplitude and phase of the transmitted pressure measured along a radius in a plane at a specified distance from the source. Referring to Figure 5, we see that these quantities are smoothly varying functions of the radius, ρ , except for relatively rapid changes associated with the nulls of the amplitude profile. So the sampling interval is determined primarily by the need to adequately determine the shape of those functions near the nulls.

The first part of the program (cards 58-69) is a linear interpolation procedure applied to both the magnitude and phase data. Uncertainties in the measurement of these quantities is likely to be greater than interpolation errors.

Partial derivatives are computed from Equations 11-13 in the text.

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1000	GFOR, SIA MAIN
0002	c
10003	C THIS IS PROGRAM PGEN2.
0004	C ITS PURPOSE IS TO PROVIDE VALUES OF THE INCIDENT PRESSURE AND ITS
0005	C DERIVATIVES FOR USE BY PROGRAM SCATTER, IT REQUIRES MEASURED VALUES
0006	C OF THE MAGNITUDE AND PHASE OF THE INCIDENT PRESSURE ON A RADIUS IN C THE MEAN SCATTERING SURFACE.
00007	C THE MEAN SCATTERING SURFACE. C IN THIS PROGRAM THE RADIAL DISTANCE IS CALLED 'X' FOR BREVITY, BUT
0000	C IT SHOULD BE UNDERSTOOD THAT, IN TERMS OF PROGRAM SCATTER, IT IS
00010	C REALLY THE QUANTITY 'RHO'.
0011	C PGEN2 INTERPOLATES TO INCREASE THE NUMBER OF VALUES, AND IT COMPUTES
0012	C THE X (RHQ) AND Z DERIVATIVES AT EACH VALUE OF X.
0013	C THE INTERPOLATED HAGNITUDES AND PHASES OF THE INCIDENT PRESSURE AND
0014	C ITS DERIVATIVES ARE PUÑCHED ON CARDS.
0015	C ••••••
0016	C Z IS THE VERTICAL DISTANCE FROM THE TRANSDUCER SURFACE TO THE PLANE
10017	C IN WHICH THE MEASUREMENTS WERE MADE.
0018	C WL IS THE ACOUSTIC WAVELENGTH.
0019	C NEARDS IS THE NUMBER OF POINTS FOR WHICH MAGY AND PHASE ARE READ IN.
00020	C AFTER INTERPOLATION, VALUES FOR NX POINTS ARE PRINTED AND PUNCHED.
0021	C NINT IS THE NUMBER OF INTERPOLATION POINTS DESIRED BETWEEN THE INPUT
0022	C POINTS. C DXD IS THE HORIZONTAL SPACING BETWEEN POINTS CORRESPONDING TO THE INPUT
0023	C DXO IS THE HORIZONTAL SPACING BETWEEN POINTS CORRESPONDING TO THE INPUT C DATA. AFTER INTERPOLATION, THE SPACING IS DX.
0025	C PO IS THE PRESSURE MAGNITUDE.
0026	C PHI IS THE PHASE OF THE MEASURED PRESSURE IN MULTIPLES OF TWOPI.
0027	C THE OUTPUT (PUNCHED AND PRINTED) PHASES ARE IN RADIANS.
0028	C DXMAG IS THE MAGN OF THE DERIVATIVE OF P WRT X.
0029	C DXPHAZ IS THE PHASE OF THE DERIVATIVE.
00030	C SIMILARLY FOR Z.
20031	6
20032	DIMENSION PO(SOD), PH1(SOD), X(SOD)
0033	DIMENSION DPODX(SOO), DPHIDX(SOO)
0034	REAL K
00035 00036	C C INPUT CARDS. ••••••••••••••••••••
0038	
20038	READ 20, Z,DXD,WL
00039	20 FORMAT (3F10+0)
00040	READ 25, NCARDS, NINT
00041	25 FORMAT(2110)
0042	<u>c</u>
0043	NX={NCARDS-1}+NINT+1
00044	<u> </u>
00045	READ 30, (PO(1), PHI(1), 1=1, NX, NINT)
00046	30 FORMAT(2F10+0)
00047 00048	C
00048	<u> </u>
00050	PRINT 35
00051	JS FORMAT(1H1, THE INPUT ARRAYS ARE', 22X, 4HPO , 8X, 3HPH1//)
00052	PRINT 40,(1,PO(1),PHI(1), 1=1,NX;NINT)
00053	40 FORMAT (30X, 15, 2F12, 4)
00054	c

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0055	1W0P1=6.283185
0056	K-TWOPI/WL
00057	
00058	C FILL UP THE PO AND PHI ARRAYS BY INTERPOLATING.
00059	<u>c</u>
00060	FINT=NINT
1900	IFIN=NINT=1
00062	JFINENX-NINT
00063	DO 50 J=1, JFIN, NINT
00064	DP=(PO(J+NINT)+PO(J))/FINT
00065	DPHI=(phi(J+NINT)-PHI(J))/FINT
00066	DO 50 1=1,1FIN
00067	
00068	PO(J+1)=PO(J)+F1+DP
00070	50 PHILJ+II=PHILJI+FI+DPHI
00071	C DX=DX0/FINT
00072	C
00073	PRINT 52, NX, DX, WL, Z
00074	57 FORMAT(14) (NX #4,15,5V (NY #4 FA # 57 ##405) FURSH -4 FA # 4
00075	52 FORMAT(1H1, "NX =", 15,5X,"DX =", F8,4,5X,"WAVELENGTH =", F8,4,5X,"Z = 1", F8,4///)
00076	PUNCH 54, NX, DX, WL, Z
00077	54 FORMAT(15,3F10,4)
00078	PRINT 58
00079	58 FORMAT(11X, 4HX , 7X, 6HPMAG , 3X, SHPHAZ , 7X, 6HDXMAG , 8H DXPHAZ , 7X
00080	1,6HDZHAG ,8H DZPHAZ //)
00081	č
00082	C ATTACH VALUES OF X TO THE PRESSURE DATA.
00083	c
00084	DO 60 Iml,NX
00085 .	FIEl
00086	60 X(I)=(FI-1.)+DX
00087	C
00088	C COMPUTE PARTIAL DERIVATIVES AND CONVERT PHASES TO RADIANS.
00089	C THE FUNCTION AINT IS AN IIGS FORTRAN INTRINSIC FUNCTION TO REMOVE
00090	C THE FRACTIONAL PART OF A REAL NUMBER.
00091	C
00092	DPHIDX(1)=0.
0009 3 00094	DPODX())=0.
00095	
00095	DO 80 (=2,)MAX
00097	DPODX(1)=(PO(1+1)-PO(1-1))/(2.+DX) BO DPHIDX(1)=(PHI(1+1)-PHI(1-1))+TWOPI/(2.+DX)
00098	DPODX(NX)=DPODX(IMAX)
00099	DPHIDX(NX)=DPHIDX(IMAX)
00100	c c
10101	DO 100 I=1,NX
00102	Q=PH1(1)
10103	AA=TWOPI+(Q-AINY(Q))
0104	5=SIN(AA)
0105	C=COS(AX)
0106	DPXRE==PO(I)+DPHIDX(I)+S+DPODX(I)+C
50107	DPXTH=PO(I)+DPHIDX(I)+C+DPODX(I)+S
0108	R=SQRT(X(1)++2+Z+Z)

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0100	
0109	DZHAG=K+PO(I)+Z/R
0111	QD=Q++25
00112	DZPHAZ=TKOPI+(QD-AINT(QD))
00113	C CONVERT REAL AND IMAG TERMS TO MACH AND BUARE
00114	C CONVERT REAL AND IMAG TERMS TO MAGN AND PHASE.
00115	DXMAG=SQRT(DPXRE++2+DPX1H++2)
00116	
00117	CALL ANGLE(DPXRE,DPXIM,DXPHAZ)
00118	PRINT 110,1,X(1),PO(1),AA,DXMAG,DXPHAZ,DZMAG,DZPHAZ
00119	100 PUNCH 110, 1,X(1),PO(1),AA,DXHAG,DXPHAZ,DZMAG,DZPHAZ
00120	110 FORMAT(15,F10.5,3(E13.6,F8.5))
00121	END
00122	C C
00123	QFOR, SIA SUBL
00124	SUBROUTINE ANGLE(X,Y,A)
00125	PI=3.14159265
00126	IF(AB5(X).GT.1.E-10 .AND. AB5(Y).GT.1.E-10) GO TO 5
00127	IF (ABS(X)+LT+1+E=10 +AND+ ABS(Y)+LT+1+E=10) 60 TO 130
00128	IF(ABS(X)+LT+1+E=10 +AND+ Y+LT+D+) GO TO 100
00129	IF (ABS(X) .LT. 1.E-10 .AND. Y.GT.D.) GO TO 110
00130	IF (ABS(Y) + LT + 1 + E + 10 + AND + K + LT + D +) GO TO 120
00132	IF (ABS(Y) . LT. I.E. IO . AND. X.GT. 0.) GO TO 130
00133	5 A=ATAN(Y/X) IF(A) 10,10,50
00134	10 1F(X) 20,20,30
00135	20 A=A+P1
00136	GO TO 200
00137	30 A=A+2PI
00138	GO TO 200
00139	50 IF(X) 20,20,200
00140	100 A=1.50PI
00141	GO TO 200
00142	110. A=.5+P1
00143	GO TO 200
00144	120 A=P1
00146	GO TO 200 130 A=0.
00147	200 RETURN
00148	END
00149	2
00150	exet
00151	50,1 ,25 ,7198
00152	01 5
00153	1. 0.
00154	•997 •001
00155	•992 •003
0156	.982 .005
00157	.965 .007
0158	,948 ,013
0159	.927 .02
0160	• 904 • 024
	.675 .036
00162	.853 .052

: :

	00163	•82	.063	
	00164	•784	•08	
	00166	•704	,12	
	00167	.665		
	00168	.62	+166	
	00169	.581	.195	
	00170	• 5 3 7	• 2 2 4	
	00171	.493	.258	
	00172	.442	• 294	· · · · · · · · · · · · · · · · · · ·
	00174	• 347	• 374	
	00175	• 3	.407	
	00176	.254	.45	
	00177	.206	.49	
	00178	•161	+518	
	00179	•121	.548	
	00160	.083	•556	
	00182	.04	.33	
	00183	.025	.345	
	00184	.053	.367	
	00185	.077	• 4	
	00186	.102	• 4 4 5	
	00187	. 23	.495	
	00188	•141 •15à	• 554	
	00190	•150	•675	
	00191	.183	•74	·
	00192	.192	.805	
	00193	.198	.872	
	00194	• 2	.942	
	00195	•201	1.007	
	00197	• 2	1.082	
	00198	.196	1.23	
-	00199	•19	1.3	
	00200	.185	1.375	
	00201	.175	1.445	
	00202	•167	1.525	
	00204	•157	1.688	
-	00205	.125	1.78	
	00206	+112	1.88	
	00207	.098	1.97	
	00208	• 084	2.055	
	00209	.075	2,125	
-	00210	•067	2,205	
	00212	.044	2.38	
	00213	.035	2.96	
	00214	•027	2.554	
	00215	• 0 2	2.632	
-	00216	.015	2.71	
	00217	.012	2.775	
	00219	+011	2.86	
	00220	.018	2.947	
	00221	.02	3,165	
	00222	.022	3.27	
	00223	.025	3.35	
	00224	•026	3.445	
	00225	.025	3.555	
	00226	.023	3.64	
	00228	• 0 2 1 • 0 2	3.71	
	00229	.019	3.81	
	00230	•017	4.24	
	00231	.017	4.36	
	00232	•016	4.53	
	00233	.017	4.72	
	00234	QFIN		

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B. Programs based on analytical expressions for acoustic scattering from a sinusoidal surface.

Two programs have been written for the purpose of evaluating Equations 35 and 39 in the text.

<u>Program PSIN1</u> -- The scattered pressure, p_m , is evaluated for a specified number of positions of the scattering surface. The surface is translated stepwise a total distance of one surface wavelength. Several quantities related to the scattered pressure are computed and printed for each position of the surface (see card 211, statement 200), and all of these quantities, except phases, are averaged over the surface wavelength (cards 215-223). One of the quantities obtained is the rms scattered pressure, p_{rms} . Another is the coherent component of the scattered pressure, $\langle p \rangle$. Values of the percentage slope correction are also computed.

The equations and definitions of terms used in this program are taken directly from Equations 15-38 in the text, and the computational procedure is clearly outlined by comment cards in the program.

Both of the subroutines in this program are useful in a variety of applications. Subroutine INTEG1, used here to compute the averages of the pressure terms, is a simple routine for the numerical integration of a function of one variable.

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00001	ØÅSG,MT 11.,T,MÅTHPK
00002	GCOPY, G 11. TPFS.
00003	DFREE 11.
00004	OFOR, SIA MAIN
00005	ç
00006	C THIS IS PROGRAM PSINI.
10007	C IT COMPUTES THE PRESSURE AND THE RNS PRESSURE SCATTERED IN THE X=2
00008	C PLANE FROM A CORRUGATED SURFACE. A SLOPE CORRECTION IS INCLUDED.
10009	C THE CORRUGATIONS ARE COSINUSUIDAL AND ARE PARALLEL TO THE VILLES
0010	C A HIGHLY DIRECTIVE TRANSHITTED ACOUSTIC BEAM AND A POINT RECEIVER
10011	C ARE ASSUMED.
0012	C THE AMPLITUDE PROFILES OF THE INCIDENT PRESSURE ARE ASSUMED TO BE
10013	C GAUSSIAN, THET ARE MEASURED ON A HURIZONTAL PLANE AT A DISTANCE RI
0014	C FROM THE SOURCE TRANSDUCER.
0015	C BWX IS THE BEAMWIDTH - FWHM/1.665 HEASURED IN THE X-Z PLANE.
0016	C BWY IS THE BEAMWIDTH MEASURED IN THE YOZ PLANE.
10017	C THE PHASE OF THE INCIDENT PRESSURE IS THAT OF A POINT SOURCE,
0018	C HODE SPECIFIES THE SCATTERING GEOMETRY.
0019	C HODE=1 SPECULAR SCATTERING.
0020	C MODE==1 BACKSCATTERING.
0021	C HODE=0 TRANSHITTER ANGLE FIXED AT ZERO. C HODED IS THE INITIAL VĂLUE OF MODE.
0023	C HODED IS THE INITIAL VALUE OF MODE.
0024	
0025	C NMODE IS THE NUMBER OF VALUES OF MODE, NORMALLY 1, 2, OR 3, C THETAD IS THE INITIAL VALUE OF THE RECEIVER ANGLE WAY VERTICAL.
0026	C THETAO IS THE INITIAL VALUE OF THE RECEIVER ANGLE WAY VERTICAL. C DTHETA IS THE INCREMENT.
0027	C NTHETA IS THE NUMBER OF VALUES USED.
0028	C OFFSET IS THE ANGULAR DISPLACEMENT OF THE TRANSDUCER WRT ITS NORMAL
0029	C POSITION FOR THE SCATTERING HODE SPECIFIED. IT IS POSITIVE IF THE
0030	C DISPLACEMENT IS IN THE + THETAL DIRECTION.
0031	C THETAL IS + "UPWIND".
0032	C THETA2 IS + 'DOWNWIND'.
0033	C ZETA IS THE AMPLITUDE OF THE CORRUGATIONS. IT MUST BE .GT. ZERO.
0034	C ZETAO IS THE INITIAL VALUE OF ZETA.
0035	C DZETA IS THE INCREMENT BY WHICH ZETA IS INCREASED.
0036	C NZETA IS THE NUMBER OF VALUES OF ZETA USED.
0037	C WE IS THE ACOUSTIC WAVELENGTH.
0038	C SAL IS THE SURFACE WAVELENGTH.
0039	C K IS THE ACOUSTIC WAVENUMBER,
0040	C SK IS THE SURFACE WAVENUMBER.
0042	C RI IS THE RADIUS OF THE ARC ON WHICH THE TRANSMITTER MOVES. C R2 IS THE RADIUS OF THE ARC ON WHICH THE RECEIVER MOVES.
0043	C M IS THE INDEX FOR HORIZONTAL DISPLACEMENT OF THE SURFACE.
0044	C XH IS THE HTH HORIZONTAL DISPLACEMENT OF THE SURFACE.
0045	C HHAX IS THE NUMBER OF VALUES OF XH. IT MUST BE ODD FOR SIMPSON'S
0046	C RULE AVERAGING.
0047	C BESSEL FUNCTIONS ARE USED IN THE FORM JINI, WHERE N IS THE ORDER AND
0048	C NHAX=16 IS THE MAXIMUM ORDER. IT CAN BE INCREASED IF NECESSARY.
0047	C A IS THE ARGUMENT OF THE BESSEL FUNCTIONS.
0050	
0051	C
0052	REAL J(100), J0
0053	DIMENSION PPRE(100), PPIM(100)
0054	DIMENSION PRE(100), PIM(100), PP2(100), P2(100)

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0055	c
0056	REAL HODED, NODE
0057	READ 40, HODED, DHODE, NHODE
0058	READ 40, BWX, BWY
0059	READ 40, THETAD, DTHETA, NTHETA
0060	READ 40, OFFSET
0061	READ 40, ZETAO, DZETA, NZETA
0062	READ 40, R1, R2
0063	READ 40, WL,SWL,MMAX
0064	40 FORMAT(2F10+0,110)
00066	PRINT 45
0067	45 FORMAT (FITHE INPUT CARDS AREY)
8400	PRINT 50, MODED, DHODE, NHODE
0069	PRINT 50, BWX, BWY
0070	PRINT 50, THETAO, DTHETA, NTHETA
0071	PRINT 50, OFFSET
0072	PRINT 50, ZETAO, DZETA, NZETA
0073	PRINT 50, R1, R2
0074	PRINT SD, WL,SWL,MMAX
0075	50 FORMAT(24X,2F10.4,110)
0076	<u>c</u>
0077	DATA NHAX/16/
0078	DATA THOP1/6.2832/
00079	REAL K
0080	K=TWOPI/WL
0081	SK # TWOP 1 / SWL
00082	A3={1./R1+1+/R2}/2. BWXD=BWX
0084	C
0085	C COMPUTE THE HAGNITUDE OF THE PRESSURE REFLECTED VERTICALLY FROM A
0086	C SHOOTH SURFACE.
00087	C
88000	AXD=SQRT(SQRT(BWX++-4+(K+A3)++2))
00089	AY0=SQRT(SQRT(BWY++=4+(K+A3)++2))
00090	PO=K/(2+R2+AXO+AYO)
00091	PRINT 60, PO
10092	60 FORMAT('OTHE PRESSURE REFLECTED FROM A SMOOTH SURFACE IS PD ="
0093	1210.51
0094	C COMPUTE DXM.
00095	
00096	C D=10/(MMAX-1)
00098	DXH=SWL+D
00099	PRINT 62, DXM
0100	62 FORMAT ('OTHE SURFACE SHIFT INCREMENT IS DAM =', F8.4)
10101	c
00102	C ESTABLISH THE GEOMETRY.
00103	c
00104	MODE=MODEC=DHODE
00105	D0 260 HD=1, NHODE
00106	HODE+DHODE
0107	C I
0108	C ESTABLISH ZETA.

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0109	6
0110	ZETA=ZETAO-DZETA
0111	D0 260 12=1,NZETA
0112	
0113	PRINT 65, ZETA
0114	65 FORMAT('1************************************
0115	C C COTARLIEU THE TRANSMITTER AND RECEIVER AND RE THERAL AND THETAS
0116 0117	C ESTABLISH THE TRANSMITTER AND RECEIVER ANGLES, THETAL AND THETA2.
0118	THETA2-THETAO-DTHETA
0119	
0120	THETA2=THETA2+DTHETA
0121	ANG2=THETA2/57.295
0122	CA2=COS(ANG2)
0123	THETA1=HODE+THETA2+OFFSET
0124	ANG1=THETA1/57,295
0125	CA1=COS(ANG1)
0126	C
0127	PRINT 70, THETAI, THETA2
0128	70 FORMAT(' THETA1 =', F6.2, ' DEGREES. '/' THETA2 ='F6.2, ' DEGREES. ')
0129	PRINT 75
0130	75 FORMAT(12X, 'PPRE ', 7X, 'PPIH ', 6X, 'PPMAG ', 6X, 'PPPHAZ', 7X, 'PRE ',
0132	17X, PIH 7,7X, PHAG 7,7X, PPHAZ 7,7X, PP2 7,7X, P2//)
0133	C COMPUTE SOME YERHS.
0134	
0135	A1={CA1++2/R1+CA2++2/R2}/2.
0136	A2=5IN(ANG1)-SIN(ANG2)
0137	A4=CA1+CA2
0138	A=K+ZETA+A4
0134	BWX07CAT
0140	AX2=SQRT(BWX++++(K+A1)++2)
0141	AX=SQRY(AX2)
0142	AY2=SQRT(BWÝ++++(K+A3)++2)
0143	AY#SQRT(#YZ) PHIX=ATAN(K+A1+BWX++2)
0145	TX=TAN(PHIX)
0146	CX=COS(PHIX)
0147	PHIY=ATAN((K+A3)+BWY++2)
0148	XY=(PHIX+PHIY)/2.
0149	GNU2=4, • AX2/(K+K+CX)
0150	CP=K+A4/(4+R2+AX+A¥)
0151	
0152	C BESJ IS A UNIVAC MATH-PAC ROUTINE TO COMPUTE BESSEL FUNCTIONS.
0153	
0154	CALL BESJ(A, D, , NMAX, J)
0155	
0157	DO 80 NN=1,NMAX 80 J(NN)=J(NN+1)
0158	
0159	C ESTABLISH XN.
0160	
0161	XMu-DXM
0142	DO 200 M=1, MMAX

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00163	XM=XH+DXM
00164	C C
00165	C COMPUTE THE APPROXIMATE SCATTERED PRESSURE, PPIH), AND THE CORRECTED
00166	C PRESSURE, P(M), FOR MHAX SURFACE POSITIONS,
00167	C THE REAL AND IMAGINARY PARTS ARE COMPUTED SEPARATELY.
00168	с
00169	C IN COMPUTING PHASES, THE CONSTANT PHID=K+(RI+R2) IS IGNORED.
00170	c
00171	Q=XY=A2+A2+TX/GNU2
00172	C=COS(q)
00173	S=SIN(q)
00174	PPREM=J0=EXP(-A2+A2/GNU2)+C
00175	PPIMH=PPREM•S/C
00176	
00177	PINM=PPINM
00178	C REAL N
00177	
00180	DO 100 NN=1,NMAX N=NN
00182	ALFP2=(A2+N+SK/K)++2/GNU2
00183	ALFM2=(A2=N+SK/K)++2/GNU2
00184	EPS=(N+5K/(K+A4))++2
00185	QP=XY-ALFP2+TX-N+SK+XM+++25+N+T#OP1
00186	QH=XY-ALFM2+TX+N+SK+XH=,25+N+TWOPI
00187	C=COS(QP)
00188	S=SIN(QP)
00189	CH=COS(QH)
00190	SM=SIN(QM)
00191	$EC=C \cdot EXP(-ALFP2) + CH \cdot EXP(-ALFM2) \cdot ((-1,) \cdot N)$
00192	ES=S+EXP(-ALFP2)+SM+EXP(-ALFH2)+((-1.)++N)
00193	PPREH=PPREH+J(NN)+EC
00194	
00195	PREM=PREM+(1.+EPS)+J(NN)+EC
00196	100 PIMH=PIMH+(1++EPS)+J(NN)+ES
00197	PPRE(H)=CP+PPREH
00198	PPIM(M) = CP = PPIM PRE(M) = CP = PREM
00200	PIM(M)=CP0PIMM
00201	PP2(H)=(PPREN++2+PPIMM++2)+CP+CP
00202	PPMAG=SQRT(PP2(M))
00203	F2(HJ=(PREM++2+F1M+++2)+CF+CF
00204	PMAG-SGRT(P2(M))
00205	C
00206	C COMPUTE THE PHASES OF PP(H) AND P(H).
00207	C
00208	CALL ANGLE(PPRE(M), PPIM(M), PPPHAZ)
00209	CALL ANGLE(PRE(M), PIN(N), PPHAZ)
00210	<u>c</u>
00211	200 PRINT 220, PPRE(H), PPIM(H), PPMAG, PPPHAZ, PRE(H), PIH(H), PHAG, PPHAZ
00212	1PP2(M),P2(M)
00213	220 FORMAT (5x,2(3E12.4,F12.3),2E12.4)
00214	<u>c</u>
00215	C AVERAGE THE REAL AND IMAGINARY PARTS OF THE PRESSURE TERMS AND FIND C THE MAGNITUDES AND PHASES OF THE AVERAGES.

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10217	<u>C</u>
0218	CALL INTEGI(PPRE, HMAX, D, APPRE)
00219	CALL INTEGI (PPIM, MHĀX, D, APPIM)
00220	CALL INTEGI(PRE,MHAX,D,APRE)
00221	CALL INTEGI (PIN,MMAX,D,APIN)
00222	CALL INTEGI(PP2,MMAX,D,APP2)
00223	CALL INTEGI(P2, HMAX, D, AP2)
00224	CALL ANGLE(APPRE, APPIM, APPPHZ)
00225	CALL ANGLE(APRE, APTM, APPHZ)
00226	APPMAG=SQRT(APPRE++2+APPIH++2)
00227	APHAGESGRTTAPRE+27APIH+421
00228	c
00229	PRINT 230
00230	230 FORMAT(56X, * + + AVERAGES + + + *)
00231	PRINT 220, APPRE, APPIN, APPHAG, APPPHZ, APRE, APIN, APHAG, APPHZ, APP2,
00233	
00235	PPRMS=SQRT(APP2)
00235	PRHS=SQRT(AP2)
00237	PPRMSN=PPRMS/PO
00238	
00239	SLPCOR=(1PPRMS/PRMS)+100.
00240	PRINT 250, PPRMS, PRMS
00241	250 FORMAT (779X, THE RMS SCATTERED PRESSURE W/D SLOPE CORRECTION IS PP
00242	IRMS =',EI3.5/ 9X,'THE RMS SCATTERED PRESSURE WITH SLOPE CORRECTION IS PP
00243	2N IS PRHS = (EI3.5)
00244	260 PRINT 265, PPRMSN, PRMSN, SLPCOR
00245	265 FORMAT (9X, THE SAME QUANTITIES DIVIDED BY PO ARE PPRKSN # ,E13.5
00246	1/48X, PRMSN = , E13.5/9X, THE CONTRIBUTION OF THE SLOPE CORRECTION
00247	210 PRASN IS SLPCOR ", F7.2, PERCENT."
00248	END
00249	C
00250	GFOR, SIA SUB1 -
00251	
00252	SUBROUTINE ANGLE(X,Y,A)
00253	(
00254	C THIS IS A SUBROUTINE TO COMPUTE THE PHASE ANGLE OF A COMPLEX NUMBER.
00255	c
00256	PI=3.14159265
00257	IF(ABS(X).GT.I.E-IO .AND. ABS(Y).GT.I.E-IO) GO TO 5
00258	C IF BOTH X AND Y ARE ZERO, A WILL BE SET = 0.
00259	IF (ABS(X) .LT.).E-10 .AND. ABS(Y).LT.1.E-10) GO TO 130
00260	IF(ABS(X)+LT+1+E=10 +AND+ Y+LT+0+) GO TO 100
0261	IF (ABS(X) +LT+1+E+10 +AND+ Y+GT+0+) GO TO 110
0262	IF(ABS(Y)+LT+1+E-10 +AND+ X+LT+0+) GO TO 120
10263	TF(ABS(Y)+LT+1+E+10 +AND+ X+GT+0+1 G0 T0 130
0264	5 AHATAN(Y/X)
0265	IF(A) 10,10,50
0266	10 IF(X) 20,20,30
0267	
0268	<u>60 TO 200</u>
0269	30 A=A+2. •P1
0270	GO TO 200

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00271	50 1F(X	1 20,20,200			
00272	100 A=1.				
00273		0 200			
00274	110 A=.5				
00275	120 A=PI	0 200			
00277		0 200			
00278	130 A=0.				
00279	200 RETU				
00280	END				
00281	C				
00282	DFOR,SIA S	082	· · · · · · · · · · · · · · · · · · ·		
00284		OUTINE INTE	G1(F,N,D,V)		
00285	C				
00286	C THIS IS	A SUBROUTI	INE TO INTEGRATE	A FUNCTION, F, VIA SIN	SON'S RULE.
00287	C C				
00288		NSION F(N)			
00289 00290		(1)+F(N) (N-1)			
00291	F3=0				
00292		=N-3			
00293	00 1	O I=2, IMAX,	2	· ·	······································
00294		2+F(1)			
00295		3+F(1+1)			
00296	RETU	/3.)+(F1+4. RN	• • • • • • • • • • • • • • • • • • • •		
00298	END				
00299	C				
00300	PXQT				
00301	0.	0.	1		
00302	5.35	5.35	2		
00304	0.	100	•		
00305	.305	0.	1	· · · · · · · · · · · · · · · · · · ·	
00306	50.1	49+2		· · · ·	
00307	.7198	8.8	49		
00308	<u>ü</u> fin				
		·····		······	
			· · · ·		
		-			
			····		
			· !		
			•		
	·				

ZETA .= +305			SAMI	ple ou	TPIIT				

	GREES.								
	GREES.								
PPRE	PPIM	PPHAG	PPPHAZ	PRE	PIN	PMAG	PPHAZ	PP2	P 2
+2257+00	3111+00	.3843+00	5.340	.2260+00	-,3120+00	.3852+60	5.339	.1477+00	,1484+00
.2169+00	2437+00	.3651+00	5.349	.2176+00	2943+00	.3660+00	5,349	.1333+00	-1340+00
•1814+00	2504+00	.3091+00	5,339	.1823+00	2506+00	.3099+00	5,341	.9557-01	•9606-01
+1092+00	1746+00	,2231+00	5,224	+1097+00-	1948+00	,2236+00	5.225		
.1903-01	1293+00	•1277+60	4,860	.1837-01	1286+00	.1299+00	4.854	.1682-01	.1687-01
-,4517-01	•.5188-01	.6879-01	3,996	4667-01-	=.5176=01	.6970-01	3,979	.4732-02	.4858-02
5221-01	.1832-01	.5549-01	2,796	5331-01	.1952-01	.5677-01	2.791	.3080-02	.3223-02
1778-01	.6276-01	.6523-01	1.847	1742-01	.6345-01	.6579-01	1.839	.4254-02	. 4329-02
.2436-02	.809/-01	.8101-01	1,541	·4118-02	.8076-01	.8106-01	1.520	.6562-02	.6571-02
4144-01	.9913-01	.1074+00	1,967	3978-01	.9857-01	.1063+00	1.954	.1154-01	.1130-01
1481+00	•1370+00	+2017+00	2,395	1479+00	·1367+DO	.2014+00	2,395	.4068-01	.4056-01
-,3086+00	+1624+00 +2032+00	•3161+UU	2,531	2623+00	.1830+00	.3199+00	2.532	+1012+00	.1023+00
2606+00	+1824+00	.3695+00	2.559	3111+00	.2043+00	.3722+00	2.540	.1366+00	+1385+00
1481+00	+1370+00	•3181+00 •2017+00	2,531 2,395	2623+00	+1830+00	.3199+00	2.532	+1012+00	.1023+00
- 4144-01	.9913-01	+1074+60	1.967	1479+00	+1367+00	.2014+00	2,395	.4068-01	.4056-01
.2437-02	.8697-01	.8101-01	1.541	,4118-02	•9857-01 •8096-01	.1063+00	1,954	.1154-01	.1130-01
1778-01	.6276-01	.6523-01	1.847	1742-01	.6345-01	•8106-01 •6579-01	1.520	,6562-02	.6571-02
5221-01	.1881-01	.5549-01	2,796	5331-01	.1952-01	+5677-01	2,791	.4254-02	+329-02
4517-01	5189-01	·6879-01	3.996	4667-01	-,5176-01	.6970-01	3.979	.4732-02	.3223-02
+1903-01	1283+00	+1297+00	4,860	.1837-01	-,1286+00	.1299+00	4.854	.1682-01	.1687-01
.1092+00	1946+00	.2231+00	5,224	+1097+00	1948+00	.2236+00	5,225	.4978-01	.4999-01
+1814+00	-+2504+00	+3091+00	5.339	.1823+00	-,2506+00	.3099+00	5,341	.9557-01	.9606-01
+2169+00	2937+00	.3651+00	5.349	.2176+00	2943+00	.3660+00	5.349	-1333+00	.1340+00
•2257+00	3111+00	.3843+00	5,340	.2260+00	-,3120+00	+3852+00	5,339	.1477+00	.1484+00
- 4404 - 00				+++ AVERAG					
6484-02	++3264-01	.3327-01	4,516	-,6484-02	3264-01	.3327-01	4.516	.5080-01	.5112-01
	,								· · · · ·
THE RHS ST	ATTERED PRES	SURE W/O SLOP	F CARDECT+	N IS BOUNE					
THE RHS SC	CATTERED PRES	SURE WITH SLO	PF CONRECTIO	UN 13 PERMS : Inn is pons .					
THE SAME O	JUANTITIES DT	VIDED BY PU A	RE PPRMSN	45548+					
			PRMSN	45491+	no				
	BUTION OF TH	E SLOPE CORRE	CTION TO PI	THSN IS SLPC	0R	ERCENT.			
ING CONTRA				· · · · · · · · · · · · · · · · · · ·					
HETAL = .00 DEG									
HETAL = .00 DEG HETAL = 20.00 DEG	GREES,								
HETAL = .00 DEG		PPMAG	PPPHAZ	PRE	PIN	PMAG	PPHAZ	PP2	P2
HETAL = ,00 DEG HETA2 = 20,00 DEG PPRE	PPIN								
HETAI = ,00 DEG HETA2 = 20,00 DEG PPRE -,1730-01	REES. PPIM .1442+00	+1452+00	1.690	-,2033-01	+1537+00	.1550+00	1.702	.2108-01	.2404-01
HETAI = .00 DEG HETA2 = 20.00 DEG PPRE 1730-01 .2162+00	REES, PPIM •1442+00 •2322-01	+1452+00	1.690	•,2033-01 •2254+00	+1537+00 +2861=01	• 1550+00 • 2272+00	1.702	•2108-01 •4726=01	•2404-01 •5161-01
(HETAL = ,00 DEG (HETA2 = 20,00 DEG PPRE -,1730-01	REES, PPIM • 1442+00 • 2322=01 • • 2735+00	•1452+00 •2174+00 •3078+00	1.690	*,2033-01 ,2254+00 ,1492+00	•1537+00 •2861-01 ••2814+00	•1550+00 •2272+00 •3185+00	1.702	•2108-01 •4726-01 •9471-01	+2404-01 +5161-01 +1015+00
HETA1 = .00 DEG HETA2 = 20.00 DEG PPRE 1730-01 .2162+00 .1410+00	REES, PPIM •1442+00 •2322-01	+1452+00	1.690 .107 5.188 4.111	*,2033-01 •2254+00 •1492+00 •2134+00	•1537+00 •2861-01 •2814+00 •3145+00	•1550+00 •2272+00 •3185+00 •3801+00	1.702 .128 5.200 4.116	•2108-01 •4726-01 •9471-01 •1357+00	•2404-01 •5161+01 •1015+00 •1445+00
HETA1 = .00 DEG HETA2 = 20.00 DEG PPRE 1730-01 .2162+00 1410+00 =.2085+00	AREES. PPIM • 1442+00 • 2322-01 • 2735+00 • 3037+00	•1452+00 •2174+00 •3078+00 •3684+00 •3684+00	1.690 .107 5.188 4.111 3.056	* • 2033-01 • 2254+00 • 1492+00 • 2134+00 * • 4024+00	+1537+00 +2861-01 ++2814+00 +-3145+00 +3376-01	•1550+00 •2272+00 •3185+00 •3801+00 •4038+00	1.702 .126 5,200 4.116 3.058	+2108-01 +726-01 +9471-01 +1357+00 +1535+00	•2404-01 •5161-01 •1015+00 •1445+00 •1631+00
(HETA] = .00 DEG (HETA2 = 20.00 DEG PPRE 1730-01 .2162+00 .1410+00 =.2085+00 3903+00	GREES. PPIM -1442+00 -2322-01 -2735+00 -3037+00 -3331-01	•1452+00 •2174+00 •3078+00 •3684+00	1.690 .107 5.188 4.111	*,2033-01 •2254+00 •1492+00 *,2134+00 *,4024+00 *,1622+00	+1537+00 -2861=01 ⇒ 2814+00 = 3145+00 +3376=01 +3668+00	.1550+00 .2272+00 .3185+00 .3801+00 .4038+00 .4038+00	1.702 .126 5,200 4.116 3.058 1.987	• 2108-01 • 4726-01 • 9471-01 • 1357+00 • 1535+00 • 1513+00	• 2404-01 • 5161*01 • 1015+00 • 1445+00 • 1631+00 • 1608+00
HETAI = .00 DEG HETA2 = 20.00 DEG PRE 1730-01 .2162+00 1410+00 3903+00 1577+00	AREES, PPIM -1442+00 -2322-01 -2735+00 -3037+00 -3331-01 -3556+00	•1452+00 •2174+00 •3078+00 •3684+00 •3718+00 •3890+00	1.690 .107 5.188 4.111 3.056 1.988	*,2033-01 ;2254+00 ;1492+00 -,2134+00 *,4024+00 *,1622+00 ;2425+00	+1537+00 +2861=01 +2814+00 =3145+00 +3376=01 +3668+00 +3031+00	.1550+00 .2272+00 .3185+00 .3601+00 .4038+00 .4010+00 .3882+00	1.702 .126 5.200 4.116 3.058 1.787 .896	• 2108-01 • 4725-01 • 9471-01 • 1357+00 • 1535+00 • 1513+00 • 1417+00	• 2404-01 • 5161+01 • 1015+00 • 1445+00 • 14631+00 • 1608+00 • 1507+00
HETA1 = .00 DEG HETA2 = 20.00 DEG PPRE 1730-01 .2[62+00 3903+00 3903+00 3590+00 .3590+00 .1114+00	AREES, PPIM -1442+00 -2322-01 -2735+00 -3037+00 -3331-01 -3556+00 -2944+00	•1452+00 •2174+00 •3078+00 •3684+00 •3918+00 •3890+00 •3765+00	1.690 .107 5.188 4.111 3.056 1.988 .898	*,2033-01 •2254+00 •1492+00 *,2134+00 *,4024+00 *,1622+00	.1537+00 .2861-01 .28614+00 .3145+00 .3376-01 .3668+00 .3031+00 .3031+00	.1550+00 .2272+00 .3185+00 .3801+00 .4038+00 .4010+00 .3882+00 .3882+00 .3771+00	1.702 .126 5.200 4.116 3.058 1.987 .896 6.085	.2108-01 .4726-01 .9471-01 .1357+00 .1535+00 .1513+00 .1417+00 .1417+00	.2404-01 .5161-01 .1015-00 .1445+00 .1631+00 .1608+00 .1507+00 .1422+00
HETAI = .00 DEG HETA2 = 20.00 DEG PPRE 1730-01 .2162+00 .1410+00 =.2085+00 3903+00 =.1577+00 .3590+00 .1114+00 2246+00	AREES. PPIM .1442+00 .2322-01 -2735+00 .3037+00 .3331-01 .3556+00 .2944+00 -7144=01 -3434+00 -2782+00	.1452+00 .2174+00 .3078+00 .3884+00 .3918+00 .3918+00 .3765+60 .3661+00	1.690 .107 5.188 4.111 3.056 1.988 .898 6.087	*,2033-01 *2254+00 *,2134+00 *,4024+00 *,1622+00 *,2425+00 *,2425+00 *,3697+00	+1537+00 +2861=01 +2814+00 =3145+00 +3376=01 +3668+00 +3031+00	• 1550+00 • 2272+00 • 3185+00 • 3801+00 • 4038+00 • 4010+00 • 3882+00 • 3771+00 • 3709+00	1.702 .126 5,200 4.116 3.058 1.987 .896 6.085 5.024	.2108-01 .4726-01 .9471-01 .1357+00 .1535+00 .1513+00 .1417+00 .1340+00 .1303+00	.2404-01 .5161+01 .1015-00 .145+00 .1631+00 .1608+00 .1507+00 .1422+00 .1376+00
HETA1 = .00 DEG HETA2 = 20.00 DEG PPRE 1730-01 .2[62+00 3903+00 3903+00 3590+00 .3590+00 .1114+00	REES. PPIM -1442+00 -2322-01 -2735+00 -3331-01 -3556+00 -2944+00 -7144-01 -3434+07	• 1452+00 • 2174+00 • 3078+00 • 3684+00 • 3918+00 • 3890+00 • 3765+00 • 3661+00 • 3610+00	1.690 .107 5.188 4.111 3.056 1.988 .898 6.087 5.026	*,2033-01 •2254+00 •1492+00 •4024+00 •1622+00 •2425+00 •3697+00 •1137+00	• 1537+00 • 2861-01 • 2814+00 • 3145+00 • 3376-01 • 3668+00 • 3031+00 • 7432-01 • 3531+00	.1550+00 .2272+00 .3185+00 .3801+00 .4038+00 .4010+00 .3882+00 .3882+00 .3771+00	1.702 .126 5.200 4.116 3.058 1.987 .896 6.085	.2108-01 .4726-01 .9471-01 .1357+00 .1535+00 .1513+00 .1417+00 .1417+00	.2404-01 .5161-01 .1015-00 .1445+00 .1631+00 .1608+00 .1507+00 .1422+00

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		E SLOPE CORRE	CTION TO PI	RMSN IS SLPCO	DK ■ 2+87 P	ERCENT.			
HETAI						·			
HETA2 = 40.00 DEG						PHAG	PPHAZ		P2
PPRE	PPIM	PPMAG	PPPHAZ	PRE	PIH	LUNAR .	EE 1186		
.5576=02		.3954-01	1.422	.6804-02		.4225-01		-1564-02	-1785-02-
+3271-01	.3178-03	.3291-01	.010	.3486-01	4780-03	.3486-01	6.269	.1083-02	+1215-02
.5019-02-				-7500=02-		.2594=01			
1304-01	1670-01	.2119-01	4.049	1318-01	1694-01	.2146-01	4.051	.4491-03	.4605-03
2475-01				2601-01	-2126-02	.2610-01	3.060		
5310-02	.3071-01	.3137-01	1.741	-,6261-02	.3280-01	.3339-01	1.759	.9838-03	.1115-02
.3338-01	.9478-02		.280		.1038-01		.284		
.9017-02	3085-01	.3214-01	4,997	.9434-02	3364-01	.3494-01	4,986	.1033-02	.1221-02
		-2510-01						+6300-03-	
.4604-02	.1633-01	.1697-D1	1.296	.5793-02	.1857-01	.1945-01	1.268	.2880-03	.3784-03
-7351=02-		-1659-01				-1909-01			
2477-01	5545-03	.2477-01	3.119	2744-01	2307-03	.2744-01	3.150	.6136-03	,7528-03
.6562-02	3250-01	3316-01		.6509=02-			1.390	-1100-02-	
.3713-01	1085-01	.3868-01	5,999	.4046-01	1142-01	.4204-01	6.008	.1496-02	.1767-02
			4.358				4.361-	-1695-02-	.2001-02-
3771-01	.1776-01	4168-01	2,701	4101-01	.1929-01	.4532-01	2,702	.1737-02	.2054-02
.2146=01	-3556=01		1.028	-2337=01		.4517-01	1.027	-1725-02	
.3341-01	2475-01	.4158-01	5.646	.3623-01	2688-01	.4511-01	5.645	.1729-02	.2035-02
		4184-01	4.025				4.024		-2038-02-
3296-01	.2615-01	.4207-01	2,471	3530-01	.2808-01	.4511-01	2,470	.1770-02	.2035-02
-2363-01-		-4245-01	.980	.2521-01	.3763-01	.4529-01	.980	.1802-02	-2051-02-
.3444-01	1977-01	.4323-01	5,808	.4101-01	2100-01	.4608-01	5.810	.1869-02	.2123-02
			4.354			.4689-01		-1928-02-	
4172-01	.1090-01	.4312-01	2.886	4459-01	.1185-01	.4614-01	2.882	.1860-02	.2128-02
	-3910-01	.3954=01		.6807=02-	.4170-01	.4225-01	1.407	.1564-02	-1785-02
			-	AVERAG	ES +++				
8220-07	.4149-06	.4230-06	1.766	Tel109=00	.4573=06	.4705-04	1.807	.1243-02	-1440-02

,						*** AVERAGES					
		2	-02 .3	978=02	4,183	-,2010-02	-,3433=02	-3978-02	4,183	.7487-01	.7935-01
	THE RMS	SCATTERED	PRESSURE	#/O SLOPE	CORRECTION	IS PPRMS =	.27362+00				
	THE RMS	SCATTERED	PRESSURE	TITH SLOP	E CORRECTIO	N IS PRHS -	.28169+00				
	THE SAM	E QUANTITI	ES DIVIDE	D BY PO AR	E PPRMSN =	•55293+01	3				
					PRISN	-56924+0	T				

							1. 200		
.5415-01		.3078+00			.3092+00				
.2277+00	.1265+00	.2604+00	.507	.2341+00	.1272+00	.2664+00	.498	.6782-01	•7099-01
.1/79+00	7858-01	.1963+00	5,871-	1826+00-			5,855		
.3362-01	1237+00	.1284+00	4.977	.3081-01	1266+00	.1303+00	4.951	.1649-01	.1699-01
			4.308				4,287		.5924-02
5402-01	5126-01	.7447-01	3,901	5252-01	-,5230-01	.7412-01	3.925	.5546-02	.5493-02
	. 6672=02	+1087+00		1110+00					
6106-01	+1411+00	.1538+00	1,979	6585-01	.1448+00	. 1591+00	1,998	.2364-01	.2530-01
.1343+00		.1975+00	.823		-1516+00-	2048+00	.834		
.2113+00	7015-01	.2220+00	5,963	.2196+00	7230-01	.2312+00	5.965	.4957-01	.5345-01
		.2098+00	4.797	-1727-01-	2182+00-	+2189+00			
1545+00	4993-01	.1624+00	3.454	1640+00	5109-01	.1718+00	3.444	.2638-01	.2950-01
		.1452+00						-2108-01	
				*** AVERAG	ES ***				
			4.183	-,2010-02-			4-183		

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THE RMS SCATTERED P	RESSURE WITH SLOPE	CORRECTION IS	PPRM5 =	.37953-01			
						1	
THE SAME QUANTITIES	DIVIDED BY PD ARE	PPRMSN	71234-01				
		PRMSN #	76696-01				
THE CONTRIBUTION OF	THE SLOPE CORRECT	TON TO PRHSN I	S SLPCOR .	7.12 PERC	ENTO		
	•						
	, 						
					,		
		1					
F1N							
			•				
RUNIDI BO9606 PROJEC	T :	USER:					
			· · · · · · · · · · · · · · · · · · ·				
LOAD HATHPK 8/3 BO9	606						
ITEM	AHOUNT	COSTIDOLLAR	e 1				
	KIIOONT	COSTUDUELAR					
PU TIME	00:00:09.778	\$0.65					
TO REQUESTS	530	\$0.63		·····			
10-AORDS TRANSFERRED	373057	\$0.22					
ORE USAGE	0.397	\$0.22				······································	
AGES PRINTED	310	\$0.15 50.39					
APE UNITS USED		SU•40					
OB CHARGE	······	50.20					
	•						
OTAL COST		\$2.86					
THE ABOVE DOLLAR AMOUNTS	S ARE APPROXIMATE	AND ARE BASED	UN RATES FI	R STANDARD	RUNS	······································	
INITIATION TIME: 0312							
TERMINATION TIME: 031	21:18-MAY 16,1973		· · · · · · · · · · · · · · · · · · ·		· · · · · · · · · · · · · · · · · · ·		
	· · · · ·						

<u>Program PSIN2</u> -- The rms scattered pressure is evaluated directly from Equation 39. A considerable simplification is obvious in comparing this program with PSIN1, but the programs are basically similar and additional comments do not seem to be necessary. One note, however; the present output format is not a particularly efficient one and should probably be changed.

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00001	BASG+MT 11.+T+NATHPK
00002	
00003	OFREE 11-
00004	afor, sia main
00005	C
.00006	
00007	c
000.08	C IT COMPUTES THE RMS - PRESSURE - SCATTERED - IN THE X-Z-PLANE FROM A
00009	C CORRUGATED SURFACE. A SLOPE CORRECTION IS INCLUDED.
00010	
00011	C A HIGHLY DIRECTIVE TRANSMITTED ACOUSTIC BEAM AND A POINT RECEIVER
00012	C ARE- ASSUKED
00013	C THE AMPLITUDE PROFILES OF THE INCIDENT PRESSURE ARE ASSUMED TO BE
00014	C GAUSSIAN, THEY ARE MEASURED ON A HORIZONTAL PLANE AT A DISTANCE R1
00015	C FROM THE SOURCE TRANSDUCER.
00016	C BWX IS THE BEAMWIDTH = FWHM/1.665 MEASURED IN THE X-Z PLANE.
-00018	C BNT IS THE BEAMWIDTH MEASURED IN THE T-2 PLANE.
00019	C MODE SPECIFIES THE SCATTERING GEOMETRY.
00020	
00021	C MODE =-1 EACKSCATTERING.
00022	C MODE=D TRANSMITTER ANGLE FIXED AT ZERO
00023	C MODED IS THE INITIAL VALUE OF MODE.
00024	
00025	C NMODE IS THE NUMEER OF VALUES OF MODE. NORMALLY 1+ 2+ OR 3-
00026	C. THETAO IS THE INITIAL VALUE OF THE RECEIVER ANGLE WAT VERTICAL.
00027	C DTHETA IS THE INCREMENT.
00028	C-NTHETA IS THE NUMBER OF VALUES USED.
00023	C OFFSET IS THE ANGULAR DISPLACEMENT OF THE TRANSDUCER WRT ITS NORMAL
00031	C DISPLACEMENT IS IN THE + THETA1 DIRECTION.
00032	C DISPLACEMENT IS IN THE + THETAL DIRECTION.
00032	C THETAZ IS + 'DOWNWIND'.
00034	C ZETA IS THE AMPLITUDE OF THE CORRUGATIONS. IT MUST BE .GT. ZERO.
00035	C ZETAO IS THE INITIAL VALUE OF ZETA.
00036	DZETA IS THE INCREMENT BY WHICH ZETA IS INCREASED.
00037	C NZETA IS THE NUMBER OF VALUES OF ZETA USED.
00038	C WL IS THE ACOUSTIC WAVELENGTH.
00033	C SWL IS THE SURFACE WAVELENGTH.
00040	CK IS THE ACOUSTIC WAVENUMBER.
00041	C SK IS THE SURFACE WAVENUMBER.
00042	CR1_IS_THE_RADIUS_OF_THE_ARC_ON_WHICH_THE_TRANSHITTER_MOVES
00043	C R2 IS THE RADIUS OF THE ARC ON WHICH THE RECEIVER NOVES.
00044	C. BESSEL FUNCTIONS ARE USED IN THE FORM JINI, WHERE N IS THE ORDER AND C NMAX=16 IS THE MAXIMUM ORDER.
00046	C A IS THE ARGUMENT OF THE BESSEL FUNCTIONS.
00047	
00048	
00049	REAL J(100),JO
00050	
00051	REAL MODED.MODE
00052	READ 40+ MODEO+DHODE-MODE
00053	READ 40+ EWX+BWY
.00054	READ 40. THETAD.DTHETA.NTHETA

-100-

00055	READ 40. OFFSET
00056	READ 40+ ZETAO DZETA NZETA
00057	READ 40+ 21-R0 00221R0 0221R
00058	READ 40, NL+SWL
00059	40 FORMAT(2F10.0+110)
00060	
00061	PRINT 45
00062	A5 FORMAT("1THE INPUT CARDS ARE")
00063	PRINT 50+ MODED+DMODE+NMODE
00064	PRINT 50+ BWX+BWY
00065	PRINT 50, THETAO, DTHETA, NTHETA
00066	PRINT 50. OFFSET
00067	PRINT 50+ ZETAO+DZETA+NZETA
83000	PRINT 50+ R1+R2
00069	PRINT 50, WL, SWL
00070	50 FORMAT(24X,2F10,4,10)
00071	
00072	DATA NHAX/16/
00073	DATA TWOPI/6.2832/
00074	
00075	KETWOPI/WL
00076	SK=TWOPI/SWL
00077	A3=(1./R1+1./R2)/2.
00078	BWADEBWX
00079	c
00080	C COMPUTE THE MAGNITUDE OF THE PRESSURE REFLECTED VERTICALLY FROM A
00081	C SMOOTH SURFACE.
00082	C
00083	AX D=SQRT (SQRT (BWX++++(K+A3)++2))
00089	AYD=SGRT(SGRT(BWY++-4+(K+A3)++2))
00085	P0=K/(2.+R2+AX0+AY0)
00086	PRINT 60, PO
00097	60 FORMAT("OTHE PRESSURE REFLECTED FROM A SMOOTH SURFACE IS PO =""
00088	
00089	C
00030	C ESTABLISH THE GEOMETRY.
00091	c
00092	MODE=HODE0-DHODE
00093	DO 260 MD=1+NMODE
00094	MODE=HODE+DHODE
00095	c
00096	C ESTABLISH ZETA.
00037	c
00098	ZETA=ZETAO-DZETA
00099	DO 260 IZ=1.NZETA
00100	ZETA=ZETA+DZETA
00101	PRINT 65. ZETA
00102	65 FORMAT(*1++++++++++++*/* ZETA =*+F5.3/* ++++++++++++++++++++++++++++++++++++
00103	c
00104	C. ESTABLISH THE TRANSMITTER AND RECEIVER ANGLES. THE TAL AND THE TAZ.
00105	C
00106	THE TA2=THE TAD-DIHETA
00107	DO 260 NT=1.NTHETA
00108	THE TA2=THE TA2+D THE TA

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00109	ANG2=THET A2/57.295
00110	CA2=COS (AN G2)
00111	THETA1=MODE + THETA2+OFFSET
00112	ANG1=THE TA1/57 •295
00113	CA1=COS(ANG1)
00114	C
00115	PRINT 70, THETA1, THETA2
00116	
00117	c
00118	C CONPUTE SOME TERNS.
00119	C C
00120	A1=(CA1++2/R1+CA2++2/R21/2.
00121	A2=SIN (ANG1)-SIN (ANG2)
00122 00123	A4=CA1+CA2 A=K+ZETA+A4
00123	BWX=BWX0/CA1
00124	AX2=SQRT(BXX*+-4+(K*A1)++2)
00125	AZ-SQRT(GWAZ)
00127	AY2=S3RT(BWY++-4+(K+A3)++2)
00128	AY=SQRT(AY2)
00129	PHIX=ATAN(K+A1+BHX++2)
00130	CX=COS (PHIX)
00131	GNU2=4•+AX2/(K+K+CX)
00132	
00133	C
00134	CBESJ IS A UNIVAC MATH-PAC-ROUTINE TO COMPUTE BESSEL FUNCTIONS.
00135	C
00136	CALL BESJ(A+0++MAX+J)
00137	JO=J(1)
-00138 00139	
00140	
00141	C COMPUTE THE RMS PRESSURE.
00142	
00143	PPSUM=J0+J0+EXP(-2.+A2+A2/GNU2)
00144	PSUM=PPSUM
00145	c
00146	REAL N
00147	DO ZOO NN=1+NMAX
00148	N=NN
00149	C
00150	ALFP2=(A2+N+SK/K}++2/GNU2
00151	ALFM2=(A2-N+SK/K)++2/GNU2
00152	E=EXP(-2.+ALFP2)+EXP(-2.+ALFM2)
00153	EPS=(N+SK/(K+A4))++2
00154	
00155	200 PSUM=PSUM+(((1.+EPS)+J(NN))++2)+E
00156	
00157	C PPRMS IS THE SMALL SLOPES APPROXIMATION OF PRMS.
00158 00159	C EPS IS THE SLOPE CORRECTION TERM.
00159 00160	C PPRMS=CP+SQRT(PPSUN)
00161	PRHS=CP+SQRT(PSUN)
00161	

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00163		IS THE PERCEN	TAGE CONTRIBUT	ION OF THE	LOPE CORRE	CTION.	
00164	C					,	
00165	SLPC	DR=(1PPRMS/	PRMS) +100+				
00166		SN=PPRMS/PO					
00167	PRMSI	N=PRMS/PO					
00168	_						
00169		T 250. PPRMS.					
00170	250 . FORM	ATT 9X+ THE	RMS_SCATTERED_	PRESSURE M/C	SLOPE CORI	CECITON IS	_ <u></u>
00171	IRMS	='+E13+5/ 9X	.THE RHS SCAT	TERED PRESS	WE WITH 20	JPE CURREC	110
		PRMS = + 113.					
00173			+PRMSN+SLPCOR		-	N -1.517	E
00174	265 FURM	LIGX THE SA	ME_QUANTITIES_ 3.5/9X. THE CO	NTOTOUTTON	E THE SLAP	C COPPECTY	0N
00175	17438	1.6KUZN1ET	R = 1 + F7.2 1 PE	DOCHT TAN	IF THE SCUP	CONNECTI	
00176	210 PI END	ANDR LA ALEU	n	∩×⊑∥⊾≞∴∕-J			
00178	C	5					
00179	axor						
00180		0.	1				
00191	6.154	6.154					
00183	0.						
			1			-	
00185	81.5	81.5	_			•	
		8.8					
00187	afin						
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C. A program for the computation of probability density functions for surface displacement: program PDF.

Much of this program is nearly identical to program <u>TRANSLATE</u>, and most of the discussion of that program is pertinent here also. The reason for this is that both programs were intended to utilize the same input (surface displacement) data. Here, a PDF for the surface displacement is computed both for the entire data set (file), as read from the digital tape, and for the 'average' data block previously described.

The basic procedure for computing the PDF of the entire file is as follows: An N-word data block is read and converted to 36-bit words by subroutine RREC. The data are placed in array IY which is common to the main program and to subroutines RREC and DSTFUN. At entry DSTF1, subroutine DSTFUN sorts the data into histogram bins. Another data block is then read and sorted. When all blocks have been read, the histogram is normalized by a call to DSTFUN at entry DSTF2. The PDF of the 'average' block is obtained by one call to each of the entry points of DSTFUN.

Much of the program between cards 124 and 246 is concerned with computing mean and rms values, and with selecting an average block. These operations are unnecessary in terms of simply obtaining a PDF for a given data set. From that point of view, the essential, and most useful, part of the program is subroutine DSTFUN.

-104-

00001	BASG, TH 17/PASSWORD BC9, REEL ID
00002	
00003	C C
00004	
00005	C IT HAS BEEN WRITTEN TO ANALYZE WAVEHEIGHT DATA THAT HAS BEEN
00006	
00007	C PLOCKS.
000012	C - THE DATA WORDS ARE CONVERTED FROM 24 BITS TO 36 BITS.
00003	C THE MEAN IS FOUND AND IS SUBTRACTED FROM EACH DATUM.
00010	C -AN RMS VALUE-IS COMPUTED FOR EACH BLOCK.
00011	C THE MEAN OF THE RMS VALUES IS CALCULATED, AND AN "AVERAGE" BLOCK
00012	C HAVINC AN RMS VALUE CLOSE TO THE MEAN RMS IS FOUND.
00013	C NORMALIZED SURFACE DISPLACEMENT PROBABILITY DENSITY SUNCTIONS (DOFICE)
00014	C ARE OBTAINED FOR THE WHOLE DATA SET AND FOR THE AVERAGE BLOCK.
00015	
00016	C - N IS THE NUMBER OF WORDS IN A DATA BLOCK.
00017	C FOR, THIS PROGRAM, N MUST RE AN INTEGER MULTIPLE OF 7
00018	C N CAN 22 CHANGED BY CHANGING THE DADAWE TED OT A TEMENT
00019	V NEILLS IS THE NUMBER OF TAPE FILES TO BE SKIPPED.
00020	C MB IS A NUMBER OF THE NUMBER OF BLOCKS (OR BECORDS) TH ANY CILE TO
00021	G DE SKIPPED.
00022	
00023	C PUBLIIVE INTEGER FORM). IYMAXIIN74 FOR A IN-ATT CTOTTTZCO.
0002 4 00025	C-N 12C 13-1HE WIDTH OF THE HISTOCRAM ATMS.
00025	C SIGN IS SET .LT. D. IF THE DATA ARE INVERTED ON THE TAPE.
00027	C SKATE 13 THE EFFECTIVE DIGTTTFFF CAMPLEND DATE
00028	C CITECTITE HEARS TOIGITIER SAMPLING RATEI+(RECORDER RECORD SPEED)
00023	C PKHT TS THE ASSUMED DEAK HETCHE OF THE HETCH HETCH
00170	C PKHT IS THE ASSUMED PEAK HEIGHT OF THE WATER WAVES. C SPEED IS THE SPEED OF THE WATER WAVES.
00031	C WLNGTH IS THE WAVELENGTH OF THE WAVES.
0032	C NB IS THE NUMBER OF NEWORD BLOCKS IN THE FILE.
00033	C IB IS THE INDEX OF THE AVERAGE BLOCK. IT CAN HAVE ANY VALUE FROM
00034	C-1 TO NE
10035	C EPS IS THE FRACTIONAL DIFFERENCE BETWEEN THE AVERAGE RMS AND THE RMS
10038	C OF THE AVERAUE FLUCK
30037	C RMGHT IS THE RMS SURFACE DISPLACEMENT IN CM FOR THE DATA IN THE
10033	C AVERAGE BLOCK.
10039	C CXP IS THE HORIZONTAL SEPARATION IN CH OF THE SURFACE DATA.
00040	
00/141	c
0042	PARAMETER N=500+IYMAX-1024
10043 1004 4	COMMON IY(N)
10044 10045	COMMON/DSF/DFIIYMAX)+NSIZE+NIY
0045	DIMENSION LABEL(7), RMS(1000)
10047	DIMENSION TEMPINIVXIIYHAX)
0048	REAL MEANING
10043	NNEN N
0050	
0051	
0052	C INPUT CARDS.
0053	
0054	READ 10+ NFILES-MB

10055		10 FORMAT(2110)
00057 10057		PEAD 10, NSTZE
0057		READ 20, SIGN
00059		
<u>, , , , , , , , , , , , , , , , , , , </u>		READ 20,PKHT, SPEED, WLNGTH
00051	с	20 FORMAT(3)10-07
10052	č	**************
00063	ç	
10064		PRINT 30, NFILES
00065		30 FORMAT(1H1+ NFILES = +12+ TAPE FILES HAVE BEEN SKIPPED. +)
10066		PRINT 35, IYAINSIZLISIGN
1006 7		40 FORMAT(1H0, THE EFFECTIVE SAMPLING RATE IS .F6.1. SAMPLES PER SEC
10068		PRINT 40, SRATE
10059		35 FORMAT(' IYMAX =',15/' NSIZE =',13/' SIGN =',F4.0)
070		10ND
00071		PRINT 50, PKHT
10072		SO FORMATI' THE PEAK HT OF THE AVERAGE WAVE IS ".F6.3." CH."
0073		PRINT 60, SPEED
10074 10075		FOPMAT(" THE SPEED OF THE WAVES IS ".F6.2." CM PER SECOND.") PRINT 70, WLNGTH
0076		TO FORMATI' THE WAVELENGTH OF THE WATER WAVES IS +F6.3. CH. +1
0077	с	TO FORMATE THE WATELENGTH OF THE WATER WAVES IS FEB-34. UN. 1
0078	č-	SKIP FILLS THAT ARE NOT OF INTEREST. LAULLS ARE FILES.
0079	c	
10030		IF (NFILES .NE. 0) CALL SKIPF (NFILES, MB)
10031	С	
0082		READ AND PRINT THE TAPE LABEL.
10037	С	IOTPIN IS AN MACC 1103 LIBRARY SUBROUTINE TO READ A DATA BLOCK FROM
0024		MAGNETIC TAPE. SEE THE REFERENCE MANUAL FOR THE 1108 COMPUTER.
00085 00086	C	IN THIS CASE THE LABEL IS STRICTLY NUMERIC AND IS RECORDED AS 24-BIT
00097		WORDS. FLD MANIPULATES THE BITS TO READ THE WORDS IN A 36-BIT (1108)
10057		
00089	C	CALL IOTPIN(17,1,1Y,N,LNZ,\$80)
0090		80 LADEL(1)=FLD(0,24,1Y(1))
0031		LABEL(2)=FLD(0+12+IY(2))
0092		LABEL (3)=FLD(12,24,1Y(2))
10033		LABEL(4)=FLC(0,24,IY(3))
0094		UABEL (5)=FUD (0+12+IY(4))
0035		LABEL(6)=FLD(12,24,IY(4))
0036		LABEL171=FLD10+24+1Y(5))
0037		PRINT 90, (LABEL(J), J=1,7)
10038		9D FORMAT (///* TAPE NUMBER*+16+10X+*DATA TAKEN ON *+12+*/*+12+*/*+12/
10093 10093	c	128X, *DATA DIGITIZED ON *+12+*/*+12+*/*+12)
0101		AN EOF SHOULD FOLLOW THE LABEL.
0102	č	FINE THE LABEL EDF.
0103	č	SUBROUTINE IOTPSP SPACES LOGICAL UNIT 17 (HERE) PAST THE NEXT DATA
0104	-	BLOCK. IT TRANSFERS CONTROL TO STATEMENT 100 IF AN EOF IS FOUND.
0105	Ċ	
0100		CALL TOTPSP(17,\$100)
0107		PRINT 95
0108		95 FORMAT(//* ***** LABEL EOF NOT FOUND. ******

00103	C
00110	C BECIN THE ANALYSIS OF THE DATA
00111	C READ THE DATA BLOCK BY BLOCK. FOR EACH BLOCK COMPUTE THE RMS AND THE
00112	
00113	C STANDARD DEVIATIONS OF THE RMS'S AND MEANS.
00114	
00115 0011 6	C (NS) WHICH CORRESPOND TO AN INTEGER NUMBER OF WAVELENGTHS.
00117	C -FINALLY, USE SUBROUTINE DSTFUN TO COMPUTE THE NORMALIZED
00119	CSUBROUTINERREC-CONVERTS-24-BIT-WORDS-TO-36-BIT_WORDSCONTROL-COES
00113	C TO STATEMENT 180 IF AN EOF IS FOUND, AND TO 160 IF A TAPE ERROR IS
01113	C TO STATEMENT 180 IF AN EUF 15 FOUNDT AND TO 180 IF A TAPE ERROR 15
00121	C AN N-WORD BLOCK IS READ BY RREC, THEN THE DATA ARE SORTED INTO
	C HISTOCRAM BINS BY SUBROUTINE DSTF1.
00123	
00124	
00125	DXP=JPEEC/SRATE
00126	SPUL=WLNGTH/DXP
00127	T = N - 1
00128	WLPBL=AINT(T+DXP/WENGTH)
00123	NS=SPWL+WLPBL+1.
00130	J=()
00131	K=0
10132	SUM2=0.
00133	SUM 3=0.
	SUM4=0.
00135 00136	SUM5=0.
00137	SUM1=0.
80139	5011-0• →=
00139	CALL RREC(NW, \$180, \$160)
00140	
00141	PRINT 105
0142	
00143	CO TO 160
00144	110 CALL DSTF1
70145	K = K + 1
00146	
00147	SUM=SUM+FLOAT(IY(I))
0148	120 SUM1=SUH1+FLOAT(IY(I))++2
00143	F=NS
00150	
70151	RMS(J)=SQRT(SUM1/F-HEAN(J)++2)
001 52 00153	
00154	SUM 3=SUM 3+MEAN (J) + + 2
00155	SUM5=SUM5+RMS(J)++2
90195 90156	
0157	150 PRINT 170, J
0158	170 FORMAT (//* ***** TAPE ERROR IN BLOCK NO.**I4** ******//)
0159	MEAN(J)=0.
10100	RMS (J)=D.
10151	GO TO 100
10162	

00163	NB=J-1
00164	AVGMN=SUM27FK
00165	SDMEAN=SQRT((SUM3-FK+AVGMN++2)/(FK-1.))
00166	AVGRMS-SUM47FK
00167	SDRMS=SQRT((SUM5-FK+AVGRMS++2)/(FK-1.))
00163	C
00169	C SUBROUTINE DSTF2 NORMALIZES THE HISTOGRAM.
00170	c
00171	CALL DSTF2
00172	PRINT 190, NS,WLPEL
00173	190 FORMAT(///* THE FIRST*+I4+* WORDS IN EACH BLOCK ARE USED TO COMPU
00174	IE THE MEAN AND THE RMS. THESE WORDS CORRESPOND TO +F5.1. WAVELE
00175	26THS.*)
00176	PRINT 200, KANGAN, SOMEAN
00177	200 FORMAT(/// THE AVERAGE OF THE MEANS OF .I4. FROR-FREE BLOCKS I
00178	1,F10.5/ THE STANDARD DEVIATION OF THE MEANS IS FID.5) PRINT 210, AVGRMS, SORMS
-00180	210 FORMAT ("OTHE AVERAGE OF THE RMSES IS" FFID. SV" THE STANDARD DEVIAT
00191	10N OF THE RMSES IS + FIG. 5)
	PRINT 220
00182 00193	220 FORMAT("ITHIS IS ARRAY MEAN."//)
-00184	PRINT 225, thean(J), J=1, NB)
00135	225 FORMAT(1H +10F1^.5)
00135	PRINT 240
00137	240 FORMAT("1THIS IS ARRAY RMS."//)
00198	PRINT 2250 (RMS(J)) J=108
00199	
-00190	- FIND A BLOCK CONTAINING DATA WITH AN RHS CLOSE TO AVORMS.
00131	
00192	EPS0=+005
00193	EPS=EPS0
00194	260 J=0
00195	270 J=J+1
00196	IF (ABS(RMS(J)/AVGRMS-1.) .LT. EPS) GO TO 300
00197	IF(J-N8) 270,280,280
00138	280 EPS=EPS+EPS0
00199	IF(EPS-10.+EPSO) 260,260,330
	300 PRCNT=100.+CPS
00201	IG=J
-00202	PRINT 320+ IB+RMS(IB)+PRCNT
00203	320 FORMAT(//' BLOCK NUMBER',14," HAS AN RMS VALUE OF', F7.2.". THIS
00204	15 WITHIN ** F4.2+ * PERCENT OF THE MEAN RMS. *}
00205	GO TO 340
00206	330 PRINT 335
00207	335 FORMAT(//* ***** CANT FIND AN RMS(J) CLOSE TO AVGRMS. GUIT. *****
00208	CALL EXIT
00209	C
00210	C REWIND THE TAPE AND RETURN TO THE BECINNING OF THE WAVE CATA.
00211	C
-00212	340-CALL_TOTPR#(17)
00213	NFP1=NFILES+1
00214	CALL SKIPF (NFP1+MB)
00215	C C

····	
00217	C AGAIN. AS A CHECK.
00219	360 IMAX=IB-1
-00220	
00221 	370 CALL IOTPSP(17)
00223	CALL RRECINW/\$400,\$440; D0 390 I=1.N
-00224	
00225	
-00226	
00227	DO 390 I=1+NS
00229	
00229	39D SUM1=SUM1+FLOAT(IY(I))++2
00230	
00231	CKME AN =SUM/F
00232 00233	
-00234	PRINT 400, CKMEAN, CKRMS
00235	400 FORMAT(//* CHECK THE HEAN AND RHS OF BLOCK IB.*/10X.*HEAN =*vF10.* 1+10X.*RMS ='+F10.5)
00236-	PRINT-428-NW
00237	420 FORMAT(//* CHECK IF NW=N. THE NUMBER OF DATA READ INTO ARRAY TEMP
-00238	
00239	GO TO 500
-00240	440 PRINT 4457 18
00241	\$45 FORMAT(//* ***** TAPE ERROR IN BLOCK IB =*,14.*. TRY ANOTHER BLO
-00242	
00243 -00244	60 TO 270 480 PRINT 490
00245	400 FRIMT 440 490 FORMAT(//* ***** EOF FOUND BY RREC. QUIT. ******)
-00246	CALL EXIT
00247	C
00248	C FIND THE LARGEST POSITIVE AND NEGATIVE VALUES IN THE AVERAGE PLOCK.
00249	C URSRCH IS AN MACC LIBRARY SUBROUTINE TO FIND THE LARGEST OR SMALLEST
- 00250	C ELEMENT IN AN ARRAY.
00251	C C
00253	500 CALL URSRCHID+N+TEMP+LP+TMPMAX+0+DUH) CALL URSRCH(1+N+TEMP+LN+TMPMIN+0+DUM)
-00254	PRINT 520, LPTTMPMAX
00255	520 FORMATI'OCATUM NUMBER'. 14." IS THE LARGEST POSITIVE DATUM IN THE
-00255	1VERACE BLOCK IT HAS THE VALUE*/F10.4)
00257	PRINT 530, LN, TMPMIN
00258	530 FORMAT(*ODATUM NUMBER*+I4+ IS THE LARGEST NEGATIVE DATUM IN THE A
00259	IVERAGE BLOCK. IT HAS THE VALUE',F10.4)
-00260	C THE CONVERSION FACTOR OF WILL ALLOW THE COMPUTATION OF ARRAY Y TH
00261	C THE CONVERSION FACTOR OF WILL ALLOW THE COMPUTATION OF ARRAY X IN
00253	
	THPETMPHIN
00265	IF(SIGN .GT. D.) THP=TMPMAX
-00266	CF=PKHT/TMP
00257	RMSHT=ABS(CF)+RMS(IB)
00268	PRINT 580, CF
00253	590 FORMAT("OTHE WAVEHT CONVERSION FACTOR IS CF ="+E10+4)

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00271	590 FORMAT(* THE RMS WAVENT OF THE AVERAGE BLOCK IS RMSHT =**F5*3** CM
00 272 00273	PRINT 500 DXP
00274	
00275	1**F6.4** CM.*)
00276 00277	C COMPUTE VALUES FOR THE ABSCISSA OF THE HISTOGRAM.
00278	C
00279	XD=FLOAT(NSIZE-IYMAX)/2.
00280	
00291	650 X(I)=(X0+FLOAT((I-1)*NSIZE))*CF
00282	C C PRINT THE X ARRAY AND THE PDF ARRAY FOR THE WHOLE DATA SET.
00293	C PRINT THE A ARRAT AND THE FUF ARRAT FOR THE WHOLE DATA SETS
00285	PRINT 670
00285	570 FORMAT (TARRAYS X AND DF. 1//)
00297	PRINT 580, $(X(I), OF(I), I=1, NIY)$
10238	680 FORMAT(1H F12.4+F14.6)
00239	c /
0290	C COMPUTE AND PRINT THE X AND PDF ARRAYS FOR THE AVERAGE BLOCK.
00291	C
00292	DO 690 I=1+NIY
00233	690 DF(I)=0.
00294	CALL DSTFI
00235	CALL DSTF2
00290	
00297	PRINT 700 700 Formatt*1Arrays X and DF for the Averace Block.*//)
00298	PRINT 680+ (X(I)+DF(I)+ I=1+NIY)
00299	END
00301	C
00302	af OR ISIA SUB1
00303	C
00301	SUBROUTINE DEFEN
00105	c c
00306	C THIS SUBROUTINE SORTS INTEGER DATA INTO BINS TO PRODUCE A
00307	C NORMALIZED HISTOGRAM. DSTF1 SORTS, DSTF2 NORMALIZES.
00308	C PARAMETER N=600+IYMAX=1024
00710	
00311	COMMON/DSF/DF(IYMAX).NSIZE.NIY
00312	C
00313	ENTRY DSTF1
00314	c
00315	CO 20 J=1.N
00316	K=IY(J)/NSIZE+1
00317	20 DF(K)=DF(K)+1.
00318	RETURN
00319	
00320	C ENTRY DSTF2
00322	د
00323	50 40 1=1.NIY
00324	40 SUMDF=SUMDF+DF(I)

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00325	DFNORM=1./(SUMDF+NSIZE)
00326	
00327	60 DF(I)=CF(I)+DFNORM
00329	END
-00330	
00331 -003 32	aFOR+SIA SUB2
00333	SUBROUTINE RREC(NWORDS+\$+\$)
-00334	C THIS IS A SUBROUTINE TO READ 1 RECORD OF TAPE FROM UNIT NO. 17 AND
00335 -80336	C THIS IS A SUBROUTINE TO READ 1 RECORD OF TAPE FROM ONLY NO. 17 AND
00337	C
00338	PARAMETER N=600 DIMENSION IX(N)
00339 80348	COMMON IY(N)
00341	c
00342	H=2+N/3 CALL IOTPIN(17+1+IX+M+LNZ+\$20+\$30)
00343 -00344	CALL INFINITIANICALUZOUSSU
00345	C GO TO 30 IF A TAPE ERROR IS FOUND.
00346	NU ORDS-LNZ/3 IM AX = N WO R CS-2
00347 -00349	1// #X = N #0 X L S = 2
00349	DO 10 I=1+IMAX+3
00350	
00351 -00352	ITTIFEFECTIVITY
00353	10 IY(I+2)=FL0(12,24,IX(J+1))
-00354 00355	20 RETURN 2
	30 RETURN 5
00357	END
0035 8 00359	afor, SIA SUBJ
-00360	C
00351	SUBROUTINE SKIPF(NFILES,MB)
00353	C THIS IS A SUBROUTINE TO SKIP TAPE DATA FILES THAT ARE NOT OF
-00364	C-CURRENT-INTEREST.
00365 	C DC 60 I=1+NFILES
00367	DO 20 J=1+MB
-00369-	
00359	40 FORMAT (///* NO EOF FOUND BY IOTPSP. GUIT.*)
00371	CALL EXIT
- 00372-	RETURN
00373 00374	
00375	C
00377	BXQT 1 250
	1 250
	i
00379	-1.
00380	400. .32 39.1 8.80
	9FIN

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