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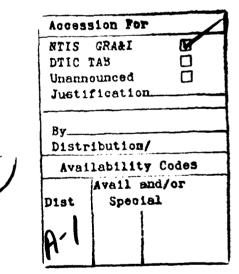
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An operational study has been made of the algorithm Station in their short base line underwater position locati have been uncovered. The issues studied include isospe profile and water layer thickness, approximate vs. exact methodology. It is shown that the practice of constant speed extrapo considerable mischief. The best remedy is to measure sp	on systems. Some important sources of systema ed vs. isogradient ray tracing, effect of the depth array tilt corrections, and ray tracing initialization plation of depth-velocity information can cause	tic error velocity
systematic errors are periodic functions of the azimuth d amplitudes of these functions are greater for the more set that reduces these errors by at least an order of magnitud	rection of the sound ray from the receiver array.	The
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1. INTRODUCTION

The purpose of this report is to discover and quantify the systematic errors in the algorithms employed by NUWES in their short base line underwater position location system. Systematic error can have a number of sources and previous works [6,7] have treated other issues. The present work deals with the algorithms used in sound ray tracing. There are a number of aspects to be treated, and some background is necessary in order to explain them.

Figure 1 contains a schematic diagram of a short base line hydrophonic array and of the signals it may receive. Sharply pulsed signals, or pings, are sent by the sound source vehicles (surface craft, submarine, torpedo) and they are received by the four transducers (called the *X*, *Y*, *Z* and C-phones) of the array. These four hydrophones form a right angled coordinate system with origins at the C-phone and arm lengths D (=30 feet) to each of the other three phones.

The ray paths from a source to the four receivers are synchronously timed with great precision (10^{-7} secs) and the differentials of arrival times are used to construct the direction of the source. But due to variability of the speed of sound at various water depths, the ray paths themselves are not straight lines. Also the paths may change from day to day as the water depth-velocity profile changes. Knowledge of the current profile allows one to perform a ray tracing computation. Recovery of the three-dimensional position of the sound source is accomplished by reconstructing the ray path and following it for the given amount of time. Each source vehicle has a phase coded "ping" so that its signals can be discriminated from those of other vehicles.

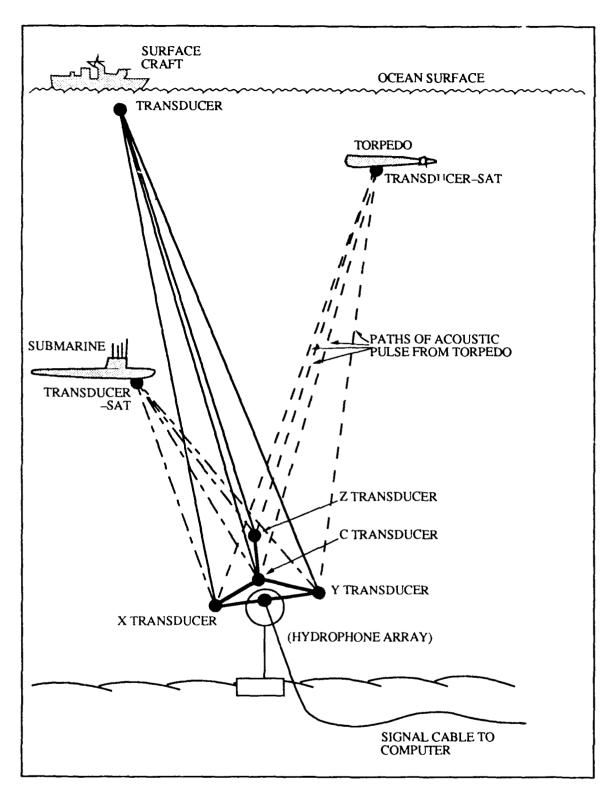


Figure 1. Short Baseline Array and Signal Sources

Each hydrophonic array is placed on the sea floor by lowering it over the side of a utility ship. Each has a self-leveling capability. When finally resting on the bottom, they are not perfectly level, but the X and Y arms have tilt meters that measure the angles that are made with the horizontal. An array surveying activity checks these angles and measures the rotation of the vertical. Thus the local coordinate system can be reconciled with the master range coordinate system.

The currently employed ray tracing methodology partitions the water column into a number of equal thickness layers and treats the speed of sound as constant in each layer. This leads to the use of isospeed ray tracing [1, 5]. Presently the layer thicknesses at Nanoose are 25 feet. In those instances for which the measured depth-velocity profile does not go as deep as the array, a constant speed extrapolation is used. In effect the thickness of the deepest layer is larger, perhaps 50 or 75 feet.

Now the issues can be detailed:

- i) How accurate is iso-speed raytracing using a 25 foot water layer increment?
- ii) What is the effect when it is necessary to use a thicker deepest layer?
- iii) How well are the ray tracing directions and transit times determined?
- iv) What effect does the various depth velocity profiles have upon the answers to i), ii) and iii)?

The treatment of these questions requires a valid sound ray construction methodology and a representative set of depth-velocity profiles. To satisfy the latter requirement we have selected twelve experimental days at the Nanoose range spanning the period May 1988 to June 1989. They are presented (in graphical form) in Appendix A. We note that the more interesting ones seem

to appear in the Spring. Two of the experimental days, 23-24 April, 1989 are consecutive. This allows us a glimpse into the question of day to day variability.

To treat the former requirement, we have developed an isogradient ray fitting algorithm. It fits a sound ray connecting two given points (in the horizontal-vertical plane) assuming direct path propagation. The depth velocity profile is partitioned into five foot equi thickness layers and within each layer the slope of sound speed vs. depth is constant. Thus the DV profile is represented as a continuous function consisting of a sequence of straight line segments. Within each layer the ray path is a circle arc because of Snell's law, [2,3,8]. The outputs of this algorithm are the angles that the ray makes with the horizontal at each of the endpoints, and the transit time of the ray from the initial point to the final one.

Section 2 of the report presents some theoretical material and formulas. The distinctions between isospeed and isogradient ray tracing are explained. Snell's law is introduced and supported.

Section 3 of the report deals with the accuracy of pure ray tracing in two dimensions, the horizontal-vertical plane. Computations are made for a number of initial angle-transit time pairs and for all twelve depth velocity profiles. Errors from this source are generally small but can be as large as a foot or more. The situation is more difficult when extrapolation of the water column is necessary. This occurs when the sounding does not extend as deep as the receiver. In such cases we cannot be definite about the nature of the errors, but several equally defensible methods lead to results that disagree by five or ten feet and even more. We state that there is a problem with

extrapolation. The soundings should be made at the deepest part of the range and to full depth. Failing that, a careful development of an extrapolation policy should be made.

Section 4 deals with the three dimensional problem of locating the position of the source based upon the transit times to each of the four hydrophones. Hence the question is one of finding the azimuth and initial elevation angles of the ray, and a matching transit time to stop the ray tracing algorithm. Also there are confounding sources of variability. The depth-velocity profile plays a role, as mentioned before. But the arrays themselves have directional properties that would interact with the algorithm even if they were fully level and aligned with the range. The fact that the arrays are tilted and rotated in a variety of ways has contributed to the puzzle of interpreting the mismatches in the array overlap areas. Taken altogether it is shown that systematic error from these sources can be as large as ten or twelve feet. Moreover errors of this magnitude are unnecessary. An alternative method is proposed which can reduce them by at least an order of magnitude.

The conclusions are summarized in Section 5. Section 6, an addendum, addresses an issue raised in the review process. Also a number of appendices are included. They hold the depth-velocity profiles, supporting mathematical details, details of the algorithms, and the source code for the FORTRAN programs.

A brief general statement of conclusions is as follows:

i) The error in iso-speed raytracing is an increasing function of horizontal range, but is seldom more than one foot.

- ii) The error due to constant speed extrapolation in the deepest layer can range up to 10 or more feet.
- iii) The error due to initializing the ray tracing is a periodic function of the azimuth direction and can be substantial for tilted arrays and at the greater horizontal ranges. The effect of the determination of azimuth is especially noticeable.

Greater details are presented as the various issues are developed.

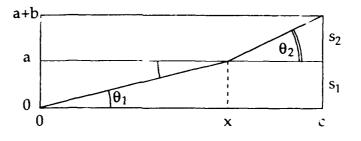
2. PERTINENT ITEMS FROM RAY TRACING

The sound source pings and sends out an isotropic wave front, which is a fixed phase point on the pressure cycle. A receiving transducer measures the time of arrival of the wave front. A ray is a path normal to the wave front and extending in space back to the source. Our goal is to trace a ray from the receiver for a fixed amount of time and thereby locate the source. To do this we must construct the azimuth and elevation angles of the ray at the receiver and then bend the ray back through the various speed layers until the measured transit time is consumed.

If the speed of sound in water were constant then the ray path would be a straight line. But it is not. Speed is a function of temperature, pressure and salinity. These variables interact in interesting ways for the water layer that affects our problem. Speed is not necessarily a monotone function of depth. Conditions change with time, and water sounding drops are made daily. They provide a depth-velocity profile which is assumed to be fixed for the entire day's exercises. Further, these values are assumed constant throughout each horizontal plane; i.e., the field is homogeneous.

Our immediate goal is to justify the use of a ray invariant in a horizontal-vertical plane and to establish the circular arc nature of ray paths in water layers for which the speed of sound is a straight line function of depth.

Let us begin with Snell's law. Consider two adjacent layers with speed s_1 in the lower layer and speed s_2 in the upper. The



ray enters the lower layer at elevation angle θ_1 and the upper layer with angle θ_2 . Given s_1 and s_2 let us find the relationship between θ_1 and θ_2 that will minimize the transit time from (0,0) to (c,a+b).

Proposition. For a ray to traverse from (0,0) to (c,a+b) in minimum time, we must have the relationship.

$$\frac{\cos(\theta_1)}{s_1} = \frac{\cos(\theta_2)}{s_2} \tag{2.1}$$

Proof. The transit time of a path from (0,0) to (c,a+b) that goes through (x,a) is given by

$$T(x) = \frac{1}{s_1}\sqrt{a^2 + x^2} + \frac{1}{s_2}\sqrt{b^2 + (c-x)^2}.$$

Further, it has derivative

$$T'(x) = \frac{1}{s_1} \frac{x}{\sqrt{a^2 + x^2}} - \frac{1}{s_2} \frac{c - x}{\sqrt{b^2 + (c - x)^2}}$$

The relationship (2.1) is a consequent of setting T'(x) = 0.

Now suppose the point c is not fixed but variable. It follows that a ray entering the lower layer at an angle θ_1 (with the horizontal) will seek the path

of minimum transit time and exit the upper layer at an angle of θ_2 ; the two angles are related by (2.1).

Next suppose that a number of layers are stacked vertically; within each layer the speed of sound is constant. The relationship (2.1) must hold for every successive pair of layers and hence the ratio

$$rv = \frac{\cos(\theta)}{s}$$
(2.2)

must be constant for the entire ray path. This value characterizes the ray path and is called the ray invariant. This is Snell's law.

Consider the vertical plane containing the source and the receiver. Call this is (h,z) plane with the depth z taken as positive downward. Our ray tracing problem is two dimensional in this plane. Given the depth velocity profile we need only the elevation angle at the receiver and the transit time to locate the source.

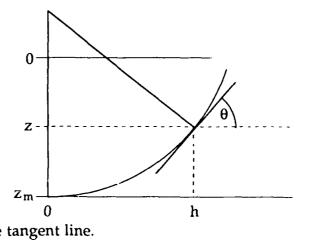
The depth-velocity profile can be approximated with a series of thin layers each having constant (internal) sound speed. Thus ray tracing can be enacted using such an approximation. This approach is called isospeed ray tracing [1].

A sharper approximation is available through the use of water layers whose sound speed structure can be represented with a linear function of depth

$$\upsilon(z) = \upsilon_0 + \upsilon_1 z \tag{2.3}$$

Proposition. If (2.3) holds then the ray path is a circle arc of radius $z_m + v_0/v_1$. If $v_1 > 0$ then the arc is a distance v_0/v_1 above z = 0.

Proof. Referring to the diagram, let (h,z) be an arbitrary point on $z_m = \frac{1}{0}$ the ray path and θ is the angle of the tangent line.



Because of Snell's law we must have

$$\frac{\cos(\theta)}{\upsilon_0 + \upsilon_1 z} = \frac{1}{\upsilon_0 + \upsilon_1 z_{\max}} = rv$$
(2.4)

and the ray path can be described parametrically in terms of

 $\{h(\theta), z(\theta)\}$ for $0 < \theta$

The radius of curvation can be found from the general formula

$$\mathbf{r} = \frac{\left\{ [\mathbf{h}'(\theta)]^2 + [\mathbf{z}'(\theta)]^2 \right\}^{3/2}}{\left| \mathbf{h}'(\theta)\mathbf{z}''(\theta) - \mathbf{h}''(\theta)\mathbf{z}'(\theta) \right|}$$

Using (2.3), $\frac{dh}{dz} = \cot(\theta)$, and implicit differentiation we find

 $z'(\theta) = -\sin(\theta)/\upsilon_1 rv$ $h'(\theta) = -\cos(\theta)/\upsilon_1 rv$

 $z''(\theta) = -\cos(\theta)/\upsilon_1 rv$ $h''(\theta) = \sin(\theta)/\upsilon_1 rv$

and hence

$$r = 1/v_1 r v = \frac{v_0}{v_1} + z_{max}$$

which is a constant, independent of θ . Hence the path is a circle arc; the radius is as specified; and the center of the circle is on a line above zero at a distance v_0/v_1 . This proof follows those found in [4,8].

If $v_1 < 0$, then the $[h(\theta), z(\theta)]$ curve is concave, still a circle arc, and the circle's center is still on the line $z = -v_0/v_1$. But now this line is below z=0. If $v_1 = 0$, the circle radius is infinite, the sound speed is constant, and the ray path is a straight line.

Now, the numerical construction of the ray path can also be accomplished by representing each layer's sound speed as a straight line segment (function of depth) and piecing together the consequent circle arcs. Since each layer has a constant speed gradient with depth, this is called isogradient ray tracing, [1].

The question of efficiency of the two methods, isospeed and isogradient ray tracing, is really a question of how well the depth-speed function in a layer can be represented by a constant on the one hand, or by a linear function on the other. For thick layers there may be oscillations that make the choice difficult. For thin layers it seems that the straight line fit should perform better.

The algorithm for isogradient ray tracing is presented in Appendix C. A corresponding algorithm for isospeed ray tracing can be extracted from it by making a number of deletions. Fortran source codes for each are shown in Appendix G. Inputs for these algorithms include the depth-velocity table; the layer depths; the depth of the receiver; the elevation (layer entrance) angle at the receiver; the ray transit time. The outputs are the horizontal and vertical end points of the ray, and the final elevation (exit layer) angle of the ray. This latter quantity is needed in the timing synchronization model, [7].

For isospeed ray tracing the pertinent formulae are

$$\Delta h = \Delta z \cot(\theta)$$
$$\Delta t = \Delta z / \upsilon \sin(\theta)$$

where θ is the angle that the ray enters the layer; Δz is the layer thickness; Δh is the horizontal distance transversed in the layer; Δt is the transit time through the layer. The next layer entrance angle is computed from the ray invariant equation (2.1).

For isogradient tracing the pertinent formulas are more complicated. For a ray that enters a layer at depth z_0 ; horizontal displacement h_0 ; and angle θ_0 we must first compute the coordinate (q_1 , q_0) of the center of the circle arc ($v_1 \neq 0$)

$$q_{2} = -v_{0}/v_{1}$$

$$q_{1} = h_{0} + (q_{2}-z_{0})\sin(\theta_{0})/\cos(\theta_{0})$$
(2.5)

and the radius of the arc

$$r = signum (q_2) (q_2 - z_0) / cos(\theta_0).$$
 (2.6)

The new horizontal displacement is

$$h_1 = q_1 - \operatorname{signum}(q_2) \operatorname{r} \operatorname{sin}(\theta_0) \tag{2.7}$$

and the increase in transit time is the line integral $\Delta t = \int \frac{ds}{v_0 + v_1 z(s)}$ along the circular path. This integral is most easily managed using $ds = \sqrt{(dz)^2 + (dh)^2} = d\theta / v_1 r v$ and $v_0 + v_1 z = \cos(\theta) / r v$ so that

$$\Delta t = \frac{1}{\nu_1} \int_{\theta_0}^{\theta_1} \frac{d\theta}{\cos(\theta)} = \frac{1}{\nu_1} \left\{ \ln\left[\frac{1+\sin(\theta_1)}{\cos(\theta_1)}\right] - \ln\left[\frac{1+\sin(\theta_0)}{\cos(\theta_0)}\right] \right\}$$

and the layer exit angle is computed from the ray invariant equation [2.1),

$$\cos(\theta_1) = \mathbf{r}\mathbf{v}\cdot\mathbf{v}(\mathbf{z}_1)$$

The layer exit angle is the entrance angle for the next layer.

The above equations form the heart of direct path ray tracing. The organizational questions that arise when developing a ray tracing algorithm are treated in Appendix C. Basically one proceeds upwards through the layers until the specified total transit time is consumed. An end correction is normally necessary because of the requirement to stop part way through a layer.

3. QUALITY OF ISOSPEED RAY TRACING

We are concerned with the quality of the currently employed isospeed ray tracing algorithm, which uses a uniform water layer thickness of 25 feet. The receiver depths range from about 1100 to 1350 feet. The elevation angle can range from 90° (directly overhead) to some rather small but positive values. Of course a variety of water columns (depth-velocity profiles) can be encountered. We have selected twelve (see Appendix A) spanning the period May, 1988 to July 1989.

Our first problem is to establish a standard ray to serve as a basis for comparison. To this end we are limited by the resolution of the water column data available. The information depicted graphically in Appendix A provides sound speed averages for every five feet. That is, at level *l* the corresponding velocity value represents

$$v_l = \frac{1}{5} \int_{\ell}^{\ell+5} v(z) dz$$
 (3.1)

and we have no information concerning the amount of variability that may exist within the layer. It is presumed small and is neglected. (A model for assessing such variability is presented in Appendix D, and this author is concerned about the issue for small entrance angles.)

The most expedient standard available is to employ the ray established by isogradient ray tracing utilizing the five foot layer thicknesses with the straight line segments as depicted (and exaggerated) in Figure 2. These rays are used to judge the rays formed by the isospeed method with 25 foot layer thickness. The corresponding depth-velocity table is formed by partitioning the $\{v_l\}$ into consecutive sets of size 5 and, within each set, average the five values.

With the above as background, the remainder of this section deals with numerical comparisons treating three issues: the computational noise generated by the processing of a large number of layers on the computer; the basic precision of the current isospeed ray tracing; the effects of extrapolation policies when the measured water column does not go sufficiently deep.

i. The adopted standard generally processes over 200 water layers (20 for each 100 feet separating source and receiver). The buildup of computational noise can be checked by using an artificial but linear depth velocity profile: An exact ray can be developed from the theory presented in the previous section and applied to a single layer of great thickness. Then the isogradient programs can attempt to match this ray by tracing through the usual five foot layers. This was done for the depth layer 150, 1200 ft. and intercepts 4250, 4800,

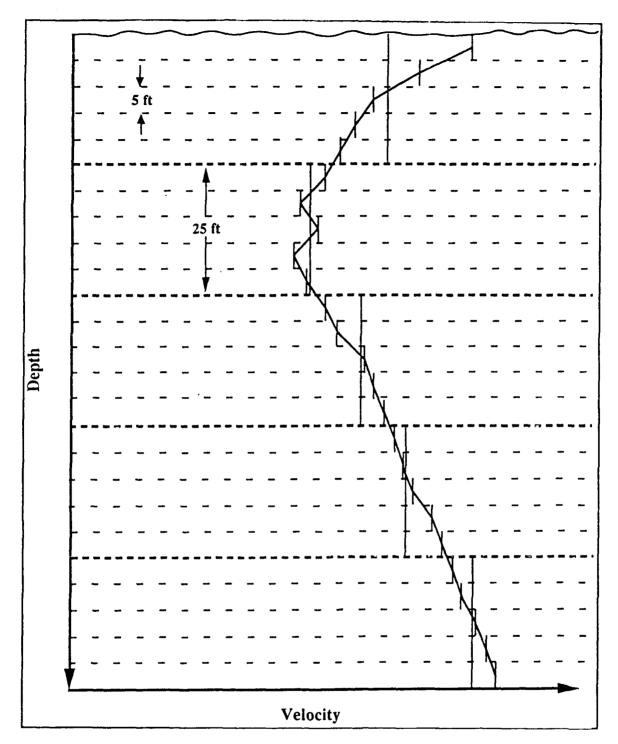


Figure 2. Schematic Diagram Comparing Isogradient and Isospeed Representations of Depth-Velocity Information

4850 ft/sec matched with slopes .1, .05, .01 ft/sec·ft respectively. The two methods produce virtually identical results. All computations are in double precision arithmetic.

ii. The ray fitting technique produces transit time, and entrance and exit angles of the ray connecting two points (source and receiver coordinates) in the horizontal-depth plane. When isogradient ray tracing is applied starting at the receiver, using the entrance angle and stopping when the transit term is consumed, then the coordinates of the source are reproduced to within some small preassigned value ϵ (we used $\epsilon = 10^{-6}$).

The accuracy of isospeed raytracing using 25 foot layer increments is compared against isogradient ray tracing using five foot-layer increments. A fixed set of entrance angle (θ_0 in radians) and transit time (t_0 in seconds) pairs have been selected. Generally they produce horizontal distances of 1000 to 6000 feet and depths of less than 50 to over 500 feet. The distance, d, between the two versions of sound source is compared for each input pair (θ_0 , t_0) and each of the twelve DV profiles. The values are recorded in Table 1 along with the source coordinates (h_c , z_c) for currently utilized isospeed methodology, and (h_s , z_s) for the standard (isogradient). An additional computation (h_g , z_g) was performed using isogradient ray tracing with 25 ft layer thicknesses. (The purpose was to provide an indication of the relative importance of layer thickness and the two types of ray tracing.) In all cases the receiver depth is 1300 feet.

The errors are computed using

$$d = \sqrt{(h_s - h_r)^2 + (z_s - z_r^2)}$$
(3.2)

with the subscript r replaced by c and g, respectively, in order to identify the current and 25 ft isogradient computations.

Examination of Table 1 shows that the errors, d, grow with range and the majority of the total error is in the vertical component. The effect of varying DV profiles is quite noticeable. The larger errors can exceed one foot. Generally isogradient ray tracing with 25-foot layer thicknesses is noticeably better than the current methods.

iii. Some measure of the effect of extrapolation techniques is given in Table 2. The same cases are developed as found in Table 1, but the sensor depth has increased to 1350 feet. The DV profiles seldom goes below 1300 feet so extrapolation of speed information is necessary. There are quite a few arrays deeper than 1300 feet and several (2, 3, 7, 13, 14 see Table B-1) considerably so. Moreover, the C-phones are even deeper, see Table B-2. (Fourteen of the C-phones are deeper than 1325 ft.) The methods of extrapolation are explained in Appendix A along with the visual effect of their use. In some instances the visual effect is appealing and in others it is not. See the insets in Appendix A. Thus the values for the standard (h_{s,Z_s}) are not always well supported. Even so, the effect is substantial and this is an important source of systematic error.

Table 2 is similar to Table 1 in the qualitative sense. The effect of the DV profile is greater and at the greater ranges the discrepancies can be quite large.

4. ERROR ASSESSMENT FOR THREE DIMENSIONAL METHODS

The ability to locate a sound source position from the transit times $(t_1, ..., t_4)$ needs to be assessed in three dimensions because of the directional properties of the array cubes. Our approach is to place the acoustic center at a

depth a_2 , and at the center of a right circular cylinder of radius h. The sound sources (k in number) are equally spaced on a section of the cylinder at depth z. Figure 3 shows a plan view. Azimuth is measured counter clockwise with zero at "3 o'clock."

The ray fitting methodology is used to construct true elevation angles (θ_1 , ..., θ_5) and true transit times (t_1 , ..., t_5) to each of the k sound sources from the four hydrophones and the acoustic center (θ_5 , t_5). We also need the azimuth from the acoustic center (origin of circle) to the sources. These latter are

$$\phi_j = 2\pi(j-1)/k \text{ for } j = 1, ..., k.$$
 (4.1)

Since the ray fitting methodology is two-dimensional, we must characterize the vertical planes connecting each sound source on the cylinder to each of the five points in the array cube. The technique for doing this is developed in Appendix B. One needs only the horizontal separations and the vertical positions of source and receiver. Thus five rays are fit to each sound source; one to each of the four phones and one to the acoustic center. Also five elevation angles are generated; but we retain only the last, the true elevation angle at the acoustic center.

During operations, the information collected consists of the set of transit times $(t_1, ..., t_4)$ from sound source to receiver array, (see Figure 1). For our purposes these four values are regarded as exact. Thus we can use the values produced in the ray fitting process. They must be converted to a ray tracing direction (azimuth angle, ϕ and elevation angle, θ) and a single transit time (t_{ac}) to stop the ray tracing. The currently employed procedure is described in [5]. It is convenient to present certain aspects , and this is done in Appendix E.

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°2°	18.47 250.85 514.42	22.28	536.89	201.90	350.72	503.51	505.42	246.66	540.54 335.43		ہ م		16.69	549.63 519.77	18.28	239.85	531.52	192.72	339.20	489.03	483.08	220.30	503.25	292.46		ъ 2		16.17	512.57	16.63	240.37	533.57	194.70	343.58	490.1U 240.52	494.53	229.63	523.77	311.49
б ц	1122.51 1015.82 930.22	2071.38	1792.54	2967.34	3018.82	3063.80	4002.02 3814.32	5001.64	4811.09 6008.74		٦ م		1122.23	03 000	2070.85	1918.23	1791.14	2965.79	3016.71	3061.33	3811 1B	4998 77	4806.86	6004.51		р Ч		1121.66 1016 70	931.30 931.30	2069.90	1920.94	1794.65	2969.54	3021.73	4006.23	3818.49	5005.79	4816.48	6014.30
ъ S	1122.47 1015.81 930.21	2071.31	1792.52	2967.31	3018.79	3063.77	4002.70 3814.28	5001.59	4811.04 6008.68		hs L		1122.17	1015.23 020 AB	2070.75	1918.20	1791.10	2965.74	3016.65	3061.26	4000.45 3811 10	4998.67	4806.75	6004.40		ъ S	10 10 11	1121.65	931.30	2069.89	1920.93	1794.64	2969.53	3021.71	4006.21	3818.47	5005.77	4816.46	6014.26
06/22/88 h _c	1122.47 1015.79 930.19	2071.29	1792.48	2967.24	3018.72	3063.70	4002.09 3814.19	5001.47	4810.92 6008.53	08/03/88	ہ م		1122.16	010.19 020.45	2070 73	1918.13	1791.04	2965.63	3016.54	3061.16	4000.30 3810 06	4998.49	4806.55	6004.17	01/13/89	ے ب		1121.61	931.26	2069.82	1920.86	1794.57	2969.42	3021.60	4006.06	3818.32	5005.57	4816.26	6014.01
бр	021	057	020	.123	152	187	281	355	475 560		δp		070	041 745	173	115	135	.245	.298	369	404	669	974	1.238		d d		210 215	014	040	036	.050	C60.	8	147	206	.230	343	407
qc	022 021 021	029	065	.126	146	1/1.	274	.375	.457 .583		ရိ	1	027	450. 620	061	260	E	.207	244	590	200.	2005	725	.789		d _c		059	850	155	8	18 0	339	397	.400 646	744	1.034	1.249	1.533
g ^z	16.05 ∠48.64 512 61	17.67	533.70	197.27	346.11	49/.00	496.08	237.31	525.07 321.84		29		17.70	511 70	21.10	242.49	534.03	198.39	345.44	495.57	402 00	236.06	519.11	314.93		βz		15./9 040 CC	512.61	18.64	241,60	534.21	198.19	347.58	490.70 247.88	498.81	241.25	529.80	328.39
5 ^S	16.04 248.63 512.60	17.62	533.64	197.16	345.97	490.89	495.80	236.96	524.60 321.28		2 _S	1	17.67	513 66	20.98	242.39	533.91	198.16	345.16	495.21	00.442 402.43	235.38	518.16	313.72		zs		15.78	512 59	18.61	241.56	534.16	198.10	347.48	490.04 247.74	498.61	241.02	529.46	327.99
2 ^C	16.02 248.61 512.58	17 58	533.58	197.04	345.83	490./2	495.53	236.59	524.15 320.71		Zc		17.65	61.002	20.92	242.31	533.80	197.97	344.92	494.93	402 00	234.81	517 44	312.94		2 _C		15.74	512.55	18.48	241.42	534.00	197.78	347.10	247.12	497.89	240.02	528.23	326.48
б _ц	1122.71 1017.35 931.35	2071.84	1794.64	2971 56	3023.32	306/.52	4000.75 3818.91	5008.96	4816.55 6017.51		рg		1122.83	010.07	2071.91	1918.85	1791.17	2966.76	3017.41	3061.49	3811 30	5000.35	4807.10	6005.85		b L		1123.99	931 72	2074.16	1922.99	1795.37	2972.64	3024.62	4010.52	3820.44	5011.13	4818.52	6020.20
s U	1122.69 1017.33 931.33	2071.81	1794.61	2971.51	3023.27	3067.47	4000.09 3818.85	5008 88	4816.47 6017.42		ь s		1122.76	920 50	2071.79	1918.80	1791.11	2966.68	3017.32	3061.39	3011 00	5000 21	4806.93	6005.62		ہ م		1123.98	931.71	2074.14	1922.98	1795.35	2972.60	5024.59	4010.48	3820.39	5011.08	4818.46	6020.13
р ц	1122.68 1017.32 931.32	2071.77	1794.59	2971.47	3023.22	300/ 42	4000.03 3818.79	5008.80	4816.40 6017.31		ິ		1122.76	30.0101	2071.78	1918.75	1791.07	2956.60	3017.24	3061.32	381118	5000 08	4806.81	6005 51		ں م		1123.93	931.67	2074.06	1922.89	1795.28	2972.48	3024.47	4010.31	3820.23	5010.85	4818.25	6019.88
Ş	30 S	84	2 4	.65	<u>5</u>	8,8	88	1.05	1.25 1.25		æ	1	35	ς Υ	3 6	45	4	.65	.65	65 7	e e	8.6	88	1.25		ء.		£.	3 K	5	45	4	.65	9 9	8 8	8	1.05	8	1.25
2/88 0 ₀ der	48.70 45.84 40.11	31.51	22.92	20.05	17.19	20.41	11.46	11.46	8.59 8.59	/88	9 ₀	deg	48.70	40.04 40.11	31.51	28.65	22.92	20.05	17.19	14.32	14.3C	11 46	8.59		/88	θo	6ep	48.70	40.04	31.51	28.65	22.92	20.05	17.19	14.32	11.46	11.46	8.59	8.59
05/12/88 θ ₀	8 8 8 8 8 8 8 8 8 8 8 8 8 8 8 8 8 8 8	50.5	3 4	.35	8	ç, y	Q 0	20	.15 .15	07/21/88	J	rad	8 8 8	8 5	22	ŝ	40	35	<u>9</u>	35	ŝ	2 2	15	.15	10/27/88	Ţ	rad	69 6	82	55	20	40	35	8	, S	2	.20	15	.15

TABLE 1 (Continued)

စီ	5	0.0	010	.036	.035	.053	.088	.047	,114	159	190	.265	411	488		őp	,	600	80	003	.018	010	600	018	2.5	520	5.6	044	045	.062		dg d		025	S S	063	064	.073	.136	.182	.191 266	294	388	524	200
qc	630	0 <u>5</u> 0	058	.167	.167	206	.368	397	.534	.684	.841	1.073	1.525	1./62		မို		043	.042	.043	.113	811.	139	256	2.0		201	748	1.031	1.221		ę		0. 7	053	94	147	171.	.318	.357	4/0/4/	738	.921	1.217	-40C
2 ²	10.03	250.89	514.27	19.20	242.32	535.59	196.24	345.06	498.67	241.33	497.79	229.88	529.48	312.42		2 ²		18.62	250.59	513.82	17.94	240.18	533.25	190.68	11.800	491.02 220.72	ARE RE	213.36	512.03	288.62		6z		18.70	514.02	19.77	240.95	533.45	192.54	340.28	492.21 233.66	487.43	217.63	511.24	11.282
zs	10.01	250.90	514.28	19.23	242.35	535.63	196.32	345.10	498.78	241.48	497.98	230.14	529.88	312.90		zs		18.62	250.59	513.82	17.93	240.17	533.24	190.67	008.00	02.124	486.82	213.31	511.98	288.56		2 _S		18.68	514.00	19.71	240.89	533.38	192.41	340.11	492.02	487.15	217.25	510.73	10.182
zc		250.86	514.24	19.09	242.21	535.45	195.98	344.72	498.26	240.82	497.16	229.09	528.38	311.16		ň		18.59	250.56	513.78	17.83	240.07	533.11	190.43	000000	930.24	486.24	212.58	510.97	287.36		20		18.65	513.96	19.59	240.76	533.23	192.11	339.77	491.57 132 BA	486.43	216.35	509.53	21.082
ۍ ۲	1110.05	1014 23	929.41	2065.08	1916.27	1791.14	2962.34	3014.71	3060.96	3996.59	3810.79	4993.81	4807.22	6000.47		եղ		1118.53	1013.06	928.56	2064.12	1914.04	1789.54	2958.77	3011.00	30.05 00	3807 15	4987.92	4802.75	5994.02		р Ч		1119.95	928.42	2066.69	1914.44	1789.16	2959.51	3011.71	305/.605	3806.57	4988.82	4801.55 5004 17	2884.11
ş	1110 06	1014 24	929.41	2065.10	1916.28	1791.17	2962.37	3014.71	3060.99	3996.63	3810.83	4993.87	4807.30	6000.56		ĥs		1118.52	1013.06	928.56	2064.11	1914.03	1789.53	2958.76 2011 60		3038.00	3807 1 A	4987.91	4802.74	5994.01		å		1119.93	928.41	2066.66	1914.40	1789.13	2959.46	3011.64	305/205	3806.51	4988.73	4801.45 5004 06	00.4880
03/22/89 h _c	110.01	1014 20	929.38	2065.01	1916.20	1791.08	2962.24	3014.60	3060.85	3996.45	3810.64	4993.63	4807.02	6000.25	04/27/89	Р С		1118.49	1013.02	928.53	2064.05	1913.97	1/89.48	2958.67	0011.00 005700	02 1005	2807.01	4987.74	4802.55	5993.78	06/06/89	ې ب		1119.89	928.37	2066.59	1914.33	1789.06	2959.35	3011.54	305/.4/	3806.34	4988.52	4801.23 5003 80	00.0880
бр	010	910	014	.051	.057	.059	.105	.132	105	509	201	.313	322	.483		đ	•	047	.021	021	.118	BSO.	190	128		3 6	040	354	470	.553		đ		660	020	860	079	088	.170	8 9 9	242	375	491	.630 756	8
မိ	760	025	032	.073	.067	.082	.162	.182	.290	.283	.415	.445	.651	/46		ဗိ		80.0	020	021	032	100	50			50		365	514	.613		ဗိ		8	80	017	003	007	005	010	015 afo	023	.015	040 040	R#0.
gz	16.43	248 70	512.35	14.77	237.94	531.28	187.69	336.12	488.35	226.76	482.24	207.98	505.16	280.89		6z		18.61	250.57	513.78	17.69	239.96	533.08	190.14		421.04 220.62	445 03	211.70	510.69	285.97		2g		18.60	513.72	18.08	239.62	532.10	189.17	336.72	488.1U	481.10	207.72	501.40	×/۵.5
zs S	16.42	248 69	512.34	14.72	237.89	531.22	187.59	336.00	488.24	226.55	482.04	207.67	504.B4	280.42		2s		18.60	250.56	513.77	17.61	239.91	533.02	190.02	330.62	10.064 0.00 40	485.65	211.35	510.22	285.43		2 _S		18.58	513.70	18.01	239.55	532.01	189.02	336.53	48/.86 226.02	480.73	207.24	500.78 277 26	67.112
zc	16.40	248.67	512.32	14.66	237.83	531.15	187.45	335.82	487.97	226.28	481.64	207.24	504.20	279.68		2		18.59	250.54	513.75	17.59	239.86	532.95	189.90	00.000	490.09 220 1 B	AR5 37	211.00	509.72	284.82		2 _C		18.58	513.70	18.01	239.54	532.01	169.01	336.52	68/.89 208.02	480.71	207.23	500.74	17.117
bu	1110 67	1014.62	929.97	2066.09	1916.99	1792.17	2963.27	3016.25	3062.59	3998.00	3812.67	4995.43	4809.68	6002.91		و م		1118.33	1012.92	928.50	2063.75	1913.78	1/89.43	2958.40	2011.64	00./COC	18.1660	4987.29	4802.42	5993.21		و م		1118.69	927.97	2064.41	1913.14	1788.33	2957.47	3010.00	3056.07	3804.60	4985.42	4799.18 5000.45	04.0220
hs	1110 60	1014 61	929.96	2066.07	1916.96	1792.15	2963.24	3016.21	3062.58	3997.94	3812.63	4995.36	4809.64	6002.82		т _s		1118.29	1012.91	928.48	2063.66	1913./5	1789.41	2958.35	2011 20	8/./coc	10.1000	4987.21	4802.34	5993.12		۲ م		1118.66	927.95	2064.35	1913.10	1788.29	2957 41	3009.94	3056.01	3804.52	4985.31	4799.07 5000 31	10,0880
ې ب	1110 50	1014 59	929.94	2066.02	1916.93	1792.11	2963.18	3016.16	3062.48	3997.87	3812.53	4995.26	4809.51	6002.68		ے ب		1118.30	1012.89	928.47	2063.68	1913.73	1789.38	2958.31	3011.13 2057.73	5/./COS	1806 72	4987.13	4802.24	5993.00		പ്		1118.66	927.95	2064.37	1913.10	1788.29	2957.41	3009.93	3056.00	3804.51	4985.31	4799.06	00.0000
e ,	20	38	5	ß	.45	4	.65	. 65	.65	8	8	1.05	8	1.25		ę.		35	8	Si i	S :	4. 5	9	ŝ	n u	e a	9 g	3 2	8	1.25		\$;	ŝ	3 8	9	4	4	.65	ເ ເ	S a	38	1.05	8 %	62.1
6/89 90	deg 46.70	45.84	40.11	31.51	28.65	22.92	20.05	17.19	14.32	14.32	11.46	11.46	8.59	8.59	(89	9 ₀	-	48 70	45.84	40.11	31.51	28.65	22.92	20.05	D	14.32	11 46	11 46	8.59	8.59	68/	9 ⁰	<u>geb</u>	48.70	40.04 40.11	31 51	28.65	22.92	20.05	17.19	14.32	11.46	11.46	8.59 e 50	0.04
03/08/89 9 ₀	rad	3	22		ß							20	5.	GL.	04/26/89	Ţ		.85										2	15	.15	05/10/89		rad	8. 28.	80	55	ß	40	.35	S, S	ų r	28	50	<u>5</u>	<u>.</u>
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Comparison
TABLE 2.

ęb	035 096 097 099 097 097 097 097 097 097 097 097		δp	0330 0320 0320 0320 0320 0320 0320 0320	δp	029 026 027 028 027 028 027 028 028 028 028 028 028 028 028 028 028
qc	189 179 500 505 505 505 1087	1.565 2.040 2.439 2.439 3.151 4.054 5.209	qc	089 081 082 0826 082 082 082 083 0894 070 11375 2183 2183 2183	qc	130 123 123 124 124 124 124 124 124 1077 1077 1077 1077 1077 1077 1077 107
6 _z	68.36 68.36 564.36 72.02 294.15 587.06 251.90 251.90	554.25 554.25 556.64 597.11 592.54 387.62	βz	66,63 66,63 67,90 67,90 67,90 289,95 582,08 542,00 286,73 536,73 536,02 546,020 546,020 546,020 546,020 546,020 546,020 546,020,020 546,0200000000000000000000000000000000000	6 _z	65.93 298.52 562.53 67.15 67.15 582.83 67.15 583.85 583.85 5847.46 547.46 546.77 546.77 546.77 546.77 546.77 546.83 577.02
S ^Z	68.34 300.71 564.33 71.94 71.94 294.07 586.97 251.70 251.70	553.97 553.97 556.20 596.52 591.85 386.64	sz	66,65 562.79 562.79 67.98 67.98 582.17 282.17 282.17 282.17 285.03 541.13 541.13 541.13 545 536.08 536.08 536.08 536.08 537.04 537.04	2 ^S	65.95 298.54 562.55 67.21 67.21 293.93 547.68 547.68 291.99 282.10 282.10 282.10 577.63 577.63
zc	68.47 68.47 564.47 72.36 294.51 587.50 252.72 252.72	555.49 555.49 558.59 599.61 595.86 391.77	2c	66.71 299.22 562.85 68.17 68.17 290.23 582.43 391.07 243.72 243.72 243.72 537.49 537.49 537.49 537.49 537.49 537.49 549.20	r K	66.04 298.63 562.64 67.51 290.84 548.13 395.34 548.75 548.75 548.75 548.75 548.75 580.36 580.36 580.36 580.36
бц	1122.41 1015.94 930.43 2071.25 1919.53 1792.94 2967.64	3064.48 3064.48 3815.17 5002.24 4812.12 6009.92	^ل م	1121.94 1015.43 929.85 2070.36 1918.57 1791.83 3017.62 3017.62 48999.66 48999.66 4809.80 6005.91	р д	1122.47 1016.95 931.58 931.58 1921.41 1795.15 2270.34 3028.21 3068.21 3068.21 3068.21 3019.77 5007.05 4817.89 6015.80
ہ s	1122.39 1015.92 930.41 2071.19 1792.90 2967.57	3064.41 3064.41 3815.08 5002.11 4812.02 6009.75	лs S	1121.96 1015.45 929.87 92070.40 1918.61 1791.87 1791.87 1791.87 1791.87 1791.87 1791.87 1791.87 1791.68 3017.68 3007.99 3007.68 3007.99 3007.99 3007.99 3007.99 3007.99 3007.55 3007.99 3007.90 3007.90 3007.90 3007.90 3007.90 3007.90 3007.90 3007.90 3007.90 3007.90 3007.90 3007.90 3007.90 3007.90 3007.90 3007.90 3007.00 30000000000	р°	1122.49 1016.97 931.60 931.60 1795.19 1795.19 2970.40 3028.27 3068.27 3068.27 3068.27 3019.85 5007.30 3819.85 5007.15
06/22/88 h _c	1122.53 1016.04 930.52 2071.45 1793.12 1793.12 2967.95	5002.76 3064.80 3815.58 5002.76 4812.66 6010.63	08/03/88 h _c	1122.02 1015.51 929.93 2070.52 1918.72 1918.72 2966.49 3017.86 3017.86 3012.92 5000.07 5000.07 6006.41	01/13/89 ħc	1122.59 1017.05 931.68 931.68 1921.61 1795.34 1795.34 3028.54 3068.54 4007.66 3820.20 5807.61 4818.45 6016.51
đ	061 057 058 058 161 183 345	497 647 768 768 1.003 1.272	đ	6.993 6.994 6.994 7.011 7.172 7.717 7.717 7.717 7.717 7.717 7.717 7.522 8.637 9.637 9.637 9.637	đ	063 062 062 062 062 168 180 1123 833 1123 833 1123 833 1123 833
ç	153 153 153 433 493 934	1.739 1.739 2.054 2.692 3.432 4.261	qc	7 143 7 143 7 176 7 646 7 642 7 692 8 679 9 149 9 149 9 149 9 149 9 149 9 149 9 149 9 149 9 149 9 149 14 247 11 280 14 245 12 268	ç	180 168 171 475 469 541 171 1031 1459 1450 2259 2259 2254 2259 2254 4668
6z	65.86 65.86 562.56 67.96 291.29 583.91 247.66	547.82 547.34 547.34 288.63 577.55 374.00	2g	67.54 299.99 563.54 70.20 70.20 292.03 247.56 394.98 545.48 545.48 563.31 14 14 363.14 1	5 ^z	65.55 562.42 562.42 562.42 569.125 581.25 548.08 548.08 548.08 548.08 548.33 558.32 578.32 578.32 578.32 578.32
2 _S	65.82 298.50 562.52 67.82 291.14 291.14 283.74 247.33	547.33 547.33 546.59 287.65 576.29 372.43	2 ^S	60.54 556.53 556.53 63.05 63.05 537.54 537.54 537.54 537.54 535.37 535.64 535.37 535.64	2 S	65.59 298.37 562.47 562.47 562.47 584.08 397.11 397.11 548.65 579.69 579.69 579.69 579.69
2c	65.93 298.60 562.64 68.19 291.53 294.20 248.20 258.20 257.	548.62 548.62 548.60 548.60 290.28 376.64	zc	67.68 300.14 563.70 70.68 547.21 396.38 547.21 547.21 547.21 547.21 547.03 574.05 574.05	zc	65.71 298.49 68.70 68.70 68.70 59.4.58 59.4.58 59.4.58 59.38 550.02 550.02 550.02 81.50 38.1.82 550.02 88.28 38.28 550.02 88.28 550.02 88.28 550.02 88.28 550.02 88.28 550.02 550.02 56.28 58.58 58.58
б _ц	1123.27 1017.57 931.58 931.58 921.58 1795.12 2972.27 2972.27	5012.37 3068.28 3819.87 5010.07 4817.86 6018.83	бц	1122.31 1015.56 929.74 2071.04 1918.83 1791.59 3017.82 3017.82 3017.82 3017.82 3017.82 3017.59 606.83 4001.75 5000.34 6006.47	б _ц	1124.00 1017.90 931.84 1923.22 1795.60 1923.22 1795.60 2973.19 3024.79 3069.11 4010.94 3820.90 5011.71 4819.11 4819.11
۶ų	1123.23 1017.53 931.55 2072.79 1922.51 1795.05 2972.16	3068 16 3068 16 4009 46 3819.71 5009 86 4817.66 6018.57	٦s	1122 22 1015 45 929 61 2070 86 1918 63 1791 33 3017 42 3061 33 3811 63 3811 63 3811 63 4999 81 4807 49 6005 67	ъs	1124.04 1017.95 931.87 1923.31 1923.31 1795.68 2933.33 3024.92 3069.25 4011.12 3821.07 5011.96 4819.34 4819.34 6020.76
ں بر	1123.35 1017.63 931.64 931.64 2073.01 1922.71 1795.24 2972.48	3068 49 3068 49 4009 90 3820 13 5010 43 4818 21 6019 27	ب د	1122.47 1015.70 929.87 2071.33 1919.10 1791.85 3018.26 3018.26 302.32 3812.78 5001.05 5001.05 6007.41	້	1124 18 1018.06 931.98 931.98 1923.53 1795.88 1795.88 3025.88 3069.62 4011.61 3021.55 6021.55 6021.52
. \$	8 8 8 8 8 8 8	88 88 1.05 1.05 1.25	\$	33 35 55 55 55 55 55 55 55 55 55 55 55 5	م	33 25 25 25 25 25 25 25 25 25 25 25 25 25
2 6 C		5 0 11 46 0 11 10 11 10 11 10 11 10 11 10 11 1	90 90	0 0 <td>2.0</td> <td>5 48 70 45 84 5 3 10 45 84 5 4 17 10 11 4 12 12 12 12 12 12 12 12 12 12</td>	2.0	5 48 70 45 84 5 3 10 45 84 5 4 17 10 11 4 12 12 12 12 12 12 12 12 12 12
05	73 200 200 200 200 200 200 200 200 200 20	<u>, </u>	60	20 20 20	10/2 rad	88233348883322 <u>45</u>

TABLE 2 (Continued)

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ືບ	.072	0690	193	270	417	.531	636	.795	1.009	1021	2.084		ືອ									-	-	-		3.245			ອີ	i	178	182	452	504	.580	1.082	1.283	2.028	2 443	3.138	4 160	120 0
đ	.261 246	250	.696	754	1.504	1.781	2 194	2.795	3.444	4 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1	7.565		ບັ			50	112	117	129	254	.299	.362	474	.568	.728	.987	12171		မိ	5.7	550	044	159	.124	.143	274	62E	207	611	.796	1.019	5/21
5 z	68.57 300.62	564.09	69.08	585.54	245.68	395.21	548.93	290.53	548.23	2/8/2 580.03	361.67		2 ₉		500 40	562 76	67.80 67.80	290.33	583.89	241.32	390.30	543.81	282.41	540.33	266.05	568.01	343.98		Ē	() () ()	100.60	563.88	68.74	290.80	583 61	242.47	390.71	282.88	538.90	268.11	563.74	343 55
2 _S	68.61 300.67	564.14	69.25	585.77	246.07	395.71	549.54	291.29	549.21	280.00	363.72		z _S		10.50	10.000	707.07	290.66	584.29	242.07	391.21	544.96	283.86	542.16	268.36	571.20	347.82		sz		00 00 JUN 72	564.02	69.14	291.24	584.15	243.48	391.93	285.84	541.28	271.17	567.83	345.45
2 ^C	68.79 200 84	564.34	69.84 202 74	586.48	247.48	397.40	551.66	293.99	552.58	284.23 587.36	371.15		7C	() ()	50.50	563 00	000.000 68.01	20076	584.41	242.31	391.50	545.31	284.32	542.72	269.07	572.18	348.99		7 ^C		50.00 76	564.05	69.26	291.35	584.28	243.74	392.25	286.32	541.88	271.94	568.84	249.75
ی م	1119.58	929.76	2066.06	1791.77	2962.86	3015.79	3062.05	3997.32	3812.11	4994.72	6001.91		б _ц	F F 0 ***	// 8111	010.43	21.626	1914 84	1790.63	2959.93	3013.11	3059.89	3993.61	3809.42	4990.01	4805.69	5996.91		Б ц		1013 54	928.80	2065.65	1914.96	1789.88	2960.28	3012.80	3993.76	3808.14	4990.19	4803.51	99.2992
лs s	1119 64 1014 46	929.80	2066.16	1791.89	2963 71	3015.97	3062.22	3997.54	3812.34 4005 01	10.0844	6002.29		h _s		1118.8/		02.636	1915.02	1790.80	2960.21	3013.40	3060.20	3994.02	3809 84	4990.54	4806.28	5997.63		٩		1013.67	928.91	2065,86	1915.20	1790.11	2960.66	3013.19	3994,29	3808.67	4990.88	4804 24	18.0550
03/22/89 ħ _c	111983	929 96	2066.53	1792.17	2963 53	3016.51	3062.79	3998.26	3813.08	40.00444 80.0444	6003.69	04/27/89	പ്		15.8111	10.0101	02.525 064 800	1915 08	1790.85	2960.30	3013.50	3060.30	3994.14	3809.96	4990.70	4806 45	5 87/665	06/06/89	ᆈ		1013 70	928.54	2065.96	1915.26	1790.16	2960.76	3013.29	3994.42	3808.80	4991.05	4804 41	01.7886
đ	085	082	227	252	490	595	214	535	611.1		2.384		dg dg		10.940	10.001	106.01	11 088	11.203	11.305	11.497	11.747	11.737	12.214	12.167	13.010	311.51		ð		510	023	050	090.	.074	53	8	247	305	380	532	624
о q	100	050.	083	105	.184	204	552	332	165	670	820		д ^с		10.01		225 11	667 11	11.615	12.089	2.445	12.930	13.251	14.100	14.549	16.287	17.052		ç	000	250	026	084	.076	.083	165	191	505	363	.473	608	765
6 2	66.21 208.58	562.33	65.16 202.27	581.92	238.78	387 69	540.92	279.25	536.57	251.98	338.11		5g	5												567 85			6z		300 54	563.77	67.95	290.10	583.09	240.55	388 93	280.42	536.35	262.75	560.42	337.18
54	66.27 208 61	562.39	65.35 766 47	582 16	239 24	388.25	541 61	280 16	537 66	253 40	340.46		NS S		54 / C		01.7CC	00.00	572 65	229.72	378.45	531.98	270.06	528.01	252.98	554.87	329.49		2 _S	1	100 F.5	563.79	67.99	290.16	583.15	240.67	389.08	280.65	536.64	263.12	560.95	337.80
,, ,	66.29 200 66	562.42	65 42 700 FF	582.25	239 41	388 45	541 86	280 48	538.04		341 26		2c		20 20		10.000 6.000	200.55	584 25	241 78	390.87	544 89	283.27	542 06	267 47	571.09	346.45		2 ^C		100 57	563.81	68.06	290.22	583.23	240.82	389.27	280.95	537.00	263.59	561.55	22 855
ഗ്	1120.32	930 47	2067 41	1793.14	2964 65	3017.78	3064 46	3999.95	3815.12	19/97 47	6006.14		<u>ق</u>		1118 53		958-10	1914.65	1790.58	2959.65	3012.84	3059.83	3993.23	3809 35	4989.54	4805.61	5996.34		бц		1012101	928.58	2064 40	1914.12	1789.52	2958.88	3011.78	3992.04	3807.23	4988 02	4802.61	2944 00
s S	1120 39	930.52	2067.53	1793.23	2964 82	2017.97	3064 65	4000.20	3815 36	4940 19	6006 58		s S		40.8111		26 976	26 F161	1790 24	2959.24	3012.33	3059.23	3992 63	3808.60	4588 80	4804 65	5995 34		ŝ		1013 12	929.60	2064 42	1914.15	1789.55	2958 92	3011.83	3992.11	3807 30	4988.10	4802.70	21 4665
ບ ເ	1:20 41	930 54	2067.57	1910 07	2964 89	3018 03	3064 71	4000.29	3815 44	494831	6006 72		ີ ເ		6/ 8111		יירא יים יירא יים	1914.83	1790.75	2959.94	3013.14	3060 15	3993.64	3809.77	4990 08	4806 20			о ц		101214	928.61	2064.47	1914.18	1789 58	2958.98	3011.89	3992.18	3807.37	4388.21	4802 81	07 1555
٩.	35	5.5	8 4	9	.65	63	3 2	<u>8</u>	85	68	3 KI		۹.	ľ	ີ ເ	2 4	9 S		9	9 9 2 9	65	65	<u>8</u> 5	ଥି	1.05	8	1 25		ጉ	Ĺ	មួទ	52	ያ	45	9	65	50	58	8	1.05	8	5
03/08/89 8°3	1 deg 48 70											04/26/89	6 ₃	6ap 1														05/10/89	φ										(211		8.59	B)
2/60	rad 85	32	47 U 47 U	7 9	35	05.	1,5 25	ci N i	8, 6 2		י י י	04/2		rad	n (0 () () (, u , u	1 G 1 G		35	8	25	25	.20	Ĵ,	5 5	4. •	05/	3		0 C	22	55	50	40	35 95		25	20	20	5	0

Finally the values (θ, t_{ac}) together with the array depth and DV information are fed into the ISOSPEED ray tracing program. The output consists of the (h,z) values. Upon combining these with the azimuth ϕ_c , the estimated position of the sound source can be computed.

A number of error calculations can be made. First the error in estimating transit time

$$tim_{er} = t_{ac} - t_5 \tag{4.2}$$

affects the rule for stopping the ray tracing algorithm. An error of 10^{-4} seconds translates to about half a foot in horizontal range, h, (i.e., $\upsilon \approx 4800$ ft./sec.). Next, the horizontal error is measured directly

$$\mathbf{h}_{\rm er} = \mathbf{h}_{\rm c} - \mathbf{h}_{\rm 1} \tag{4.3}$$

The values timer and her should be strongly correlated.

In a similar fashion, the error in the elevation angle is correlated with the error in the vertical component

$$\theta_{\rm er} = \theta_{\rm c} - \theta_{\rm o} \tag{4.4}$$

$$Z_{\rm er} = Z_{\rm c} - Z_{\rm I}$$

but, because of ray bending, the relationship is non-linear. The value z_{er} is measured directly in feet. The value θ_{er} is more difficult to interpret. It is a function of the water layers involved. See [7].

Finally we have the error in azimuth

$$\phi_{\rm er} = \phi_{\rm c} - \phi_{\rm o}. \tag{4.5}$$

These values are in radians and, when multiplied by h_1 , measure how far off the mark (see Figure 3) in feet, along the cylinder perimeter. It is a rather incipient fact that our errors are periodic functions of the true azimuth. This is illustrated in Figure 4. We suspect that this is at the root some earlier attempts to treat possible causes of systematic error, [6,7].

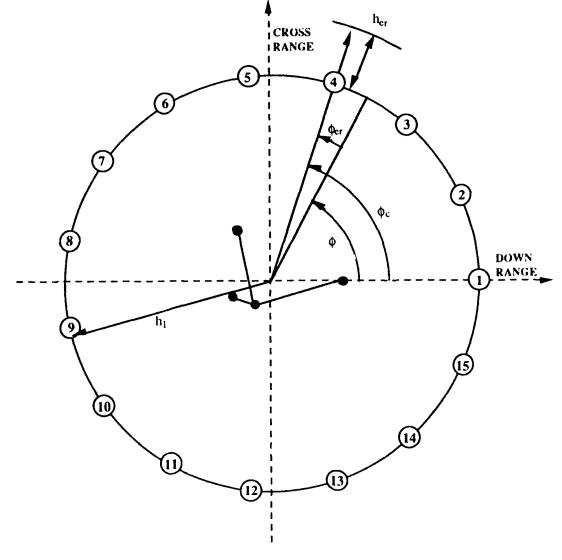


Figure 3. Cylinder Cross Section Illustrating Sound Source and Receiver Positions

Many tabulations of these errors have been made. Four sets, each with 15 points around the cylinder and four radii ($h_1 = 2000, 3000, 40000, 5000$ ft.),

Horizontal and Vertical Errors h_{er} ſ Feet er 3000 4000 5000 ŏ 2π Radians $2.5 \int_{-3}^{\times 10^{-3}}$ Azimuth and Elevation Angle Errors 3000 ¢ er 2 4000 5000 1.5 Radians θ er 0.5 0 -0.5L 2π Radians 3^{x 10-4} Errors in Transit Time 2 tim _{er} Seconds 0

Figure 4. Periodic Nature of Errors as a Function of Azimuth Case: DV of 3/22/89 and Array No. 1. Ref. Table 3.

3000 4000 -1 5000 -2L 0 2π Radians

have been selected for display and they appear in Table 3. The basis of selection was to choose two of the more interesting water columns (10/27/88) and 3/22/89, see Appendix A) each matched with two of the more severely tilted arrays: Number 1 and 56, see Table B-1. The tilts for array number 1 have opposite signs while those for array 56 are of the same sign.

Let us make a sample calculation in order to aid in the interpretation of these errors. The total error offset is essentially

$$d = \sqrt{z_{er}^2 + h_{er}^2 + (h_1 \phi_{er})^2}.$$

Thus for the case of Array No. 1 on 10/27/88 at $h_1 = 5000$ ft. and azimuth 0.418879 radians (i.e., 24 deg. north of east) we have $\sqrt{(2.45)^2 + (.28)^2 + (5 \cdot 2.3525)^2} = 12.02$ ft. The dominant portion of the error is in the azimuth, which (measured along the arc) contributes 2.3525 ft. of error for every 1000 ft. of horizontal range. A scan of Table 3 shows generally that this condition persists, although the periodic nature of the errors can make the effect small in some directions.

It seems that the most severe errors are associated with the arrays having the larger tilts. We suspect that the causes lie in the use of the array center as the acoustic center and the assumption of constant sound speed for the entire array. The z-phone is about 30 ft higher than the others. Also, the tilt correction method is but approximate (see Appendix B.)

We are disinclined to attach great importance to the other source, the isospeed ray tracing. The former was considered in Section 3 and cannot

Table 3. Error Structure - Current Methodology

0 1831	Zer	-2.44	-2.01	-1.42	-1.22	-1.62	-2.37	-2.90	-2.70	-1.71	40	.52	.59	19	-1.35	-2.25	0	= 1.0477325	Zer	-2.77	-2.45	-1.96	-1.84	-2.31	-3.09	-3.58	-3.26	-2.11	61	.47	.62	18	-1.43	14.40
pth = 250 t ₅ = 0.6521831	her	41	.15	.62	.75	.48	05	62	-1.00	-1.08	94	74	66	75	87	79	th = 25	; = 1.04	her	20	.28	.64	.72	.49	.05	42	80	98	.98	88	80	78	76	on
ce Dej	tim _{er}	8060000.	.0001704	.0002208	.0002320	.0002073	.0001551	.0000816	0000056	0000911	0001541	0001806	0001710	0001336	0000746	.0000029	Sound Source Depth = 250	200409 E	tim _{er}	.0000842	.0001670	.0002188	.0002297	.0002041	.0001511	.0000762	0000139	0001035	0001703	0001989	0001892	0001502	0000887	
Sound Sour $\theta = 0.334327$	θ_{er}	.0007312	.0005428	.0003178	.0002425	.0003906	.0006718	.0008889	.0008694	.0005861	.0001824	0001161	0001448	.000093	.0004578	.0007135	Sound	$\theta = 0.200409$	θ_{er}	.0004691	.0003869	.0002773	.0002508	.0003501	.0005176	.0006315	.0005880	.0003756	0060000.	0001223	0001542	0000012	.0002374	CC2+000.
1308.76 h ₁ = 3000	þer	.0018177	.0023335	.0020193	.0010944	.0001812	0001053	.0004353	.0014525	.0022728	.0023491	.0016246	.0005677	0001358	0000311	.0008061	1308.76	$h_{1} = 5000$	φer	.0018424	.0023525	.0020293	.0010939	.0001705	0001244	.0004109	.0014270	.0022506	.0023341	.0016193	.0005731	0001206	0000087	N70000.
11	Zer	-1.87	-1.48	93	71	-1.04	-1.68	-2.13	-1.95	-1.08	.05	.84	.87	.14	90	1.70	11	1453	Zer	-2.60	-2.21	-1.66	-1.49	-1.93	-2.71	-3.24	-2.99	-1.91	50	.51	.62	17	-1.39 7 25	CC.7.
Acoustic Center Depth $t_5 = 0.4639118$	h_{er}			69.	.87						70	41	34	54	80	- 83	Acoustic Center Depth	= 0.8482453	her	27	.24								96	82	74		08	
Acousti 83579 t ₅	timer	9960000.	.0001706	.0002181	.0002293	.0002062	.0001561	.0000859	- 00000039	0000750	0001326	0001566	0001476	0001127	0000574	.0000150	Acousti	52155 t ₅	tim _{er}	.0000869	.0001686	.0002201	.0002311	.0002058	.0001530	.0000785	0000107	- 0660000'-	0001644	0001922	0001825	0001441	0000835	r c n n n n n
Array #1 30 θ=0.4	θ_{er}	.0009728	.0006887	.0003597	.0002388	.0004279	.0008080	.0011156	.0011180	.0007739	.0002686	0001035	0001271	.0001992	.0006640	.0009805	Array #1) = 0.2	θ_{er}	.0005733	.0004483	.0002923	.0002464	.0003658	.0005795	.0007354	.0007016	.0004601	.0001263	0001214	0001523	.0000373	.0003244	1 20000
Arr h ₁ = 2000	\$er	.0017878	.0023100	.0020063	.0010939	.0001933	0000828	.0004645	.0014834	.0023000	.0023678	.0016317	.0005621	0001532	0000576	.0007752	Arr	$h_1 = 4000$	φer	.0018328	.0023452	.0020255	.0010941	.0001747	0001171	.0004203	.0014369	.0022592	.0023399	.0016214	.0005710	0001265	0000174	N770NNN.
10/27/88	Ð	0000000	.4188790	.8377580	1.2566370	1.6755160	2.0943951	2.5132741	2.9321531	-2.9321531	-2.5132741	-2.0943951	-1.6755160	-1.2566370	8377580	4188790	10/27/88		Ð	0000000	.4188790	.8377580	1.2566370	1.6755160	2.0943951	2.5132741	2.9321531	-2.9321531	-2.5132741	-2.0943951	-1.6755160	-1.2566370	8377580	

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Table 3. (Continued)

50 3758	Zer	-1.39	-1.66	-1.45	74	.14	.74	.73	.11	79	-1.50	-1.70	-1.42	98	78	97	0	42213	Zer	-1.20	-1.39	-1.07	23	.76	1.43	1.41	.71	31	-1.16	-1.47	-1.26	86	66	83
pth = 250 t ₅ = 0.6463758	her	.58	.14	19	33	30	23	24	31	35	20	.13	.56	.93	1.07	.94	th = 25	$t_5 = 1.0442213$	her	.62	.24	09	29	36	36	36	37	31	11	.23	.61	.90	1.02	.92
ce De	tim _{er}	.0002035	.0001355	.0000557	0000168	0000682	0000939	0000946	0000695	0000173	.0000562	.0001358	.0002030	.0002456	.0002602	.0002464	Sound Source Depth = 250		timer	.0001765	.0001071	.0000247	0000508	0001045	0001313	0001320	0001057	0000513	.0000252	.0001074	.0001759	.0002187	.0002334	.0002197
Sound Sour $\theta = 0.307561$	θ_{er}	.0004311	.0005548	.0005262	.0003282	.0000586	0001297	0001258	.0000693	.0003432	.0005422	.0005690	.0004414	.0002745	.0001984	.0002692	Sound	$\theta = 0.183359$	θ_{er}	.0003250	.0003753	.0003278	.0001768	0000117	0001395	0001356	0000000	.0001922	.0003448	0003909.	.0003366	0002491	.0002066	.0002430
1218.84 h ₁ = 3000	þer	0018021	0014151	0006730	0000670	0000006	0005216	0012916	0018094	0017387	0011301	0003888	0000055	0002298	0009067	0015818	1218.84	$h_1 = 5000$	φer	0018125	0014285	0006870	0000792	0000089	0005245	0012886	0018010	0017265	0011162	0003755	.0000048	0002241	0009067	0015875
Acoustic Center Depth = 1218.84 $t_5 = 0.4556447$ $h_1 =$	Zer	,	-1.70	-1.52	90	11	.42	.42	12	92	-1.55	-1.73	-1.47	-1.08	90	-1.07	Acoustic Center Depth = 1218.84	= 0.8 4 38392	Zer	'	-1.55	-1.29	52	.42	1.06	1.05	.38	58	-1.36	-1.61	-1.36	93	72	91
stic Cer t ₅ = 0.4	her	.52	.05	26	32	19	06	06	21	33	27	.04	.51	.93	1.10	.94	tic Cen	$t_5 = 0.8$	her	.60	.20	14	32	34	32	32	36	33	15	.18	.59	16.	1.04	.92
006	tim _{er}	.0002299	.0001651	0000899	.0000218	0000263	0000505	0000512	0000275	.0000212	.0000902	.0001654	.0002296	.0002710	.0002853	.0002716	Acous	211	tim _{er}	.0001872	.0001182	.0000367	0000378	0000906	0001171	0001178	0000919	0000383	.0000372	.0001186	.0001867	.0002295	.0002442	.0002304
Array #56 00 0 = 0.447	θ_{er}	.0005318	.0007249	.0007152	.0004745	.0001303	0001141	0001104	.0001405	.0004885	.0007297	.0007372	.0005403	.0002988	.0001914	.0002946	Array #56	$\dot{\theta} = 0.231$	θ_{er}	.0003668	.0004465	.0004064	.0002365	.0000153	0001367	0001328	.0000261	.0002518	.0004231	.0004615	.0003779	.0002587	.0002028	.0002529
Arr h ₁ = 2000	þer	0017890	0013984	0006555	0000518	9600000.	0005179	0012953	0018198	0017540	0011477	0004055	0000185	0002368	0009066	0015747	Arr	$h_1 = 4000$	φer	0018086	0014235	0006818	0000747	0000058	0005234	0012897	0018041	0017311	0011214	0003805	6000000.	0002262	0009067	0015854
10/27/88	Ð	0000000	.4188790	.8377580	1.2566370	1.6755160	2.0943951	2.5132741	2.9321531	-2.9321531	-2.5132741	-2.0943951	-1.6755160	-1.2566370	8377580	4188790	10/27/88		Ð	0000000	.4188790	.8377580	1.2566370	1.6755160	2.0943951	2.5132741	2.9321531	-2.9321531	-2.5132741	-2.0943951	-1.6755160	-1.2566370	8377580	4188790

Table 3. (Continued)

																				Y		•												
0 40366	Zer	-2.46	-2.00	-1.38	-1.14	-1.52	-2.27	-2.80	-2.60	-1.62	32	.58	.62	19	-1.37	-2.27	0	90690	Z _{CT}	-2.75	-2.38	-1.83	-1.66	-2.10	-2.88	-3.37	-3.07	-1.92	43	.62	.73	11	-1.41	-2.45
pth = 250 $t_5 = 0.6540366$	her	41	.15	.64	.78	.51	01	58	96	-1.05	91	72	65	75	- 88	80	oth = 25	$t_5 = 1.0506906$	h_{er}	19	.29	68	LL.	.54	.10	37	75	94	94	84	77	76	75	58
ce De	tim _{er}	7060000.	.0001700	.0002203	.0002315	.0002069	.0001548	.0000815	0000054	000000	0001535	0001799	0001704	0001331	0000743	.0000031	Sound Source Depth = 250		tim _{er}	.0000842	.0001667	.0002184	.0002293	.0002039	.0001509	.0000762	0000136	0001030	0001696	0001980	0001884	0001495	0000882	0000073
Sound Sour 9 = 0.332511	θ_{er}	.0007400	.0005436	.0003088	.0002241	0003650	.0005419	.0008578	.0003399	.0005607	.0001637	0001263	0001457	.0001065	.000-4701	.0007264	Sound	$\theta = 0.197561$	θ_{er}	.0004768	.0003860	.0002656	.0002295	.0003224	000-4871	.0006012	.0005597	.0003505	.00008698	0001355	0001587	.0000028	.0002477	.0004357
1308.76 h ₁ = 3000	φer	.0018168	.0023317	.0020174	.0010933	.0001817	0001031	.0004386	.0014554	.0022741	.0023484	.0016227	.0005661	0001364	0000310	.0008061	1308.76	$h_1 = 5000$	φer	.0018406	.0023503	.0020274	.0010931	.0001714	0001218	.0004147	.0014305	.0022525	.0023338	.0016177	.0005714	0001217	0000093	.0008311
Acoustic Center Depth = t ₅ = 0.4652331	Zer	-1.81	-1.40	83	60	91	-1.55	-2.00	-1.82	96	.16	.94	.95	.21	85	-1.64	Acoustic Center Depth =	= 0.8506490	Zer	-2.61	-2.18	-1.58	-1.37	-1.79	-2.56	-3.09	-2.85	-1.78	38	.61	.68	15	-1.40	-2.37
stic Cent t ₅ = 0.46	her	- 44	.18	.75	.93	.64	.02	60	95	92	64	35	30	50	TT	80	tic Cent	$t_5 = 0.8t$	h_{cr}	27	.25	.67	.78	.54	.06	45	83	98	93	79	72	76	80	66
Acous 82345 t	timer	.0000963	.0001701	.0002175	.0002287	.0002057	.0001559	.0000859	.0000040	-0000747	0001321	0001561	0001471	0001123	0000572	.0000150	Acous		tim _{er}	.0000868	.0001683	.0002197	.0002307	.0002055	.0001528	.0000785	0000105	0000985	.0001637	0001915	0001818	0001435	0000830	0000033
Array #1 $\theta = 0.482345$	θ_{er}	.0009788	.0006881	.0003511	.0002221	.0004044	.0007800	.0010857	.0010891	.0007490	.0002503	0001134	0001284	.0002048	.0006737	.0009902	Array #1	$\theta = 0.249797$	θ_{er}	.0005823	.0004488	.0002822	.0002266	0665000.	.0005492	.0007046	.0006725	.0004349	.0001070	0001327	0001545	.0000436	.0003366	.0005523
Ап h ₁ = 2000	\$er	.0017867	.0023079	.0020040	.0010925	.0001937	0000804	.0004680	.0014866	.0023015	.0023674	.0016301	.0005606	0001538	0000575	.0007751	Arı	$h_1 = 4000$	φει	.0018316	.0023433	.0020237	.0010932	.0001753	0001147	.0004237	.0014399	.0022606	.0023393	.0016196	.0005694	0001272	0000175	.0008217
3/22/89	Ð	0000000	.4188790	.8377580	1.2566370	1.6755160	2.0943951	2.5132741	2.9321531	-2.9321531	-2.5132741	-2.0943951	-1.6755160	-1.2566370	8377580	4188790	3/22/89	0	o	0000000	.4188790	.8377580	1.2566370	1.6755160	2.0943951	2.5132741	2.9321531	-2.9321531	-2.5132741	-2.0943951	-1.6755160	-1.2566370	8377580	4188790

Table 3. (Continued)

	ŏ	Array #56 30	6542	tic Cent $t_5 = 0.45$	Acoustic Center Depth = 1218.84 t t ₅ = 0.4569723 h ₁ =	= 1218.84 h ₁ = 3000	Sound Sour $\theta = 0.305570$	ce De	pth = 250 t ₅ = 0.6482545	0 2545
¢er		θ_{er}	timer	her	Zer	þer	θ_{er}	tim _{er}	her	Zer
0017923		.0005007	.0002303	.47	-1.57	0018058	.0003725	.0002037	.50	-1.61
0014023		.0006844	.0001653	.01	-1.80	0014196	.0004854	.0001356	.08	-1.85
0006594		.0006649	0060000.	29	-1.60	0006775	.0004460	.0000556	24	-1.61
0000549		.0004158	.0000217	35	95	0000708	.0002393	0000171	38	88
.000000		.0000659	0000266	22	15	0000030	0000358	0000687	34	10.
0005183		0001808	0000509	08	.38	0005220	0002263	0000945	27	.62
0012941		0001759	0000515	- 00	.37	0012901	0002212	0000952	28	.61
0018174		.0000794	0000278	24	18	0018065	0000216	0000700	36	02
0017509		.0004347	.0000211	37	99	0017350	.0002596	0000177	40	94
0011442		.0006852	.000000	31	-1.64	0011261	.0004683	.0000561	26	-1.67
0004020		.0007023	.0001656	00	-1.83	0003848	.0005056	.0001359	.06	-1.91
0000155		.0005136	.0002300	.45	-1.59	0000021	.0003874	.0002032	.48	-1.66
0002351		.0002773	.0002715	.87	-1.21	0002278	.0002267	.0002459	.84	-1.24
0009067		0001200.	.0002858	1.03	-1.03	0009068	.0001520	.0002605	96.	-1.04
0015765		.0002706	.0002721	.88	-1.20	0015840	.0002188	.0002467	.85	-1.22
4	Ę	Arrav #56	Acoust	tic Cent	Acoustic Center Depth	= 1218.84	Sound	Sound Source Depth = 250	oth = 25	0
$h_1 = 4000$		$\dot{\theta} = 0.22$	8642	$t_5 = 0.84$	= 0.8462835	$h_1 = 5000$	$\theta = 0.180284$		$t_5 = 1.0472331$	2331
¢er		θ_{er}	tim _{er}	her	Zer	þer	θ_{er}	tim _{er}	her	Zer
0018133		.0002785	.0001874	.51	-1.69	0018185	.0002049	.0001765	.50	-1.78
0014291		.0003471	.0001182	.11	-1.88	0014356	.0002445	.0001069	.13	-1.91
0006874		.0002964	.0000365	21	-1.58	0006942	.0001869	.0000243	19	-1.54
0000795		.0001182	0000382	39	78	0000855	.0000287	0000514	39	67
0000089		0001077	0000912	41	.17	0000130	0001638	0001052	45	.33
0005240		0002614	0001177	38	.82	0005254	0002927	0001321	- 44	1.00
0012878		0002563	0001184	39	.80	0012862	0002877	0001328	45	.98
0018004		0000935	0000925	43	.12	0017961	0001498	0001065	47	.26
0017262		.0001387	0000387	41	86	0017202	.0000489	0000518	42	77
0011161		.0003191	.0000370	23	-1.67	0011093	.0002095	.0000249	22	-1.65
0003753		.0003680	.0001185	60.	-1.96	0003688	.0002655	.0001073	.11	-2.01
.0000052		.0002942	.0001867	.49	-1.75	.0000102	.0002208	.0001758	.48	-1.86
0002239		.0001815	.0002297	.80	-1.35	0002212	.0001400	.0002188	LL.	-1.49
0009070		.0001272	.0002444	.93	-1.15	0009071	.0000992	.0002335	88.	-1.31
0015882		.0001732	.0002306	.81	-1.32	0015912	.0001315	.0002198	.78	-1.45

Table 4. Error Structure - Proposed Methodology

10/27/88	Array h ₁ = 2000	#1 Acou: $\theta = 0.489930$	stic Cent $t_4 = 0.4$	Acoustic Center Depth = 0.4654736	1308.76 h ₁ = 3000	Sound Source Depth = 250 $\theta = 0.339157$ $t_4 = 0.65328$	Depth = $t_4 = 0.6$	epth = 250 t ₁ = 0.6532862
	фег	θ_{er}	her	Zer	¢er	θ_{er}	her	Zer
	.00002258	.00006163	07	12	.00001590	.00006335	07	19
	.00003532	.00004005	04	08	.00002490	.00004109	04	12
	.00004159	76600000.	01	02	.00002924	09600000.	01	03
	.00004007	00002200	.02	.04	.00002802	n0002369	.03	.07
	.00003132	00004904	.05	.10	.00002171	00005135	.06	.15
	.00001754	00006587	.07	.13	.00001201	00006775	.07	.20
	.00000186	00007002	.08	.14	.00000126	00007047	.08	.21
	00001258	00006218	.07	.12	00000820	00006092	.07	.18
	00002356	00004547	.05	60.	00001486	00004348	.05	.13
	00003016	00002363	.03	.05	00001853	00002295	.03	.07
	00003232	.00000056	00.	00	00001982	00000153	00.	00 [.]
	00002989	.00002534	03	05	00001885	.00002127	02	06
	00002235	.00004820	05	10	00001469	.00004464	05	13
	00000073	.00006478	07	13	00000662	.00006356	07	19
	.00000632	00007019	08	14	.00000441	.00007116	08	21
	Arrav	#1 Acous	stic Cen	Acoustic Center Depth =	1308.76	Sound Source	Depth = 250	250
	$h_1 = 4000$	$\theta = 0.255966$	$t_4 = 0.8$	= 0.8490844	$h_1 = 5000$	$\theta = 0.203535$	$t_4 = 1.0$	$t_4 = 1.0484023$
	ber	θ_{er}	her	Zer	þer	θ_{er}	her	Zer
	.00001262	.00006511	07	26	00001070	.00006692	08	34
	12610000.	.00004209	05	17	.00001660	.00004306	05	22
	.00002304	.00000926	01	04	.00001930	.00000895	01	04
	.00002197	00002525	.03	.10	.00001832	00002670	.03	.13
	.00001693	00005352	.06	.21	.00001407	00005557	90.	.28
	.00000928	00006961	.08	.28	.00000768	00007145	.08	.36
	76000000.	00007102	.08	.28	.00000081	00007167	.08	.36
	00000606	00005967	.07	.24	00000482	00005837	.07	.29
	00001058	00004127	.05	.17	00000807	00003877	<u>.</u> 0	.19
	00001271	00002212	.02	60.	00000923	00002101	.02	.11
	00001351	00000407	00.	.02	00000967	00000727	.01	.04
	00001335	.00001629	02	07	00001008	.00001000	01	05
	00001095	.00004049	04	16	00000881	.00003551	04	18
	00000511	.00006230	07	25	00000425	.00006100	07	31
	.00000350	.00007226	08	29	.00000299	.00007353	08	37

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250 73052	Zer	.14	.07	01	10	17	20	19	12	04	.04	.11	.17	.20	.20	.18	250	47819	Zer	.26	.16	.03	13	27	34	30	17	01	.12	.22	.29	.33	.34	.32
Depth = 250 t ₄ = 0.6473052	her	.05	.02	00	03	06	07	06	04	01	.01	.04	.06	.07	.07	90.	Jepth =	t ₄ = 1.0447819	her	.05	.03	0.	03	06	07	06	03	00	.02	.05	90.	.07	.07	.07
Sound Source Depth = 250 $\theta = 0.312071$ $t_4 = 0.64730$	θ_{er}	00004670	00002438	.00000353	.00003280	.00005670	.00006774	.00006179	.00004134	.00001359	00001428	00003791	00005528	00006540	00006759	00006136	Sound Source Depth = 250	$\theta = 0.186264$	θ_{er}	00005136	00003192	00000539	.00002573	.00005426	.00006832	.00005969	,00003339	.00000255	00002389	00004379	00005772	00006597	00006807	00006336
= 1218.84 h ₁ = 3000	þer	00001609	00002056	00002197	00001934	00001235	00000218	.00000845	.00001658	.00002058	.00002052	.00001732	.00001195	.00000519	00000227	00000965	= 1218.8 4	$h_1 = 5000$	þ er	00000852	00001126	00001265	00001181	00000795	00000143	.00000551	.00001017	.00001169	.00001090	.00000882	.00000598	.00000259	00000116	00000498
Acoustic Center Depth : 1929 t ₄ = 0.4569741	h _{er} z _{er}		.02 .04		0407		0713		0409	0204	.01 .02	.03 .07	.05 .11	.06 .13	.07 .13	.06 .12	Acoustic Center Depth	t4 = 0.8445435	h _{er} z _{er}			00. 00.	0312	0622	0727	0624	0415		.02 .08	.04 .16	.06 .23	·	•	.06 .25
$#56 Acous \theta = 0.453929$	θ_{er}	00004423	00002062	.00000761	.00003575	.00005766	.00006756	.00006258	.00004438	.00001822	00000987	00003494	00005397	00006515	00006738	00006026	#56 Acous	$\theta = 0.234753$	θ_{er}	00004909	00002816	0000080	.00002949	.00005559	.00006799	.00006086	.00003774	.00000842	00001896	00004086	00005653	00006568	00006782	00006239
Array i h ₁ = 2000	ber	00002534	00003175	00003311	00002836	00001774	00000311	.00001211	.00002430	.00003109	.00003191	.00002755	.00001929	.00000841	00000365	00001542	Array	$h_1 = 4000$	¢er	- 00001139	00001482	00001625	00001471	00000963	00000172	.00000662	.00001263	.00001513	.00001461	.00001207	.00000824	.00000357	00000158	00000674
10/27/88	Ð	0000000	.4188790	.8377580	1.2566370	1.6755160	2.0943951	2.5132741	2.9321531	-2.9321531	-2.5132741	-2.0943951	-1.6755160	-1.2566370	8377580	4188790	10/27/88		Ð	000000	.4188790	.8377580	1.2566370	1.6755160	2.0943951	2.5132741	2.9321531	-2.9321531	-2.5132741	-2.0943951	-1.6755160	-1.2566370	8377580	4188790

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epth = 250 t ₄ = 0.6551362	Zer	26	17	04	60.	.21	.28	.29	.25	.18	60.	.01	08	18	26	29	250	$t_4 = 1.0513520$	Zer	- 44	28	06	.17	.36	.47	.47	.38	.25	.14	.07	03	21	39 - 48
Depth = t4= 0.6	her	.00	06	01	.03	.08	.10	11.	60.	.06	.03	00.	03	07	60	11	Depth =	$t_4 = 1.0$	her	10	06	01	.04	.08	.11	.11	60.	.06	.03	.02	01	05	- 00. - 11
Sound Source Depth = $250 \theta = 0.337225 t_{4} = 0.65513$	0er	.00008688	.00005637	.00001347	00003181	00006968	00009251	00009663	00008341	00005939	00003159	00000317	.00002745	.00005986	.00008666	1770000.	Sound Source Depth = 250	$\theta = 0.200474$	θ_{er}	.00008844	.00005702	.00001270	00003348	00007149	00009323	00009423	00007645	00004999	00002791	00001347	.00000667	.00004166	.00007855 00009699
= 1308.76 h ₁ = 3000	¢er	.00002163	.00003395	.00003989	.00003830	.00002986	.00001668	.00000193	00001116	00002024	00002516	00002692	00002586	00002045	00000943	.00000584	= 1308.76	$h_1 = 5000$	Øer	.00001385	.00002162	.00002516	.00002403	.00001871	.00001048	.00000135	00000623	00001044	00001169	00001226	00001333	00001224	00000630
Acoustic Center Depth = 452 t ₄ = 0.5438750	her Zer	ſ	0809	0202	.04 .05		.13 .15		.13 .14		.05 .05	0000	0606		1315		Acoustic Center Depth =	$t_4 = 0.8514820$	her Zer	1	0623	0105	.04 .13		.10 .37			.06 .22	.03 .12	.01 .03			0933
#1 Acous $\theta = 0.711452$	θ_{er}	.00006864	.00004455	.00001161	00002346	00005365	00007333	00007943	00007193	00005335	00002743	.00000209	.00003173	.00005753	.00007479	00007919	#1 Acous	$\theta = 0.253446$	θ_{er}	.00008781	.00005682	.00001303	00003288	00007091	00009314	0000956+	00008011	00005503	00002991	00000760	.00001845	.00005177	.00008284 .00009744
Array h ₁ = 2000	Ø er	.00003985	.00006200	.00007302	00001069	.00005579	.00003169	.00000343	00002363	00004518	00005863	00006280	00005718	00004173	00001775	.00001127	Array	$h_1 = 4000$	¢er	.00001680	.00002631	.00003080	.00002948	.00002294	.00001279	.00000157	00000806	00001408	00001677	00001784	00001801	00001520	00000736 .00000443
3/22/89	Ð	0000000	.4188790	.8377580	1.2566370	1.6755160	2.0943951	2.5132741	2.9321531	-2.9321531	-2.5132741	-2.0943951	-1.6755160	-1.2566370	8377580	4188790	3/22/89)	Ð	0000000	.4188790	.8377580	1.2566370	1.6755160	2.0943951	2.5132741	2.9321531	-2.9321531	-2.5132741	-2.0943951	-1.6755160	-1.2566370	8377580 4188790

Table 4. (Continued)

3/22/89	Array	#56 Acou	Acoustic Center Depth	II	1218.84	Sound Source I	Depth = 250	250
	$h_1 = 2000$	$\theta = 0.452549$	$t_4 = 0.4583009$	33009	$h_1 = 3000$	$\theta = 0.309991$	$t_4 = 0.6491808$	91808
Ð	þer	θ_{er}	her	Zer	þer	θ_{er}	her	Zer
0000000	00003537	00006070	.06	.12	00002153	00006076	90.	.18
.4188790	00004398	00002798	.03	.06	00002708	00003140	.03	60.
.8377580	00004567	.00001082	01	02	00002868	.00000469	00	01
1.2566370	00003907	.00004953	05	10	00002512	.00004263	04	13
1.6755160	00002430	00007979	08	16	00001584	.00007385	07	22
2.0943951	00000389	.00009328	09	19	00000226	.00008781	09	26
2.5132741	.00001726	.00008583	09	17	.00001180	.00007868	08	24
2.9321531	.00003398	.00005998	06	12	.00002219	.00005037	05	15
-2.9321531	.00004307	.00002352	02	05	.00002695	.00001363	01	04
-2.5132741	.00004398	00001501	10.	.03	.00002653	00002208	.02	.07
-2.0943951	.00003788	00004933	.05	.10	.00002229	00005202	.05	.16
-1.6755160	.00002639	00007547	.07	.15	.00001530	00007429	.07	.22
-1.2566370	.00001128	18060000	60.	.18	.00000635	00008739	60.	.26
8377580	00000553	00009361	60.	.19	00000361	00008975	60.	.27
4188790	00002183	00008334	.08	.17	00001332	00008074	.08	.24
3/22/89	Array	#56 Acou	Acoustic Center Depth	11	1218.84	Sound Source Depth = 250	Depth = 3	250
	$h_1 = 4000$	$\theta = 0.235058$	$t_4 = 0.846981$	5981 7	$h_1 = 5000$	$\boldsymbol{\theta}=0.183019$	$t_4 = 1.0477857$	77857
Ð	фег	θ_{er}	h_{er}	Zer	¢er	θ_{er}	her	Zer
0000000	00001438	00005913	.06	.24	00000981	00005569	90.	.28
.4188790	00001806	00003375	.03	.13	00001218	00003487	.04	.17
.8377580	00001942	00000166	00.	.01	00001316	00000825	.01	.04
1.2566370	- 00001739	.00003401	03	14	00001201	.00002314	02	12
1.6755160	00001110	.00006516	07	26	00000769	.00005246	05	26
2.0943951	00000123	00620000.	08	32	00000038	.00006500	07	33
2.5132741	.00000893	.00006768	07	27	.00000701	.00005078	05	25
2.9321531	.00001574	.00003703	04	15	.00001123	.00001837	02	09
-2.9321531	.00001805	.00000136	00	01	.00001187	00001400	0.	.07
-2.5132741	.00001701	00002969	.03	.12	.00001057	00003790	.04	.19
-2.0943951	.00001394	00005396	.06	.22	.00000846	00005498	.06	.27
-1.6755160	.00000944	00007185	.07	.29	.00000567	00006792	.07	.34
-1.2566370	.00000373	00008260	.08	.33	.00000203	00007617	.08	.38
8377580	00000270	00008440	60.	.34	00000220	00007731	.08	.39
4188790	00600000	00007648	.08	.31	00000634	00007033	.07	.35

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support errors of this magnitude. (The calculations in this section apply the polynomial extrapolation prior to the development of isospeed layers. In this way the DV table extension is not an issue in the present comparisons.)

Accordingly, the author has proposed and tested a method that treats the two main issues. The acoustic center is moved to the C-phone. (Visualize Figure 3 with the C-phone on the axis of the cylinder.) In this way the transit time that stops the ray tracing is t_4 , a directly measured quantity. Also the sound speed in the layer containing the (X, Y, C) phones (Z-phone omitted) is assumed constant in order to determine (ϕ_p , θ_p), the azimuth and elevation angles of the proposed method. In this way the change in sound speed at the Z-phone is taken out of the computation. Details appear in Appendix F.

Computations using this technique appear in Table 4. The improvement is dramatic. Let us repeat the earlier exemplary computations 10/22/88, Array No. 1, $h_1 = 5000$ and $\phi = 0.418879$. This time the inputs come from Table 4: specifically $\sqrt{(.22)^2 + (0.5)^2 + (5 \cdot 0.04306)^2} = 0.31$ ft., a value dramatically smaller than 12 ft. A scan of Table 4 and comparison with Table 3 shows that these improvements are quite consistent.

These computations also have the advantage that the five-foot layer isogradient ray tracing was used together with the exact tilt corrections, eq. (B.5.)

5. CONCLUSIONS

Generally there are a number of sources of systematic error. In some the effects are small, e.g., isogradient vs. isospeed raytracing, 5 foot versus 25 foot layer thicknesses. None the less they are systematic and certainly no longer necessary. The presence of systematic errors tends to build up idiosyncracies

that frustrate the use of standard statistical methodology when troubleshooting other aspects of the data.

The discovery of the periodic nature of the errors was a surprise. Its presence adds another dimension to the interpretation of the data. No doubt it is the source of some deception.

The major sources of error related to ray tracing (see pages 5, 6) are believed to be associated with

- (i) the conversion of transit times to the direction of the sound source
- (ii) the constant speed extrapolation of the DV profile below 1300 ft.
- (iii) the use of approximations for tilt corrections.

Of course the effects of these errors are directional because of their periodic nature. We have documented errors of 12 ft. or so due to (i). We can speculate 10 or more feet because of (ii). Moreover the combined effects could be additive. The effect of (iii) increases as the tilts increase.

It is further noted that the above errors apply to single arrays. At this point we have no comment about how they may combine to produce mismatches in the array overlap regions. It would be more comfortable to treat this issue by introducing the changes and then collecting more data.

6. ADDENDUM

It came to the author's attention during the final editing phase of this report, that the approximate tilt corrections, eq. (B.4), are not the ones currently employed by the software. Rather, the system is using

$$\begin{cases} X(1) \\ X(2) \\ X(3) \end{cases} = \begin{cases} c_2 & 0 & -s_2 \\ 0 & c_1 & -s_1 \\ s_2 & s_1 & c_1c_2 \end{cases} \begin{cases} X_0(1) \\ X_0(2) \\ X_0(3) \end{cases}$$
(6.1)

Accordingly the author made some specialized computations for purposes of indicating the effect. Referring to Tables 3 and 4, we have chosen to treat the exemplary case: water column of 10/27/88, array number 1, $h_1 = 5000$ ft, $z_1 = 250$, and $\phi = 0.418879$ radians. The table below contains the four errors and total offset for each of four algorithms:

- (i) the isospeed method (from Table 3)
- (ii) the isospeed method using (6.1) vice (B.4)
- (iii) the isospeed method using (B.5) vice (B.4)
- (iv) the proposed method (from Table 4)

TABLE 5. EFFECT OF TILT CORRECTION METHOD

10/27/88, Array #1, $h_1 = 5000$ ft, $z_1 = 250$ ft, $\phi = 0.418879$ rad., $a_2 = 1308.76$ ft.

	\$ er	θer	her	zer	d
(i)	.0023525	.0003869	.28	-2.45	12.02
(ii)	.0021174	0001567	1.47	2.95	11.09
(iii)	0000161	0000077	1.26	2.19	2.53
(iv)	.0000166	.0000430	05	- 22	0.31

The results of Table 5 suggest the following: there is little distinction between the use of (B.4) and (6.1) for the tilt correcting (compare (i) and (ii)). However case (iii) suggests that there is much to be gained by use of the exact tilt correct (B.5) even if isospeed ray tracing is used with 25 ft layer thickness. Finally case (iv) suggests that considerable gains are available if, in addition, we use the proposed methodclogy.

All of these systematic errors are mathematical, i.e., due to choice of algorithms. There is no longer any reason not to use the best.

APPENDIX A

The twelve depth velocity profiles used in this study are recorded here in graphical form. They include the profiles used in [7]. Generally the measured values stop at a depth of about 1300 feet and it is necessary to extrapolate, in several instances to depths greater than 1350 feet (see Table B-2).

The insets of the graphs illustrate two different extrapolation schemes. The constant value extrapolation is the one currently in use, [5]. But there is a slight deception. Current methodology utilizes isospeed profiles, see Figure 2, and 25 foot layer thicknesses. The constant value extrapolation is the value of the deepest 25 foot layer appearing in an isospeed profile.

The other extrapolation is based upon fitting a second order polynomial to the deepest hundred feet of the original profile (five foot layer thicknesses). There are two steps in this process. First is fitting the curve by least squares. Because of the equally spaced depth increments, the fitting takes an especially simple form. Using the equation,

$$v = a + b(u - \overline{u}) + c(u - \overline{u})^2 \tag{A.1}$$

1

where v is velocity, u is layer depth, and \overline{u} = average depth, the normal equations take the form

$$\Sigma v = \Sigma a - \Sigma b(u - \overline{u}) - c\Sigma (u - \overline{u})^{2}$$
$$\Sigma v(u - \overline{u}) = \Sigma a(u - \overline{u}) - \Sigma b(u - \overline{u})^{2} - c\Sigma (u - \overline{u})^{3}$$
$$\Sigma v(u - \overline{u})^{2} = \Sigma a(u - \overline{u})^{2} - \Sigma b(u - \overline{u})^{3} - c\Sigma (u - \overline{u})^{4}$$

Using the notation $S_v = \Sigma v$, $S_{vu} = \Sigma v (u - \overline{u})$, $S_{vuu} = \Sigma v (u - \overline{u})^2$. $S_{u2} = \Sigma (u - \overline{u})^2$, $S_{u4} = \Sigma (u - \overline{u})^4$ and recognizing $\Sigma (u - \overline{u}) = \Sigma (u - \overline{u})^3 = 0$ because of the uniform spacing, the above equations assume the reduced form

$$Sv = na + cS_{u2}$$

$$S_{vu} = bS_{u2}$$

$$S_{vuu} = aS_{u2} + cS_{u4}$$
(A.2)

and we solve for the coefficient of the quadratic term

$$c = (nS_{\upsilon uu} - S_{\upsilon}S_{u2}) / (nS_{u4} - S_{u2}^2)$$
(A.3)

(This done, values for a and b come easily from the first two equations.)

We note in passing that if n = 2K+1, an odd number, and Δ is the spacing between consecutive values of {u}, then \overline{u} is an integer and

$$S_{u2} = \frac{\Delta^2}{3} K(K+1)(2K+1),$$

$$S_{u4} = \frac{\Delta^4}{2} [K(K+1)]^2,$$

$$nS_{u4} - S_{u2}^2 = \Delta^4 K^2 (K+1)^2 (2K+1)(7-4K) / 18$$

(A.4)

Such formulas expedite the computation of c in (A.3).

The second step deals with the issue of how to use this quadratic for extrapolation purposes. Generally, the direct use of the equation (A.1) would lead to a visual discontinuity between the last measured value and the first extrapolated one. We have chosen not to do this. Instead we recognize that Δ ·c represents the first difference of the gradient sequence {v₁(i)}, see Figure 2. So to preserve continuity, the extrapolation is enabled by using successive updates

$$v_1(i) = v_1(i-1) + \Delta \cdot c$$

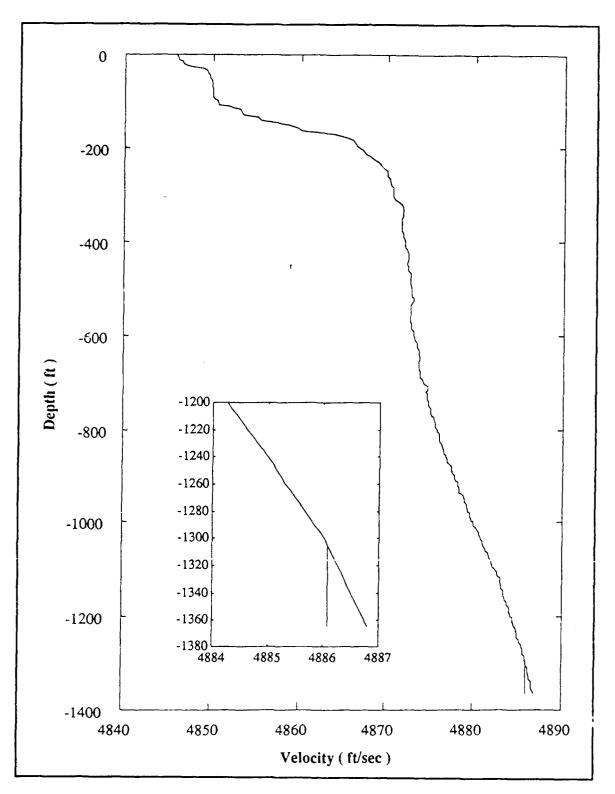
$$v(i+1) = v(i) + \Delta v_1(i)$$

$$v_0(i) = \left[u(i+1)v(i) - u(i)v(i+1)\right] / \Delta$$

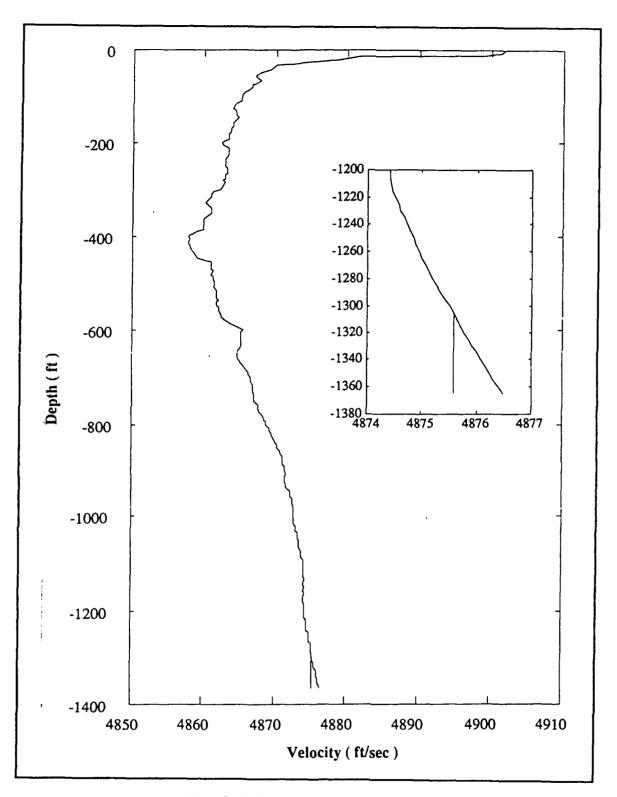
where v(i) is the estimated sound speed at u(i); $v_0(i)$ and $v_1(i)$ are the intercept and slope values of the straight line fit for the ith layer. It is this continuity preserving step that explains the occasional appearance of "crooked" extrapolations in the insets.

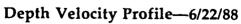
The use of isogradient ray tracing with 25 foot layer thicknesses was used in Section 3. If extrapolation of the water column was required, then a different second order polynomial method was used, and after the conversion to 25 foot layers. Basically the quantity $\Delta \cdot c$ was estimated by averaging the difference of the last five values of v_1 (last 100 feet), and then proceeding in the same way as stated above. No graphs showing the effect of this have been prepared. But the choice does have an effect upon the d_g values computed for Table 2.

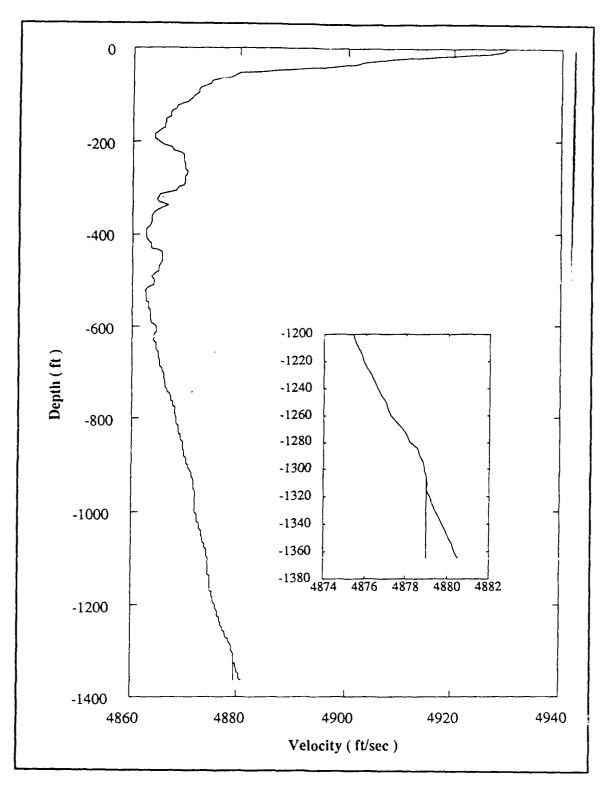
We take this opportunity to note that the visual appeal of the quadratic extrapolation is rather good in instances 1, 2, 3, 4, 6 and 11. The others are easily challenged. This information may influence the reader's interpretation of some values appearing in Table 2.



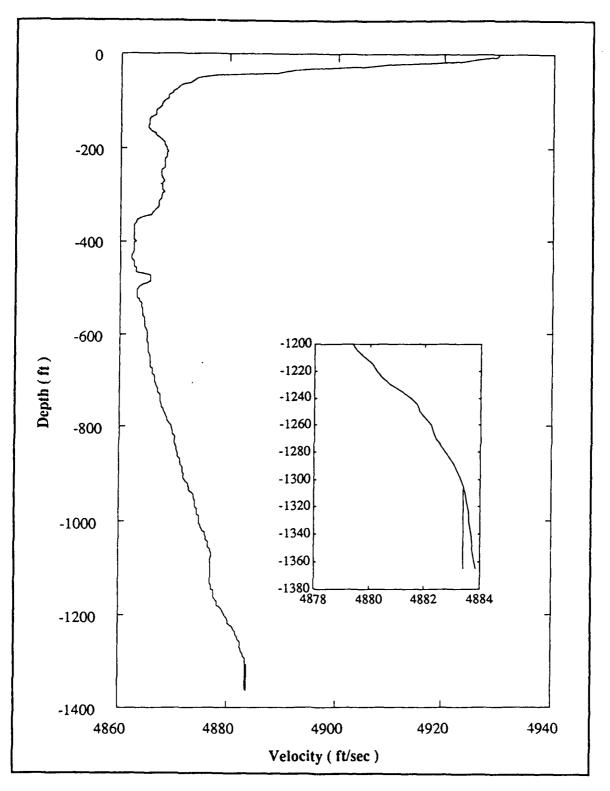
Depth Velocity Profile-5/12/88



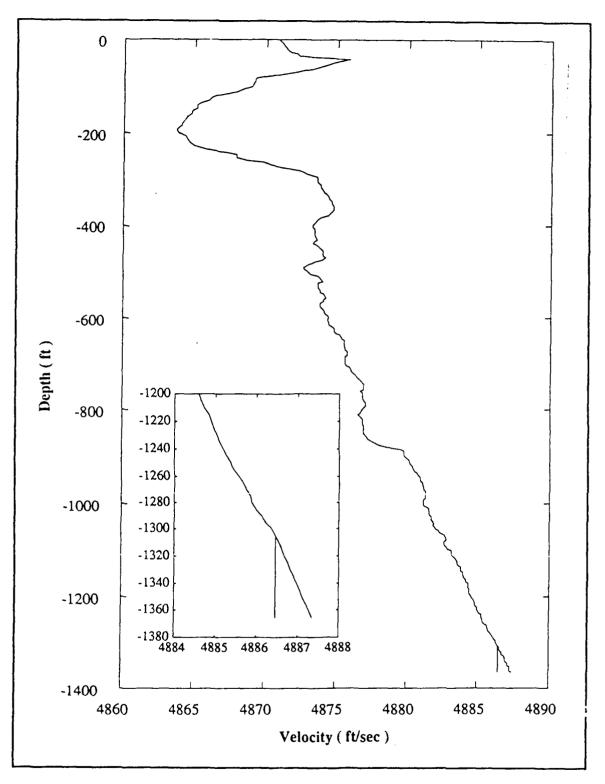


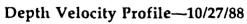


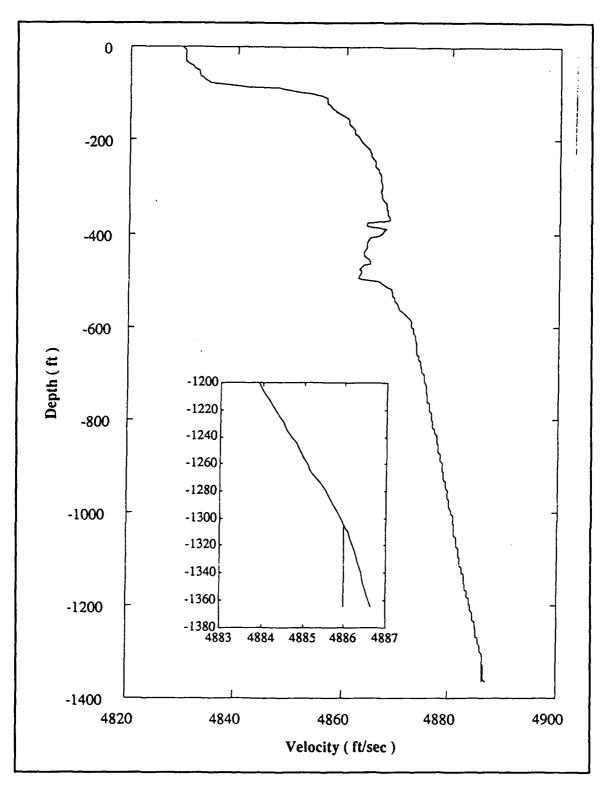
Depth Velocity Profile-7/21/88



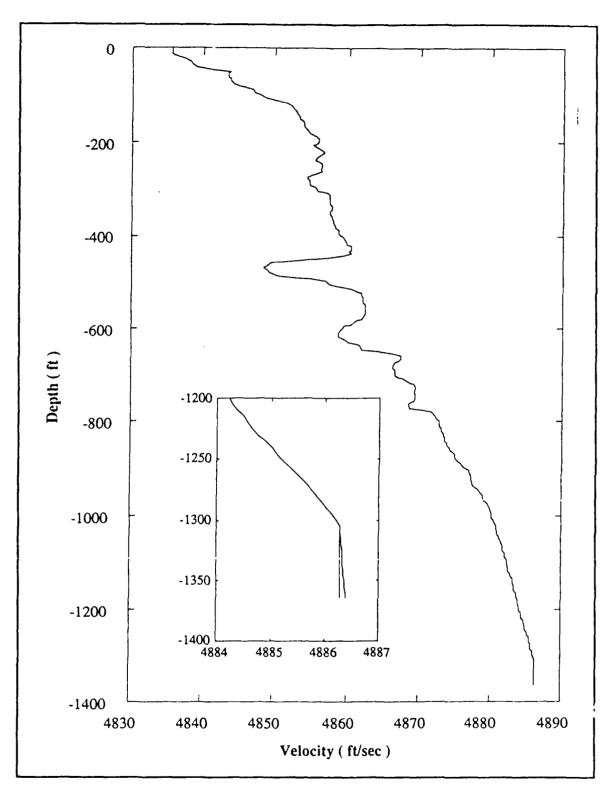
Depth Velocity Profile-8/03/88



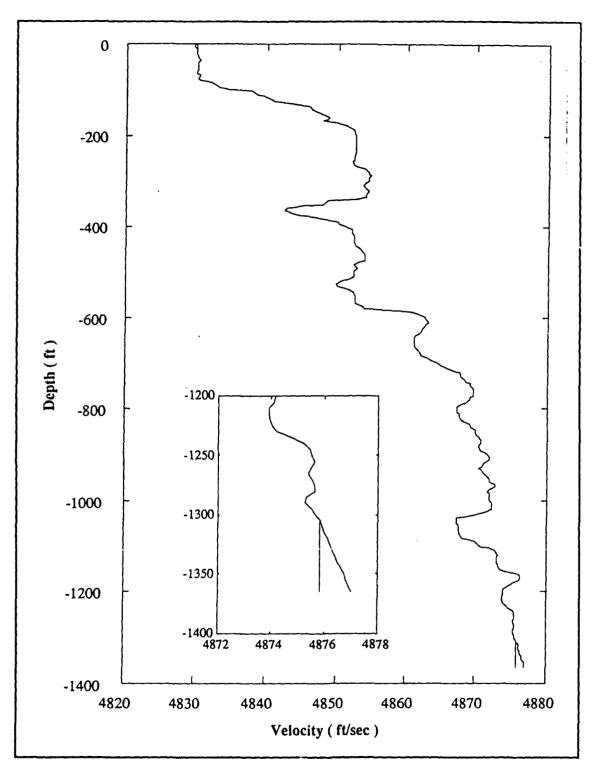




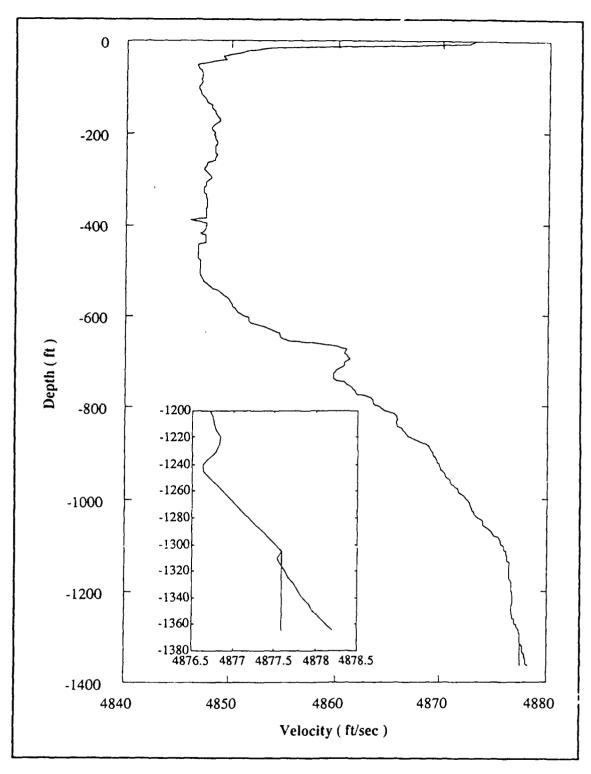
Depth Velocity Profile-1/13/89



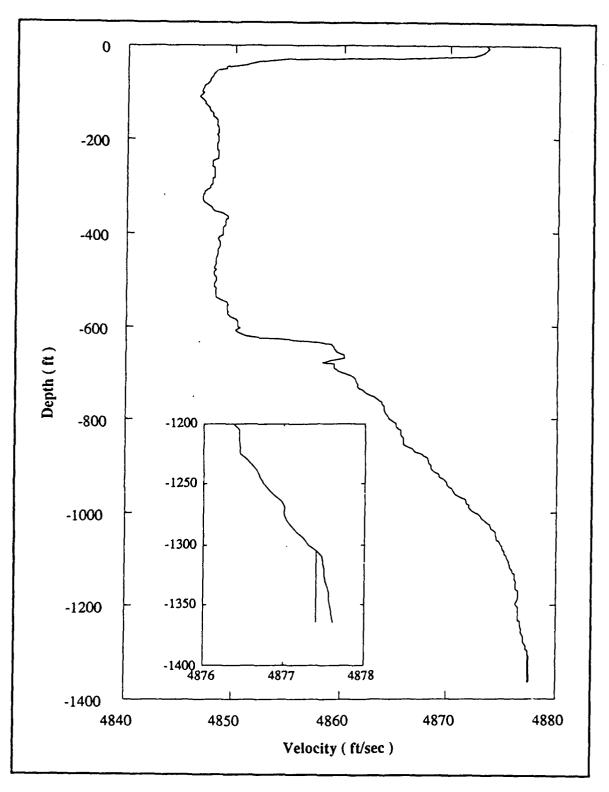
Depth Velocity Profile-3/08/89



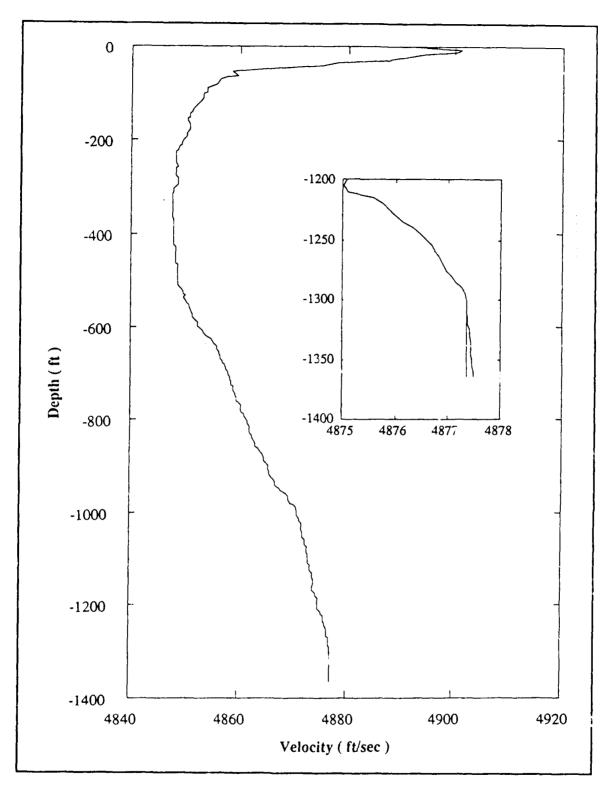
Depth Velocity Profile-3/22/89



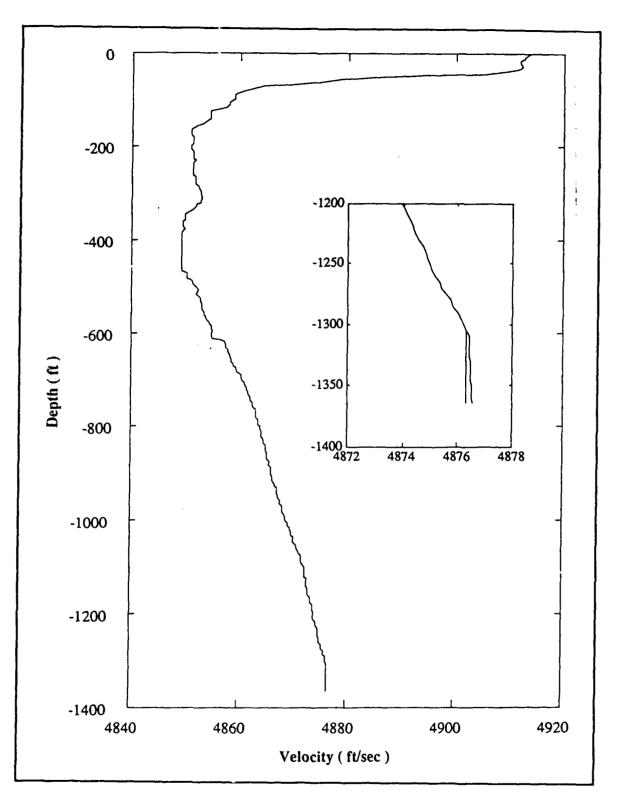
Depth Velocity Profile-4/26/89



Depth Velocity Profile-4/27/89



Depth Velocity Profile-5/10/89



Depth Velocity Profile-6/06/89

i.

APPENDIX B

COORDINATE SYSTEMS

The developments that follow deal with several coordinate systems and we need efficient ways to distinguish among them. First of all is the right handed system defined by the X, Y, Z and C hydrophones of the array. All incoming transit times must be interpreted in this system and we call it cs(a), or the coordinate system of the array. The origin is at the c hydrophone.

Since this system is generally tilted with respect to a "flat earth surface" system there is need to rotate cs(a) into alignment with a common coordinate system for all arrays, or a range coordinate which has horizontal directions consistent with the earth's surface. Such a resultant system will be called cs(b) and, for convenience of terminology, the X arm of cs(a) is rotated to a position called *east*; the Y arm to a position called *north*, and the Z arm to a position called vertical or zenith. The origin is still at the c hydrophone.

The ray tracing methodology of Ref [5] attempts to locate the sound source in relation to a specific point termed the acoustic center. This center is the geometric center of the array cube. The resulting coordinate system is a translation of cs(a), and will be called cs(ac). In this system the four hydrophone positions are specified by D/2 times the vectors

$$\begin{pmatrix} 1 \\ -1 \\ -1 \\ -1 \end{pmatrix} \begin{pmatrix} -1 \\ 1 \\ -1 \\ -1 \end{pmatrix} \begin{pmatrix} -1 \\ -1 \\ 1 \end{pmatrix} \begin{pmatrix} -1 \\ -1 \\ -1 \\ -1 \end{pmatrix}$$
(B.1)

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respectively, whereas in cs(a) these four vectors would be the three unit basis vectors and the origin. Conversion from cs(ac) to cs(a) is a direct translation.

The conversion of a point located in cs(a) to the cs(b) system requires a three-dimensional rotation based upon the tilt angles and the "ZROT" horizontal direction correction. It is convenient to describe this in terms of the three Euler angles ϕ_1 , ϕ_2 , ϕ_3 , or roll, pitch and yaw. Letting

$$s_i = sine(\phi_i)$$
 and $c_i = cosine(\phi_i)$ $i = 1, 2, 3$ (B.2)

we define three successive rotations which, when applied sequentially to a point in cs(a), will in the end describe it in cs(b). First hold the X arm fixed and rotate the Y-Z plane through an angle ϕ_1 ; the matrix of this transform is

$$\rho_1 = \begin{bmatrix} 1 & 0 & 0 \\ 0 & c_1 & -s_1 \\ 0 & s_1 & c_1 \end{bmatrix}$$

Next hold the (current position of) Y-arm fixed and rotate the X-Z plane through an angle of ϕ_2 ; this transformation has matrix

$$\rho_2 = \begin{bmatrix} c_2 & 0 & -s_2 \\ 0 & 1 & 0 \\ s_2 & 0 & c_2 \end{bmatrix}$$

Finally hold the (current position of) Z-arm fixed and rotate the X-Y plane through an angle of ϕ_3 ; the transform is

$$\rho_3 = \begin{bmatrix} c_3 & -s_3 & 0 \\ s_3 & c_3 & 0 \\ 0 & 0 & 1 \end{bmatrix}$$

The successive applications of these three rotations is a unitary transformation, (i.e., its inverse is equal to its transpose),

$$B = \rho_3 \rho_2 \rho_1 \tag{B.3}$$

and if a is a three vector in cs(a), then b = Ba is the same vector referenced in cs(b).

The determination of the Euler angles is accomplished as follows. The submerged arrays have tilt indicators on the X and Y arms which, individually, measure the angles that these arms make with the horizontal. An accounting for these tilts must be made when the ray trace azimuth and elevation angles are converted to a horizontal based coordinate system. The technique currently in use takes the apparent position, X_0 , and applies the transformation

$$\begin{cases} X(1) \\ X(2) \\ X(3) \end{cases} = \begin{cases} 1 & 0 & -\sin(XTILT) \\ 0 & 1 & -\sin(YTILT) \\ \sin(XTILT) & \sin(YTILT) & 1 \end{cases} \begin{cases} X_0(1) \\ X_0(2) \\ X_0(3) \end{cases}$$
(B.4)

so that the new apparent position, X, is referenced in a plane level with the earth. This transformation is an approximation which simply rotates the two arms to the horizontal as if they were separate unconnected arms and the rotation of one does not affect the rotation of the other. That is, the first two rows of the coefficient matrix are not orthogonal. The result is an approximation whose success depends upon the smallness of the tilt angles.

The exact way to accomplish this goal involves the direct replacement of the coefficient matrix with the product rotation $\rho_2\rho_1$:

$$\begin{cases} X(1) \\ X(2) \\ X(3) \end{cases} = \begin{cases} c_2 & -s_1 s_2 & -c_1 s_2 \\ 0 & c_1 & -s_1 \\ s_2 & s_1 c_2 & c_1 c_2 \end{cases} \begin{cases} X_0(1) \\ X_0(2) \\ X_0(3) \end{cases}$$
(B.5)

7.

Upon comparing these two coefficient matrices, one sees there is choice in identifying the two tilt angles with these two Euler angles. We have chosen to match the first two elements of the third row:

$$s_2 = \sin(XTILT)$$
 $s_1 = \sin(YTILT) / c_2.$ (B.6)

The geometric interpretation is as follows. First hold the X arm fixed and rotate the plane of the Y-Z arms so that the Y arm is horizontal. This is not a vertical projection. The division by c₂ shows that one must rotate through an angle greater than YTILT in order to maintain orthogonality of the coordinate system when making the Y arm parallel to the earth's surface. This done, we next hold the Y arm fixed in its new position and rotate the plane of the X-Z arms so that the X arm is horizontal. Since the new Y arm is already horizontal this second rotation is a vertical projection through an angle of XTILT.

This latter method is exact. The nature of the original approximation can be assessed by comparing the two coefficient matrices, using numerical inputs. The effect is not great for most of the tilts present at Nanoose.

To complete the conversion of cs(a) to cs(b) we must find the third Euler angle in terms of ZROT, the rotatic of the horizontal plane determined by array survey. In [5], ZROT is defined as the angle from the horizontally projected X arm to the range center line (east). The comparison of this definition with that of ϕ_3 (see ρ_3) leads to

$$s_3 = -\sin(ZROT) \tag{B.7}$$

The subsequent application of ρ_3 to $\rho_2\rho_1$ will bring the point a into east-north orientation.

Finally the position location system prefers to specify an object's position in terms of east, north, and depth (positive) below the sea surface. This system is a left-handed one with origin on the sea surface directly above the acoustic center of the array.

Table B-1 contains the positions of the Nanoose arrays at the time of their most recent survey dates. These are the positions of the acoustic centers in the range coordinate system. It is both useful and instructive to use our coordinate system superstructure and locate the hydrophones of each array in the range coordinate system.

Let α be the 3-vector locating the acoustic center of an array in the range coordinate system, see Table B-1. Let a be the location of one of the phones in cs(a) and let f,

$$f = (D/2) \begin{pmatrix} 1 \\ 1 \\ 1 \end{pmatrix}$$
 (B.8)

be the position of the acoustic center in cs(a). In the system cs(b) these vectors are Ba and Bf, and the location of the phone relative to the acoustic center is B(a-f). The position p of the phone in the range coordinate system is

$$p = \alpha + B(a - f) \tag{B.9}$$

Since the C-phone is the origin in cs(a), its range coordinate position is α – Bf. the result of this computation is in Table B-2.

TABLE B-1. NANOOSE ARRAY COORDINATES AND ORIENTATION ANGLES

				. 0	-		
Surv	ey Date	X	Y	Z	XTILT	YTILT	ZROT
AR 0	6/20/85	12188.01	-131.52	-1295.33	0.002909	0.014835	-0.208183
AR 1	6/20/85	19463.16	-174.99	-1308.76	0.061523	-0.036070	1.362579
AR 2	7/12/85	26991.39	-109.83	-1323.25	0.000145	0.005236	2.670336
AR 3	1/7/88	34505.10	-80.76	-1323.32	0.027925	-0.011345	2.928139
AR4	10/24/88	42005.19	-55.17	-1318.28	0.001164	-0.040288	-2.315877
4 D.C	< 100 /0E	10105 00	05.00	1015 50	0.000001	0.004050	4 (()505
AR 5	6/20/85	49497.00	-25.23	-1315.58	-0.000291	-0.004072	1.668535
AR 6	6/20/85	56972.28	-21.21	-1308.50	0.013817	0.041161	-0.703420
AR 7	7/30/85	64680.66	15.33	-1353.39	0.034907	0.022835	-0.574144
AR 8	11/16/88	71969.73	-29.28	-1300.89	-0.005963	-0.012217	-1.577341
AR 9	5/7/84	3.00	3.00	1.00	0.000000	0.000000	0.000000
AR 10	3/12/84	47100.00	-3600.00	-1300.00	0.000000	0.000000	0.000000
AR 11	7/18/85	23173.89	-6488.40	-1312.09	-0.004654	0.000436	2.784376
AR 12	6/20/85	30731.25	-6553.05	-1312.90	0.002036	0.001745	-3.042179
AR 13	6/20/85	38213.61	-6640.77	-1323.05	0.000291	0.006254	1.373522
AR 14	6/20/85	45647.07	-6513.18	-1324.78	0.001309	0.002327	-2.348044
	0, 20, 00	10017.07	0010.10	1021.70	0.001007	0.002027	2.010011
AR 15	6/19/85	53249.43	-6354.60	-1316.66	0.003345	0.004509	0.581544
AR 16	9/13/85	60859.74	-6356.07	-1313.42	0.014835	0.036943	2.303276
AR 17	6/16/87	68217.93	-6524.10	-1313.43	0.008290	0.034761	2.158449
AR 54	2/2/88	38029.95	5401.98	-1212.69	0.007709	-0.003782	-1.056919
AR 55	6/20/85	45645.75	6369.66	-1188.12	0.027634	0.039415	-0.728553
		50100 10	(115.07	1010.04	0.005505	0.040140	
AR 56	7/30/85	53180.13	6417.96	-1218.84	0.037525	0.048142	-1.392651
AR 57	7/30/85	60745.71	6419.40	-1088.24	0.006981	0.001891	-3.108606
AR 23	6/20/85	41605.14	-12150.18	-1268.23	002182	0.003200	-1.845214
AR 24	4/17/89	49572.00	-12966.00	-1300.00	-0.007272	0.055269	-1.343904
AR 25	10/24/88	56993.79	-12999.33	-1205.48	0.000291	-0.002182	-0.593726
AR 26	8/8/88	64442.94	-12971.04	-1255.35	-0.014835	-0.012654	3.134192
AR 27	7/15/80	22119.60	-15908.70	83.00	0.000000	0.000000	0.000000
AR 28	5/4/83	45000.00	1500.00	-1350.00	0.000000	0.000000	0.000000
AR 29	2/2/79	4 5000.00 0.00	0.00	0.00	0.000000	0.000000	0.000000
1 MAX 27		0.00	0.00	0.00	0.00000	0.00000	0.00000

(Distances in feet; angles in radians)

TABLE B-2. LOCATIONS OF THE C-HYDROPHONES

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AR	Date	XC	YC	ZC
0	6/20/85	12170.190	-143.316	-1310.105
1	6/20/85	19473.791	-192.189	-1325.075
2	7/12/85	27011.563	-103.354	-1338.287
3	1/7/88	34522.511	-69.103	-1338.681
4	10/24/88	42004.318	-33.398	-1332.413
5	6/20/85	49513.397	-38.633	-1330.630
6	6/20/85	56950.933	-23.551	-1323.123
7	7/30/85	64659.408	10.557	-1367.552
8	11/16/88	71954.918	-13.999	-1315.793
9	5/7/84	-12.000	-12.000	-14.000
10	3/12/84	47085.000	-3615.000	-1315.000
11	7/18/85	23193.258	-6479.598	-1327.004
12	6/20/85	30744.657	-6536.662	-1327.956
13	6/20/85	38225.375	-6658.513	-1337.943
14	6/20/85	45646.877	-6492.002	-1339.829
15	6/19/85	53245.085	-6375.441	-1331.552
16	9/13/85	60880.699	-6357.757	-1328.681
17	6/16/87	68238.605	-6528.792	-1328.448
54	2/2/88	38009.399	5407.725	-1227.510
55	6/20/85	45624.165	6367.887	-1202.470
56	7/30/85	53162.159	6429.360	-1233.742
57	7/30/85	60760.102	6434.858	-1103.370
23	3/20/85	41594.799	-12131.725	-1283.312
24	4/17/89	49554.120	-12955.612	-1315.728
25	10/24/88	56972.961	-13003.338	-1220.483
26	<u>9/8/99</u>	64458.271	-12955.966	-1269.935
27	7/15/80	22108.374	-15890.700	68.000
28	5/4/83	44985.000	1485.000	-1365.000
29	2/2/79	-15.000	-15.000	-15.000

APPENDIX C

2. RAY FITTING AND RAY TRACING

A "ping" sound source produces a wave front that travels through the water and is detectable by the receiving transducers. A ray is the path generated by the normal to the wave front. Ray tracing is the activity of following a ray from a receiver for a fixed amount of time in order to locate the sound source. Ray fitting is the activity of recreating the ray path from the positions of the source and the receiving sensor. The latter is needed to study the error performance of the former.

The speed of sound in water is assumed to vary with depth, but remain homogeneous in the horizontal. Thus a single depth-velocity profile is valid throughout the field.

We begin with some preliminaries followed by the development of a ray fitting algorithm, which may be viewed as an inverse to ray tracing. It is certainly more difficult. All paths are direct paths. That is, no provision is made for reflections or refractions that produce non monotone rays connecting source and receiver. Also our interest lies in the greater ranges and no adjustments are presented for sources that are directly above the receiver.

The notation is chosen to be consistent with that used in the Fortran source codes. We are given a speed-depth profile in the form of pairs lm_i and vel_i.

 $lm_i = depth of ith water layer, positive down.$ $vel_i = speed of sound at the depth lm_i.$ Digression. The water speed processing system at NUWES produces average velocity values for a large number of equally spaced points. These averages are intended to serve as constant values for the entire layer, eq. (3.1). Thus the information is provided in the form, for i = 1, ..., m

l_i = depth of water layer boundary
 vel_i = velocity constant for layer (l_i, l_{i+1})

The values used for lm_i above are the layer midpoints:

 $lm_i = .5 \times (l_i + l_{i+1})$ for i=1, ..., m-1

The algorithms developed below do not require layers of equal thickness. Thus they can accommodate the user who wants to use thin layers at depths of rapid change and thick layers at depths of slow changes. All computations should be in double precision arithmetic.

Occasionally as application leads to a receiver whose depth is greater than the largest {lm_i}. For these cases we have been using an extrapolation of the sound speed profile that adjoins a sufficient number of layers, all of the same thickness as the deepest of the original layers. The corresponding velocities are extrapolated using a second order polynomial calculated from a fixed second difference whose value is the average of the four deepest second differences. (Use of the coefficient of the quadratic term in a least square fit has also been employed as an option.)

Since ray fitting is a rather delicate operation we use the isogradient technique. I.e., straight lines are fitted to the speed for each layer; the constant gradient of slope is computed; the profile itself is then a continuous function of connected straight line segments. See Figure 2. I.e., if

$lm_i \le z \le lm_{i+1}$

then

$$vel(z) = v_0(i) + v_1(i)z$$

where

$$v_0(i) = lm_{i+1} \cdot vel_i - lm_i \cdot vel_{i+1}$$
$$v_1(i) = (vel_{i+1} - vel_i)/dz_i$$

 $dz_i = lm_{i+1} - lm_i$

Typically the values of $\{v_0(i)\}\$ are positive and large compared to the $\{v_1(i)\}\$ which are small and of either sign or zero. Snell's law comes into play and the ray invariant will be denoted rv. The ray path in each layer will be a circle arc $(v_1(i)\neq 0)$ or a straight line $(v_1(i)=0)$.

Now we are positioned to describe our ray fitting algorithm. It is two dimensional (horizontal, vertical) and given the endpoints

(a1,a2) receiver(p1,p2) sound source

the goal is to compute

- θ_0 entrance angle at the receiver
- θ_1 exit angle at the sound source
- t transit time of sound from source to receiver

There is no loss in placing the origin on the water surface directly over the receiver. Thus, $a_1 = 0$ and $p_2 < a_2$ (depth of receiver).

The algorithm is an iterative one that operates as follows. Initialize the entrance angle θ_0 . Use this angle to ray trace upward through the layers until the vertical value of p_2 is achieved. Compute the current horizontal value h,

and compare it with p_1 . If this is within a preassigned small number, ε , stop

Figure C-1. Details of single layer processing

```
\begin{aligned} \mathbf{v}(z) &= \mathbf{v}_0 + \mathbf{v}_1 z \quad (\text{linear profile}) \quad z_1 < z < z_0 \\ \text{Given: } \theta_0, z_0, z_1, \mathbf{v}_0, \mathbf{v}_1, \text{ and} \\ & \text{the ray invariant } \mathbf{rv} = \frac{\cos(\theta_0)}{\mathbf{v}(z_0)} \end{aligned}\begin{aligned} \text{Case } \mathbf{v}_1 &= 0 \\ & \text{d}z = z_1 - z_0 \\ & \text{d}w = \text{d}z/\sin(\theta_0) \\ & \text{h}_1 &= \text{h}_0 + \text{d}w \cdot \cos(\theta_0) \\ & \text{d}t &= \text{d}w/\mathbf{v}_0 \end{aligned}
```

 $\cos(\theta_1) = \mathbf{r}\mathbf{v} \cdot \mathbf{v}(z_1)$

Case $\mathbf{v}_1 \neq \mathbf{0}$

$$q_{2} = -v_{0}/v_{1}$$

$$s = \sin(\theta_{0})$$

$$c = \cos(\theta_{0})$$

$$q_{1} = h_{0} + (q_{2} - z_{0})s/c$$

$$r = \operatorname{signum}(q_{2})(q_{2} - z_{0})/c$$

$$h = q_{1} - \operatorname{signum}(q_{2}) r \cdot \sin(\theta)$$

$$dt = \frac{1}{v_{1}} \int_{\theta_{0}}^{\theta_{1}} \frac{d\theta}{\cos(\theta)}$$

$$= \frac{1}{v_{1}} \ln\left(\frac{1 + \sin(\theta)}{\cos(\theta)}\right)/_{\theta_{0}}^{\theta_{1}}$$

$$\cos(\theta_1) = \mathbf{r}\mathbf{v} \cdot \mathbf{v}(z_1)$$

and compute the transit time. Otherwise, adjust the entrance angle θ_0 according to the ratio of the current rise over run, $\frac{h}{a_2 - p_2}$, to the desired one, $\frac{p_1}{a_2 - p_2}$, and repeat the algorithm using the new initialization.

ALGORITHM RAYFIT

Initialize

(i) Determine the layers that contain the source and receiver. Choose j, n so that

$$lm_j \leq p_2 < lm_{j+1}$$

 $lm_n \leq a_2 < lm_{n+1}$

(ii) Make thickness corrections for the extreme layers

 $dz_j = lm_{j+1} - p_2$

 $dz_n = a_2 - lm_n$

(iii) Compute the sound speed at depths a_2 , p_2

 $\mathbf{va}_2 = \mathbf{v}_0(\mathbf{n}) + \mathbf{v}_1(\mathbf{n}) \cdot \mathbf{a}_2$

 $\mathbf{vp}_2 = \mathbf{v}_0(\mathbf{j}) + \mathbf{v}_1(\mathbf{j}) \cdot \mathbf{p}_2$

(iv) Unless a "previous" value for θ_0 is available, fit a straight line through the depth-velocity profile in the range (p₂, a₂) and use θ_0 corresponding to the circle arc (or line) of that approximate profile. I.e.,

If $va_2 = vp_2$ then $\theta_0 = tan^{-1}((a_2 = p_2)/p_1)$, else

$$q_2 = \frac{va_2 \cdot p_2 - vp_2 \cdot a_2}{(a_2 - p_2)}$$

$$q_1 = .5 \cdot p_1 + .5 \cdot (p_2 - a_2)(p_2 + a_2 - 2q_2)/p_1$$

$$\theta_0 = \tan^{-1} \{q_1/(q_2-a_2)\}$$

endif

Set initial values for iteration

A.
$$i = n$$
, $s = sin(\theta_0)$, $c = \sqrt{1-s^2}$
 $rv = c/va_2$, $h = 0$, $z = a_2$

Main raytracing code

If $v_1(i) = 0$, then Β. $dw = dz_i/s$ $\mathbf{h} = \mathbf{h} + \mathbf{c} \cdot \mathbf{d}\mathbf{w}$ else $q_2 = -v_0(i)/v_1(i)$ $q_1 = h + (q_2 - z) \cdot s/c$ $\mathbf{r} = \sqrt{(h - q_1)^2 + (z - q_2)^2}$ $c = rv \cdot vel(i)$ $s = \sqrt{1-c^2}$ $h = q_1 - signum(q_2) \cdot r \cdot s$ endif goto TEST If i = j, $z = lm_i$ i = i - 1goto B TEST: $\theta_1 = \cos^{-1}(\mathbf{rv} \cdot \mathbf{vp}_2)$ If $v_1(j) \neq 0$ $h=q_1-\text{signum}(q_2) \cdot r \cdot \sin(\theta_1)$ If $|h-p_2| < \epsilon$, goto FINI

Re-estimate θ_0

$$\theta_o = \tan^{-1} \{ \tan(\theta_0) \cdot h/p_1 \}$$

goto A

FINI:

 $ang(j) = \theta_i$ $ang_{n+1} = \theta_0$ $ang_i = \cos^{-1}(rv \cdot vel(i)) \text{ for } i = j+1, ..., n$ Compute transit time

$$\begin{split} t &= 0\\ Do \ TC \quad i = j, \ n\\ If \ v_1(i) &= 0\\ t &= t + dz_i / (v_0(i) \cdot sin(ang_i))\\ else\\ t &= \frac{1}{v_1(i)} \ ln \Big\{\!\! \frac{\cos(ang_{i+1})(1+sin(ang_i))}{(1+sin(ang_{i+1}))\cos(ang_i)}\!\! \}\\ end if \end{split}$$

TC continue

Remove extreme layer thickness corrections

```
dz_i = lm_{j+1} - lm_jdz_n = lm_{n+1} - lm_nend
```

This algorithm has been quite useful to us. Using $\epsilon = 10^{-6}$ we typically have 8 to 10 iterations through A.

Raytracing algorithms are less sensitive, but since a good one is readily available by merely modifying the above, let us do that. The process is an inverse one in that the goal is to compute p_1 , p_2 given θ_0 and t_0 , where t_0 is the transit time.

ISOGRAD

Initialize by locating the layer containing the receiver, establishing the ray invariant, etc.

```
Choose n so that lm_n \le a_2 < lm_{n+1}

i = n, \quad h = 0, \quad z = a_2, \quad t = 0

s = sin(\theta_0), \quad c = cos(\theta_0)

va_2 = v_0(i) + v_1(i) \cdot a_2

rv = c/va_2

dz_n = a_2 - lm_n
```

AA: If $v_1(i) = 0$, then

$$dw = dz_i/s$$
$$dt = dw/v_0(i)$$
$$h = h + c \cdot dw$$

else

$$q_{2} = -v_{0}(i) / v_{1}(i)$$

$$q_{1} = h + (q_{2} - z) \cdot s / c$$

$$r = \sqrt{(h - q_{1})^{2} + (z - q_{2})^{2}}$$

$$cp = rv / v_{1}(i)$$

$$sp = \sqrt{1 - cp^{2}}$$

$$dt = \frac{1}{v_{1}(i)} \ln \left\{ \frac{c}{1 + s} \frac{1 + sp}{cp} \right\}$$

$$h = q_{1} - signum(q_{2}) sp$$

en

$$t = q_1 - signalit(q_2) sp$$

dif
$$t = t + dt$$

If $t \ge t_0$ goto FINAL
$$z = lm_i$$

$$s = sp$$
 $c = cp$ $i = i-1$ goto AA

.

FINAL

 $dt = t_0 + dt - t$ If $\mathbf{v}_1(\mathbf{i}) = 0$ $dz_i = v_0(i) \cdot dt$ $dw = dz_i/sp$ $p_1 = h + dw \cdot cs$ $p_2 = z - dz_i$

else

$$x = \exp\{dt \cdot v_1(i)\}(1+s)/c$$

$$cp = 2 \cdot x/(1+x^2)$$

$$sp = \sqrt{1-cp^2}$$

$$p_1 = q_1 + r \cdot sp$$

$$p_2 = q_2 + r \cdot cp$$

endif Restore layer integrity

 $dz_i = lm_{i+1} - lm_i$ $dz_n = lm_{n+1} - lm_n$

Compute exit angle

 $\theta_1 = \cos^{-1}(cp)$

APPENDIX D. EFFECT OF SOUND SPEED OSCILLATION WITHIN A LAYER OF WATER

Suppose a water layer of thickness Δ_z has an average sound speed of v feet per second. Suppose further that the speed profile within the layer is an oscillation of frequency K and amplitude δ . We address the question of how this profile can affect the nominal calculation of the ray's horizontal distance and transit time through the layer.

The question is most easily treated using isospeed ray tracing and modeling the oscillations as follows: Partition the layer into 2K equithick sublayers and assume the sound speed alternates between the values $\nu+\delta$ and $\nu-\delta$ as we move through these layers. Let θ_1 be the elevation angle for the $\nu+\delta$ speed layers and θ_2 for the $\nu-\delta$ layers.

According to isospeed ray tracing formulation, the horizontal distance advanced in the layer is

$$H = \frac{\Delta_z}{2K} \sum_{j=1}^{2K} \cot(\theta_j) = \frac{\Delta_z}{2} [\cot(\theta_1) + \cot(\theta_2)]$$

and the transit time

$$T = \frac{\Delta_z}{2K} \sum_{1}^{2K} \frac{1}{\upsilon_j \sin(\theta_j)}$$
$$= \frac{\Delta_z}{2} \left[\frac{1}{(\upsilon + \delta)\sin(\theta_1)} + \frac{1}{(\upsilon - \delta)\sin(\theta_2)} \right]$$

The nominal values are found by using v as the speed and θ_1 as the angle throughout the layer. Thus the error in these two values is

$$\Delta H = \frac{\Delta_z}{2} \left[\cot(\theta_2) - \cot(\theta_1) \right],$$

$$\Delta T = \frac{\Delta_z}{2} \left[\frac{1}{(\nu + \delta) \sin(\theta_1)} + \frac{1}{(\nu - \delta) \sin(\theta_2)} - \frac{2}{\nu \sin(\theta_1)} \right],$$

and θ_1 and θ_2 are related by the ray invariant equation

$$\frac{\cos(\theta_1)}{\upsilon+\delta} = \frac{\cos(\theta_2)}{\upsilon-\delta}$$

The error ΔH is affected by velocity only through this equation. Notice that the errors do not depend upon the frequency K. Using δ/v as the proportion of the speed appearing in the amplitude, we can rewrite

$$\Delta T = \frac{\Delta_z}{2\upsilon} \left[\frac{1}{(1+\delta/\upsilon)\sin(\theta_1)} + \frac{1}{(1-\delta/\upsilon)\sin(\theta_2)} - \frac{2}{\sin(\theta_1)} \right]$$
$$= \frac{\Delta_z}{2\upsilon} \left[\frac{1-\delta/\upsilon}{\sin(\theta_1)} + \frac{1+\delta/\upsilon}{\sin(\theta_2)} - \frac{2}{\sin(\theta_1)} \right]$$
$$= \frac{\Delta_z}{2\upsilon} \left[\left(\frac{1}{\sin(\theta_2)} + \frac{1}{\sin(\theta_1)} \right) (1+\delta/\upsilon) \right]$$

since the proportion δ/υ is believed small.

Some speculative calculations appear in Table (D-1) for $\Delta_z = 5$ feet and $\upsilon = 4800$ feet/second.

The calculation is not very sensitive to values of v, but quite responsive to the elevation angle. It should be noted that the signs of ΔH , ΔT change if the modeled oscillations are in reverse order. This is equivalent to replacing with $-\delta$.

5

This author is not qualified to judge the reality of the suggested values of δ/υ . Certainly the question deserves more attention. It is common to process

some 200 of these 5 foot layers in a ray computation. The error buildup, even if the signs change randomly, could be significant, perhaps 15 times ΔH .

TABLE D-1. EFFECT OF OSCILLATION ($v = 4800$ feet/second; $\Delta_z = 5$ feet)					
δ/υ	θ1	ΔН	ΔΤ		
.0001	0.1	0.48	0001006		
	0.15	0.15	0000301		
	0.2	0.06	0000127		
	0.3	0.02	0000037		
	0.4	0.01	0000015		
.0005	0.1	2.18	0004517		
	0.15	0.70	0001432		
	0.2	0.30	0000615		
	0.3	0.009	0000181		
	0.4	0.04	0000074		
.001	0.1	3.87	0008032		
	0.15	1.31	0002699		
	0.2	0.58	0001189		
	0.3	0.18	0000357		
	0.4	0.08	0000147		

The support for this calculation proceeds as follows: Suppose X is a random variable uniformly distributed on the interval ($-\Delta H, \Delta H$). Then the mean of X is zero and the variance is $(\Delta H)^2/3$. If there are n layers to be r ocessed then the error in determining the horizontal distance is the sum of N independent and identical such X's. It will have mean zero and standard deviation $\Delta H\sqrt{n/3}$; zero plus or minus two standard deviations could be a significant amount, especially for small elevation angles.

APPENDIX E. CONVERSION OF FOUR TRANSIT TIMES TO INPUTS FOR THE RAY TRACING ALGORITHM

The method in current use for converting the four transit times $(t_1, ..., t_4)$ into an azimuth angle, ϕ_c , an elevation angle, θ_c and t_{ac} an estimated transit time from the source to the acoustic center is outlined and critiqued below.

The two angles ϕ_c and θ_c are generated from a description that assumes a constant value v for the speed of sound for all points in the array. Then the concept of an "apparent position," (X,Y,Z) relative to the acoustic center for the sources in a constant speed median is utilized for purposes of estimating the two angles. The apparent position must satisfy the system of equations

$$(X + D/2)^{2} + (Y - D/2)^{2} + (Z - D/2)^{2} = v^{2}t_{1}^{2}$$

$$(X - D/2)^{2} + (Y + D/2)^{2} + (Z - D/2)^{2} = v^{2}t_{2}^{2}$$

$$(X - D/2)^{2} + (Y - D/2)^{2} + (Z + D/2)^{2} = v^{2}t_{3}^{2}$$

$$(E.1)$$

$$(X - D/2)^{2} + (Y - D/2)^{2} + (Z - D/2)^{2} = v^{2}t_{4}^{2}$$

This is a system of four equations in three unknowns. Unless the values of t_1 , ..., t_4 are singularly coherent, there are an infinity of solutions for X, Y, Z.

The current policy is to obtain a unique solution by subtracting the fourth equation successively from each of the other three, leaving a three by three system remaining. I.e.,

$$2XD = v^{2}(t_{1}^{2} - t_{4}^{2})$$

$$2YD = v^{2}(t_{2}^{2} - t_{4}^{2})$$

$$2ZD = v^{2}(t_{3}^{2} - t_{4}^{2})$$

(E.2)

and a unique solution. Other rationales supporting this approach can be found in [5].

At this point an adjustment is made under the name of a direction cosine correction (DCC). The quantity D/2 is added to each component of the apparent position (in effect making the position relative to the C-phone) and the length of the new vector is computed and divided by vt_4 . Call this

$$DCC = \sqrt{(X + D/2)^2 + (Y + D/2)^2 + (Z + D/2)^2} / vt_4$$
(E.3)

The translated values are normalized by DCC prior to returning the origin to the acoustic center:

$$X_{c} = \frac{X + D/2}{DCC} - D/2$$

$$Y_{c} = \frac{Y + D/2}{DCC} - D/2$$

$$Y_{c} = \frac{Y + D/2}{DCC} - D/2$$
(E.4)

Next come the tilt corrections, required because the corrected apparent position is in the coordinate system cs(ac), see Appendix B. This is accomplished by applying the approximation of eq. (B.4), i.e., that coefficient matrix (instead of the exact one (B.5)) to the vector (X_c , Y_c , Z_c): The Z rotation can also be applied at this point, i.e., multiplying by ρ_3 . Let us call the result of all this (X,Y,Z) and form the functions

$$\sin(\theta_{c}) = Z / \sqrt{X^{2} + Y^{2} + Z^{2}}$$

$$\sin(\phi_{c}) = Y / \sqrt{X^{2} + Y^{2}}$$

$$\cos(\phi_{c}) = X / \sqrt{X^{2} + Y^{2}}$$

(E.5)

and, from these, the angles θ_c and ϕ_c can be found.

Finally, there is need to construct a transit time to the acoustic center because it is not measured. Calling the values t_{ac} , the proportionality adjustment with t_4 is used.

$$t_{ac} = t_4 \sqrt{X^2 + Y^2 + Z^2} / \sqrt{(X + D/2)^2 + (Y + D/2)^2 + (Z + D/2)^2}$$
(E.6)

Ś.

APPENDIX F. ALTERNATIVE INITIALIZATION OF RAY TRACING

Current methodology utilizes the four measured ray transit times (t_1 , t_2 , t_3 , t_4) from source to the X, Y Z, C hydrophones (respectively), and converts them to ϕ_c , θ_c , and t_{ac} , the azimuth and elevation angles and an estimated transit time from source to the acoustic center. It is seen in Section 5 that there can be considerable error in this, especially for ϕ_c . Our alternative is to shift the acoustic center to the C-phone and to ignore t_3 . The resulting angles will have much smaller error, and the transit time t_4 is used directly.

We proceed to develop the ray azimuth (spherical coordinate longitude) and elevation (spherical coordinate latitude) angles from times t_1 , t_2 , t_4 , in the range coordinate system. (See Appendix B for a general discussion of coordinate systems.) The conversion of those times is influenced by the array orientation.

Let s_i,c_i be the sine and cosine of the ith Euler angle in the orientation of the 3–D sensor array j; i = 1, 2, 3. The conversion of XTILT, YTILT, and ZROT into Euler angles is explained in Appendix B, together with methodology for locating the positions of the phones at the ends of the array arms.

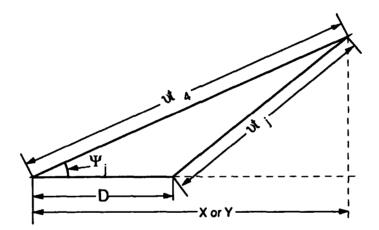
Having moved the acoustic center to the C-phone, we are using the coordinate system cs(a). Letting B_j represent the jth column of the matrix B, eq. (B.3), the locations of the X,Y,Z phones are, respectively,

$$D \cdot B_1$$
 $D \cdot B_2$ $D \cdot B_3$

and, of course, the C-phone is at the origin. These locations in the range coordinate system can be found by adding the location of the C-phone as indicated in Table B-2.

Typically there is very little difference in the depths (3^{rd} component) of the X,Y, and C-phones. Hence there is small variability if the speed of sound at these three depths and the use of a constant common value, say v, is more tenable than the previous use.

Again we adopt the concept of an apparent position (X,Y,Z) of the sound source and the first two direction cosines can be found by applying the law of cosines (see the figure) for j = 1, 2.



where t_1 , t_2 , t_4 are the signal transit times to the X, Y and C phones, respectively.

For j = 1, the cosine of the marked angle is $X/\upsilon t_4$ and by the law of cosines

$$(v \cdot t_1)^2 = D^2 + (v t_4)^2 - 2D \cdot t_4 \cos(\Psi_1)$$

It follows that

$$X = \frac{D}{2} + (\upsilon \cdot t_4 - \upsilon \cdot t_1)(\upsilon \cdot t_4 + \upsilon \cdot t_1)/(2D)$$

and a like equation for j = 2, leading to

$$Y = D/2 + (vt_4 - vt_2) (vt_4 + vt_2) / 2D.$$

The third direction cosine is obtained from

$$\cos(\Psi_3) = \sqrt{1 - \cos^2(\Psi_1) - \cos^2(\Psi_2)} = \sqrt{1 - (X^2 + Y^2) / \upsilon^2 t_4^2}$$

Next we rotate these values into alignment with the test range coordinate system, i.e., apply the matrix B, eq. (B.3)

$$\begin{bmatrix} X_c \\ Y_c \\ Z_c \end{bmatrix} = B \begin{bmatrix} X \\ Y \\ Z \end{bmatrix}$$
(F.1)

The proposed ray elevation (θ) and azimuth (ϕ_p) angles can be found from

$$\cos(\theta_p) = Z_c / \sqrt{X_c^2 + Y_c^2 + Z_c^2}$$
 (F.2)

and

$$\sin(\phi_p) = Y_c / \sqrt{X_c^2 + Y_c^2}$$

$$\cos(\phi_p) = X_c / \sqrt{X_c^2 + Y_c^2}$$
(F.3)

Comment. This technique was designed to treat the speculated shortcomings of the procedure currently in use. The results are quite successful, especially in reducing azimuth error. It has the curious property of not using information provided by the Z-phone. It seems wasteful not to use this information. Yet any system that does use it must be required to perform at least as well as this one.

Appendix G.

SUBROUTINE ERRCUR(H1,Z1,K,A2,XTILT,YTILT,ZROT,D,L,VEL,M, * THR.THC.THER.PHR.PHC.PHER.HC.HER.ZC.ZER,TEE,TIMER) C* С 08/10/90 С COMPUTES THE TRUE ELEVATION ANGLES AND THE ESTIMATED ELEVATION С ANGLES FOR K SOUND SOURCE DISTRIBUTED EQUALLY ON A CIRCLE OF RADIUS С H1 AT DEPTH Z1 FOR A RECIEVING ARRAY AT DEPTH A2 BUT OTHERWISE AT THE С CENTER OF THE CIRCLE. THE ACCOUSTIC CENTER IS THE GEOMETRIC CENTER OF THE ARRAY CUBE. THE CURRENTLY USED METHODOLOGY IS APPLIED. C C \mathbf{C}^{\prime} ************************* **INPUTS: RADIUS OF CIRCLE IN FEET** H1: DEPTH OF SOUND SOURCE IN FEET Z1: NUMBER OF SOURCES ON THE CIRCLE K: DEPTH OF THE C-HYDROPHONE A2: XTILT, YTILT, ZROT: ORIENTATION INFORMATION ABOUT THE SENSING ARRAY (RADIANS) D: LENGTH OF ARRAY EDGES. DEPTH OF LAYER BOUNDARIES. Ŀ NUMBER OF RECORDS IN THE VELOCITY DEPTH PROFILE. **M**: AVERAGE SPEED OF SOUND IN THE LAYERS. VEL: **OUTPUTS:** THR: **ACTUAL ELEVATION ANGLE (THETA)** THETA ESTIMATES AT ACCOUSTIC CENTER, CURRENT METHOD THC: THER: THETA ERROR AT THE ACCOUSTIC CENTER, CURRENT METHOD PHR: ACTUAL AZIMUTH ANGLE (PHI). PHC: PHI ESTIMATES AT ACCOUSTIC CENTER, CURRENT METHOD. PHER: PHI ERROR AT ACCOUSTIC CENTER. HORIZONTAL ESTIMATE, CURRENT METHOD. HC: HER: HORIZONTAL ERROR. VERTICAL ESTIMATE, CURPENT METHOD. ZC: ZER: VERTICAL ERROR. TRANSIT TIME TO ACCOUSTIC CENTER. T: TIMC: ESTIMATE OF TIME TO ACCOUSTIC CENTER, CURRENT METHOD. TIMER: TRANSIT TIME ERROR. *********** DIMENSION B(5,3), DZ(300), THER(30), HD(5), HC(30), ZC(30) DIMENSION L(300), LM(300), LL(300), PX(30), PY(30), T(5), TEE(30) DIMENSION THC(30),THR(30),V0(300),V1(300),VEL(300),VV(300) DIMENSION HX(5), HY(5), HER(30), ZER(30), TIMER(30) DIMENSION X0(3).PHC(30).PHR(30).PHER(30),TIMC(30) REAL*8 A1,A2,A2M,ANG,B,C1,C2,C3,D,DP,DW,DZ,THER,GG REAL*8 H,H0,H1,HD,L,LM,LL,LPZ,PIE,P1,P2,PX,PY,S1,S2,S3 REAL*8 T,T0,THEC,TH0,TH1,THC,THR,HC,ZC,HER,ZER REAL*8 V,V0,V1,VEL,VV,X,X0,XTILT,Y,YTILT,Z,Z0,Z1,ZROT REAL*8 CC1,CC2,VV1,VV0,HX,HY,SC,DCC,RAC,RC,TEE

PIE = 3.14159265359D0

REAL*8 CAZ,SAZ,PHC,PHR,PHER,LH,C,T3P,TIMC,TIMER REAL*8 SR,SRER,LB,SV,SVU2,SVU,SU2,SU4,G1,G IEST = 0M1 = M

C DISTRIBUTE THE K SOURCES EQUALLY AROUND THE CIRCLE COUNTER C CLOCKWISE FROM THE EAST. DO 10I = 1.KANG = 2*PIE*(I-1)/KPX(I) = H1*DCOS(ANG) $PY(I) = H1^*DSIN(ANG)$ PHR(I) = ANGIF(ANG.GT.PIE) PHR(1) = ANG - 2*PIE **10 CONTINUE** C FORM SINES AND COSINES OF ALL THE EULER ANGLES: ROLL, PITCH, YAW S2 = DSIN(XTILT) $C2 = DSORT(1 - S2^{++}2)$ S1 = DSIN(YTILT)/C2 $C1 = DSORT(1 - S1^{**}2)$ S3 = -DSIN(ZROT)C3 = DCOS(ZROT)C IN THE COORDINATE SYSTEM HAVING CENTER AT THE C-HYDROPHONE C AND POSITIVE-UPWARD, THE LOCATIONS OF THE FOUR HYDROPHONES (RELATIVE TO THE ARM LENGTH D) ARE DEVELOPED NEXT. THIS IS C THE TRANSPOSE OF THE MATRIX B IN APPENDIX B. $B(1,1) = C2^*C3$ B(1,2) = C2*S3B(1,3) = S2B(2,1) = -S1*S2*C3 - C1*S3B(2,2) = -S1*S2*S3 + C1*C3 $B(2,3) = S1^{*}C2$ B(3,1) = -C1*S2*C3 + S1*S3B(3,2) = -C1*S2*S3 - S1*C3B(3,3) = C1*C2C LIKE NOTATION WILL BE USED TO LOCATE THE C-HYROPHONF AND THE C ACCOUSTIC CENTER (AC). DO 12 J = 1.3B(4,I) = 0.0D0 $B(5,J) = 0.5^{*}(B(1,J) + B(2,J) + B(3,J))$ 12 CONTINUE A1 = 0.0D0P2 = Z1C LOCATE THE HYDROPHONE HORIZONTAL COMPONENTS IN THE COORDINATE C SYSTEM CENTERED AT AC. DO 14 J = 1.5 $HX(J) = D^{*}(B(J,1) - B(5,1))$ $HY(J) = D^*(B(J,2) - B(5,2))$ 14 CONTINUE C DETERMINE THE DEPTHS OF THE FOUR HYDROPHONES AND THE AC. $HD(1) = A2 + D^{*}(B(5,3) - B(1,3))$ $HD(2) = A2 + D^{*}(B(5,3) - B(2,3))$ $HD(3) = A2 + D^{*}(B(5,3) - B(3,3))$ $HD(4) = A2 + D^{*}(B(5,3) - B(4,3))$ HD(5) = A2

```
C FIND THE DEEPEST HYDROPHONE
     A2M = 0.D0
     DO 51 J=1,4
        IF(HD(J).GT.A2M) A2M = HD(J)
    CONTINUE
51
C FORM THE SET OF LAYER MIDPOINTS.
     DO 105 I = 1, M-1
        LM(I) = .5^{*}(L(I) + L(I+1))
105 CONTINUE
     LM(M) = LM(M-1) + L(M) - L(M-1)
   FORM DEPTH INCREMENTS, AND ALL SOUND VELOCITY SLOPL
С
   INTERCEPTS.
С
     DO 110 I=1,M-1
         DZ(I)=LM(I+1)-LM(I)
         VO(I) = (LM(I+1)*VEL(I) - LM(I)*VEL(I+1))/DZ(I)
         V1(I)=(VEL(I+1)-VEL(I))/DZ(I)
110 CONTINUE
С
     IF(A2M.LT.LM(M-2)) GOTO 126
C IF A2M IS DEEPER THAN THE LAST LAYER MIDPOINT, THEN WE EXTRAPOLATE
C THE SOUND VELOCITY PROFILE BY USING A QUADRATIC FUNCTION OVER THE
C DEEPEST 100 FEET.
C FIRST COUNT THE NUMBER OF LAYERS (OF THICKNESS DZ(M-2)) TO
C BE ADIOINED. ALSO MUST EXTEND THE L ARRAY.
     K0 = 2 + MAX(0,NINT((A2M-LM(M-1))/DZ(M-1)))
     MC = 21
C FIND THE AVERAGE DEPTH FOR THE LAST 100 FEET
     LB = 0.0D0
     DO 200 J = M+1-MC,M
         LB = LB + LM(I)
200
     LB = LB/MC
C FORM SUM OF POWERS AND PRODUCTS.
     SV = 0.0D0
     SVU2 = 0.0D0
     SVU = 0.0D0
     SU2 = 0.0D0
     SU4 = 0.0D0
     G1 = 0.0D0
     DO 210 J = M+1-MC,M
         U = LM(J) - LB
         SV = SV + VEL(J)
         SVU = SVU + U^*VEL(J)
         SVU2 = SVU2 + U^{**2} * VEL(J)
         SU2 = SU2 + U^{**2}
         SU4 = SU4 + U^{**}4
         G1 = G1 + V1(j)
210 CONTINUE
     G1 = G1/MC
     G = SVU/SU2
     GG = (MC^{*}SVU2 - SU2^{*}SV)/(SU4^{*}MC - SU2^{**}2)
```

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79
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IF(GG.LT.0.0D0) GG = 0.0D0IF(V1(M-1).LT.0.0D0) V1(M-1) = G1C PERFORM THE EXTRAPOLATION. DO 125 I=M,M+K0 $V1(I) = V1(I-1) + GG^*DZ(M-1)$ LM(I+1) = LM(I) + DZ(M-1)VEL(I+1) = VEL(I) + DZ(M-1)*V1(I)VO(I) = (LM(I+1)*VEL(I) - LM(I)*VEL(I+1))/DZ(M-1)L(I+1) = L(I) + DZ(M-1)DZ(I) = DZ(M-1)125 CONTINUE C UPDATE M, THE NUMBER OF LAYERS M = M + K0126 CONTINUE C THE OUTER LOOP WILL PERFORM COMPUTATIONS FOR THE K SOUND C SOURCES. C ADJUST DV TABLE TO 25 FT. INCREMENTS. CALL VELMOD(L, VEL, M1, LL, VV, MM) C RAYFITTING C THE INNER LOOP WILL FIT RAYS TO THE FOUR HYDROPHONES C AND THE AC IN THE ORDER X.Y.Z.C AND AC AND THE AC IN THE ORDER X,Y,Z,C AND AC. DO 501 = 1.KWRITE(*,*)' OUTER LOOP I = ',I,' K = ',KDO 35 J = 1,5 $P1 = DSQRT((PX(I)-HX(J))^{**}2 + (PY(I)-HY(J))^{**}2)$ Z0 = HD(I)CALL RAYFIT1(A1,Z0,P1,P2,M,VEL,LM,DZ,V0,V1,T0,TH0, TH1.IEST) C COLLECT THE FIVE TRANSIT TIMES. T(J) = TOC IN THIS PROGRAM WE KEEP ONLY THE TRUE ELEVATION ANGLE AT AC. THR(I) = THO35 CONTINUE C INNER LOOP COMPLETED. C LOCATE THE WATER LAYER. N. CONTAINING THE ARRAY. USE THIS C TO DEVELOP THC, THE CURRENTLY USED ESTIMATE OF THO. N = MMDO 37 J = 2,MM IF((LL(J-1),LE,A2),AND.(LL(J),GT,A2)) N = J-137 CONTINUE С V = VV(N)C USE THE FOUR TRANSIT TIMES TO PRODUCE ESTIMATES OF THE C ENTRANCE ANGLE. CALCULATE THE PRE-TILT CORRECTED APPARENT C POSITION AND INCLUDE THE DIRECTION COSINE CORRECTION. DO 40 J = 1.3

 $X0(J)=((V^{T}(4)-V^{T}(J))^{*}(V^{T}(4)+V^{T}(J)))/(2^{D})$ $X0(J) = 0.5^{D} + X0(J)$ NITINUE

40 CONTINUE

 $SC = V^*T(4)$ DCC = (DSQRT(X0(1)**2 + X0(2)**2 + X0(3)**2))/SC DO 41 J = 1,3 X0(J) = (X0(J)/DCC) - 0.5*D

- 41 CONTINUE
- C NEXT MAKE TILT CORRECTIONS $X = X0(1) - X0(3)^*DSIN(XTILT)$ $Y = X0(2) - X0(3)^*DSIN(YTILT)$ $Z = X0(3) + X0(1)^*DSIN(XTILT) + X0(2)^*DSIN(YTILT)$

RAC = DSQRT(X**2 + Y**2 + Z**2) RC = DSQRT((X + D/2)**2 + (Y + D/2)**2 + (Z + D/2)**2) T0 = T(4)*RAC/RC TIMC(I) = T0

- C PERFORM Z ROTATION IN THE (X,Y) PLANE. XO(1) = C3*X - S3*YXO(2) = S3*X + C3*Y
- C COMPUTE THEC: THE ESTIMATE OF THETA CURRENTLY IN USE. THEC = DASIN(Z/DSQRT(X**2 + Y**2 + Z**2))
- C COMPUTE SINES AND COSINES OF AZIMUTH. $SAZ = X0(2)/DSQRT(X^{*2} + Y^{*2})$ $CAZ = X0(1)/DSQRT(X^{*2} + Y^{*2})$ PHC(I) = DATAN2(SAZ,CAZ) PHER(I) = PHC(I) - PHR(I) IF(ABS(PHER(I)).GT.PIE) PHER(I) = PHC(I) + PHR(I)IFG = 0
- C RAYTRACE BY THE ISOSPEED METHOD. CALL ISOSPEED(A1,Z0,T0,THEC,LL,VV,MM,H,Z,TH1)

HC(I) = HZC(I) = Z

C NOW FINISH THE OUTPUT. THC(I) = THEC

C AND THE ERRORS TEE(I) = T(5) THER(I) = THC(I) - THR(I) TIMER(I) = TIMC(I) - T(5) HER(I) = HC(I) - H1 ZER(I) = ZC(I) - Z1 SR = DSQRT(H1**2 + (A2 - Z1)**2) SRER = DSQRT(HC(I)**2 + (A2 - Z1)**2) - SR 50 CONTINUE

C OUTER LOOP COMPLETED!

RETURN

- 100 FORMAT(3(5X,E13.6))
- 120 FORMAT(3(5X,F15.12))
- 130 FORMAT(3(F12.8,2X))
 - END

C ***** ********* С LIBRARY FILE: LIB12.FOR 2/22/90 Č C **************** ****************** SUBROUTINE ISOGRAD1(A1.A2.T0.TH0.N.LM.VEL.V0.V1.DZ.H.Z.TH1) C* 00000000 09/25/89 T0 : TRANSIT TIME (SEC). TH0: ELEVATION ANGLE AT SENSOR (RAD). HORIZONTAL COORDINATE OF SENSOR. A1 : A2 : VERTICAL COORDINATE OF SENSOR, POSITIVE DOWN. V0,V1 : ARRAYS CONTAINING SOUND VELOCITY PARAMETERS. LM: ARRAY CONTAINING LAYER MIDPOINTS. INDEX OF DEEPEST LAYER USED. N : С C* DIMENSION LM(300),V0(300),V1(300),DZ(300),VEL(300) REAL*8 T0,H,H0,Z,A1,A2,TH0,TH1,VEL REAL*8 LM, V0, V1, DZ, Q1, Q2 REAL*8 VA2, R, VP2, TH, RV, DW, DT, X, T **INTEGER N,IS** I = NT = 0.0D0TH=TH0 H0 = A1VA2=V0(I)+V1(I)*A2RV=DCOS(TH)/VA2 Z = A2DZ(N) = Z - LM(N)50 IF(V1(I).EQ.0.0) THEN DW = DZ(I)/DSIN(TH)DT = DW/V0(I)H = H0 + DW*DCOS(TH)TH1 = THELSE Q2 = -V0(I)/V1(I)IF (Q2) 51,52,53 51 IS = -1GOTO 54 52 IS = 0GOTO 54 53 IS = 154 CONTINUE Q1=H0 + (Q2-Z)*DTAN(TH)R=DSQRT((Q2-Z)**2 + (Q1-H0)**2) TH1=DACOS(RV*VEL(I)) DT=DLOG((DCOS(TH)/(1+DSIN(TH)))*((1+DSIN(TH1))/DCOS(TH1)))/V1(I) H=Q1 - IS*R*DSIN(TH1) **ENDIF** T=T+DTIF (T.GE.T0) GOTO 60 Z=LM(I)H0 = HTH=TH

I=I-1 **GOTO 50**

```
DT=T0+DT-T
60
     IF(V1(I).EQ.0.0) THEN
        DW = V0(I)^*DT
        DZ(I) = DW^*DSIN(TH1)
        H = H0 + DW^*DCOS(TH1)
        Z = Z - DZ(I)
     ELSE
        X = (EXP(DT^*V1(I)))^*(1+DSIN(TH))/DCOS(TH)
        TH1=DACOS((2*X)/(1+X**2))
        H = Q1 - IS^*R^*DSIN(TH1)
        Z = Q2 - IS*R*DCOS(TH1)
     ENDIF
C RESTORE THE END LAYERS.
     DZ(I) = LM(I+1) - LM(I)
     DZ(N) = LM(N+1) - LM(N)
     RETURN
     END
                        ***************
C**
C++++
     SUBROUTINE ISOSPEED(A1,A2,T0,TH0,L,VEL,M,H,Z,TH1)
C*
C
C
                                                     08/09/89
    This is a 2-D ray tracing algorithm that mimics the one in
С
  Proceedure 5181. It utilizes the assumption that the speed
С
   of sound in water is constant for the entire layer encompassed
Č
C
  by the layer boundaries. A fixed ray invariant is used
   throughout the entire migration.
INPUTS:
     A1,A2 - POSITION OF SENSOR (A2>0 DOWN)
     T0
           - TRANSIT TIME
           - ELEVATION ANGLE (OF THE RAY AT THE SENSOR,
     TH0
             ALSO CALLED THE ENTRANCE ANGLE)
     L
           - ARRAY CONTAINING LAYER BOUNDARIES
     VEL
           - ARRAY CONTAINING SOUND VELOCITY AT THE
             THE LAYER MIDPOINTS
  OUTPUTS:
           - POSITION OF TARGET (SOUND SOURCE)
     HZ
С
           - ELEVATION ANGLE AT TARGET
     TH1
C
     DIMENSION L(300), VEL(300)
     REAL*8 A1,A2,C.S,T,DT,DW,TH,DZ,RV,TH1,Z,H,L,VEL,TH0,T0
     Z = A2
     H = A1
     T = 0.0
C CHOOSE N SUCH THAT L(N) <= A2 < L(N+1). IF A2 IS DEEPER
С
```

THAN LOWEST LAYER THEN N = M, THE INDEX OF THE DEEPEST LAYER

С BOUNDARY

N = M

DO 5 I = 2,M IF ((L(1-1).LE.A2).AND.(L(I).GT.A2)) N \approx I - 1 5 CONTINUE RV = DCOS(TH0)/VEL(N) J = N TH = TH0 S = DSIN(TH0) C = DSQRT(1 - S**2) 10 DZ \approx Z - L(J)

- C COMPUTE THE INCREMENTAL SLANT RANGE DW = DZ/S
- C COMPUTE THE INCREMENTAL TRAVEL TIME DT = DW/VEL(J)
- C ACCUMULATE TOTAL TRAVEL TIME AND TEST T = T + DT IF (T.GE.T0) GOTO 50
- C UPDATE THE HORIZONTAL AND VERTICAL ACCUMULATIONS H = H + DW*C Z = Z - DZ
- C USE SNELL'S LAW TO UPDATE THE LAYER ENTRANCE ANGLE AND

C THE TRIG FUNCTIONS. J = J - 1 C = RV * VEL(J) S = LSQRT(1 - C**2)GOTO 10

- 50 T = T DT DT = T0 - T DW = VEL(J)*DT DZ = DW*S H = H + DW*C Z = Z - DZ TH1 = DASIN(S)
 - RETURN
 - END

SUBROUTINE RAYFIT1(A1,A2,P1,P2,M,VEL,LM,DZ,V0,V1,T0,TH0, TH1, IEST) ***** С 09/12/89 NEW SUBROUTINE TO REPLACE TGEN, RAYTRACING ALGORITHM. *********** INPUTS: A1,A2 - POSITION OF SENSOR (A2 > 0 DOWN) P1,P2 - POSITION OF SOUND SOURCE (P2 > 0 DOWN) LM - ARRAY CONTAINING LAYER MIDPOINTS - NUMBER OF LAYER MIDPOINTS Μ VEL - ARRAY CONTAINING SOUND VELOCITY AT THE LAYER MIDPOINTS.

- C C **V**0 - SPEED INTERCEPT VALUES
- V1 - SPEED SLOPE VALUES
- DEPTH INCREMENTS DZ
- **IEST FLAG FOR INITIALIZING THE ANGLE**
- CCCCCCC **OUTPUTS:**
- TRANSIT TIME T0
- ELEVATION ANGLE AT THE SENSOR TH0
- ELEVATION ANGLE AT THE SOUND SOURCE TH1
- C ****

DOUBLE PRECISION VEL(300), DZ(300), LM(300), V0(300) DOUBLE PRECISION V1(300), ANG(300), G(300) REAL*8 A1, A2, P1, P2, T0, TH0, TH1, EP, S, C REAL*8 H,H0,DW,VA2,VP2,GG,R,Z,TH,RV,Q1,Q2 **INTEGER M, IS**

EP = 1D-6

- C DETERMINE LAYERS INVOLVED IN RAY FITTING N = MI = M
 - DO 30 I=1.M 1 IF ((LM(I).LE.A2).AND.(LM(I+1).GT.A2)) N=I IF ((LM(I).LE.P2).AND.(LM(I+1).GT.P2)) J=I **30 CONTINUE**
- C MAKE END CORRECTIONS FOR THE LAYERS DZ(N) = A2 - LM(N)DZ(J) = LM(J+1) - P2
- С COMPUTE SPEED OF SOUND AT A2 AND P2 $VA2 = V0(N) + V1(N)^*A2$ VP2 = V0(J) + V1(J)*P2IF(IEST.NE.0) GOTO 50
- C INITIALIZE THE ELEVATION ANGLE AT THE SENSOR, THO, BY C FITTING A STRAIGHT LINE SPEED PROFILE BETWEEN P2 AND A2.

IF(VEL(N).EQ.VEL(J)) THEN TH0 = DATAN((A2-P2)/(P1-A1))ELSE $Q2 = (VEL(N)^*LM(J) - VEL(J)^*LM(N)) / (VEL(N)-VEL(J))$ $O1 = 0.5^{(P1+A1)+(0.5^{(P2-A2)^{(P2+A2-2^{Q2}))/(P1-A1)}$ TH0 = DATAN((Q1-A1)/(Q2-A2))**ENDIF**

C OUTER LOOP: SET UP RAY FITTING FOR TH0 = ELEVATION ANGLE 50 S = DSIN(TH0) $C = DSQRT(1.0 - S^{**}2)$

I = NRV = C/VA2H0 = A1Z = A260 IF(V1(I).EQ.0.0) THEN DW = DZ(I)/S

 $H = H0 + DW^*C$ ELSE Q2 = -V0(I)/V1(I) ŝ.

IF (Q2) 61,62,6361 IS = -1 GOTO 64 62 IS = 0 GOTO 64 63 IS = 1 64 CONTINUE

Q1 = H0 + (Q2-Z)*S/C R = DSQRT((Q2-Z)**2 + (Q1-H0)**2) C = RV*VEL(I) S = DSQRT(1.0-C**2) H = Q1 - IS*R*SENDIF

IF (I.EQ.J) GOTO 80 H0 = H Z = LM(I)I = I - 1 GOTO 60

80 TH1 = DACOS(RV*VP2)

- C FRACTIONAL LAYER CORRECTION IF(V1(J).NE.0.0) H = Q1 - IS*R*DSIN(TH1) IF (ABS(H-P1).LT.EP) GOTO 90
- C RE-ESTIMATE TH0 TH0 = DATAN(DTAN(TH0)*H/P1) GOTO 50
- C PREPARE FOR COMPUTATION OF TRANSIT TIME.
- C COLLECT EXIT AND ENTRANCE ANGLES. 90 ANG(I) = TH1
- ANG(N+1) = TH0 DO 95 I = J+1,N ANG(I) = DACOS(RV*VEL(I)) 95 CONTINUE

C COMPUTE TRANSIT TIME T0 = 0.0D0DO 100 I = J,N IF(V1(I).EQ.0.0) THEN T0 = T0 + DZ(I)/(V0(I)*DSIN(ANG(I)))ELSE T0=T0 + DLOG((DCOS(ANG(I+1))*(1+DSIN(ANG(I))))/* ((1+DSIN(ANG(I+1)))*DCOS(ANG(I))))/V1(I) ENDIF 100 CONTINUE C REMOVE THE END CORRECTIONS. DZ(I) = IM(I+1) + IM(I)

SUBROUTINE VELMOD(L, VEL, M, LL, VV, MM) C*** С 02/22/90 THIS PROGRAM TAKES THE VELOCITY SOUND PROFILE GIVEN IN FIVE (5) С FOOT INCREMENTS AND CONVERTS IT INTO TWENTYFIVE (25) FOOT INCREMENT C C PROFILE. С C CCCCCCCC **INPUT: DEPTH IN 5 FT INCREMENTS** Ŀ VEL: SOUND VELOCITY IN 5 FT. INCREMENTS **M**: NUMBER OF ELEMENTS IN DEPTH ARRAY OUTPUT LL: **DEPTH IN 25 FT INCREMENTS** SOUND VELOCITY IN 25 FT INCREMENTS VV: С MM: NUMBER OF ELEMENTS IN DEPTH ARRAY DIMENSION L(300), LL(300), VEL(300), VV(300) REAL*8 L,LL,VEL,VV,VS MM = INT(M/5) $MREM = M - 5^*MM$ VS = 0.0D0DO 10 J = 1,MM $LL(J) = L(5^*J-4)$ $VV(J) = 0.2^{+}(VEL(5^{+}J-4) + VEL(5^{+}J-3) + VEL(5^{+}J-2) +$ VEL(5*J-1) + VEL(5*J))CONTINUE 10 DO 20 J = 1,MREMVS = VS + VEL(M+1-J)20 CONTINUE LL(MM+1) = L(5*MM + 1)VV(MM+1) = VS/MREMRETURN **END** SUBROUTINE ERRPROP(H1,Z1,K,A2,XTILT,YTILT,ZROT,D,L,VEL,M, THR.THCER.TH1ER.PHR.PHER.PH1ER.H2ER.Z2ER.T4) C*** 08/22/90 С POSITION ERROR ANALYSIS WHEN THE ORIGIN IS OVER THE C-HYDROPHONE С AND THE PROPOSED SYSTEM IS APPLIED. COMPUTES THE TRUE ELEVATION С ANGLES AND ESTIMATED ELEVATION ANGLES FOR K SOUND SOURCES С DISTRIBUTED EQUALLY ON A CIRCLE OF RADIUS H1 AT DEPTH Z1 FOR A С RECIEVING ARRAY AT DEPTH A2 BUT OTHERWISE AT THE CENTER OF THE C CIRCLE. A2 IS DEPTH OF THE GEOMETRIC CENTER OF THE ARRAY CUBE. THE С ACCOUSTIC CENTER IS THE C-HYDROPHONE. BECAUSE OF DIRECT С C MEASUREMENT THERE IS NO TRANSIT TIME ERROR. C **************************** Ċ С С INPUTS: **RADIUS OF CIRCLE IN FEET** С H1:

د

С	Z1:	DEPTH OF SOUND SOURCE IN FEET			
С	K:	NUMBER OF SOURCES ON THE CIRCLE			
С	A2:	DEPTH OF THE CENTER OF THE ARRAY			
C XTILT.YTILT.ZROT: ORIENTATION INFORMATION ABOUT THE					
С		SENSING ARRAY (RADIANS)			
C	D:	LENGTH OF ARRAY EDGES.			
С	Ŀ	DEPTH OF LAYER BOUNDARIES.			
С	M :	NUMBER OF RECORDS IN THE VELOCITY DEPTH PROFILE.			
С	VEL:	AVERAGE SPEED OF SOUND IN THE LAYERS.			
С					
	C OUTPUTS:				
С	THR:	ACTUAL ELEVATION ANGLE (THETA)			
C THONE: THETA ESTIMATE, PROPOSED METHOD					
C THER: THETA ERROR, PROPOSED METHOD					
С	PHR:	• • • •			
С	PH1:				
C PH1ER: PHI ERROR, PROPOSED METHOD.					
С	HC:				
С	HER:				
С	ZC:	VERTICAL ESTIMATE, CURRENT METHOD.			
С	ZER:	VERTICAL ERROR.			
С	T:	TRANSIT TIME TO THE C-HYDROPHONE.			
С					
Č**************					

DIMENSION B(5,3),DZ(300),HD(5),HX(5),HY(5),ITH2(20),L(300) DIMENSION LM(300),PH1(20),PH1ER(20),PHC(20),PHER(20) DIMENSION PHR(20),PX(20),PY(20),T(4),TH1ER(20),TH2(20) DIMENSION TH2ER(20),THC(20),THCER(20),THONE(20),THR(20) DIMENSION V0(300),V1(300),VEL(300),X0(30),PZ(2,2),PC(2) DIMENSION V0(300),V1(300),VEL(300),X0(30),PZ(2,2),PC(2) DIMENSION PP(2,3),E(3),TH3(20),TH3ER(20),TT3(3) DIMENSION IFL(2),SSC(2),SP(2),H2(20),Z2(20),H2ER(20) DIMENSION Z2ER(20),T4(30)

REAL*8 B,DZ,HD,HX,HY,L,LM,PH1,PH1ER,PHC,PHER,PHR,PX REAL*8 PY,T,TH1ER,TH2,TH2ER,THC,THCER,THONE,THR,V0,V1 REAL*8 VEL,X0,PONE,T4

REAL*8 A1,A2,A2M,ALPH1,ANG,C,C1,C2,C3,CAZ,CT,CX,CX0,CY,CY0 REAL*8 CZ,CZ0,D,DR,DR1,DP1,DP2,DTDS,EP5,F,FT,GG,H1,HH1,P1 REAL*8 P2,PIE,Q1,Q2,Q1P,R0,R1,S,S0,S1,S2,S3,SAZ,ST,SS,T0,T3 REAL*8 T3P,TH0,TH1,THEC,V,VV0,VV1,X,XTILT,Y,Y0,Y1,YTILT REAL*8 Z,Z0,Z1,ZROT,ZZ1,SC,DT

REAL*8 PP,E,DEL,A,TH3,TH3ER,TT3,SSC,SP,PZ,PC REAL*8 A11,A12,A21,A22,B1,B2,H2,H2ER,Z2,Z2ER,THE REAL*8 LB,SV,SVU,SU2,SU4,SVU2,G,U INTEGER ITH2,IFL

PIE = 3.14159265359D0 Et'S = 1D-6 IEST = 0 M1 = M

C DISTRIBUTE THE K SOURCES EQUALLY AROUND THE CIRCLE COUNTER C CLOCKWISE FROM THE EAST.
 DO 10 I = 1,K
 ANG = 2*PIE*(I-1)/K

PX(I) = H1*DCOS(ANG)PY(I) = H1*DSIN(ANG)PHR(I) = ANGIF(ANG.GT.PIE) PHR(I) = ANG - 2*PIE 10 CONTINUE C FORM SINES AND COSINES OF ALL THE EULER ANGLES: ROLL, PITCH, YAW S2 = DSIN(XTILT) $C2 = DSORT(1 - S2^{**}2)$ S1 = DSIN(YTILT)/C2 $C1 = DSORT(1 - S1^{++}2)$ S3 = -DSIN(ZROT)C3 = DCOS(ZROT)C IN THE COORDINATE SYSTEM HAVING CENTER AT THE C-HYDROPHONE C AND POSITIVE-UPWARD, THE LOCATIONS OF THE POUR HYDROPHONES C (RELATIVE TO THE ARM I FNICTURE) ARE DOWN OF THE POUR HYDROPHONES B(1,1) = C2*C3B(1,2) = C2*S3B(1,3) = S2B(2,1) = -S1*S2*C3 - C1*S3B(2,2) = -S1*S2*S3 + C1*C3B(2,3) = S1*C2B(3,1) = -C1*S2*C3 + S1*S3 $B(3,2) = -C1^*S2^*S3 - S1^*C3$ B(3,3) = C1*C2C LIKE NOTATION WILL BE USED TO LOCATE THE C-HYROPHONE AND THE C ARRAY CENTER. DO 12 J = 1.3B(4,J) = 0.0D0 $B(5,J) = 0.5^{*}(B(J,1) + B(J,2) + B(J,3))$ 12 CONTINUE A1 = 0.0D0P2 = Z1C LOCATE THE HYDROPHONE HORIZONTAL COMPONENTS IN THE COORDINATE C SYSTEM CENTERED AT C-HYDROPHONE. DO 14 J = 1.5 $HX(J) = D^*B(J,1)$ $HY(I) = D^*B(I,2)$ 14 CONTINUE C DETERMINE THE DEPTHS OF THE FOUR HYDROPHONES AND THE ARRAY CENTER. $HD(1) = A2 + D^{*}(B(5,3) - B(1,3))$ $HD(2) = A2 + D^{*}(B(5,3) - B(2,3))$ $HD(3) = A2 + D^{*}(B(5,3) - B(3,3))$ $HD(4) = A2 + D^*(B(5,3) - B(4,3))$ HD(5) = A2**C** FIND THE DEEPEST HYDROPHONE A2M = 0.D0DO 51 J=1,4 IF(HD(J).GT.A2M) A2M = HD(J)51 CONTINUE

C FORM THE SET OF LAYER MIDPOINTS.

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DO 105 I = 1, M - 1
        LM(I) = .5^{*}(L(I) + L(I+1))
105 CONTINUE
   FORM DEPTH INCREMENTS, AND ALL SOUND VELOCITY SLOPES AND
С
С
   INTERCEPTS.
     DO 110 I=1,M-2
        DZ(I)=LM(I+1)-LM(I)
        VO(I) = (LM(I+1)*VEL(I) - LM(I)*VEL(I+1))/DZ(I)
        V1(I)=(VEL(I+1)-VEL(I))/DZ(I)
110 CONTINUE
     LM(M) = LM(M-1) + DZ(M-2)
С
     IF(A2M.LT.LM(M-1)) GOTO 126
C IF A2M IS DEEPER THAN THE LAST LAYER MIDPOINT, THEN WE EXTRAPOLATE
C THE SOUND VELOCITY PROFILE BY USING A QUADRATIC FUNCTION OVER
C THE DEEPEST 100 FEET.
C FIRST COUNT THE NUMBER OF LAYERS (OF THICKNESS DZ(M-2)) TO
C BE ADJOINED. ALSO MUST EXTEND THE L ARRAY.
     K0 = 2 + MAX(0,NINT((A2-LM(M-1))/DZ(M-2)))
C FIND AVERAGE DEPTH OF LAST 100 FEET.
     LB = 0.0D0
     DO 43 I = M-21.M-1
        LB = LB + LM(I)
43 CONTINUE
     LB = LB/21
C FORM SUMS OF POWERS AND PRODUCTS.
     SV = 0.0D0
     SVU2 = 0.0D0
     SVU = 0.0D0
     SU2 = 0.0D0
     SU4 = 0.0D0
     DO451 = M-21.M-1
        U = LM(I) - LB
        SV = SV + VEL(I)
        SVU = SVU + U*VEL(I)
        SVU2 = SVU2 + U^{**2} * VEL(I)
        SU2 = SU2 + U^{**2}
        SU4 = SU4 + U^{**}4
45 CONTINUE
     G = SVU/SU2
     GG = (21*SVU2 - SU2*SV)/(SU4 - SU2**2)
     V1(M-1) = G
C PERFORM THE EXTRAPOLATION.
     DO 125 I=M,M+K0
        V1(I-1) = V1(I-2) + GG^*DZ(M-1)
        LM(I) = LM(I-1) + DZ(M-2)
        VEL(I) = VEL(I-1) + DZ(M-2)*V1(I-1)
        VO(I-1) = (LM(I)*VEL(I-1) - LM(I-1)*VEL(I))/DZ(M-2)
        L(I+1) = L(I) + DZ(M-2)
        DZ(I-1) = DZ(M-2)
```

125 CONTINUE

- C UPDATE M, THE NUMBER OF LAYERS M = M + K0
- 126 CONTINUE
- C LOCATE THE WATER LAYER, N, CONTAINING THE ARRAY. N = MDO 37 J = 2,M
 - IF((LM(J-1).LE.HD(4)).AND.(LM(J).GT.HD(4))) N = J-1
- 37 CONTINUE С
 - $V = V0(N) + V1(N)^{*}HD(4)$
- C THE OUTER LOOP WILL PERFORM COMPUTATIONS FOR THE K SOUND C SOURCES.
- **C** RAYFITTING
- ć THE INNER LOOP WILL FIT RAYS TO THE FOUR HYDROPHONES
- C IN THE ORDER X,Y,Z, AND C.
 - DO 50I = 1.KWRITE(*,*)' OUTER LOOP I = ',I,' K = ',KIVV1 = 0DO35I = 1.4 $P1 = DSQRT((PX(I)-HX(J))^{**}2 + (PY(I)-HY(J))^{**}2)$ Z0 = HD(I)CALL RAYFIT1(A1,Z0,P1,P2,M,VEL,LM,DZ,V0,V1,T0,TH0,
 - TH1, IEST)
- C COLLECT THE FOUR TRANSIT TIMES. T(J) = TO
- C IN THIS PROGRAM WE KEEP ONLY THE TRUE ELEVATION ANGLE AT THE
- C C-HYDROPHONE.

THR(I) = THOT4(I) = T(4)

- 35 CONTINUE
- C CALCULATE THE PRE-TILT CORRECTED APPARENT POSITION DO 40 = 1.3 $X0(I)=(D^{**2} + (V^{*}T(4)-V^{*}T(I))^{*}(V^{*}T(4)+V^{*}T(I)))/(2^{*}D)$
- 40 CONTINUE

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- C COMPUTE DIRECTION COSINES. $CX0 = X0(1)/(V^*T(4))$ $CY0 = X0(2)/(V^*T(4))$ $CZ0 = DSORT(1 - CX0^{**}2 - CY0^{**}2)$
- C PERFORM EXACT TILT CORRECTIONS AND THE ROTATIONAL ALIGNMENT.

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CX = B(1,1)*CX0 + B(2,1)*CY0 + B(3,1)*CZ0CY = B(1,2)*CX0 + B(2,2)*CY0 + B(3,2)*CZ0CZ = B(1,3)*CX0 + B(2,3)*CY0 + B(3,3)*CZ0IF(IVV1.EQ.1) THEN $TH_2(I) = 0.5*PIE - DACOS(CZ)$ **GOTO 49 ENDIF** $SAZ = CY/DSQRT(CX^{**}2 + CY^{**}2)$ $CAZ = CX/DSQRT(CX^{**2} + CY^{**2})$

 $\begin{array}{l} PH1(I) = DATAN2(SAZ,CAZ) \\ PH1ER(I) = PH1(I) - PHR(I) \\ IF(ABS(PH1ER(I)).GT.PIE) PH1ER(I) = PH1(I) + PHR(I) \\ THONE(I) = 0.5*PIE - DACOS(CZ) \\ TH1ER(I) = THONE(I) - THR(I) \end{array}$

50 CONTINUE

C OUTER LOOP COMPLETED!

RETURN

- 100 FORMAT(3(5X,E13.6))
- 120 FORMAT(3(5X,F15.12))
- 130 FORMAT(4(F10.8,2X),F13.8,1X,F12.8,2X,2(F12.8,2X),F12.8)
- 140 FORMAT(5X,'The transit time to the z-phone is not bracketed') END

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