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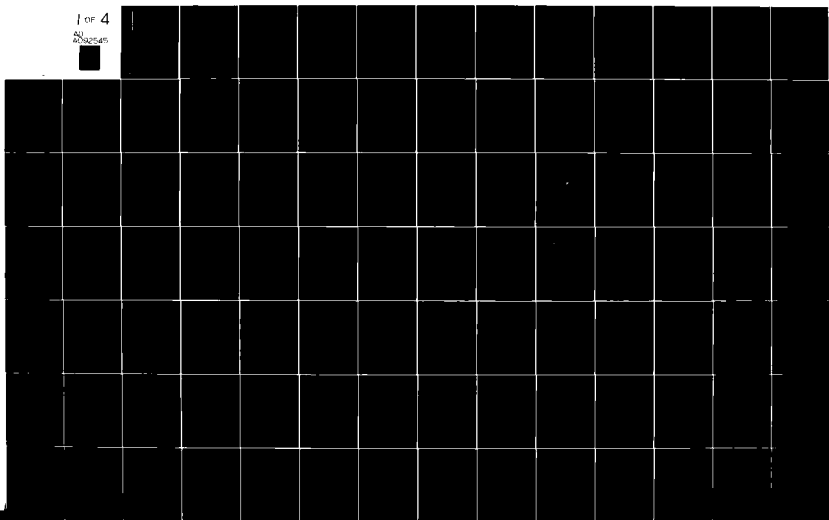
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Los Angeles

Relativistic Electron Precipitation:
An Observational Study

A dissertation submitted in partial satisfaction
of the requirements for the degree Doctor of Philosophy
in Atmospheric Sciences

by

Leopold Joseph Andreoli

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1980

Dedicated to Kathleen, David, Michael and Donna.

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	of precipitating electrons from 235 keV to highest energy channel at which strong precipitation occurred, $J(p/cm^2 s str) =$ integral flux of strong diffusion precipitation of protons per energy channel if any, 80 p ⁺ L-VAL = L value of 80 keV proton precipitation if any, strong precip 12, 33, L-VAL = check if 12, 33 keV electron channels had strong precip and L value of 12 keV channel UT (sec) = time at which 235 keV electron precip occurred.	251
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ABSTRACT OF THE DISSERTATION

Relativistic Electron Precipitation:
An Observational Study

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Leopold Joseph Andreoli

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The object of this dissertation is to analyze satellite data to investigate the morphology of strong diffusion relativistic electron precipitation (REP) events with the ultimate purpose of gaining insights into the precipitation mechanism and determining atmospheric consequences. Analysis of fourteen (14) months of magnetic electron spectrometer and proton telescope data from the Air Force S3-3 satellite from July 1976 to September 1977 is presented. Analysis techniques and data catalogues are presented in detail. Three distinct classes of REPs were discovered: an overwhelming majority (96.5%) exhibited precipitation in a large range of energies below 1 MeV for electrons with concurrent proton precipitation; the smallest class (1.3%) also exhibited concurrent electron and proton precipitation but the electron

and proton precipitation but the electron precipitation occurred only at energies above 1 MeV; a third class (2.2%) had no associated proton precipitation and electrons precipitated only at energies below a few hundred keV. In addition, all REPs have a 10-20% frequency of occurrence that linearly increases with magnetic activity, show seasonal variation with equinoxial maxima and solstice minima, have mostly small latitudinal (~0.5L) and large longitudinal (1000-10,000km) extents, and occur most frequently (9 out of 10) in the local nighttime sector.

A theoretical approach to the REP mechanism is presented. Interaction of the radiation belt electron and ring current proton populations with electrostatic and electromagnetic ion cyclotron and whistler mode chorus waves can account for the three classes of precipitation, respectively. Examples of each class are presented in detail.

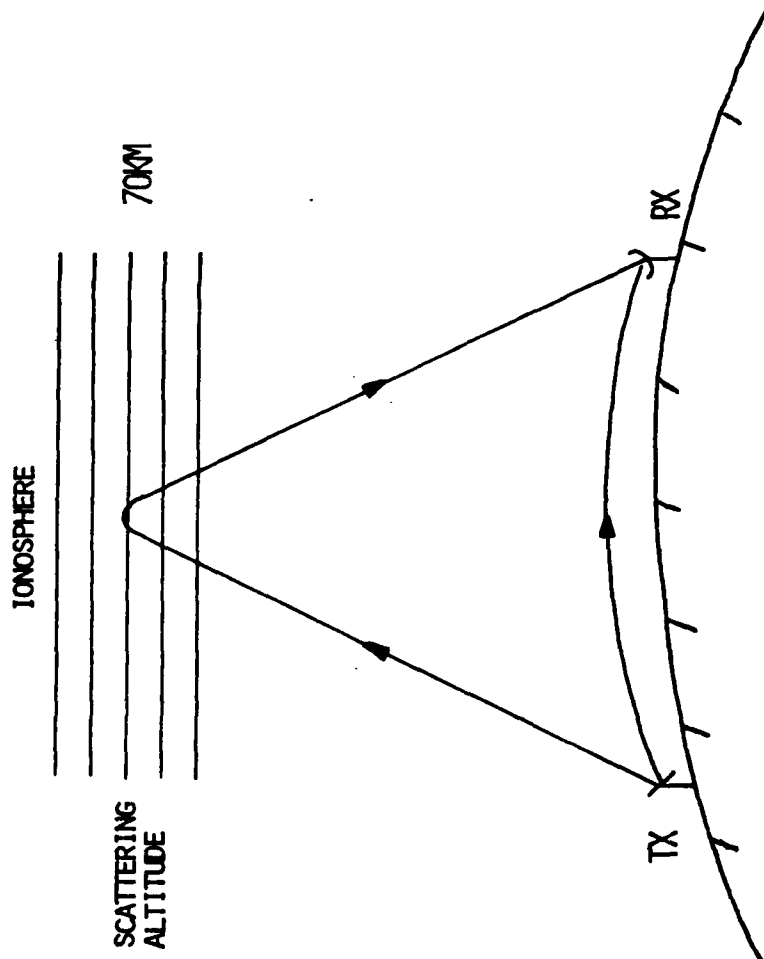
Finally, an investigation of the energy deposition rates from these events and their ultimate effect on the ionized and neutral atmosphere with emphasis on ozone modification is presented. The local hydroxyl concentration can be increased up to a factor of 10 and the local ozone concentration can be decreased by 10-30% by REP events.

Introduction

HF radio waves can propagate from transmitter to receiver via refractive interaction with the ionosphere (see Fig. 1). This propagation is known as forward scatter or skywave propagation. The received signal can be diminished (in power) if the electron density below the scattering altitude is enhanced in any way. Signal absorption occurs during the day due to the natural increase of free electrons below the scattering altitude by solar ionizing radiation (1000-1100⁰A). Absorption can also be increased by precipitating energetic particles (e.g., solar protons) that reach altitudes below the scattering altitude. One well known class of absorption event is the Polar Cap Absorption (PCA) events which occur over the polar region following the outburst of energetic protons during flare activity on the sun. A lesser known but more frequent class of events involves the precipitation of relativistic electrons and these will be the subject of this dissertation.

Relativistic Electron Precipitation (REP) events were first observed as anomalous daytime absorption of forward scatter (more correctly - ionospheric refraction) HF signals by Bailey and Pomerantz (1965). The detection system used by Bailey and Pomerantz consisted of three paths each in Alaska and Antarctica operating near 23 MHz with skywave scattering points over the invariant latitude (Λ) range 58-77⁰ ($L = 3.6-20.0$). The scattering altitude was estimated to be ~70 km. Peak signal absorption typically occurred between $\Lambda = 62-66^0$ ($L = 4.5-6.0$); namely, within the region of long term trapping for

FIG 1. Illustration of one-hop skywave propagation via the
Ionosphere.



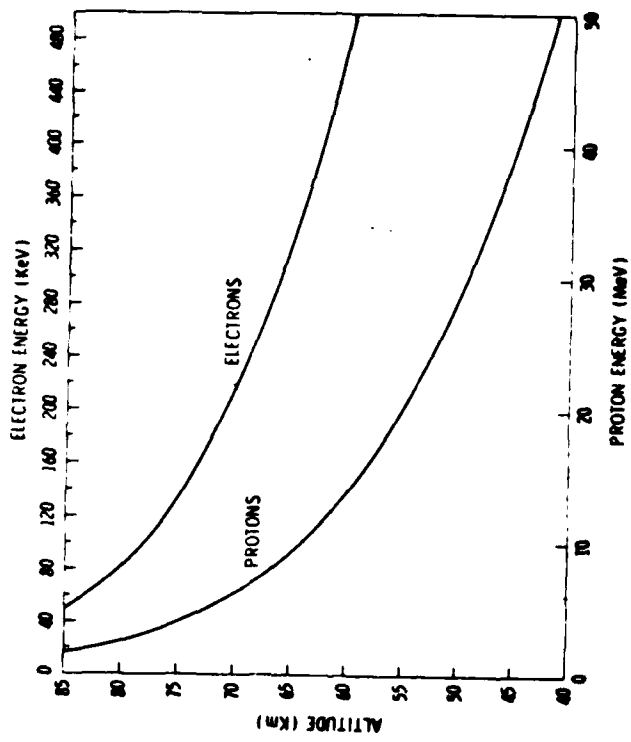
radiation belt particles.

Since solar protons could be ruled out due to the absence of absorption over the higher latitude polar cap region ($\Lambda > 72^\circ$; $L > 10$), absorption was attributed to a locally precipitated source of particles. These particles dissipate most of their energy at altitudes determined by their range in the atmosphere and thus by their energy (Aikin and Bauer, 1965). As shown in Fig. 2, protons require at least several MeV while electrons require only a few hundred keV of energy to penetrate below 70 km. This, however, immediately excludes the main bulk of auroral electrons ($\sim 1-10$ keV) as the precipitating particles causing absorption and indeed visible auroral activity (monitored by all sky cameras) was found to have no correlation with dayside HF absorption (Bailey et al., 1970). Since protons of several MeV energy are tenuous beyond ($L \sim 3.5$) and do not exhibit outer zone integral flux variations of the order estimated for strong ionospheric absorption ($\sim 40 \text{ cm}^{-2} \text{ str}^{-1} \text{ sec}^{-1}$ for protons), Bailey and Pomerantz concluded that trapped electrons at energies greater than 400 keV were the cause. In situ satellite measurements have also demonstrated that relativistic electrons trapped from $\Lambda = 55-60^\circ$ ($L = 3-4$) can exhibit order of magnitude flux decreases during disturbed times (Frank, 1965). Furthermore, since the maximum observed absorption intensities occur at higher L values than the normally expected peak intensity of trapped electrons, Bailey and Pomerantz (1965) concluded the mechanism operates more efficiently near the trapping boundary.

Since Bailey and Pomerantz only observed the events during the daytime, they assumed, inconclusively, that the precipitating mechanism is more intense on the dayside. It must be noted that this type of measurement is highly frequency dependent. This is supported by the fact that as the forward scatter observing frequency changed from 50 MHz in the early 1950's to 35 MHz pre-IQSY (Int'l Quiet Sun Year) and finally to 23 MHz post IQSY, ability to detect less intense events improved. The day/night occurrence probability of REPs will be further discussed as we review subsequent measurements.

Other types of observations were soon employed to verify inferences based on the forward scatter results. Riometers which integrate the effect of precipitating particles (and radiation) of all energies were used to confirm signal absorption (Bailey et al., 1970). All sky cameras (observing auroral activity), ionosondes (for Es, sporadic E layer, detection), magnetometers (Bailey et al., 1966) and balloon X-ray measurements (for energy spectrum information) (Barcus and Rosenberg, 1966) basically confirmed signal absorption differences between daytime REPs and nighttime Es (absorption by precipitating 5-20 keV electrons). In addition, estimates of precipitating electron spectra for REP and Es conditions for quiet to severely disturbed magnetic conditions were constructed from X-ray and forward scatter measurements (Bailey et al., 1966; Bailey et al., 1970). Comparison of X-ray and forward-scatter-deduced spectra for the same event revealed the marked inability of the forward scatter system alone to detect the presence of relativistic electrons due

FIG 2. Atmospheric penetration depths for protons and electrons assuming vertical incidence (adopted from Aikin and Bauer, 1965).



to the large uncertainty in scattering altitude (Rosenberg et al., 1972). Spectra derived from various sources will be compared later.

Another drawback to estimating REP spectra from forward scatter measurements alone is the uncertainty of absolute absorption. The signal power received is enhanced by increases in the electron density at the scattering altitude as well as diminished by increases in electron density along the signal path below the scattering altitude, thus an accurate measure of absorption is difficult.

REPs (and PCAs) appear as absorption on forward scatter records by day. At night, although the scattering altitude is higher (by 10-20 km), there are insufficient free electrons below the scattering altitude to produce signal absorption and even the strongest precipitation events appear as signal enhancements. Riometers, however, continue to show absorption during such nighttime events (Bailey, 1968).

Two Alaskan paths at $\Lambda = 65^\circ$ and 62° ($L = 5.7$ and 4.6) were chosen for long term statistical study (Bailey, 1968). For the period 1 March 1964 to 31 December 1966, REPs producing absorption ≥ 8 db occurred on 319 days with maxima at equinoxes and minima at solstices. The 93 medium and large REPs (absorption ≥ 14 db) for the same period are not associated with sudden commencement (SC) geomagnetic storms, but the events mostly occurred on magnetically disturbed days, lasted from 1-6 hours (with some seeming to last 24 hrs) and were mainly confined to the 12 hour period centered on local noon.

After 1967, the system detection capability became erratic due to the increasing solar activity. The system was terminated for various additional reasons early in 1973 (Sullivan, 1979). Attempts were made to obtain some 1969 forward scatter data for correlation and calibration with OV1-19 satellite spectrometer data (see Appendix A) but all of the forward scatter data except for the first five months of 1969 had been lost (Lincoln, 1978).

In 1968-1971, the Norwegian Defense Research Establishment (NDRE) undertook a program entitled "Omega Research in Norway" (Larsen, 1973). This program monitored the VLF (12.3 kHz) short hop (433 km) path between the transmitter at Aldra, $\Delta = 63.7^\circ$ (L ~ 5.1), and the receiver at Tromsø, $\Delta = 66.5^\circ$ (L ~ 6.3), both in Norway, for PCA, substorm, REP, and stratospheric warming research. Phase recordings of the VLF signal perpendicular (to propagation plane) and parallel magnetic components provide sensitive detection of ionization variations below the scattering altitude of about 60-65 km daytime and 80-90 km nighttime (Larsen, 1973). A diurnal effect is seen as an advance shift in baseline of 60-80 μ sec from local night to day. REPs (and PCA/solar flare events) are readily observed both day and night as phase advances beyond the normal undisturbed levels.

Thorne and Larsen (1976) compared the Norwegian VLF and Alaskan HF observations of REPs and Auroral Electrojet (AE) Indices (for substorm identification) for the period January-May, 1969. They indicate "...a direct correlation between the onset of intense nighttime REP events with substorms while dayside REPs are far less frequent,

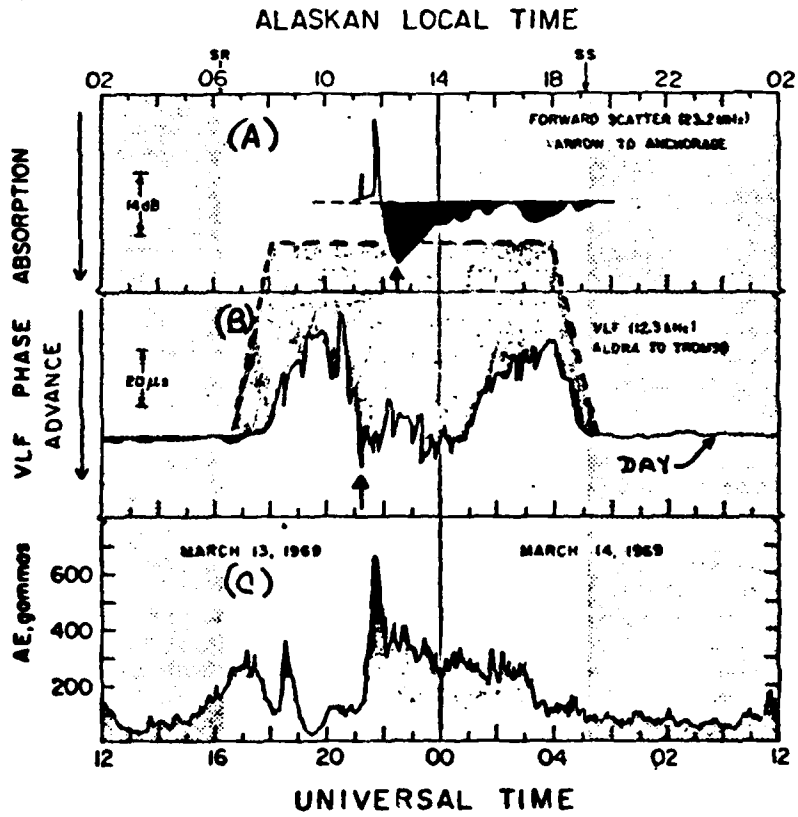
are shorter in duration and are often delayed by as much as several hours from the onset of substorm activity" (Thorne and Larsen, 1976). Figure 3 shows for example, a simultaneous increase in AE and VLF phase advance in Norway at 2100 UT followed 1-1 1/2 hrs later by the HF absorption in Alaska. This sequence agrees with the observation of the 25-26 August 1967 substorm by Hones et al., (1971) but not their inference of gradient drifting electrons injected at substorm onset since the gradient drift lifetimes of relativistic electrons are on the order of minutes (Thorne and Larsen, 1976). In addition, only a select class of substorms can trigger dayside REPs (Rosenberg et al., 1972; Thorne and Larsen, 1976) and substorm intensity is only weakly correlated in the selection of which substorm will spawn a REP (Thorne and Larsen, 1976).

Many observers using riometers (Brown, 1964; Jelly and Brice, 1967) and chains of riometers (Ansari, 1965; Lukkari et al., 1975; Lukkari and Kangas, 1976; Lukkari et al., 1977) have observed similar relations between subauroral energetic electron precipitation and slowly varying absorption (SVA), substorm onsets, and magnetic pulsations. Riometers observe total absorption occurring both above and below the scattering altitudes of VLF and HF systems and alone cannot determine where (in altitude) the absorption occurs nor yield a precise precipitation spectrum (Bailey, 1968).

Balloon-borne measurements of X-rays during REP events (usually with uncollimated NaI detectors near 32 km altitude) have the advantage over riometer measurements in sensitivity, time resolution, and the fact that they provide an

1

FIG 3. Time sequence of (A) HF absorption (dark shaded area) in Alaska; (b) VLF phase advance (shaded) from normal nighttime trace (dashed line); (c) AE index for 13-14 March 1969 from Thorne and Larsen, 1976.



energy spectrum (Bewersdorff et al., 1968). In addition to the disadvantages of imprecise balloon location (~ 100 km horizontally) and proton contamination of X-ray scintillators, the theoretically predicted X-ray spectrum (the scintillator measures mostly secondaries) is very sensitive to the assumed incident electron spectrum, usually in exponential form ($\propto \exp E/E_0$), and atmospheric depth (Bailey et al., 1970). Only six or seven X-ray measurement deduced electron energy spectra during energetic electron precipitation events have published and range from spectra constructed by the addition of soft ($E_0 = 9$ KEV) and hard ($E_0 = 40$ KEV) components (Bailey et al., 1970) to very hard ($E_0 = 180-200$ KEV) spectra (Rosenberg et al., 1972; Parks et al., 1979).

Rockets have carried spectral and integral electron spectrometers (Matthews and Simons, 1973; Vij et al., 1975; Matthews et al., 1976) as well as having boosted parachute-borne X-ray scintillators (Vij et al., 1975) during REP and auroral events. Simultaneous electron (rocket to 275 km apogee) and X-ray (rocket boost to 60 km) observations show fair agreement at least in the much "softer" auroral events (Vij et al., 1975). The spectrometers yield a more precise electron energy spectrum and, depending on instrument orientation and flight configuration, add pitch angle distribution information. An obvious drawback of this type of experiment for REP research is the lack of morphological statistics due to prohibitive logistics expense.

Satellite X-ray spectrometer experiments provide simultaneous measurements over a wide range of local times because of their extensive fields of view (Imhof et al., 1974) as well as opportunities for coordinated studies with particle spectrometers for improved spatial/temporal resolution (Imhof et al., 1975). Comparison of X-ray and particle measurements show agreement to only a few hundred keV and that impulsive (< 5 min duration) events near the energetic trapping boundary were substorm correlated (Imhof et al., 1978). Coordinated two satellites X-ray and particle measurements show significant variations in local time dependence of energetic electron precipitation though fluxes seemed greater in late morning than early morning.

The majority of satellite REP measurements have been made with integral (Brown and Stone, 1972; Larsen and Thomas, 1974; Morfill et al., 1975) and spectral (Vampola, 1969; Vampola, 1971 a band b; Imhof et al., 1973; Imhof et al., 1975) particle spectrometers. REP precipitation is observed to be a narrow ($\Delta L \sim .5-1.5$) phenomenon. Figure 4 shows an invariant latitude versus magnetic local time plot of electron precipitation spikes with $E > 425$ keV from the last five months of 1967 from the low polar orbit satellite OGO 4 (Brown and Stone, 1972). Notice a relative scarcity of events between 06-12 MLT. Also, the events extend from approximately the middle of the outer electron zone to the limit of closed field lines with the highest density clustered about the limit of stable trapping. A general shift to lower invariant latitudes is observed

with increasing geomagnetic activity. Event time scales range from several minutes to several hours. Approximately five times as many events occur at 2200 MLT as at 1000 MLT. This study was extended by the use of the > 500 keV measurements on board another low polar orbit satellite ESRO 4 with similar results (Morfill et al., 1975).

Another investigation of magnetic electron spectrometer data of precipitating electrons in the energy range $225 \text{ keV} < E < 2.47 \text{ MeV}$ aboard polar orbit satellite OV3-3 for the period August 1966-September 1967 show common occurrences of precipitation at local evening and midnight in a narrow region ($\Delta L \sim 0.5$) at the local limit of trapping (Vampola, 1971). Rare events (3) at local evening just inside the plasmopause where higher energy electrons ($> 1 \text{ MeV}$) are preferentially precipitating were also found.

Measurements of precipitating electrons from $> 48 \text{ keV}$ to $> 434 \text{ keV}$ on the polar orbit satellite ESRO 1A during the REP of 2 February 1969 at local morning at $\Delta L = 63^{\circ}-70^{\circ}$ ($L = 4.9-8.5$) have been reported by Larsen and Thomas (1974). The energy spectrum for this event is a bit "harder" than those previously obtained by rocket measurements (Matthews and Simons, 1973) but much harder than those deduced from balloon X-ray measurements (Bailey et al., 1970).

Figure 5 compares REP energy spectra from various sources previously discussed with typical discrete and diffuse auroral spectra (Lui et al., 1977). Satellite-borne spectrometers, particularly spectral spectrometers, have measured the hardest spectra in comparison with non satellite spectra and represent the most direct means of examining REPs of all intensities.

FIG 4. Latitude-local time plot of all 425 keV electron spike observations from Brown and Stone, 1972.

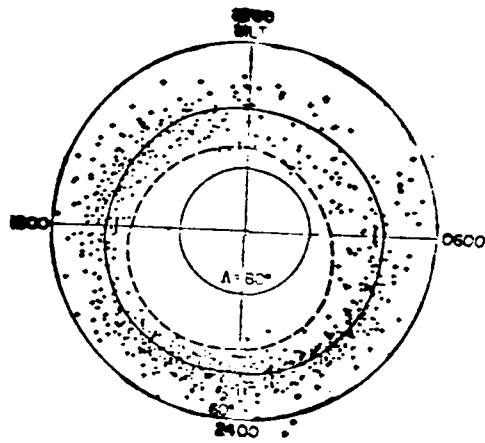
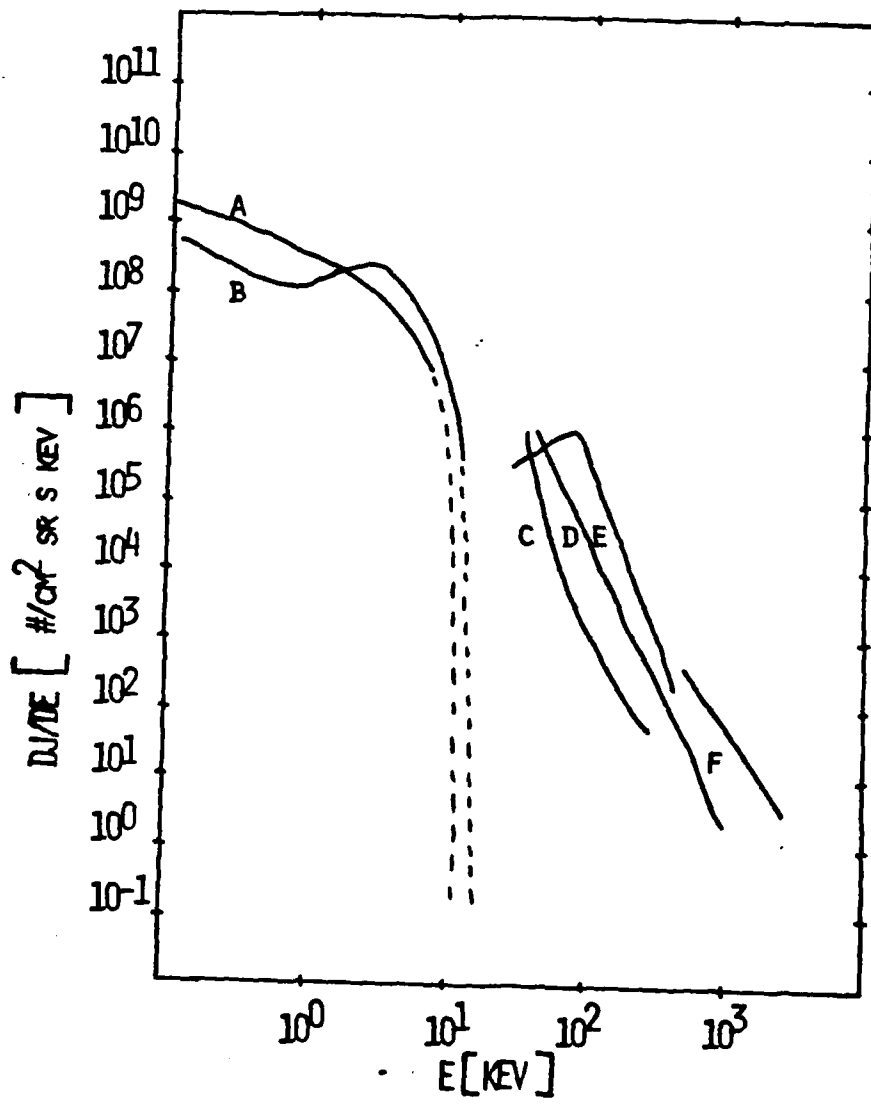


FIG 5. Differential energy spectra of precipitating electrons of the following types: (A) Diffuse aurora (Lui et al., 1977); (B) Discrete aurora (Lui et al., 1977); (C) Balloon X-ray (Bailey et al., 1970); (D) Rocket (Matthews and Simons, 1973); (E) Satellite (Larsen and Thomas, 1974); (F) Satellite (Vampola, 1971).



It was the purpose of this dissertation to determine the mechanisms responsible for REP events, their relative importance and atmospheric consequences through the study of satellite data. Magnetic electron spectrometer data on the low polar orbit satellites OV1-19 and S3-3 were graciously provided by the instrument builder and prime investigator, A. L. Vampola, Aerospace Corp., El Segundo, Ca. An exhaustive study of the OV1-19 data set was performed but for various reasons the data set yielded only marginally significant results. This study is described in detail with results in Appendix A. The S3-3 data set has few of the drawbacks of the OV1-19 data set and is the main source of study here. The S3-3 satellite orbit and orientation as well as particle and wave measurements are described in Appendix B.

This thesis is arranged with four main chapters. In Chapter 1, details of the S3-3 REP event morphology are presented. A description of the objective and subjective analysis of the data needed to identify events is contained in Appendix C.

Chapter 2 contains a general background of the types of mechanisms previously postulated and others newly discovered capable of explaining REPs. The mechanisms are compared with the classes of events found in Chapter 1. In addition, a case study of each type is presented with the view of advancing the knowledge of each mechanism.

Chapter 3 compares the atmospheric energy deposition profiles for REP and REP X-rays with profiles from other sources.

Profiles for REPs of various classes and intensities are presented as well as determinations of their importance and consequences in ionization production and ozone chemistry. The electron and X-ray deposition calculations are presented in Appendix D.

Chapter 4 summarizes the findings and discusses the conclusions and outlooks.

CHAPTER 1. EVENT MORPHOLOGY

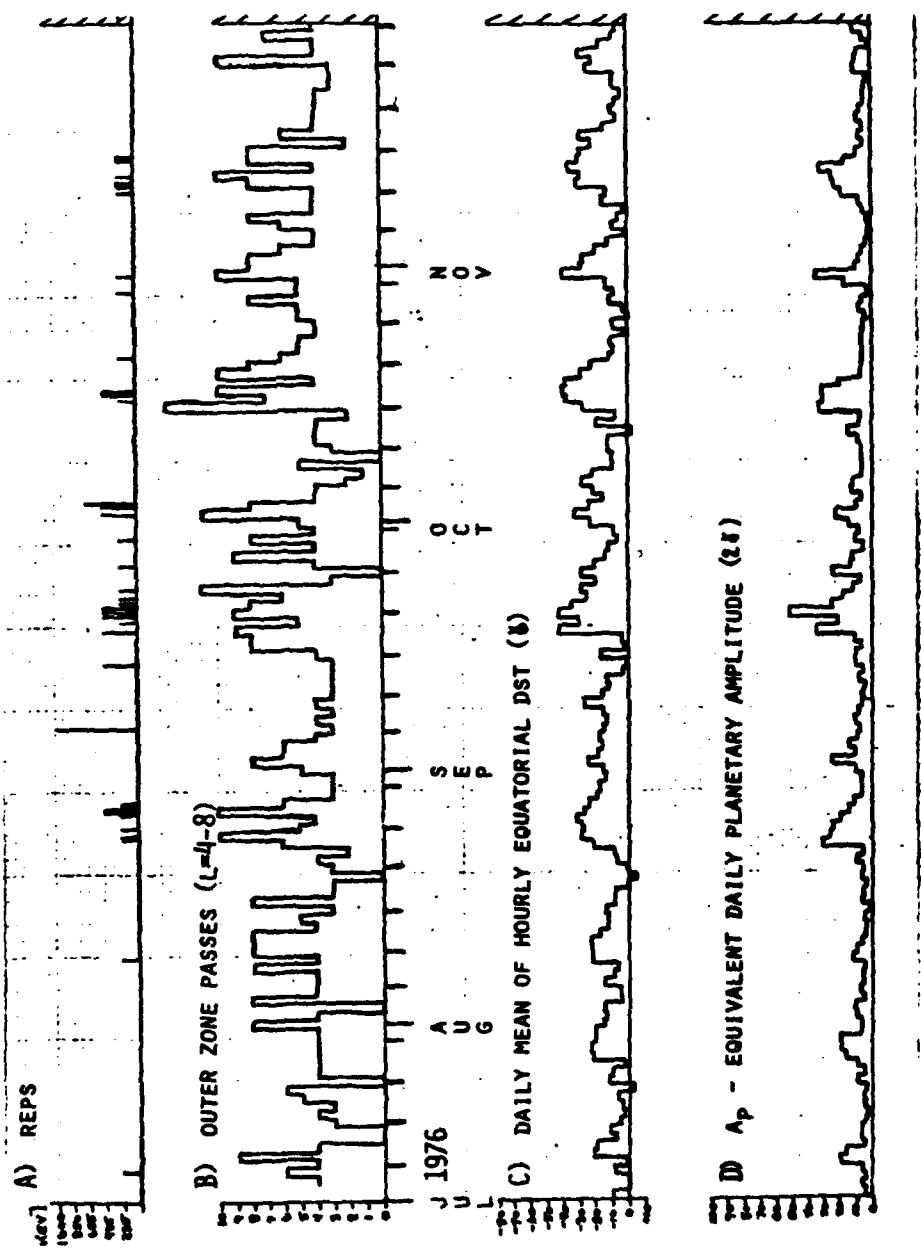
1.1 Results

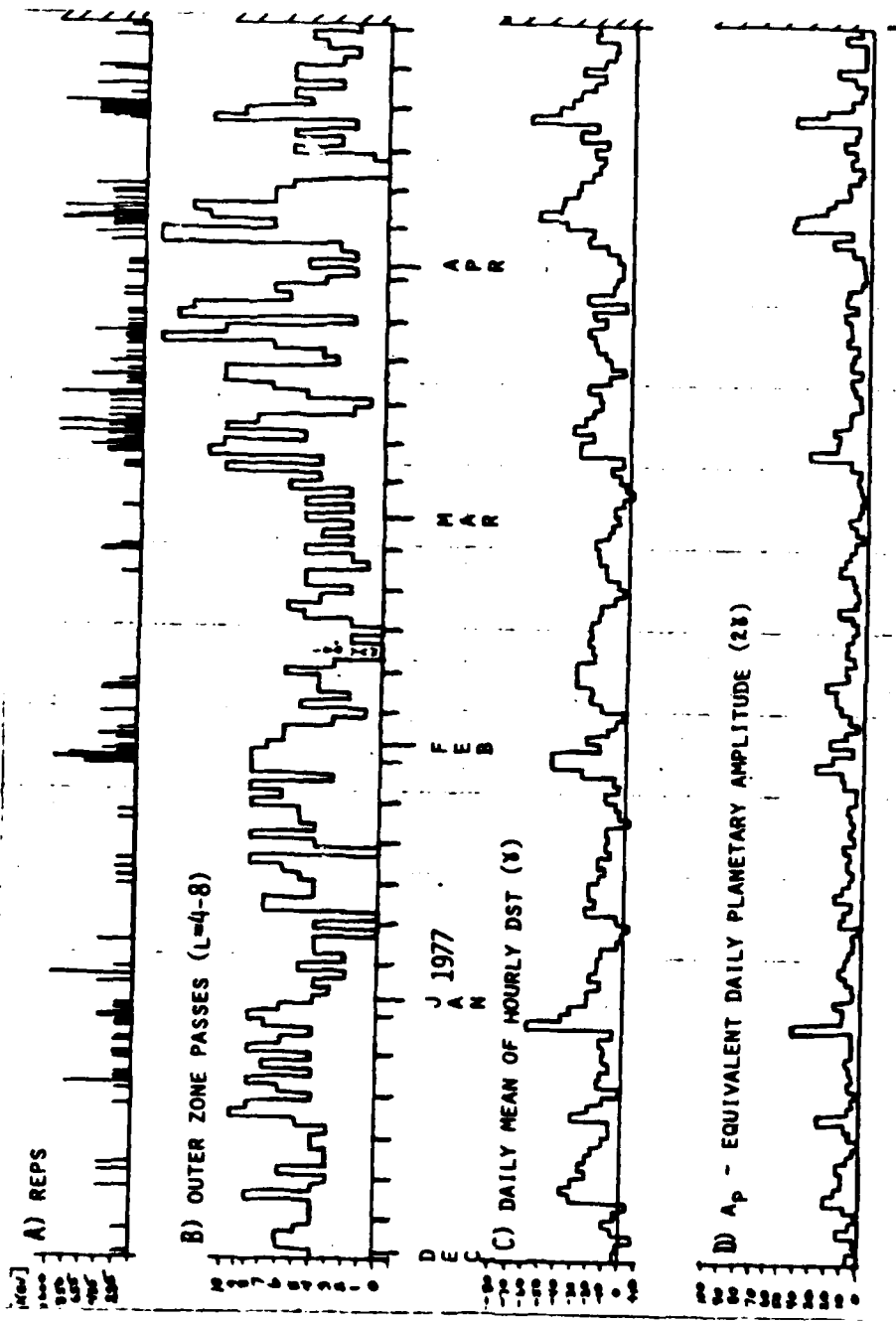
We present here the statistical data on REP occurrence measured on S3-3 during a 14 month period from 11 July 1976 to 29 September 1977. The S3-3 data collection specifications are described in Appendix B. The methodology of selecting REPs as well as the data base description is provided in Appendix C. Fig. 1.1, a-c, shows a summary of the S3-3 data set; the plot includes daily values of the highest energy channel (in keV) exhibiting strong diffusion precipitation (Panel A); the daily number of outer zone (L=4-8) passages (Panel B); the daily mean of the hourly equatorial DST in gamma units (Panel C); and the daily A_p , equivalent daily planetary magnetic amplitude in 2 gamma units (Panel D). Outer zone coverage (Panel B) was fairly uniform throughout the data set (although 29 days had no coverage) averaging a little over 5 outer zone passages per day. Coverage was a function of tape recorder capacity, power limitations, microfiche availability, etc. DST and A_p are presented to represent the relative magnetic condition of the magnetosphere since AE, the Auroral Electrojet Index, was not yet available for this time period. AE indices for five stations, AE (5), for seven 2 day periods in 1977, however, were available and are discussed in the next section. (Allen, private communication, 1979.)

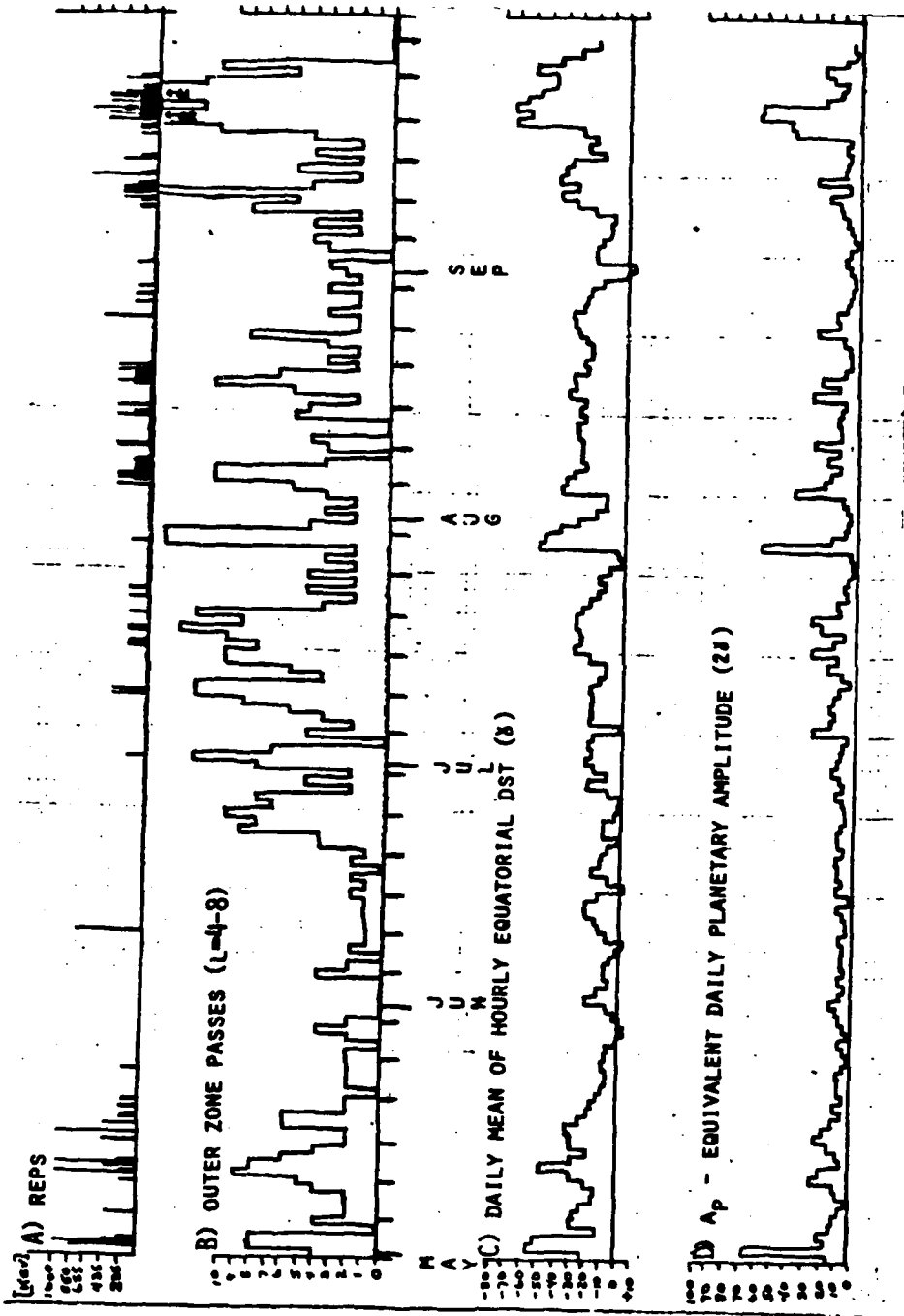
1.2 Frequency of Occurrence and Magnetic Dependence

The monthly average frequency of occurrence of REP events observed during the satellite transit of the outer radiation zone (L=4-8)

FIGURE 1.1a,b,c. Daily values from 11 July 1976 to 29 September 1977. Plots are given of REPs with highest energy channel exhibiting strong diffusion (Panel A); the total number of outer zone passings, L=4-8 (Panel B); daily mean of the hourly Dst, γ (Panel C); A_p , equivalent daily planetary magnetic index, 2γ (Panel D).







is represented in Fig. 1.2. Maximum occurrence is found near both (1976, 1977) autumnal equinoxes, at and slightly beyond vernal equinox, and near winter solstice while pronounced minima are evident near both (1976, 1977) summer solstices. This is generally consistent with the equinoxial maxima and solstice minima found by Bailey (1968). In this sense, the winter maxima is anomalous, but approximately five strong magnetically disturbed periods took place in Dec 1976-Jan 1977 (see Fig. 1.1b) probably accounting for the rise in activity. The magnetic dependence is apparent when the same monthly percentage of REPs normalized to outer zone coverage is plotted for days when $A_p > 30$ (Panel A), $20 \leq A_p < 30$ (Panel B), $10 \leq A_p < 20$ (Panel C), and $A_p < 10$ (Panel D) in Fig. 1.3. The anomalous winter maximum in Dec is apparent for $30 < A_p < 20$ while the Jan maximum is seen only for $20 < A_p < 30$. The remaining maxima and minima are fairly consistent with magnetic conditions.

Plotting the REP occurrence frequencies for all 14 months for different ranges of A_p (magnetospheric disturbance) (Fig. 1.4), one sees a nearly linear increase in REP occurrence with magnetic disturbances.

Fig. 1.5 shows S3-3 coverage of seven 2 day periods by AE index (Allen, private communication, 1979). This data is merely presented to support the fact that not all substorms trigger REPs (note particularly Panels A and E) in agreement with the findings of Rosenberg et al. (1972) and Thorne and Larsen (1976). The local times of collections for both Panels A (1 outer zone pass per rev) and E (2 outer zone passes

FIGURE 1.2 Monthly REP event percentage of occurrence normalized to outer zone coverage for July 1976-September 1977.

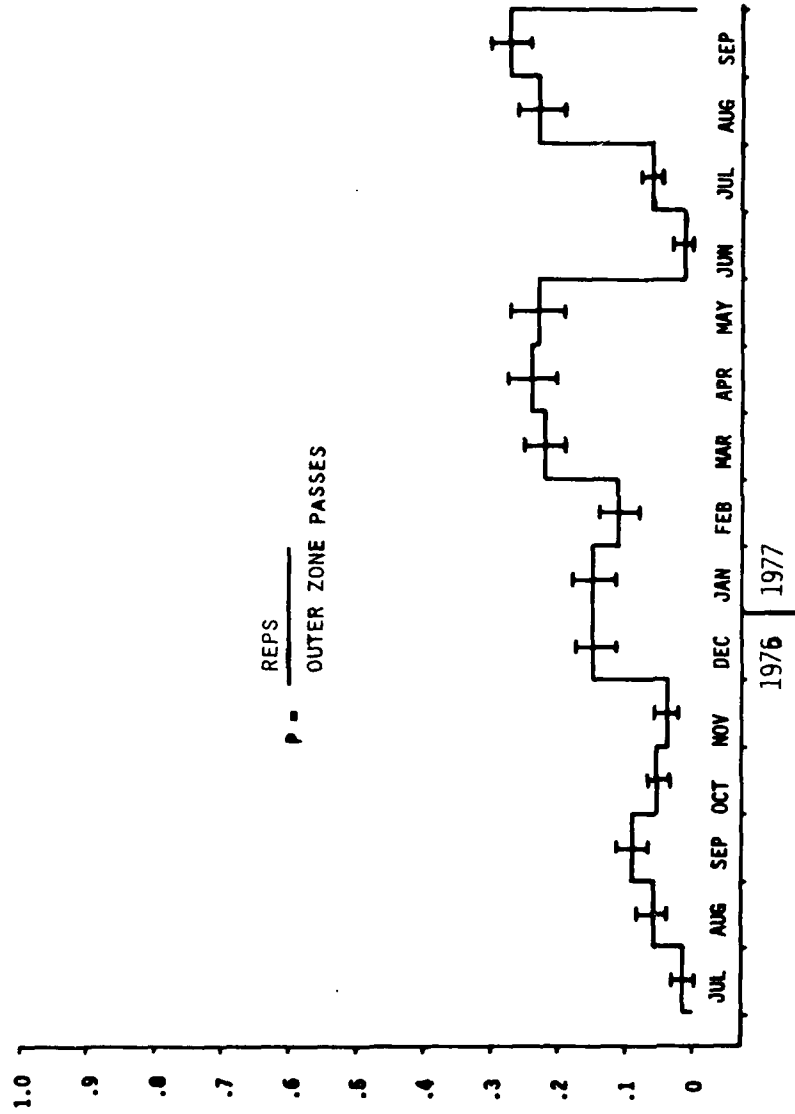


FIGURE 1.3 Monthly percentage of REPs normalized to outer zone crossings ($L=4-8$) for magnetic conditions: $A_p > 30$ (Panel A); $20 \leq A_p < 30$ (Panel B); $10 \leq A_p < 20$ (Panel C); $A_p < 10$ (Panel D). Values for JAN and FEB in Panels A and B overlap. $P=0$ for both JAN and FEB in Panel A and $P=0.86$ and 0.39 for JAN and FEB respectively in Panel B.

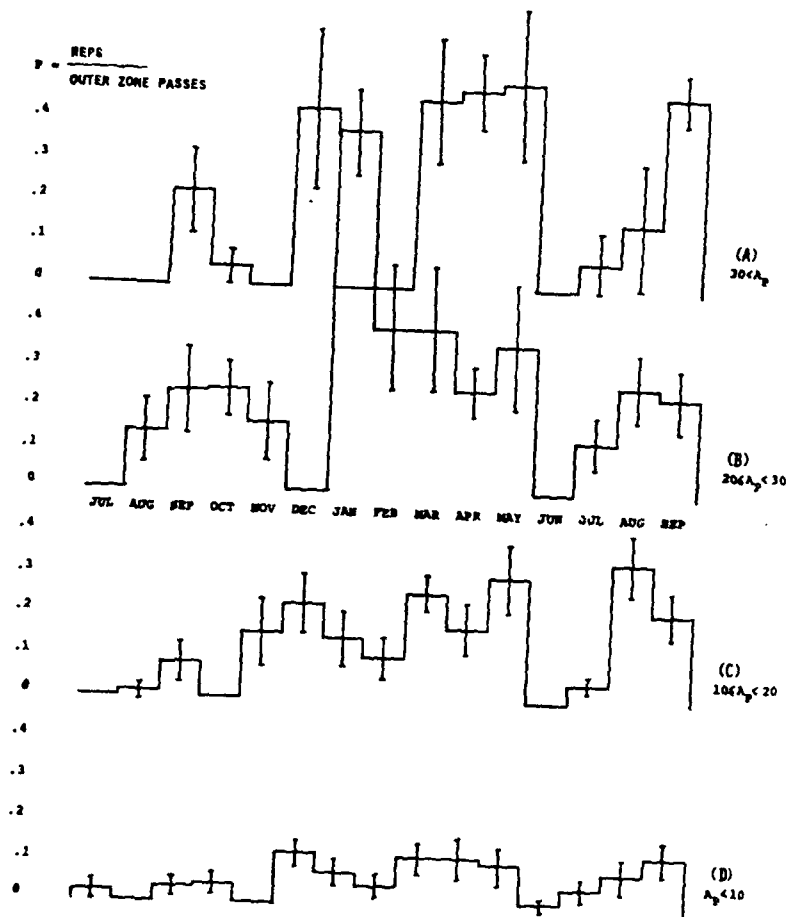


FIGURE 1.4 REP event percentage of occurrence normalized to outer zone coverage (L=4-8) versus magnetic conditions for all months JUL 1976-SEP 1977. Total number of outer zone passes for each category are in parentheses.

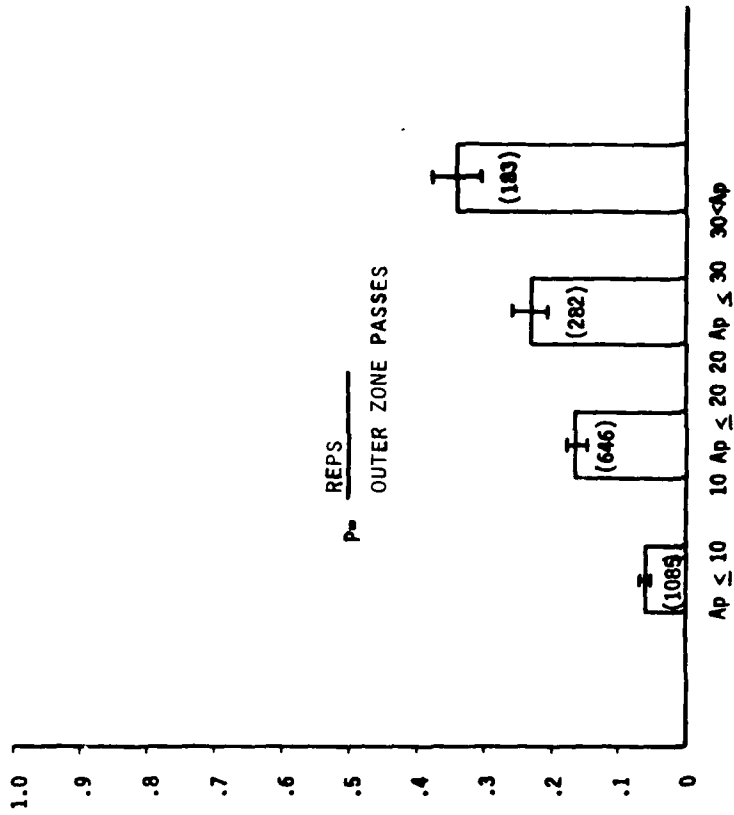
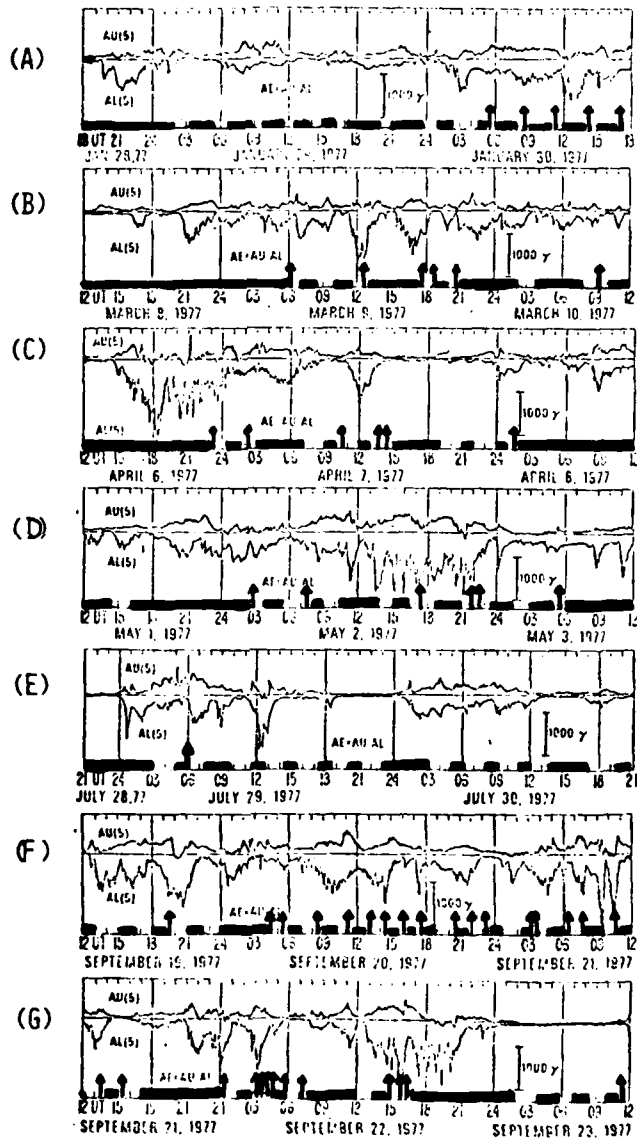


FIGURE 1.5 Seven 2 day periods of AE index (γ) throughout 1977 with REPs indicated by arrows (\uparrow) and hours of noncoverage by shaded horizontal bar (■). AE = AU (Upper Trace) - AL (Lower Trace).

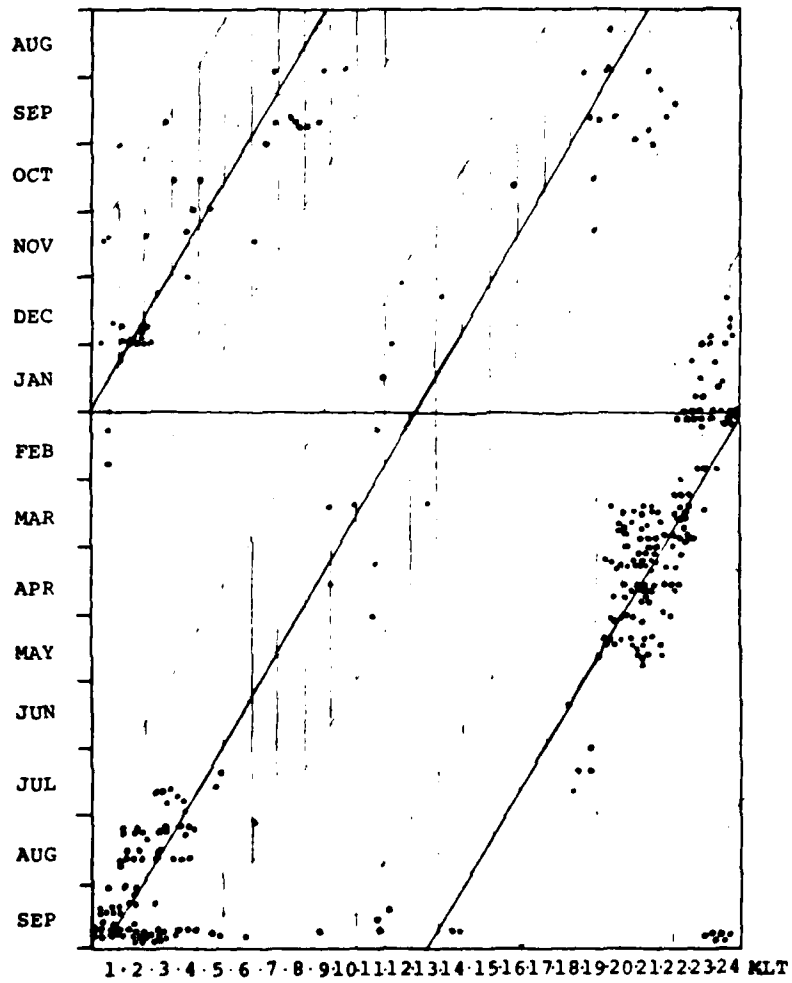


per rev) occurred near midnight where one would expect highest probability of detection (see Fig. 1.7 for example). Many of the outer zones, particularly the outer edge, in the passes of Panels A and E were very irregular showing trapped flux variations to nearly an order of magnitude though no precipitation was evident at the time of the pass. Also noticeable are times when substorms with large values ($\sim 900\gamma$) of AE do not spawn REPs (03-04L, 30 Jan 77, Panel A; 12-14L, 29 Jul 77, Panel E; 03-11L, 30 Jul 77, Panel E; 13-18L, 19 Sep 77, Panel F) while substorms corresponding to smaller AE may be associated with REPs (08-10L, 10 Mar 77, Panel B; 01-02L, 8 Apr 77, Panel C; 05-06L, 3 May 77, Panel D; 04-06L, 20 Sep 77, Panel F; 11-12L, 23 Sep 77, Panel G). This agrees again with Thorne and Larsen's (1976) finding of weak correlation between AE index and REP observations. An exception to this correlation is the REP activity at 03-04L, 22 Sep 77, Panel G, where a 1500 γ storm has associated REP activity for all four outer zone passes on the same revolution (see event catalogue, Appendix C, REV P 3569). This orbit was oriented in the 0200 and 1400 hrs MLT plane and could infer a longitudinal extent of nearly 10,000 km.

1.3 Local Time Dependence and Latitudinal Extent

To examine the local time morphology, REPs (Fig. 1.1, Panel A) were plotted for a given date versus magnetic local time (MLT) in Fig. 1.6. In order to test for any bias in the excursion of the orbit plane with time (skewed solid lines), we have sketched (hatched areas) the span of local times over which REPs could be observed in the range $4 \leq L \leq 8$ for a given local time orientation of this orbit plane. A

FIGURE 1.6 REPs by date versus MLT (hrs). Hatched area indicates regions of coverage ($4 \leq L \leq 8$) around the local time of the orbit plane (skewed solid line).



possible apogee-perigee bias in the orbit plane was investigated but after careful analysis (see Appendix C) the bias was found to be insignificant. The most striking feature is the pronounced night-day asymmetry with the greatest number of REPs occurring in the night sector. A plot of the ratio of day to night REPs is presented in Fig. 1.7 as a function of magnetically disturbed conditions. Overall there is a 9:1 preference for nighttime REP events. This local time asymmetry is less for highly disturbed ($A_p > 30$) conditions.

The actual local time morphology of REP events is shown in Fig. 1.8 (dots indicate very narrow localized events and lines are drawn for extended events) where they are plotted against invariant latitude for varying magnetic conditions. Most events show narrow latitudinal extends $\sim 0.5L$ (dots) while a few showed excursions up to 4-5L, mostly in the 235, 435 keV electron channels. Readily apparent is the relative lack of dayside REPs particularly during magnetically quiet days ($A_p \leq 10$). Furthermore, the scarcity of REPs over the local time range 1300-1900 remains apparent for $A_p \geq 10$. As can be seen from Fig 1.6 the orientation of the orbit plane for these MLTs covered roughly all months except the Mar-May maxima. Both autumnal equinoxes and the anomalous Dec-Jan maxima were covered so we believe that the statistics for the local afternoon hours are not biased by the orientation of the orbit plane: thus the tenuous afternoon statistics appear to be real. Also, for highly disturbed conditions ($A_p \geq 20$) the event location moves towards lower invariant latitudes; this is most noticeable for $A_p > 30$.

FIGURE 1.7. Ratio of day (MLT 06-18) to night (MLT 18-06) for varying magnetic conditions.

$$\text{RATIO} = \frac{\text{DAY (MLT 06-18) REP}}{\text{NIGHT (MLT 18-06) REP}}$$

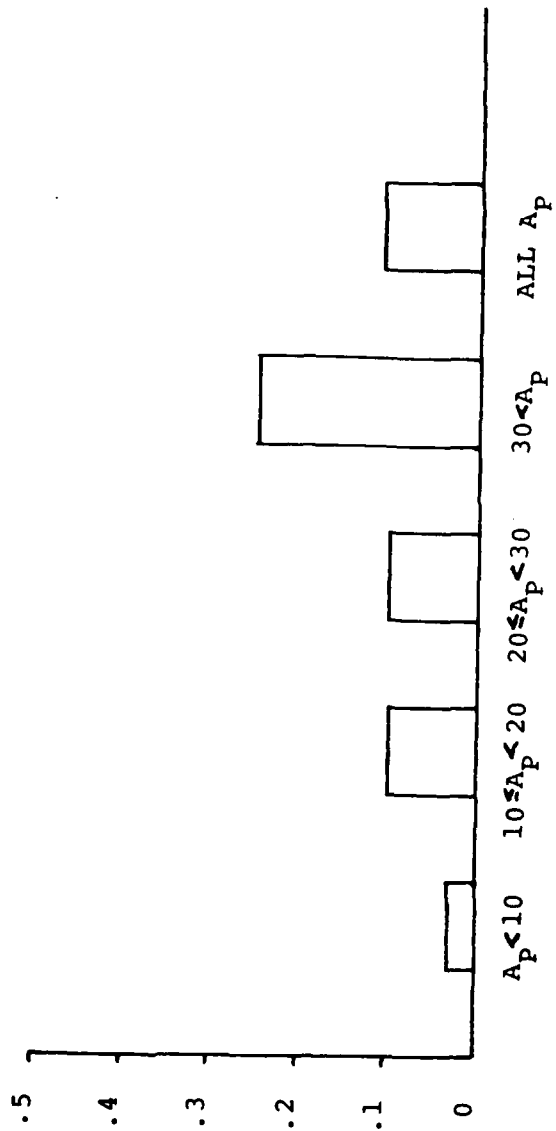
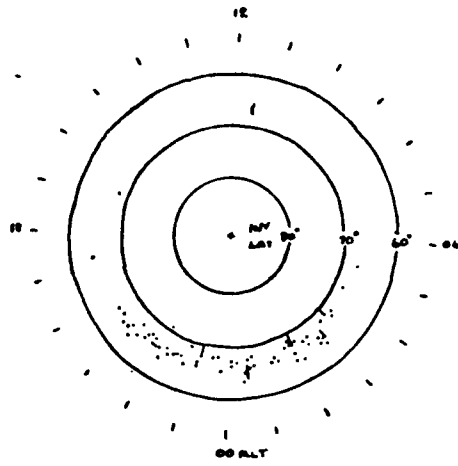
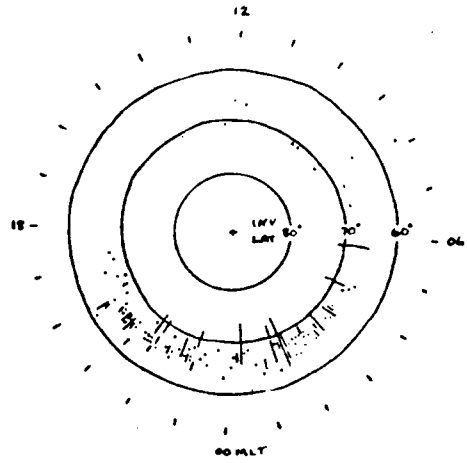


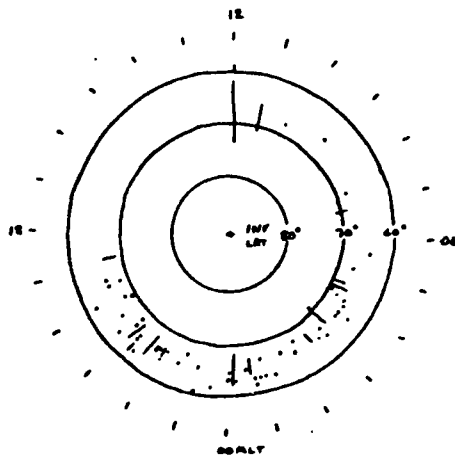
FIGURE 1.8 REPs displayed by invariant latitude (degrees) versus MLT (hrs) for varying magnetic conditions. Lines indicate excursions of events in invariant latitude.



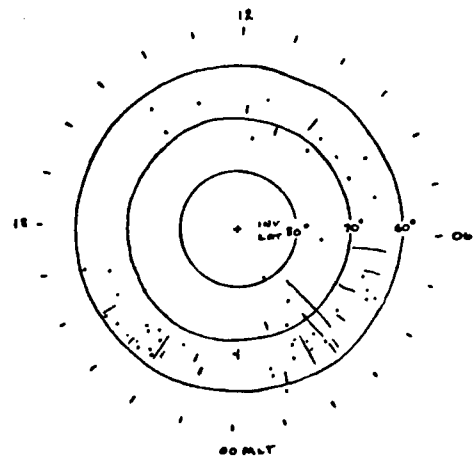
$A_p \leq 10$



$10 < A_p \leq 20$



$20 < A_p \leq 30$



$A_p > 30$

1.4 Longitudinal Extent

The slow (1.5 hr/month) excursion of the orbit plane in local time (see Appendix B) does not allow direct analysis of longitudinal extent. Estimates, however, can be made from previously published ground based events (e.g. Bailey, 1968) though these would probably be a lower limit. In addition, general limits could be estimated on an individual case basis from the satellite revisit times (~3 hrs) where consecutive coverage exists.

Duration of ground based REP events was observed to be 1-6 hrs with some events lasting for 24 hrs (Bailey, 1968). Several S3-3 REPs seemed to last for 24 hrs but continuity of these events is uncertain since events can occur over a period of many substorms (e.g. Fig 1.5, Panel F). The maximum longitudinal extent at 60° latitude is approximately 20,000 km. Bailey's 1 hr lower limit translates into 850 km which roughly agrees with Larsen's (1973) estimate of REPs extending greater than 1000 km. The 6 hour duration estimate translates to a 5,000 km longitudinal extent. For very disturbed times, all four outer zone crossings on the same 3 hour rev were on strong diffusion (see Fig 1.5, Panel (G), 22 Sep 1977, 0300-0500 UT) indicating longitudinal extents of 10000-15000 km.

1.5 REP - Proton Precipitation Correlation

In scanning through the data set it was immediately apparent that there was a high correlation between electron and proton precipitation (see catalogue remarks, Appendix C). Particularly noticeable was the similarity between precipitating fluxes in the 235 and 435 keV

electron and the 80 keV proton channels. An example of this feature is presented in Fig 1.9 which shows the 12, 33, 235 keV electron channel and 80 keV proton channels for 30 Jan 1977. The 12 keV electrons are precipitating from 19590 UT to 20565 UT. Precipitation features for 33 and 235 keV electrons and >80 keV proton are evident at 19646-19700 UT, 19815 UT, and 20475 UT (proton precipitation began at 19580 UT). This suggests the precipitation mechanism for both electrons and protons may be either related or identical. Protons did, however, precipitate much more frequently than electrons and their zone of strong diffusion precipitation generally extended to lower latitude. This frequency difference can be readily seen in Fig 1.10 where the revs (1-4 outer zone passes per rev) with strong electron ($E_{\text{precip}}^- > 235 \text{ keV}$) and proton ($E_{\text{precip}}^+ > 80 \text{ keV}$) precipitation normalized to the number of revs is plotted for increasingly disturbed magnetic conditions. Note that proton precipitation is seen on 65% of revs while electrons on 22% of revs for all magnetic conditions. The high occurrence of proton precipitation in the energy range (100-200 keV) has been previously observed (e.g. On ESRO 1A, Lindalen et al., 1971).

1.6 Classification of REP Events

Due to the strong correlation of REPs with proton (>80 keV) precipitation, the data set was analyzed for REPs without concurrent strong proton precipitation. Only 7 of 313 outer zone passes fell in this category. In addition, all 7 cases had concurrent 33 keV electrons precipitating on strong diffusion and were found at altitudes below 1500 km.

FIGURE 1.9 Flux versus UT plot of 12, 33, 235 keV electrons and
80 keV protons for rev 1656 30 Jan 1977

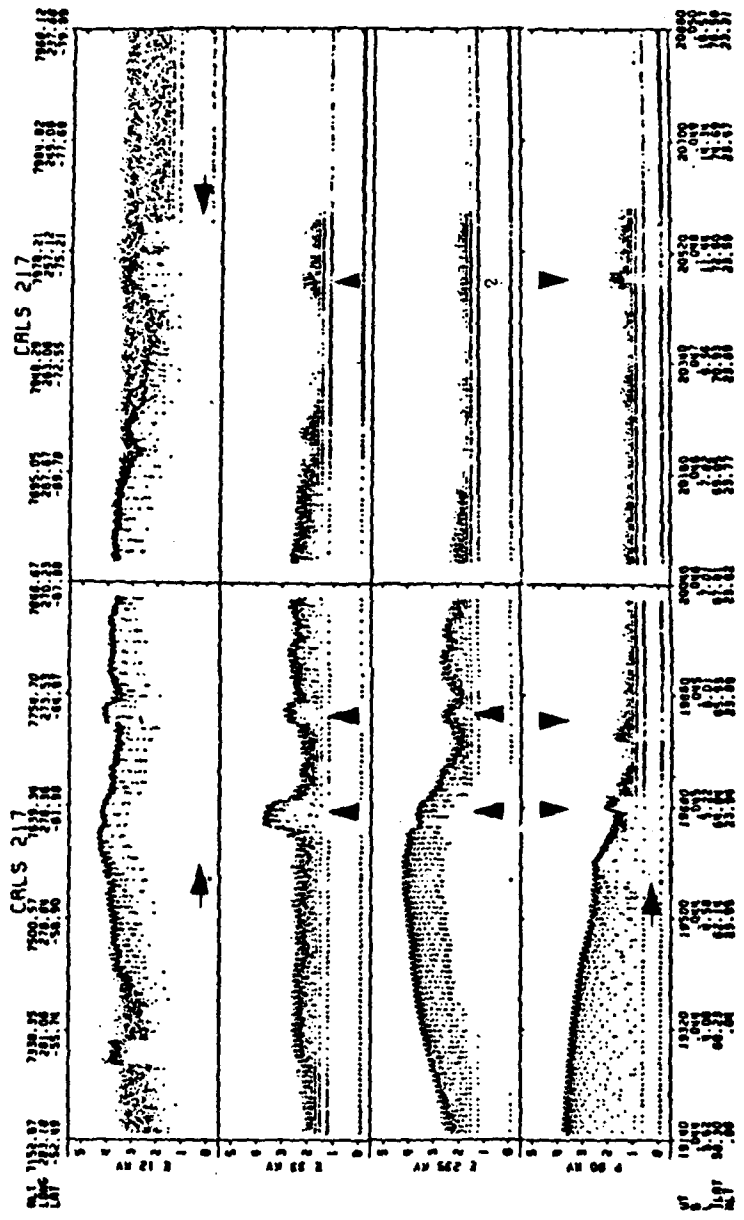
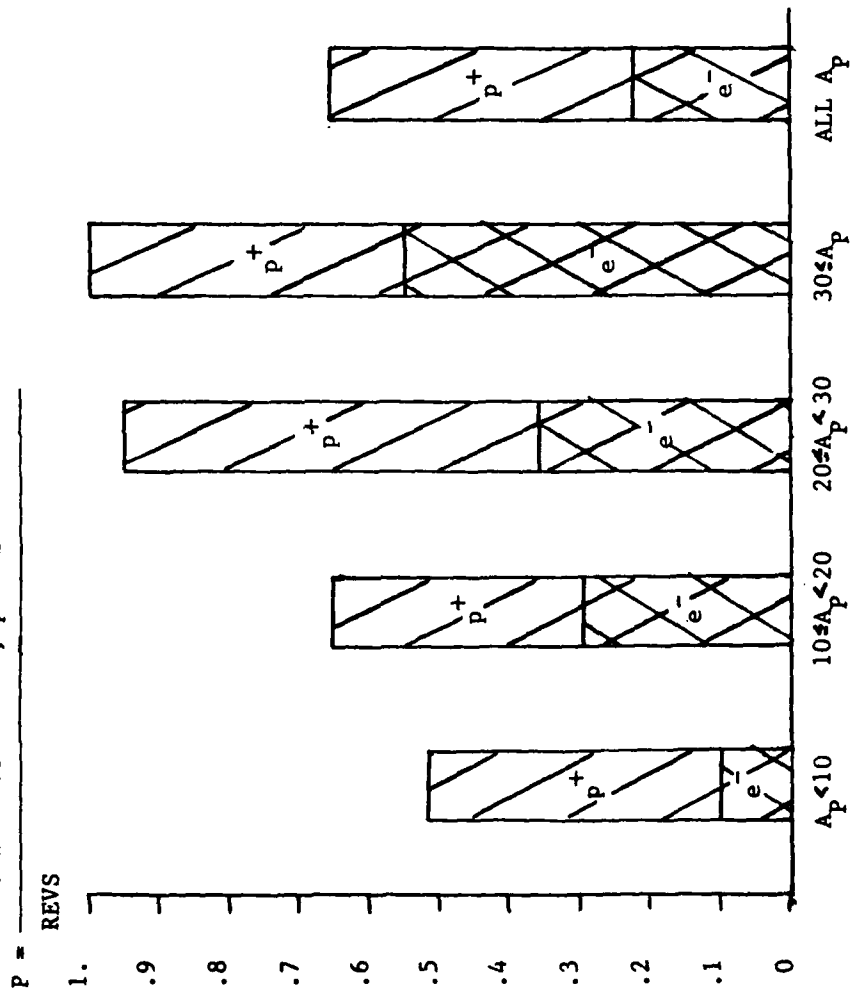


FIGURE 1.10 Percentage of revs (1-4 outer zone passes) with strong e^- or p^+ precipitation normalized to revs number by magnetic conditions.

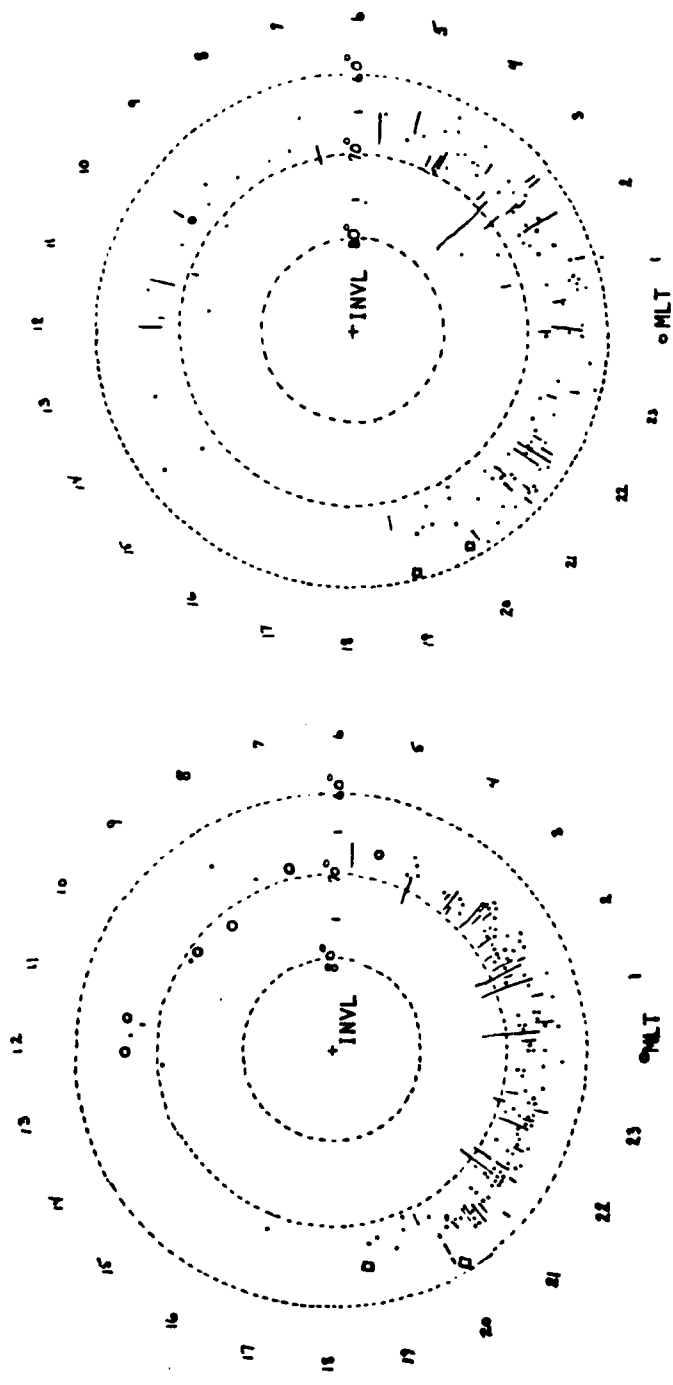
REVS WITH STRONG e^- , p^+ PRECIPITATION



Analysis of REP energy dependence revealed that the overwhelming majority of REPs had energies in the range 235-1000 keV on strong diffusion. Only 4 events had spectra such that only energies above 1 MeV were on strong diffusion. These same 4 cases were different in another way. They occurred equatorward of the outer edge of the outer zone; the remaining cases except for the seven above all occurred right at the outer edge or extended inward from the outer edge.

The four cases with strong precipitation only above 1 MeV were not among the seven cases which showed no concurrent proton precipitation. REPs could thus be divided into three distinct classes: those with strong diffusion precipitation below (302) and only above (4) 1 MeV with concurrent protons on strong diffusion and those with strong diffusion precipitation below 1 MeV but without strong proton precipitation (7). These classes are displayed in invariant latitude versus MLT in Fig 1.11 and will be discussed in detail in Chapter 2.

FIGURE 1.11 REP events displayed according to invariant latitude and MLT for the following classes: dots and lines (. |) for strong precipitation at energies ≤ 1 MeV with concurrent proton precipitation; squares (\square) for strong precipitation at energies ≥ 1 MeV with concurrent proton precipitation; and circles (o) for strong precipitation at energies $\ll 1$ MeV without concurrent proton precipitation.



$A_p > 20$

$A_p \leq 20$

CHAPTER 2. THEORY AND CASE STUDIES

2.1 Background

To account for the observed subauroral radio signal absorption, Bailey (1968) estimated that the precipitating electrons at energies exceeding several hundred keV required fluxes $J_p \geq 10^3$ electrons/cm²s str. Such intense subauroral (closed field lines) precipitation requires pitch angle scattering of trapped radiation belt electrons near the strong diffusion limit. This leads to an isotropic pitch angle distribution, i.e., precipitated flux \approx trapped flux (Kennel, 1969). The rapid scattering can occur during resonant interaction between plasma waves and particles in the magnetosphere. Either of the first two adiabatic invariants (μ and J) can be violated during the pitch angle diffusion and this can provide a major loss mechanism for geomagnetically trapped particles (Thorne, 1975).

The condition for first order resonant scattering with relativistic electrons is:

$$1 - \mu_{||} \beta_{||} = \frac{\Omega_{-}}{\omega \gamma} \quad (1)$$

where $\mu_{||} = \frac{k_{||} c}{\omega}$, parallel (to ambient magnetic field) refractive index

$\beta_{||} = \frac{v_{||}}{c}$, normalized parallel electron velocity

ω = wave frequency

$\frac{\Omega_{-}}{\gamma}$ = relativistic electron gyrofrequency

$\gamma = (1 - \beta^2)^{-1/2}$, relativistic mass enhancement factor

c = speed of light in vacuo

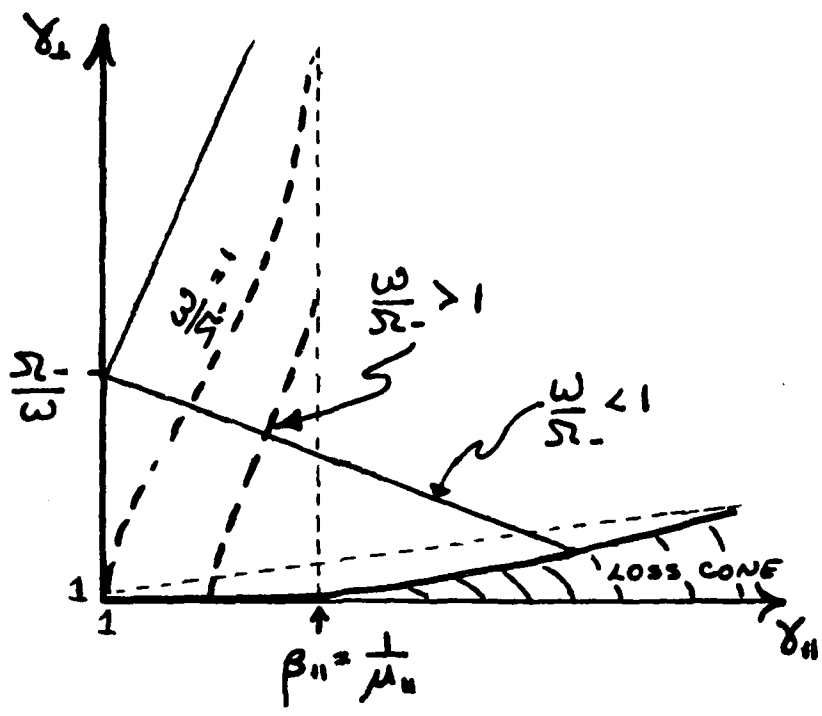
k_{\parallel} = parallel propagation constant

(wave number)

Thorne (1977) has scaled the general morphology of velocity space surfaces for resonance with a wave of a given frequency and refractive index (Fig 2.1). Waves with $\mu_{\parallel} > 1$ generally require $\omega < \Omega_{-}$ to resonate with relativistic electrons because only this surface ($\frac{\omega}{\Omega_{-}} < 1$) reaches the loss cone at relativistic energies. Waves with $\omega > \Omega_{-}$ even with $\mu_{\parallel} < 1$ as in the case of Auroral Kilometric Radiation (Kurth et al., 1975) do not usually resonate with relativistic electrons. A brief review of plasma waves follows.

Southwood (1978) has reviewed the plasma waves found in the magnetosphere. Four distinct classes of waves have been identified in the outer radiation zone $4 \leq L \leq 8$ where REP events are observed. The first class are electron electromagnetic waves at frequencies below the electron gyrofrequency; waves typically occur in the frequency band $.1 \leq \frac{\omega}{\Omega_{-}} \leq .5$. These so-called right hand polarized (Burton and Holzer, 1974) whistler mode "chorus" waves are generated at high altitudes (near the magnetic equator) outside the plasmasphere with typical wide band wave amplitudes ~ 10 m γ (Dunckel and Helliwell, 1969, Anderson and Maeda, 1977). The waves propagate down the field lines to low altitudes where they are the most frequently occurring emission outside the plasmopause (Kennel and Thorne, 1967; Thorne and Kennel, 1967; Gurnett et al., 1969; Dunckel and Helliwell, 1969; Tsurutani and Smith, 1974). Dunckel and Helliwell (1969) examined emission

FIGURE 2.1 Velocity space surfaces for resonance with a wave of frequency ω and parallel refractive index μ_{\parallel} for γ_{\perp} versus γ_{\parallel} .



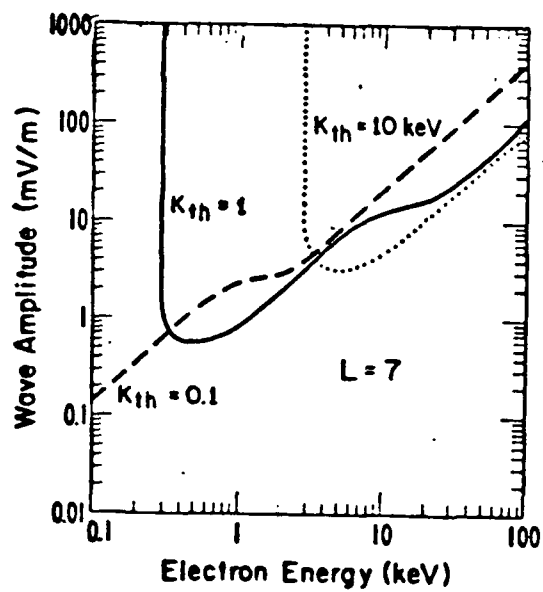
intensities for 300-500 Hz whistler mode chorus and found that they peak near 1000 MLT. Also, just outside the plasmopause, intensities increase abruptly at 0600 MLT and decrease gradually after 1400 MLT in agreement with OV3-3 1.3 kHz low altitude emissions (McPherson and Koons, 1970). A one year study of OGO-5 search coil magnetometer data show that equatorial chorus is strongly substorm correlated with two peaks: one at post midnight (L=5-8) and a second in the dawn to noon sector (L=7-11) (Tsurutani and Smith, 1977). The source of the waves is generation by loss cone instability associated with the pitch angle anisotropy of substorm injected electrons (Kennel and Petschek, 1966). Association of this wave phenomenon with precipitation of energetic electrons (10-100 keV) has been made by several authors (Dungey, 1963; Cornwall, 1964; Kennel and Petschek, 1966; Vampola et al., 1971; Thorne, 1974; Tsurutani and Smith, 1977).

The second class of waves are electron electrostatic waves found at frequencies above the electron gyrofrequency (Kennel et al., 1970). These so-called "n + 1/2" waves ($\frac{\omega}{\Omega_e} \simeq n + 1/2$) are found throughout the magnetosphere outside the plasmopause (Kennel et al., 1970; Shaw and Gurnett, 1975) and are thought to be generated by $\frac{\partial f}{\partial v} > 0$ electron distributions (Fredricks, 1971). The waves are localized about the geomagnetic equator from L=4-10 a majority of the time in the morning sector (Kennel et al., 1970; Fredericks and Scarf, 1973). Wave amplitudes range from 1-10 mV/m (Kennel et al., 1970) with excursions to 100 mV/m during substorms (Scarf et al., 1973). Theoretical parametric studies yielding realistic spatial growth rates

for these waves have been accomplished (Young et al., 1973; Karpman et al., 1975; Ashour-Abdalla and Kennel, 1976) and computations of the resonant electron population show energies of 1-10 keV (possibly accounting for the diffuse aurora, Lyons, 1974) or higher (~ 100 keV) during substorms (Scarf et al., 1973). However, this class of waves does not resonate effectively with relativistic electrons in strong diffusion, requiring wave amplitudes in excess of 100 mV/m to cause strong diffusion of 100 keV electrons (see Fig. 2.2, Lyons, 1974). This class will not be considered further in this study.

Analogous to the right hand electron electromagnetic waves which cut off at Ω_- , there is a third class, the left hand ion electromagnetic waves below Ω_+ (e.g., Cornwall, 1965). In the micropulsation literature these are sometimes referred to as "PCI or IPDP micropulsations" or periodic emissions: PCI ("Pearls") are dispersive waves propagating along field lines thought to be generated by the loss cone distribution of ring current protons just inside the plasmapause (Cornwall, 1965; Cornwall et al., 1970; Williams and Lyons, 1974; Kaye et al., 1979); "IPDP," intervals of pulsations of diminishing period are non-dispersive waves propagating near dusk during substorm expansion (McPherson et al., 1968; Lukkari and Kangas, 1976; Bossen et al., 1976). In a statistical study by Bossen et al. (1976) EMIC occurred mostly about the dusk meridian with frequencies $\frac{\omega}{\Omega_+} \sim .1-.2$ and wave amplitudes 1-7 γ during the main phase of geomagnetic storms. PCI turbulence has also been observed in detached plasma regions (Kivelson et al., 1972) prevalent in the late morning to dusk sector (Chappell,

FIGURE 2.2 Minimum wave amplitudes required for electron strong diffusion scattering at L=7 versus electron energy for electron thermal energies 0.1, 1, 10 keV (Lyons, 1974)



1974). Theoretical calculations show that these waves are candidates for hard REPs (precipitating electron energies ≥ 1 MeV) in the evening plasmopause (Thorne and Kennel, 1971) and possibly during the day in detached plasmas (Thorne, 1974).

The last class of waves to be discussed here are the electrostatic ion cyclotron waves predicted by Kennel et al. (1970) after the analogous electrostatic electron wave discovery and measured by Gurnett and Frank (1977) and Kintner et al., (1978). These waves have frequencies $\Omega_+ < \omega < \Omega_{LHR}$, amplitudes on the order of a few 10's of mV/m, and are found at all local times at $L > 6$ (Gurnett and Frank, 1977; Kintner et al., 1978). Generation of the waves is thought to be by current driven instability low on the field line to yield resonant Landau growth (Kindel and Kennel, 1971; Ashour-Abdalla and Thorne, 1978) and/or equatorial anisotropic pitch angle instability due to distributions of hot protons (Coroniti et al., 1972; Ashour-Abdalla and Thorne, 1977; Ashour-Abdalla and Thorne, 1978). Observations have indeed correlated REPs with intense ELF electrostatic waves (Koons et al., 1972) and shown that some cases can be explained only as a low altitude phenomenon (Vampola, 1977).

2.2 Resonant Populations

Since the observational data covers the electron energy range 12 keV-1.6 MeV and proton range > 80 keV, > 150 keV etc., it is important to determine the electron and proton resonant populations for each wave phenomena if any success at discrimination (between wave mechanism) is to be accomplished. We proceed with calculations for

each wave type.

2.2.1 Whistler Mode Chorus

The cold plasma dispersion relation for parallel whistler mode chorus propagation assuming $\Omega_+ \ll \omega \ll \Omega_-$ and $\mu_{||} \gg 1$ ($\omega_p \gg \Omega_-$) is given by Kennel and Petschek (1966):

$$\mu_{||}^2 \approx \frac{\omega_p}{\omega \Omega_-} \left(1 - \frac{\omega}{\Omega_-}\right)^{-1} \quad (2)$$

where ω_p = plasma frequency.

Using this dispersion relation and the condition for first order cyclotron resonance with relativistic electrons (eq. 1) assuming $\beta_{||} \approx \beta$ (valid near the loss cone), Thorne (1974) has derived an expression for resonant electron energies:

$$\left(\gamma^2 - 1\right)_{res} \approx \frac{2 E_m^- \Omega_-}{E_0^- \omega} \left(1 - \frac{\omega}{\Omega_-}\right)^3 \quad (3)$$

where $E_m^\pm = \frac{B^2}{8\pi N^\pm}$, magnetic energy per particle

B = local magnetic field value

N^\pm = particle number density

$E_0^\pm = m^*c^2$, particle rest energy.

Thus the expression for the resonant electron population with whistler mode chorus waves of frequency ω is:

$$E_r^- = \left[\left(\frac{2 E_m^- \Omega_-}{E_0^- \omega} \left(1 - \frac{\omega}{\Omega_-}\right)^3 + 1 \right)^{1/2} - 1 \right] E_0^- \quad (4)$$

where E_r^- is the resonant electron energy.

Since $B \approx \frac{B_0}{L^3}$ at the equator where $B_0 = .311$ Gauss and N varies from approximately $10^3/\text{cm}^3$ in the plasmasphere ($L=4$) to $1-10/\text{cm}^3$ in the plasma trough region ($L=5-8$) for a $K_p \sim 2-3$ (Chappell et al., 1970), we have a range of E_m^- of .5 keV ($L=4$, $N=10^3$) to 150 keV ($L=5$, $N=1$) at the equator. Down on the field line (away from the equator) where B increases and N is less than ionospheric densities ($\leq 10^4$) the value of E_m^- can reach a few 100 keV. Values of N not greater than 10^4cm^{-3} are used since whistler mode chorus propagating down the field line rapidly becomes perpendicular to the local magnetic field line and reflects at frequencies near the lower hybrid resonance (Thorne and Kennel, 1967). [For propagation at angles to the field line, the cumulative contribution from parallel velocity resonances is appropriate since higher order resonances become important (Stix, 1962; Kennel and Petschek, 1966) although this is not the case for electromagnetic waves with $k_{\perp} \approx 0$.] So for $.1 \leq \frac{\omega}{\Omega_-} \leq .5$ (Tsurutani and Smith, 1977) and $E_m^- = .5-150$ keV (both typical equatorial values), we have $E_r^- = 1-800$ keV. Off the equator (down the field line where $\frac{\omega}{\Omega_-}$ is lower and E_m^- is higher) E_r^- is expected to reach values of nearly 10^3 keV.

For proton precipitation we use the resonance condition for first order ($n=1$) cyclotron resonance with protons

$$1 - \mu \beta_{\parallel} = \frac{\Omega_+}{\omega} \quad (5)$$

and the dispersion relation for whistlers to obtain

$$E_r^+ \approx \left(\frac{m^+}{m^-} \right) E_m^+ \left(\frac{\omega}{\Omega_-} \right) \left(1 - \frac{\omega}{\Omega_-} \right) \quad (6)$$

where $\frac{m^+}{m^-} = 1.837 \times 10^3$, proton to electron mass ratio. We take $E_m^+ \approx E_m^-$, since we assume the densities of electrons to be the same as protons (DeForest and McIlwain, 1971). Using the same range of values for $\frac{\omega}{\Omega_-}$ we can readily see that E_r^+ will typically be greater than an MeV (except in the dense plasmasphere) in agreement with Kennel and Petschek (1966).

Resonance energies for both electrons and protons are shown in Fig. 2.3. One readily sees that electrons can be resonant at energies up to several hundred keV while protons are resonant only at energies above 1 MeV.

2.2.2 Electromagnetic Ion Cyclotron (EMIC) Waves

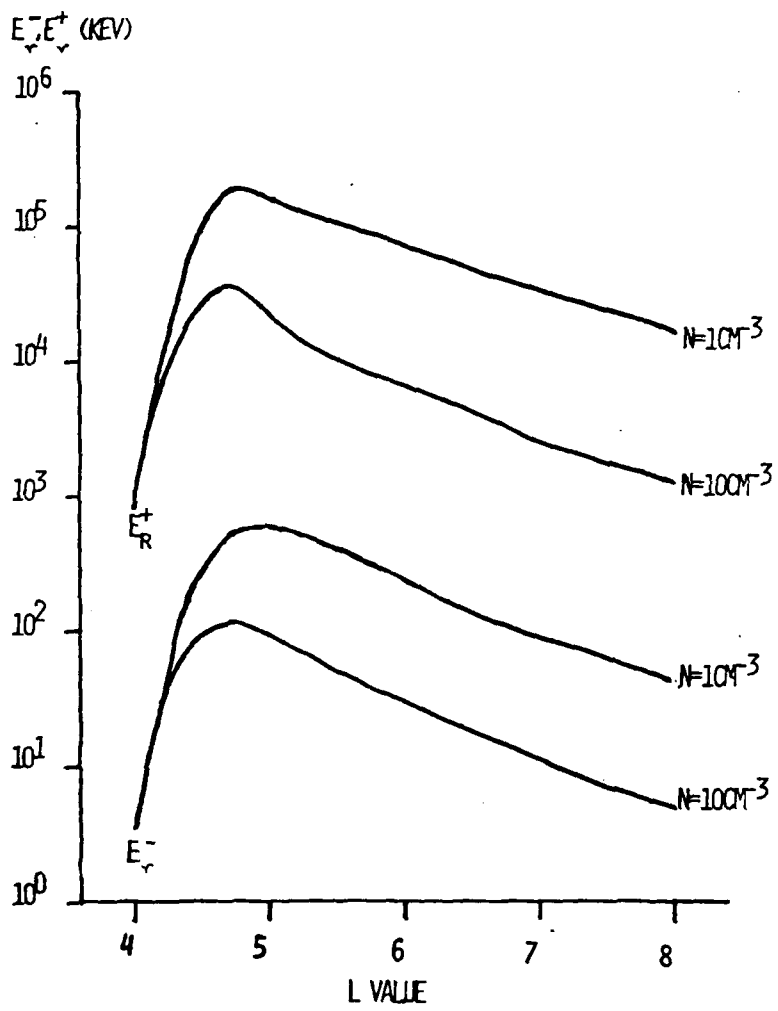
The cold plasma dispersion relation for parallel propagating EMIC waves below Ω_+ is given by Kennel and Petschek (1966):

$$\mu_{||}^2 = \frac{\omega_p^2}{\Omega_- \Omega_+} \left(1 - \frac{\omega}{\Omega_+} \right)^{-1} \quad (7)$$

Thorne and Kennel (1971) and Thorne (1974) have used equations (1) and (7) to derive the equation for resonant electrons with EMIC:

$$\left(\gamma^2 - 1 \right)_{res} \approx \frac{2 E_m^-}{E_o} \left(\frac{m^+}{m^-} \right) \left(\frac{\Omega_+}{\omega} \right)^2 \left(1 - \frac{\omega}{\Omega_+} \right) \quad (8)$$

FIGURE 2.3 Electron and proton resonant energies for the whistler mode chorus mechanism ($.1 \leq \frac{\omega}{\Omega} \leq .5$) [in keV] versus L for $N=10^4 \text{ cm}^{-3}$ for $L=4$ and $N=1, 10 \text{ cm}^{-3}$ for $L \geq 5$.



and we see that (8) becomes

$$E_r^- = \left\{ \left[\frac{2 E_m^- m^+ \Omega_+^2}{E_o^- m^- \omega^2} \left(1 - \frac{\omega}{\Omega_+} \right) + 1 \right]^{1/2} - 1 \right\} E_o^- \quad (9)$$

Using the previous value of E_m^\pm at $L=4$, $N=10^3 \text{ cm}^{-3}$, and $\frac{\omega}{\Omega_+} \approx .1-.5$ (Bossen et al., 1976) we see that $E_r^- \sim$ few MeV, see Fig 2.4. Even during intense injection events where EMIC waves may be generated outside the plasmasphere to $L=8$ due to enhanced ion fluxes and larger loss cone distributions (Coroniti et al., 1976), E_r^- for $L=8$, $N=10$, and $\frac{\omega}{\Omega_+} \sim .2$ is 5 MeV.

The equation for EMIC resonance with protons is given by Kennel and Petschek (1966):

$$E_r^+ = E_m^+ \left(\frac{\Omega_+}{\omega} \right)^2 \left(1 - \frac{\omega}{\Omega_+} \right)^3 \quad (10)$$

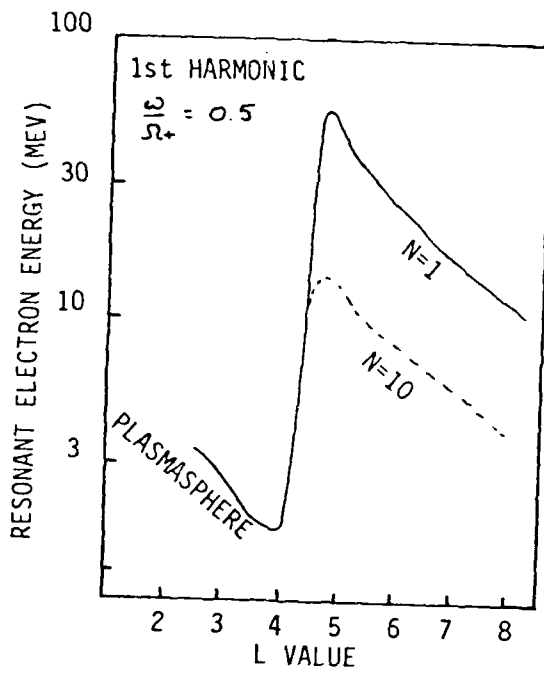
Using the same E_m^\pm values and $\frac{\omega}{\Omega_+} = .1-.2$ (Bossen et al., 1976) we see $E_r^+ \approx 10\text{s}-100\text{s}$ keV near the outer edge of the plasmasphere.

In summary, EMIC waves resonate with electrons only at energies $\gg 1$ MeV and only near 1 MeV in the relatively dense plasmasphere or detached plasma regions where N is relatively large and at the equator where B is lowest. The resonant proton population for the same wave mechanism is on the order of 10's and 100's keV.

2.2.3 Electrostatic Ion Cyclotron (ESIC) Waves

Since the frequency of ESIC waves is near the Ω_+ , we can

FIGURE 2.4 Resonant electron energies versus L for the ion cyclotron ($\frac{\omega}{\Omega} = 0.5$) wave mechanism for densities $N=1, 10 \text{ cm}^{-3}$ outside the plasmasphere and $N=10^3 \text{ cm}^{-3}$ $L=4$ (Thorne and Kennel, 1971).



eliminate the unity term in (1) to obtain:

$$\gamma^2 \mu_{ii}^2 \beta_{ii}^2 = \left(\frac{\Omega_-}{\omega} \right)^2 \quad (11)$$

Using $\gamma^2 \beta_{ii}^2 \simeq \gamma^2 - 1$ (valid for particles near the loss cone) in (11) we have:

$$\gamma^2 - 1 = \mu_{ii}^{-2} \left(\frac{\Omega_-}{\omega} \right)^2 \quad (12)$$

By writing the wave refractive index $\mu_{ii}^2 = \frac{c^2 k_{ii}^2}{\omega^2} \frac{f_+^2 \Omega_+^2}{v_{th}^{+2}}$, where Ω_+ is the ion gyroratio, v_{th} is the ion thermal speed, and $f_+ = v_{th}/\Omega_+$, eq. (12) becomes:

$$\left(\gamma^2 - 1 \right)_{res} \simeq \left(\frac{\Omega_-}{\Omega_+} \right)^2 \left(\frac{v_{th}^+}{c} \right)^2 \left(\frac{1}{k_{ii} f_+} \right)^2 \quad (13)$$

Using $\left(\frac{v_{th}^+}{c} \right)^2 = 2 \frac{E_{th}^+}{E_0^+}$, where E_{th}^+ is the ion thermal energy, in (13) we have:

$$\left(\gamma^2 - 1 \right)_{res} = \frac{2 E_{th}^+}{E_0^+} \left(\frac{m^+}{m^-} \right)^2 \left(\frac{1}{k_{ii} f_+} \right)^2 \quad (14)$$

The resonant electron energy for ESIC waves is thus:

$$E_{\nu}^- = \left\{ \left[\frac{2 E_{th}^+}{E_0^+} \left(\frac{m^+}{m^-} \right)^2 \left(\frac{1}{k_{ii} f_+} \right)^2 + 1 \right]^{1/2} - 1 \right\} E_0^- \quad (15)$$

Ashour-Abdalla and Thorne (1977 and 1978) have solved the electrostatic dispersion relation numerically and obtained k space contours of temporal and convective growth rates normalized to f_+ for typical equatorial and ionospheric conditions. If waves are excited by the current driven instability in the topside ionosphere with electron drifts just above the Kindel and Kennel (1971) marginal stability limit, 1 Re above the earth for $L \simeq 10$, $E_{th} \simeq 2\text{ev}$, the first three harmonics (typical of wave observations, Kintner et al., 1978) have typical values for $k_{||} f_+ \simeq .26-.1$ (Ashour-Abdalla and Thorne, 1978). This yields values of $E_r^- \simeq 50-300 \text{ keV}$. For excitation in the equatorial plasmashet by the loss cone instability, typical equatorial parameters are $E_{th} \simeq 1 \text{ keV}$, and $k_{||} f_+ \simeq .4-.5$ for the first five harmonics. This gives $E_r^- \simeq 100 \text{ keV} - 2 \text{ MeV}$. Again using the first order cyclotron relation for resonance with protons (eq. 5) we have:

$$1 - \mu_{||} \beta_{||} = \frac{\Omega_+}{\omega}$$

Rewriting, we have

$$\beta_{||}^2 = \frac{1}{\mu_{||}^2} \left(1 - \frac{\Omega_+}{\omega}\right)^2$$

or

$$\beta_{||}^2 = \left(\frac{\Omega_+}{\omega}\right)^2 \left(1 - \frac{\omega}{\Omega_+}\right)^2 \frac{1}{\mu_{||}^2}$$

Substituting for $\mu_{||}^2$, we have:

$$\beta_{||}^2 = \frac{v_{||}^2}{c^2} = \frac{\Omega_+^2}{c^2 k_{||}^2} \left(1 - \frac{\omega}{\Omega_+}\right)^2 \quad (16)$$

and

$$\frac{1}{2} m^+ v_i^2 = E_r^+ = \frac{v_{th}^2 m^+ c^2}{2 k_{||}^2 f_+^2 c^2} \left(1 - \frac{\omega}{\Omega_+}\right)^2 \quad (17)$$

thus

$$E_r^+ = E_{th}^+ \left(\frac{1}{k_{||} f_+}\right)^2 \left(1 - \frac{\omega}{\Omega_+}\right)^2 \quad (18)$$

Since $\left(\frac{1}{k_{||} f_+}\right)^2 \left(1 - \frac{\omega}{\Omega_+}\right)^2$ is of the order one (1), $E_r^+ \approx E_{th}^+$, i.e., 2 eV - 1 keV. It must be noted, however, that higher order resonances must be taken into account since these electrostatic ion cyclotron waves have polarizations $\frac{k_{\perp}}{k_{||}} \approx 10$ (Kintner et al., 1978). It will be shown in the next section that, unlike electromagnetic waves, resonant electrostatic wave amplitudes required for strong diffusion scattering decrease with increase in energy of the resonant population ($\gamma \rightarrow \infty$). Thus resonant energies for ions can easily reach a few 100 keV (Ashour-Abdalla and Thorne, 1978). So in summary, electrons with $E_r^- \approx 50$ keV-2MeV and ions with $E_r^+ \approx 1 \approx 100$'s keV can resonate with electrostatic ion cyclotron waves.

2.3 Wave Amplitudes

Now that the resonant electron and proton population is known for each wave mechanism, it is useful to determine the resonant wave amplitudes required for strong diffusion. We will then relate the ability of observed wave amplitudes to

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populations on strong diffusion but the equations will help later when examining real cases.

The condition for strong diffusion, i.e., very rapid scattering, is

$$\tau_D \ll \tau_{1/4B} \quad (19)$$

where τ_D = diffusion loss time

$\tau_{1/4B}$ = 1/4 bounce time

This permits the particle distribution to become isotropic even over the loss cone ($\Rightarrow J_p \approx J_t$). The loss time approaches a minimum value under strong diffusion and the diffusion coefficient, D , is larger than $1/\tau_{\min}$, where τ_{\min} is the minimum rate of diffusion loss. $\tau_{\min} \approx \frac{2 \tau_{1/4B}}{\alpha_L^2}$ where α_L is the pitch angle of the loss cone.

2.3.1 Electromagnetic Waves

The action of electromagnetic waves on particles is mainly via pitch angle scattering. The pitch angle scattering coefficient, $D_{\alpha\alpha}$, is given by Kennel and Petschek (1966):

$$D_{\alpha\alpha} = \frac{(\Delta\alpha)^2}{2 \Delta t} \approx \pi \Omega - \left(\frac{b}{B_0}\right)^2 \quad (20)$$

where b = wave amplitude

B_0 = ambient magnetic field value

and for strong diffusion scattering one requires

$$D_{\alpha\alpha} \approx D_{SD} \approx \frac{\alpha_L^2}{2 \tau_{1/4B}} \quad (21)$$

where α_L = loss cone pitch angle

D_{SD} = coefficient for strong diffusion.

Equation (20) can be rewritten including the relativistic mass correction and bounce averaging to obtain:

$$D_{aa} \approx \frac{\pi \Omega_-}{\gamma} \left(\frac{b}{B_0} \right)^2 f \quad (22)$$

where f = is the fraction of the electron bounce orbit spent in resonance with waves.

For equatorial loss cones, $\alpha_L^2 \approx \frac{1}{2L^3}$. For electrons, $\tau_{1/4B} \approx \frac{LR_0}{v^-} - \frac{LR_0}{c}$ where $R_0 = 1$ earth radius. Equating (22) and (20) with substitutions yields a fluctuating field wave amplitude for strong electron diffusion:

$$b_{SD}^- = \left(\frac{B_0^2 c \gamma}{4\pi L^4 R_0 \Omega_- f} \right)^{1/2} \quad (23)$$

and

$$b_{SD}^- = \frac{100 \gamma^{1/2}}{L^{7/2}} \text{ (gammas) for } f = \frac{1}{5} \quad (24)$$

For ions, $\tau_{1/4B} = \frac{LR_0}{v^+} = \frac{LR_0}{(2E^+/E_0^+)^{1/2} c}$, $\Omega_- = \frac{m^+}{m^-} \Omega_+$, and $\gamma \sim 1$ where E^+ and E_0^+ are the ion energy and ion rest energy respectively. Likewise, the fluctuating field wave amplitude for strong ion diffusion is

$$b_{SD}^+ = \left[\frac{m^+}{m^-} \left(\frac{2 E^+}{E_0^+} \right)^{1/2} \frac{B_0^2 c}{4 \pi L^4 R_0 \Omega_e} \right]^{1/2}$$

and

(25)

$$b_{SD}^+ = \frac{1.65 E^+{}^{1/4} (\text{KeV})}{L^{7/2}} \text{ (gammas)}$$

Strong electron diffusion ($E_\gamma^- \sim 1/2 - 1$ MeV; $\gamma \sim 2-3$) thus requires amplitudes of 1 gamma at L=4 and 100 m γ at L=8. For ion strong diffusion ($E_\gamma^+ \sim 50$ keV), amplitudes of 3.5 gammas at L=4 and 300 m γ at L=8 are required. These values are within the range of amplitudes for both electrons and ions for EMIC waves (1-7 γ) found by Bossen et al. (1976) but only the most intense chorus events with amplitudes approximately 100 m γ (Tsurutani and Smith, 1977) meet the criteria for electrons at L=8.

2.3.2 Electrostatic Amplitudes

Electrostatic waves scatter electrons primarily by energy diffusion. Scattering by energy in the perpendicular (to the magnetic field line) direction at the edge of the loss cone requires a diffusion coefficient as follows:

$$D_{\perp\perp} = \frac{\langle \Delta v_{\perp} \rangle^2}{\Delta t} \quad (26)$$

For strong diffusion which can maintain isotropy over the loss cone one requires

$$D_{\perp\perp} \geq D_{SD} \approx \frac{V_{res}^2 \alpha_{\perp}^2}{\tau_{1/4B}^2} \approx \frac{V_{res}^2}{\tau_{1/4B}^2} \left(\frac{B_0}{B_A} \right) \quad (27)$$

where $B_A = B(100 \text{ km})$.

An approximate expression for $D_{\perp\perp}$ as a function of electrostatic wave amplitude can be derived (e.g. Ashour-Abdalla and Thorne, 1978):

$$D_{\perp\perp}^{\pm} = \frac{\Omega_{\pm}^2 c^2 \mathcal{E}_w^2}{\gamma^{\pm} B_0^2} \quad (28)$$

where \mathcal{E}_w is the electric field amplitude of the electrostatic wave. Equating eqs. (27) and (28) and using $\tau_{1/4B} = \frac{LR_0}{v} \sim \frac{LR_0}{c}$, we have for the fluctuating wave amplitude for electron strong diffusion

$$\mathcal{E}_w^- = \left[\frac{c}{LR_0} \left(\frac{B_0}{B_A} \right) \frac{B_0^2 \gamma^-}{\Omega_-} \right]^{1/2} \quad (29)$$

At the equator $\frac{B_0}{B_A} \approx \frac{1}{2L^3}$ and $\mathcal{E}_{w_{eq}}^- = [3.7 \times 10^8 \frac{\gamma^-}{L}]^{1/2} (\frac{mV}{m})$. For $\gamma^- = 2-3$, we have $\mathcal{E}_w^- = 20-250 \frac{mV}{m}$ for $L=8-4$. In the topside ionosphere near 1 Re, $\frac{B_0}{B_A} \approx \frac{1}{10}$, $B_0 \approx .05$, and $\mathcal{E}_{w_{1Re}}^- = [1.4 \times 10^7 \frac{\gamma^-}{L}]^{1/2} (\frac{mV}{m})$. For $\gamma^- = 2-3$, we have $\mathcal{E}_w^- \approx 1-3 \frac{V}{m}$ for $L=8-4$.

For the concomitant ion strong diffusion using eq. (27) and

$$D_{SD}^+ = \Omega_+ c^2 \frac{\mathcal{E}_w^+{}^2}{B_0^2} \quad (30)$$

we have

$$\mathcal{E}_w^+ = \left[\frac{c}{LR_0} \left(\frac{B_0}{B_A} \right) \left(\frac{B_0^2 \gamma^+ m^+}{\Omega^- m^-} \right) \left(\frac{2 E^+}{E_0^+} \right) \right]^{1/2} \quad (31)$$

and

$$\mathcal{E}_w^+ = \left[9.4 \times 10^6 \left(\frac{B_0}{B_A} \right) \frac{B_0 E^+ (\text{keV})}{L} \right]^{1/2} \left(\frac{\text{mV}}{\text{m}} \right) \quad (32)$$

Again for the equator $\frac{B_0}{B_A} \approx \frac{1}{2L^3}$ and $\gamma^+ = 1$ we have $\mathcal{E}_w^+ = 6 \frac{\text{mV}}{\text{m}} - 66 \frac{\text{V}}{\text{m}}$ at $L=8-4$ for a 50 keV ring current proton. At the 1 Re, $\frac{B_0}{B_A} \approx \frac{1}{10}$ $B \approx .05$ and $\gamma^+ = 1$, we have $\mathcal{E}_w^+ = 550 \frac{\text{mV}}{\text{m}} - 750 \frac{\text{mV}}{\text{m}}$ for a 50 keV ring current proton.

Electrostatic ion cyclotron waves have electric field amplitudes typically on the order of 10-100's $\frac{\text{mV}}{\text{m}}$ (Gurnett and Frank, 1977 and Kintner et al., 1978). Required wave amplitudes for both ions and electrons are easily met by these waves in the equatorial regime but not in the topside ionosphere.

Table 2.1 lists the range of parameters for strong ion and electron diffusion.

2.4 Case Studies

An example of each of the classes described in Section 1.6 is presented. Each example was selected as being most representative with amount of peripheral data available as an added factor.

2.4.1 Case 1, ESIC (rev 479, day 250, 6 SEP 1976).

The flux versus time plots for electron channels 12-1600 keV

TABLE 2.1 Theoretical REP mechanisms and characteristics.

THEORETICAL REP MECHANISMS

<u>WAVE</u>	<u>FREQ</u>	E^- RES	E^+ RES	<u>MIN AMP S.D.</u> ELECTRONS	<u>MIN AMP S.D.</u> PROTONS
WHISTLER MODE CHORUS	$.1 < \omega < \omega_{ce1}$	10-100s KEV	1-10 MEV	100mV- γ	300mV-3 γ
EM ION CYCL	$\omega < \omega_+$	1-10 MEV ($N=10^3$)	10 KEV-1 MEV	100mV- γ	300mV-3 γ
ES ION CYCL	$\omega_+ < \omega < \omega_{ce}$ LHR	10 KEV-2 MEV	100 EV-100s KEV	10s-1000mV/M	5-750mV/M

and protons 80-770 keV for three outer zone passes are shown in Fig 2.5 a,b,c, respectively. The low altitude outer zone pass at local evening (Fig 2.5a) shows an extended region of 12 keV precipitation (isotropic) at $L \approx 5-10$. This is actually an inverted V event which is bounded by 235 keV precipitation. At $L=5$ the precipitation extends to nearly 1 MeV electrons and 80 and 150 keV protons. This agrees well with the criteria of ESIC (Table 2.1). Fig 2.5b shows the outer zone pass at high altitude (~ 8000 km) near local dusk. Twelve (12) keV electrons are precipitating from $L \approx 7-9$ with two large spikes near $L \approx 8.4$. The two spikes are associated with two electrostatic shocks (Kintner, private communication) with perpendicular electric field $\mathcal{E}_\perp \approx 500$ mV/m. In addition > 80 and > 150 keV protons are precipitating from $L \approx 6$. Although there is significant trapped 235 keV electron flux from $L \approx 6.8-7.8$, there is no precipitation. The next pass (Fig 2.5c) is a medium altitude (~ 5000 km) pass near local morning. Twelve (12) keV electron precipitation is evident from $L \approx 6-13$. Extensive 33 keV is coincident although not on strong diffusion from $L \approx 9-12$. The 235 keV channel shows "moderate" precipitation at $L \approx 6-8$. Significant fluxes of > 80 keV protons are present to roughly $L \approx 12$ without any precipitation. This fits the criteria of whistler mode chorus in Table 2.1.

Pitch angle plots of the low altitude local evening outer zone pass of rev 479 are shown in Fig 2.6. Panel (A) at 5962 UT shows a normal pitch angle distribution both down (0°) and up (180°) the local magnetic field line. Panel (B) shows the next distribution towards

FIGURE 2.5a Flux (counts/sec, every fourth data point) versus L value of low altitude (700 km) outer zone pass near local evening. Vertical arrows (\uparrow) indicate singular precipitation and horizontal arrows ($\rightarrow \leftarrow$) indicate limits of extended precipitation.

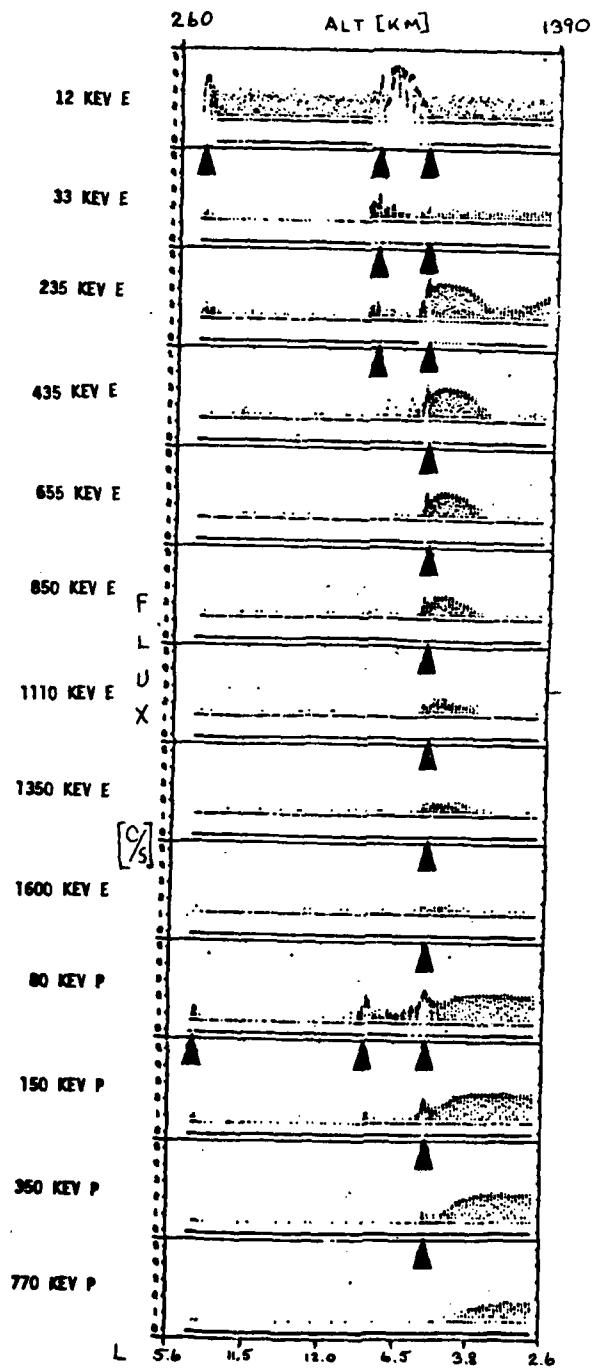


FIGURE 2.5b Same plots for high altitude (8000 km) passes at local dusk.

ALT [KM] 7600

8000

7900

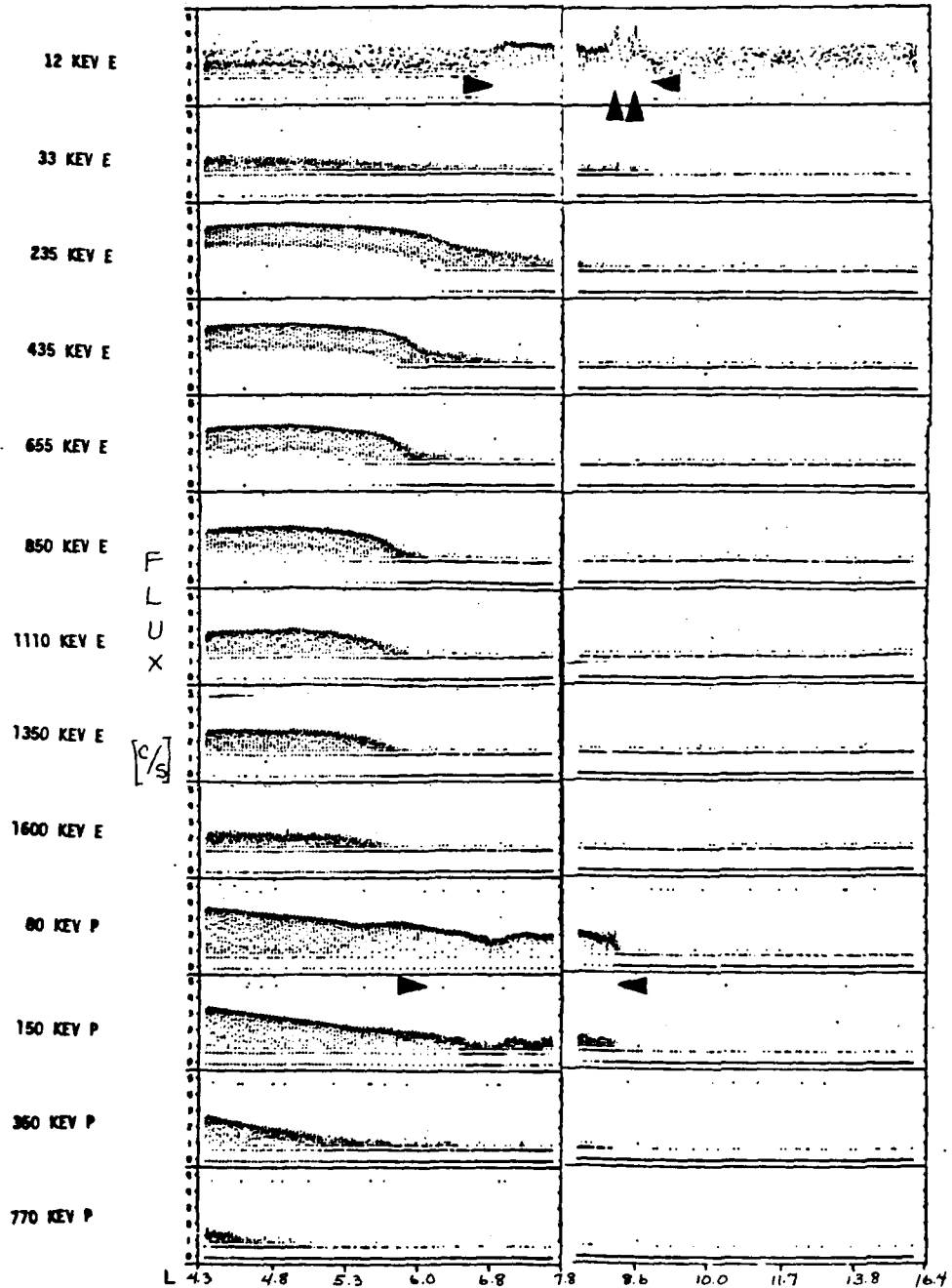


FIGURE 2.5c Same for medium altitude (5000 km) passes at
local morning.

ALT [KM] 6250

4550

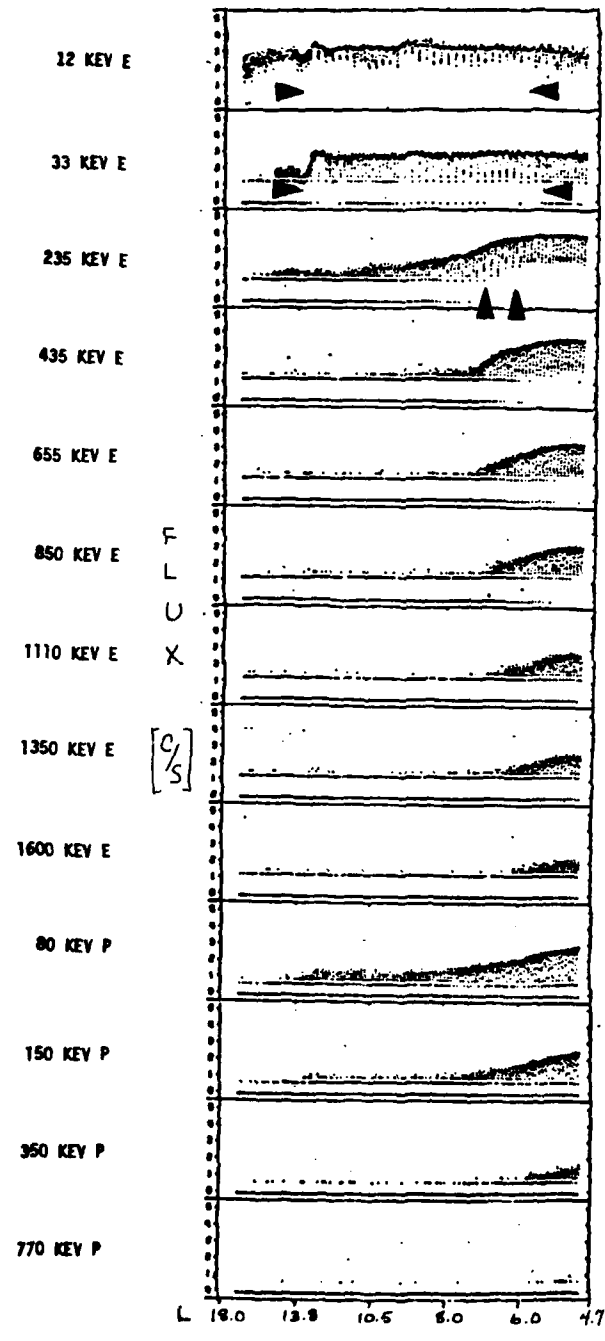
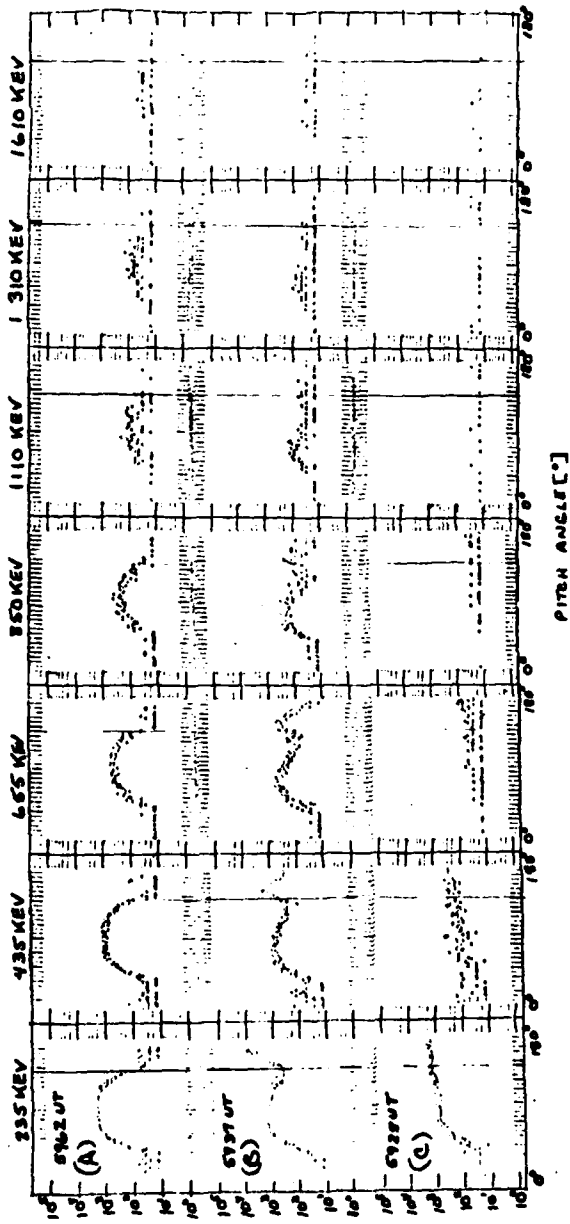


FIGURE 2.6 Pitch angle versus flux (counts/sec) for 235-1610 keV electron channels for 3 times in the local evening outer zone of rev 479. The upward loss cone (L.C. = 50°) is indicated by the vertical line.



J
[neutrons/cm²-sec]

the outer edge of the outer zone with a large spike in the upward viewing loss cone. The very next distribution Panel (C) shows isotropic flux in the upward loss cone. Isotropic fluxes are evident to ~ 1 MeV.

Since the precipitating flux is higher than the locally trapped flux, one can estimate the altitude at which pitch angle scattering occurred (Vampola, 1977). This is accomplished by tracing the precipitation flux to the shape of the outer zone as illustrated in Fig 2.7. At L=4 the trapped flux would be roughly 2×10^4 counts/sec up the field line where the precipitation occurred. Fig 2.5b shows fluxes near L=4 of approximately 2×10^4 counts/sec. The altitude of this pass is ~ 7600 km. It must be noted that this "spikey" event is very rare with the overwhelming majority of precipitation cases showing flat isotropic fluxes.

Resonance with 1 MeV electrons at ~1 Re would require $\frac{1}{k_{\parallel} \beta_+}$ \approx .008, slightly smaller than values for first three harmonics calculated by Ashour-Abdalla and Thorne (1978). In addition, the wave amplitude is $\mathcal{E}_w \approx 350 \frac{mV}{m}$. Both these calculations indicate that higher harmonics may play a very important role in the mechanism.

Fig 2.8a shows time spectrograms for the filtered plasma wave data (frequency .03-100 KHz) and low energy electron (.2-33 keV) and protons (.1-3.9 keV), 235 keV electrons and > 80 keV protons of the low altitude local evening pass. The 80 keV proton and 235 keV electrons are seen to precipitate (upward loss cone filled) at 5815 UT and 5950 UT (the spike occurred here). Between these times is the inverted V electron event. Plasma turbulence is also evident but not well

FIGURE 2.7 Enlargement of portions of Figure 2.5a.

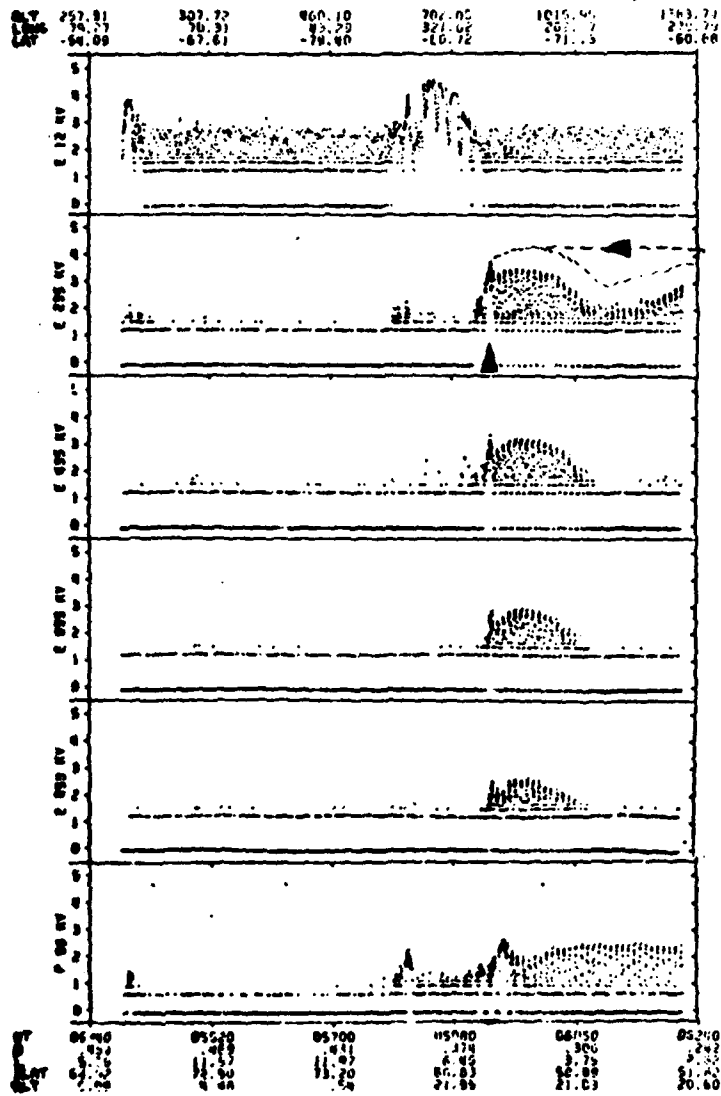
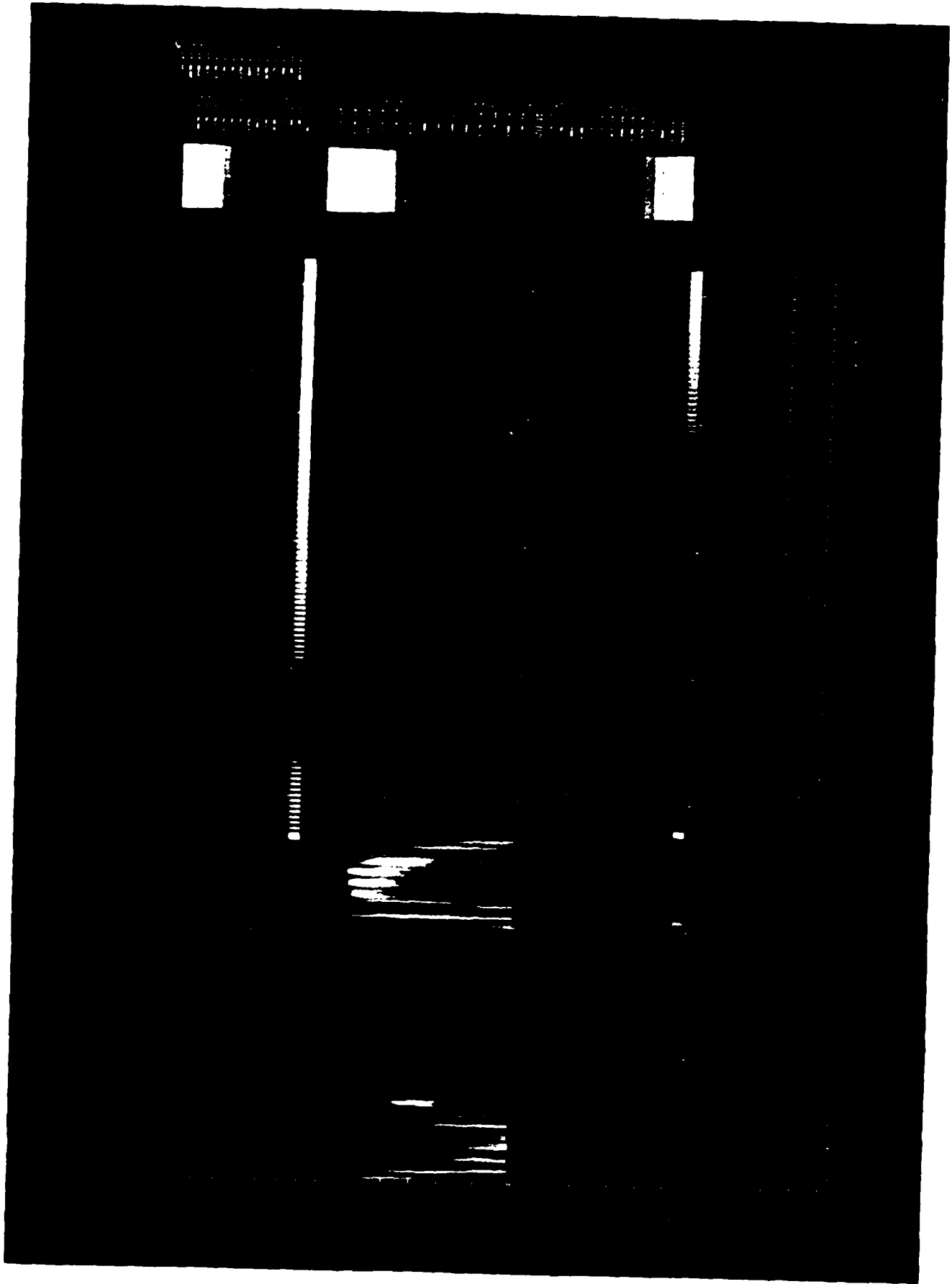


FIGURE 2.8a Spectrogram of plasma waves, 235 keV electrons, low energy electrons, low energy protons, 80 keV protons for the low altitude local evening pass.



defined. Fig 2.8b shows the high altitude local dusk pass with the inverted V electrons evident from 10500 UT to 10700 UT with ion beams also evident between these times. Plasma turbulence is also evident between these times but the 235 keV electrons have been depleted.

Fig 2.9 shows the raw magnetogram for Kiruna ($67^{\circ} 50'N$, $20^{\circ} 25'E$) for 4-6 September 1976. The period is relatively quiet from 0200 UT to 2200 UT 5 September 1976. A small substorm occurs shortly before midnight on 6 September and the estimate (to be shown later) of the extent of the precipitation in time is indicated by the arrows. In addition, the geosynchronous satellite, ATS-6, showed two small excursions in the 140-800 keV electron channel (see Fig 2.10).

The S3-2 satellite has a magnetic spectrometer similar to S3-3 (but only to electron energies of 400 keV). These satellites were coplanar at this time of year. The satellite tracks are shown in Fig 2.11. The X's indicate regions of precipitation for either satellite. The precipitation is first seen on S3-3 at 1.6 UT at 21.5 MLT (spike event). A similar spike event is then seen at 2.1 UT (both 200 and 400 keV channels) by S3-2 in the same location and altitude. Neither satellite observed precipitation electrons (protons were precipitating, Fig 2.5b) at local dusk. Both satellites observed precipitation near 200 keV at local morning at 3.05 UT (S3-2) and 3.8 UT (S3-3). The S3-2 satellite continues in orbit till precipitation at 200 keV is observed again at 3.65 UT at 3.0 MLT, 3.7 UT at 2.3 MLT, and 3.8 UT at 22.5 MLT. The extent of precipitation is thus ≈ 2.2 hrs. The actual S3-2 flux versus times are shown in Fig 2.12a,b,c.

FIGURE 2.8b Same for the high altitude local dusk pass.

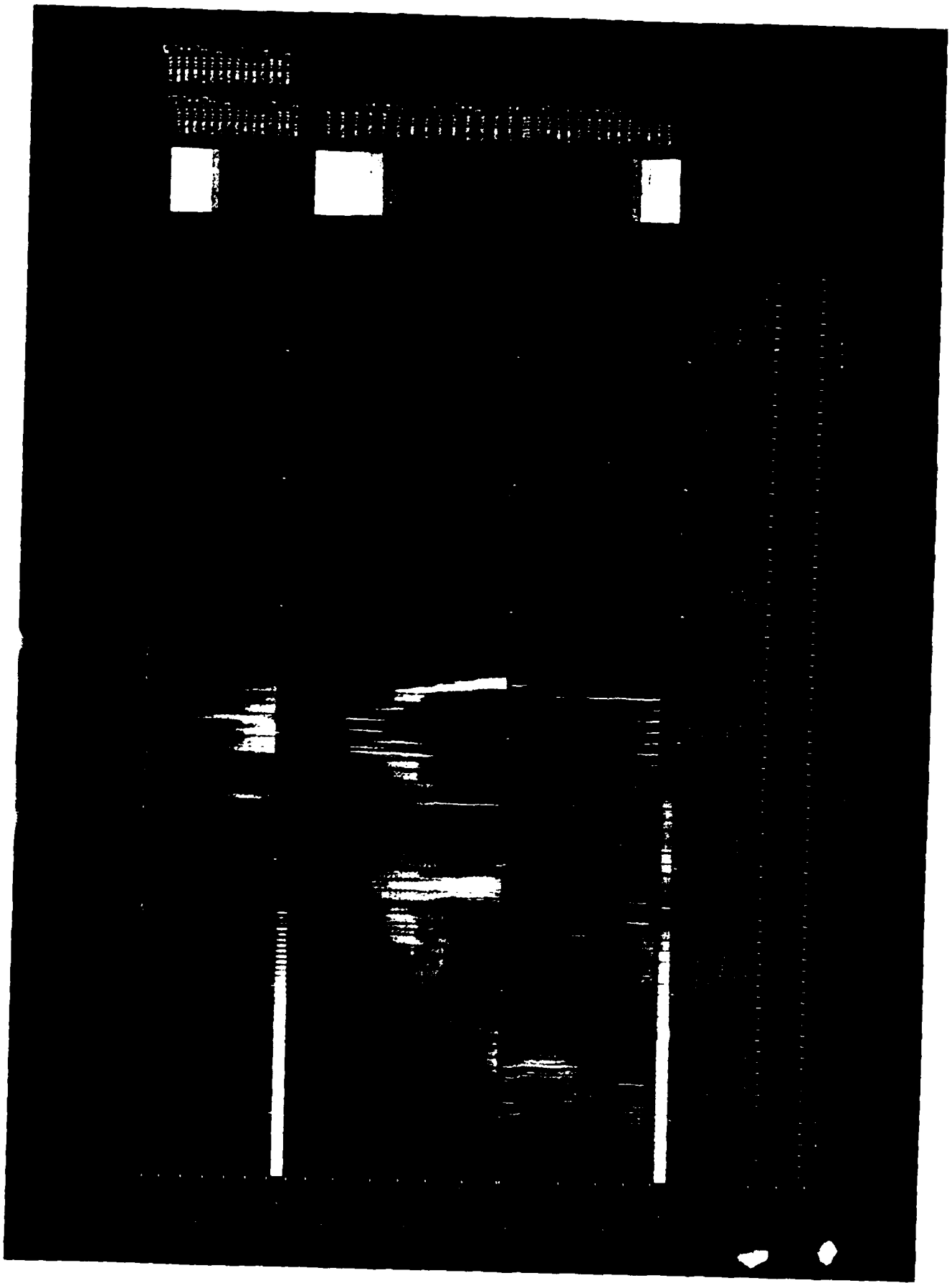


FIGURE 2.9 Kiruna raw magnetogram for 4-6 September 1976.

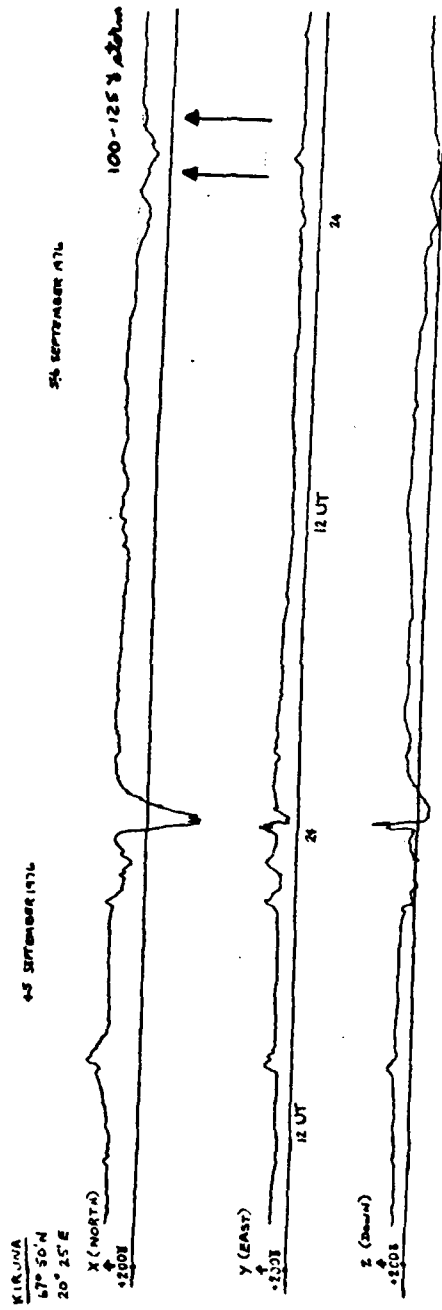


FIGURE 2.10 ATS-6 for 6 September 1976. Arrows indicate beginning and ending of event.

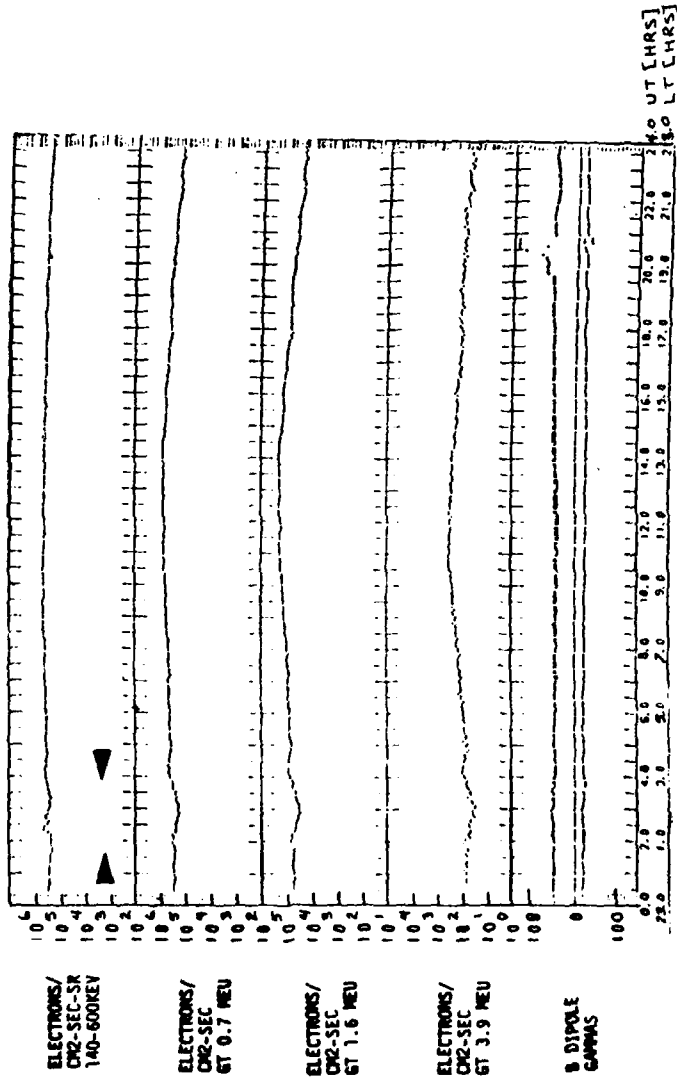


FIGURE 2.11 S3-2 (1, 2, 3) and S3-3 satellite (a, b, c) tracks for
6 September 1976 in invariant latitude versus MLT.

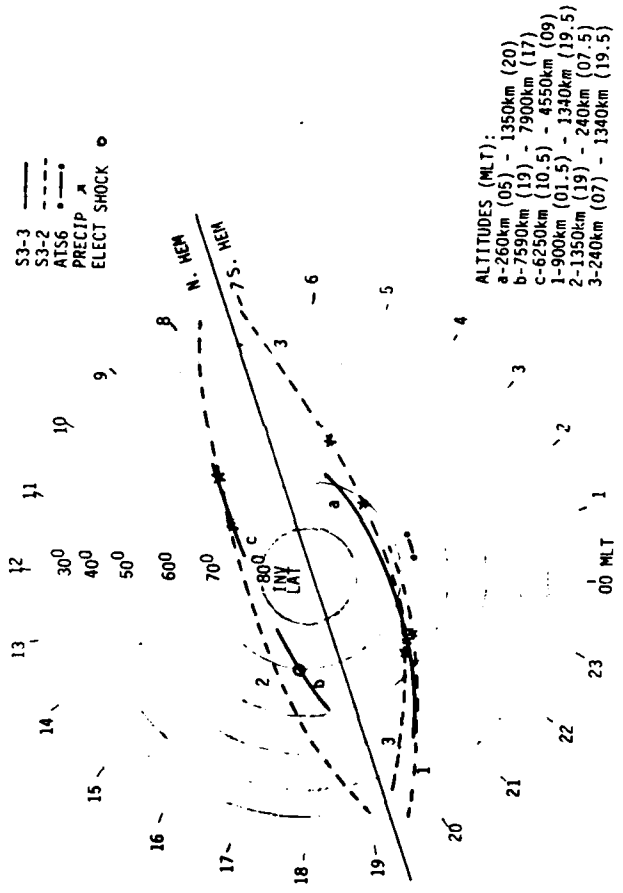


FIGURE 2.12a S3-2 flux versus time plots of pass 1, Figure 2.11.
Arrows indicate precipitation.

ALT 451.08
 LONG 626.20
 LAT 55.52
 59.49

626.20
 46.94
 -78.08

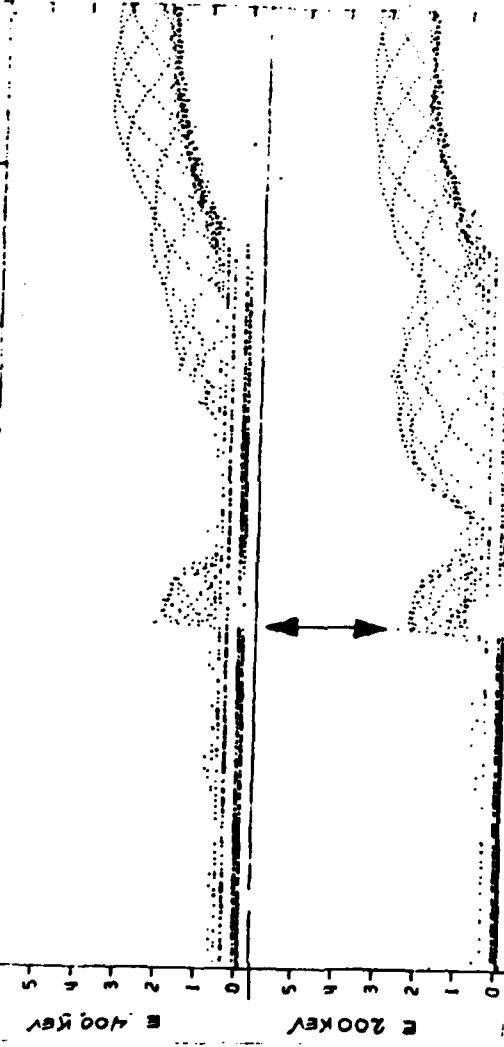
504.02
 294.23
 -79.74

976.59
 268.52
 -63.72

1129.55
 261.28
 -47.40

1252.24
 257.06
 -31.53

1337.12
 253.78
 -16.07



UT 06360
 G 419
 L 6.98
 ILAT 67.76
 MLT 5.36

07260
 401
 12.67
 73.69
 1.45

07560
 365
 5.83
 65.88
 21.54

07860
 304
 2.84
 53.61
 20.51

08160
 243
 1.82
 46.11
 20.03

08460
 182
 1.41
 32.57
 15.72

08760
 124
 1.00
 20.24
 13.44

FIGURE 2.12b Same for pass 2, Figure 2.11.

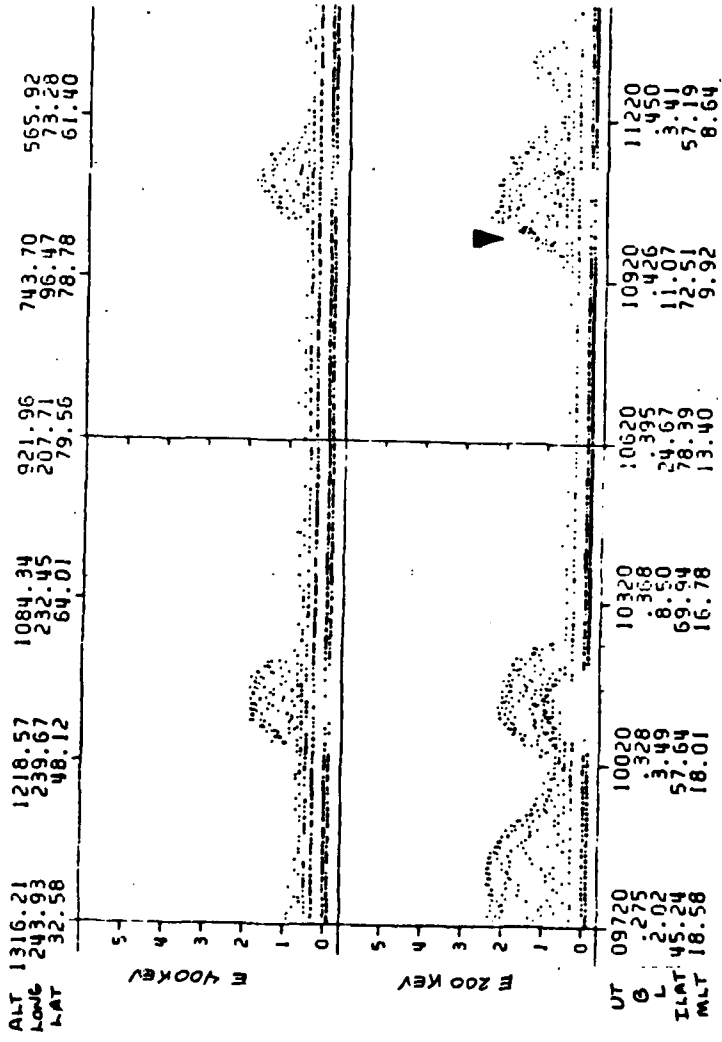
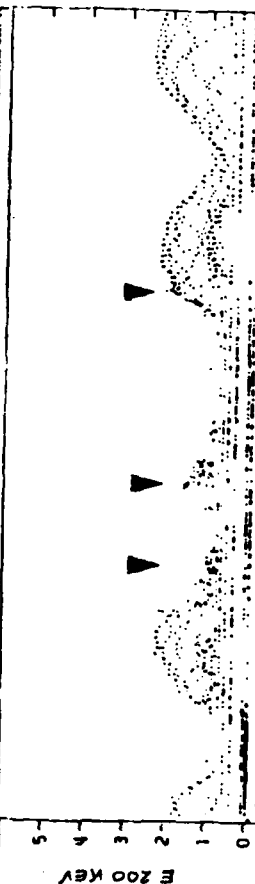


FIGURE 2.12c Same for pass 3, Figure 2.11.

ALT 328.96 461.00 626.19 804.08 976.65 1129.62
 LONG 50.16 43.34 22.21 269.71 243.38 236.06
 LAT -40.01 -59.78 -77.89 -79.91 -63.93 -47.61



UT 12720 13020 13320 13620 13920 14220
 G 352 348 379 381 345 290
 L 2.38 4.76 9.50 6.52 3.38 2.01
 ILAT 49.56 62.71 69.94 66.94 57.06 45.70
 MLT 6.28 5.19 2.28 22.48 20.95 20.20

In summary this event fits the criteria for ESIC mechanism at local evening with whistler mode chorus at local morning. This particular event seems to be a low altitude ($\sim 1.2 R_0$) phenomenon correlated with a 120γ substorm and inverted V event.

2.4.2 Case 2, EMIC (rev 2484, day 132, 12 May 1977)

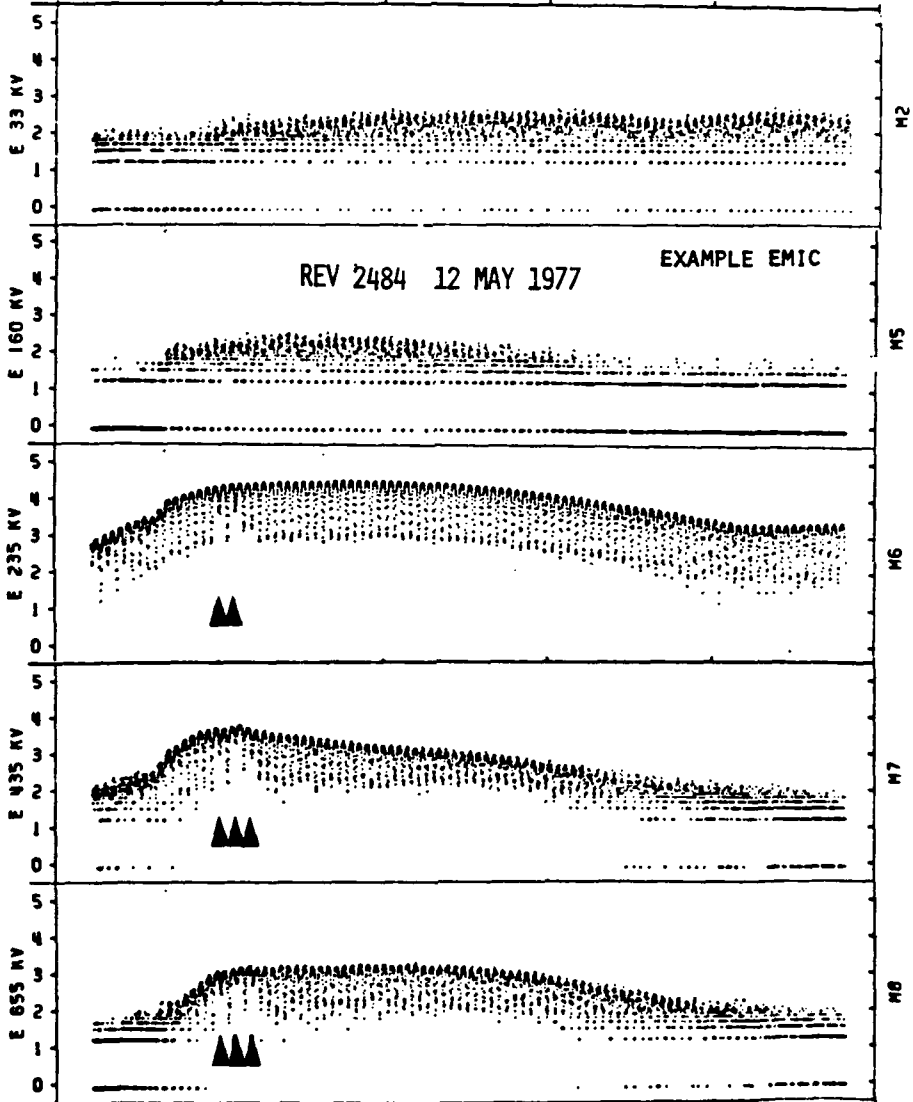
The flux versus time plot for the EMIC example is shown in Fig 2.13a,b. No precipitation is seen in the 12 or 33 keV electron channels for this local evening, high altitude (~ 7800 km) pass. In the 235 keV channel, however, the flux at $L \approx 4.3$ falls an order of magnitude into the loss cone where normal loss cones are nearly 2 orders of magnitude deep. Proceeding up in energy, the loss cone at this same L value begins to fill until at 850 keV, the loss cones are nearly full (Fig 2.13b). Isotropic distributions are evident in the 1.1, 1.3, and 1.6 MeV channels. In addition, the > 80 keV protons are on strong diffusion from just equatorward of the electron precipitation and continuing poleward. The loss cone filling with increasing energy is evident in Fig 2.14 where pitch angle plots of flux for 160 keV to 1600 keV electron energies are shown. The upward observing loss cone is at 0° in these plots. This spectrum and concurrent proton precipitation is indicative of the EMIC mechanism.

The spectrogram for this pass is shown in Fig 2.15. Low energy electron and proton activity is evident poleward of the REP and only the > 80 keV is coincident. No plasma data was available for this pass.

No direct means of observing the plasmopause was available for

FIGURE 2.13a Flux versus time plots of electrons for local evening
pass of 12 May 1977.

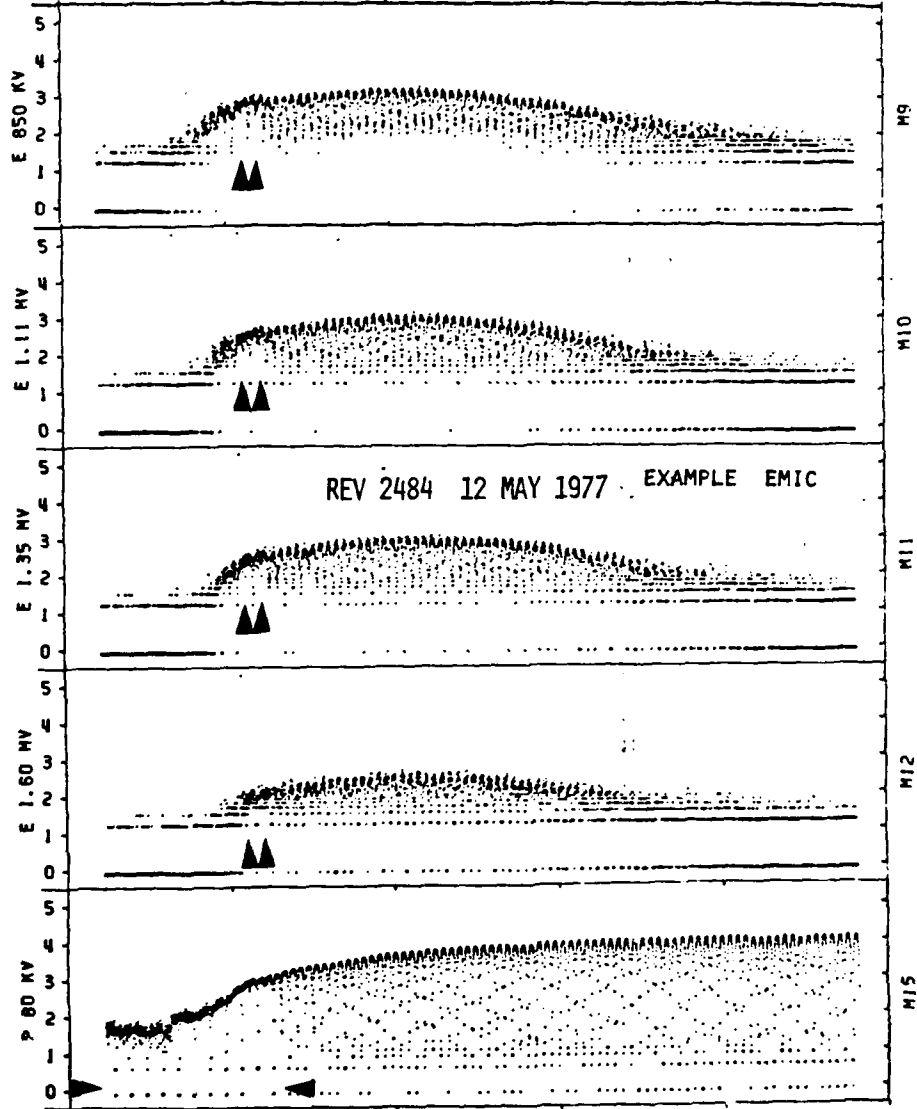
ALT	7735.28	7805.36	7852.12	7875.63	7675.93	7853.01
LONG	241.89	240.47	239.10	237.77	236.49	235.23
LAT	40.78	37.73	34.71	31.69	28.68	25.66



UT	16440	15620	16800	16900	17160	17340
B	.047	.045	.043	.041	.039	.038
L	4.87	4.35	3.93	3.58	3.30	3.06
ILAT	63.05	61.34	59.69	58.10	56.58	55.12
MLT	20.17	20.18	20.18	20.19	20.20	20.20

FIGURE 2.13b Same for higher energy electrons and protons.

ALT	7735.28	7805.36	7852.12	7875.63	7875.93	7853.01
LONG	241.89	240.47	239.10	237.77	236.49	235.23
LAT	40.78	37.73	34.71	31.69	28.68	25.66



UT	16440	16620	16800	16980	17160	17340
B	.047	.045	.043	.041	.039	.038
L	4.87	4.35	3.93	3.58	3.30	3.06
ILAT	63.05	61.34	59.69	58.10	56.58	55.12
MLT	20.17	20.18	20.18	20.19	20.20	20.20

FIGURE 2.14 Pitch angle versus flux plots for electrons with energies 160-1600 keV for 12 May 1977.

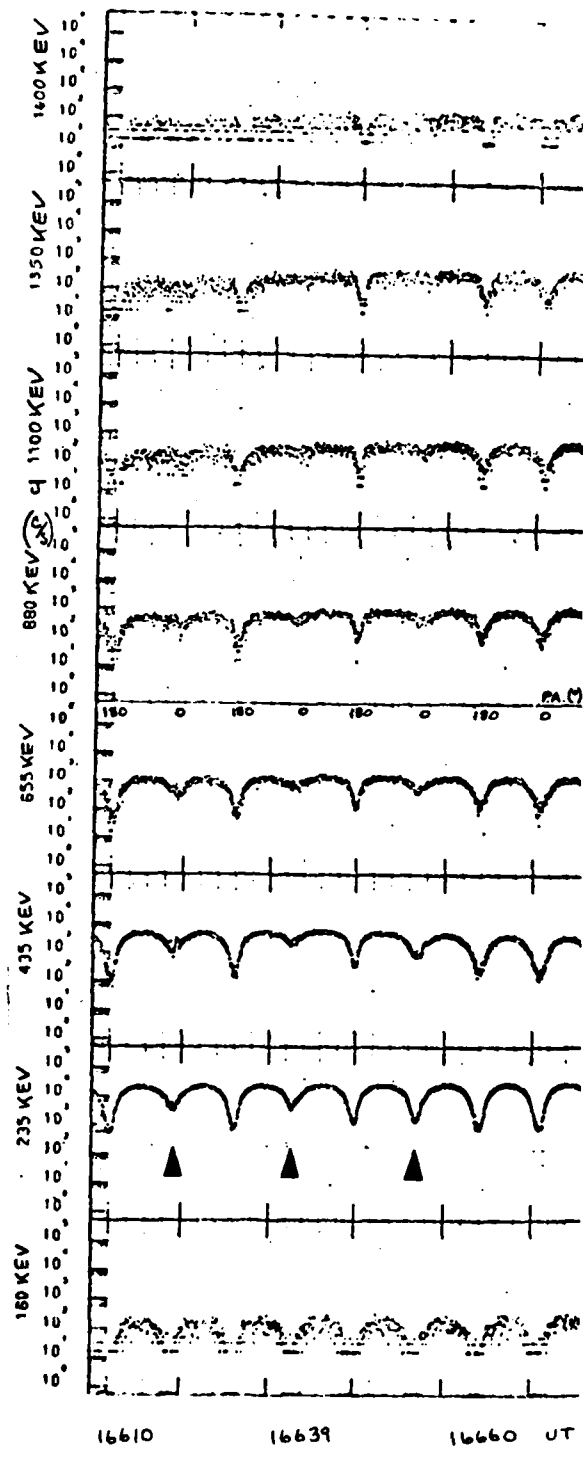
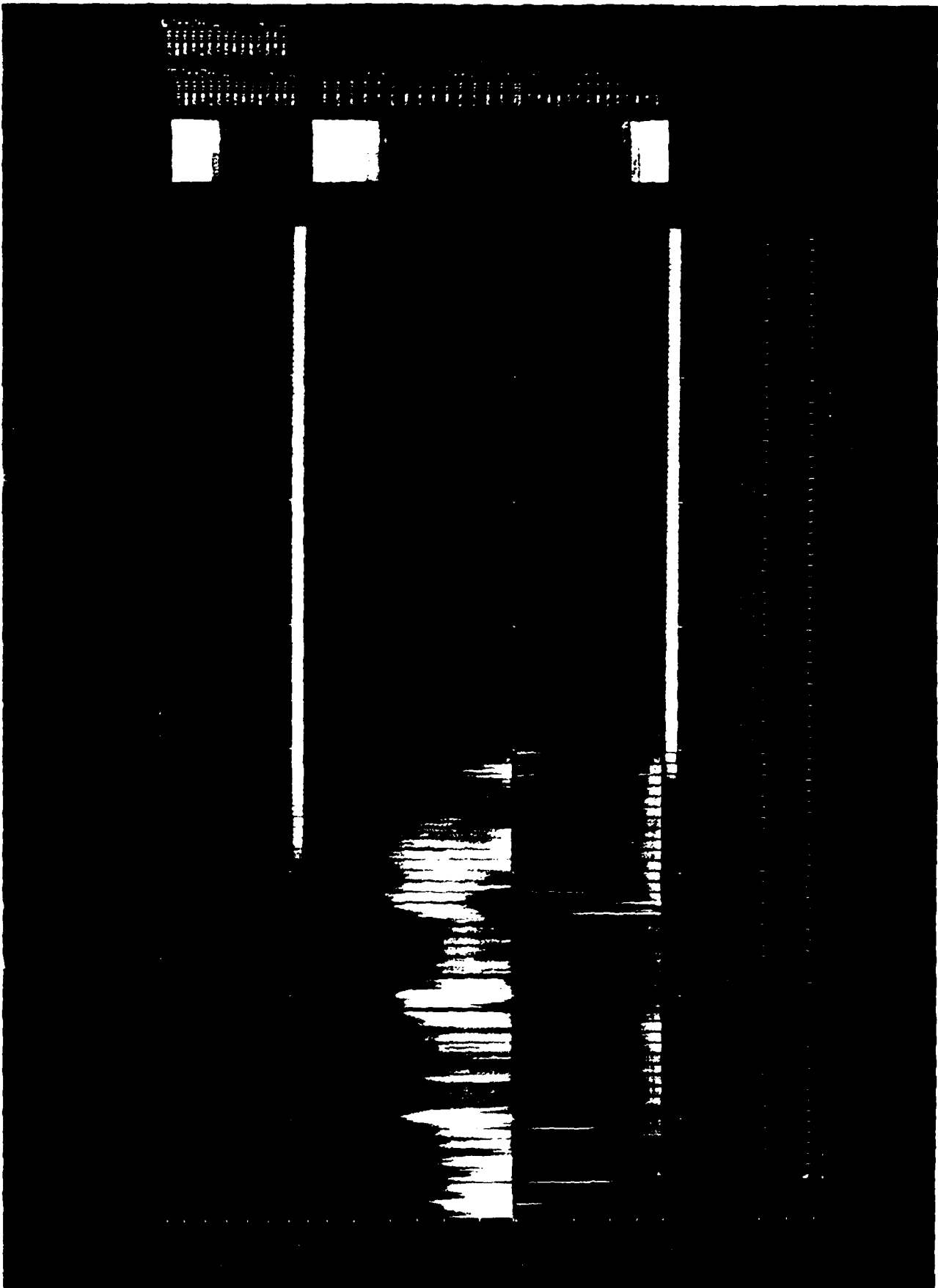


FIGURE 2.15 Spectrogram for 12 May 1977.



this pass. However Mizera (1974) has shown that the midnight plasma-pause is ~ 8 L equatorward of the energetic proton precipitation boundary. The electron and proton precipitation are nearly coincident at local evening. Since the plasmasphere bulges poleward from midnight to dusk, the REP may indeed be just within the plasmopause. Another rough approximation to the L value of the midnight-dawn plasmopause location is given empirically by Carpenter and Park (1973) from ATS-5 data:

$$L_{pp} = 5.7 - 0.47 K_p \quad (29)$$

where K_p is the maximum geomagnetic planetary three-hour-range indices 12 hours previous to the event. The same formula may be used with K_p as the actual value at the time of the event. For UT of 4.6 hrs on 12 May 1977 the K_p at the time of event was ~ 4.5 and the max 12 hrs previous was 5.0. These two values yield $L_{pp} = 3.3-3.9$ as the midnight-dawn-noon plasmopause and approximately twice that value for the dusk plasmopause. These are sketched in Fig 2.16. For these rough approximations, the location of the precipitation of rev 2484 is clearly within the plasmasphere.

Finally, Fig 2.17 shows the ATS-6 electron data on 12 May 1977. The REP occurred during a highly disturbed period as indicated in the 140-600 keV electron channel. In addition, S3-3 showed precipitation (ESIC) at energies up to 435 keV on the preceding and following passes at the same local time. Only four of these events ($E_p^- \geq 1$ MeV) were observed for the entire data set.

FIGURE 2.16 Plasmopause relation to rev 2484, 12 May 1977 based on rough estimates.

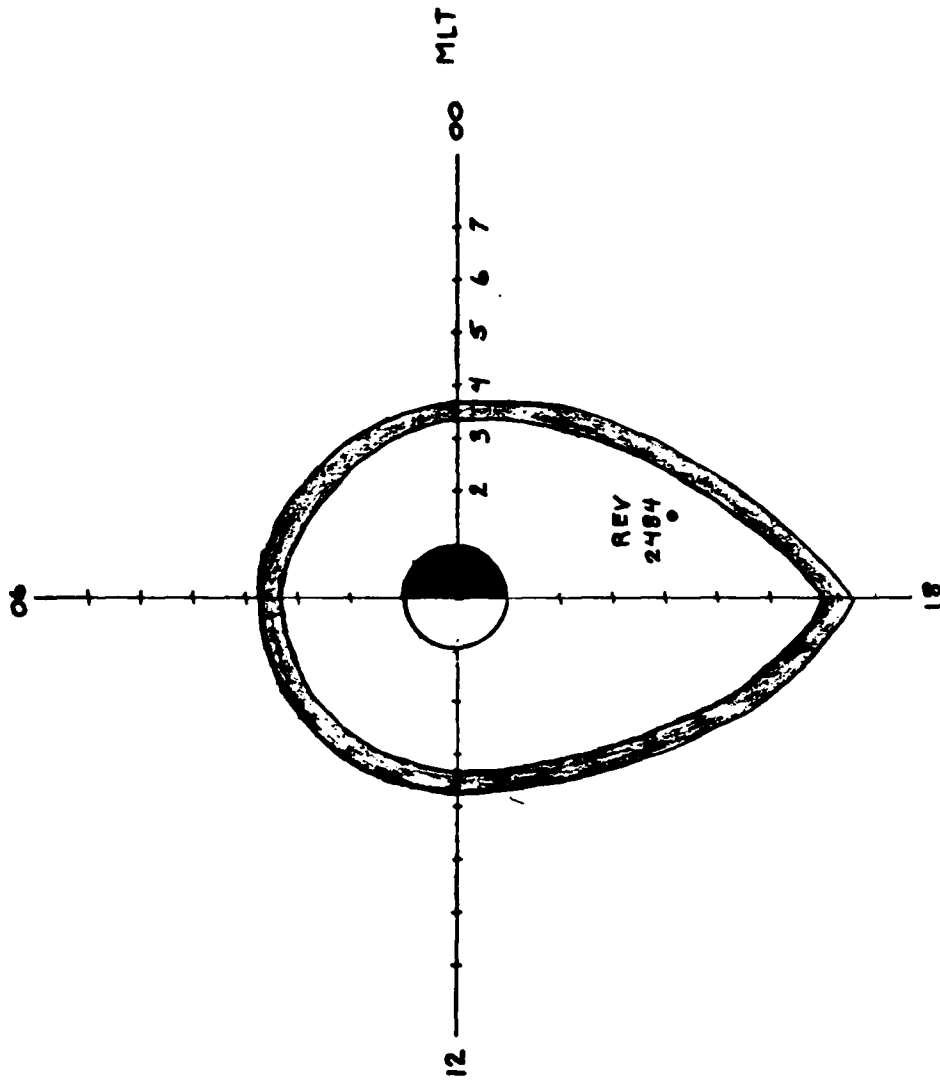
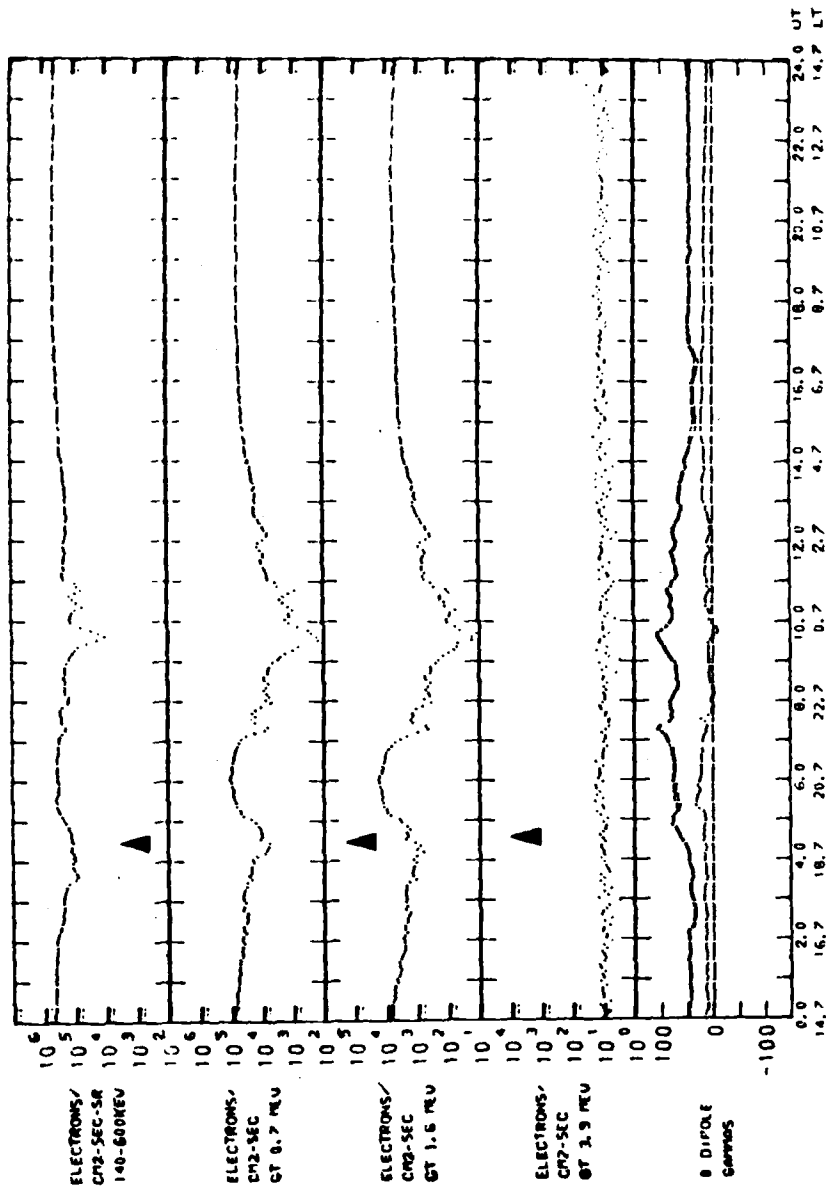


FIGURE 2.17 ATS-6 for 12 May 1977.



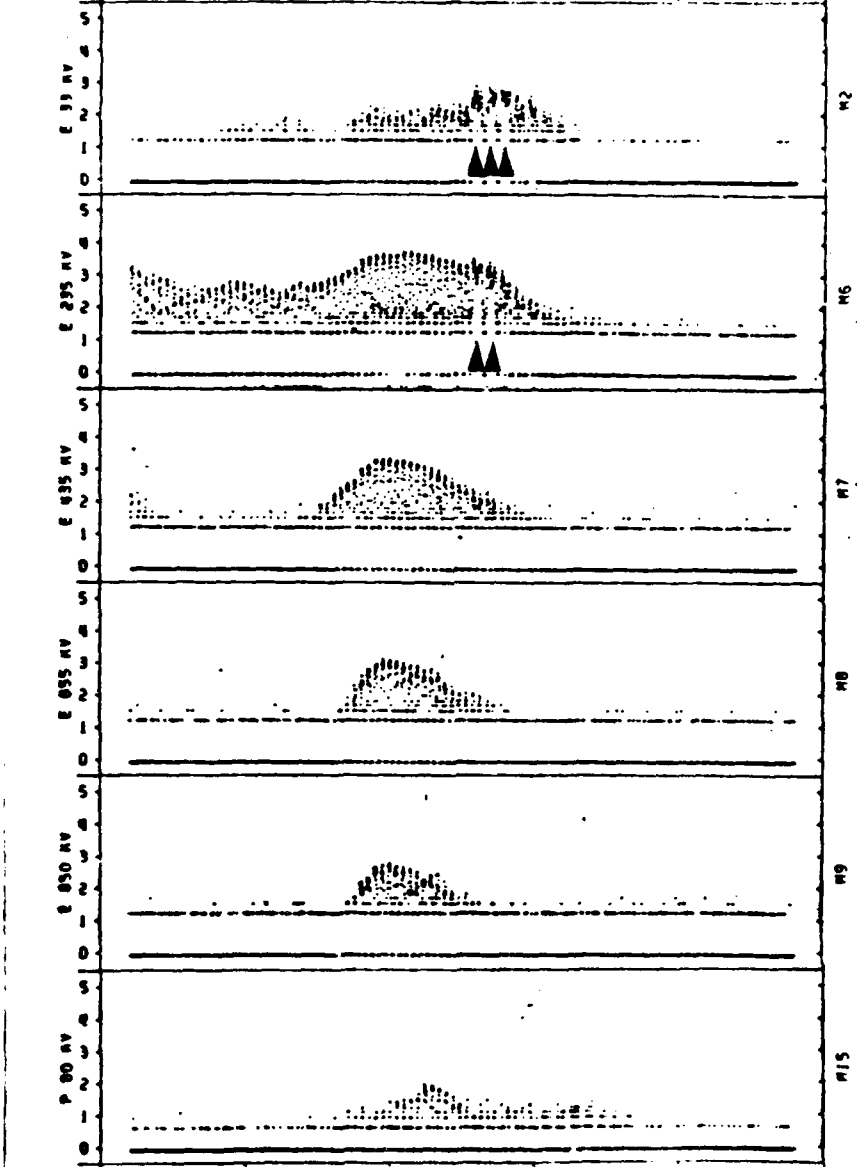
2.4.3 Case 3, Chorus (rev 1413, day 366, 31 December 1976)

The flux versus time plot for 31 December 1976 is shown in Fig 2.18. Only 33 and 235 keV electron precipitation is evident just prior to local noon at low altitude (550 km). No protons were precipitating on strong diffusion though they were present. The pitch angle plot for this event is shown in Fig 2.19 where the loss cone $\approx 55^\circ$. Normal distributions are evident at 12812 UT but at the later times, 33 and 235 keV channels are isotropic in the upward viewing loss cone.

The spectrogram (Fig 2.20) shows evidence of plasma noise just above 30 Hz at and poleward of the REP (12820 UT). A faint band near 10 kHz is also evident.

FIGURE 2.18 Flux versus time plots for 31 December 1977. Arrows indicate strong diffusion precipitation

ALT	1567.20	1179.00	836.03	556.72	356.78	257.58
LONG	154.81	150.49	143.79	126.84	55.69	2.99
LAT	49.67	59.61	64.83	70.17	82.15	72.62



HT	12360	12540	12720	12900	13080	13260
B	1.244	3.19	4.00	4.64	4.87	4.77
L	1.86	3.44	3.92	4.50	17.73	8.64
ILAT	42.80	50.24	59.60	69.94	76.26	71.23
MLT	13.11	12.79	12.29	11.95	9.31	6.16

FIGURE 2.19 Pitch angle distributions for 33, 235, 435 keV electron channels for 31 December 1976.

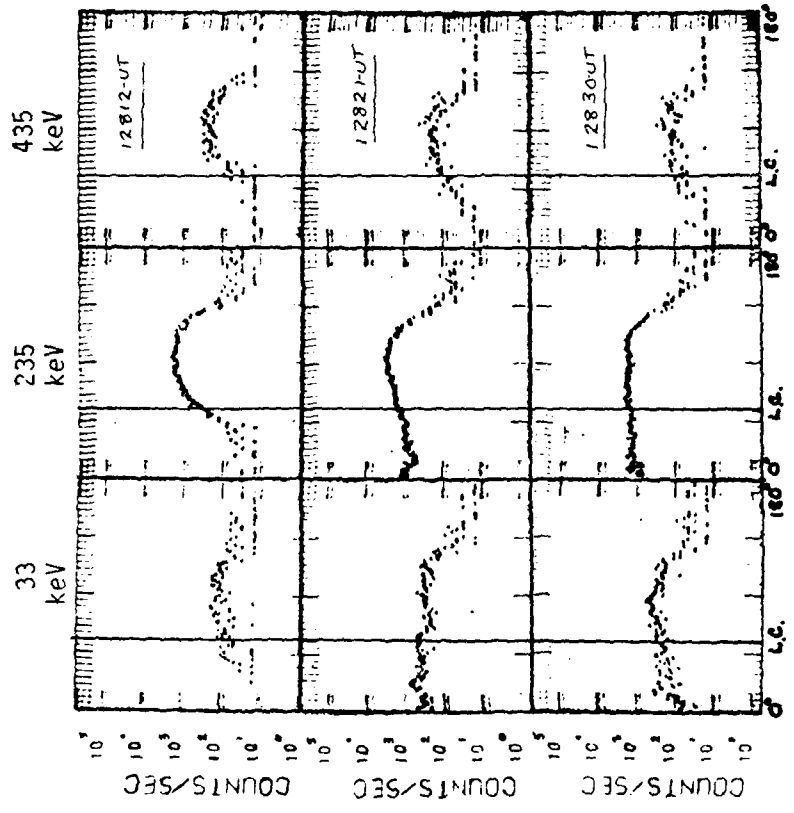
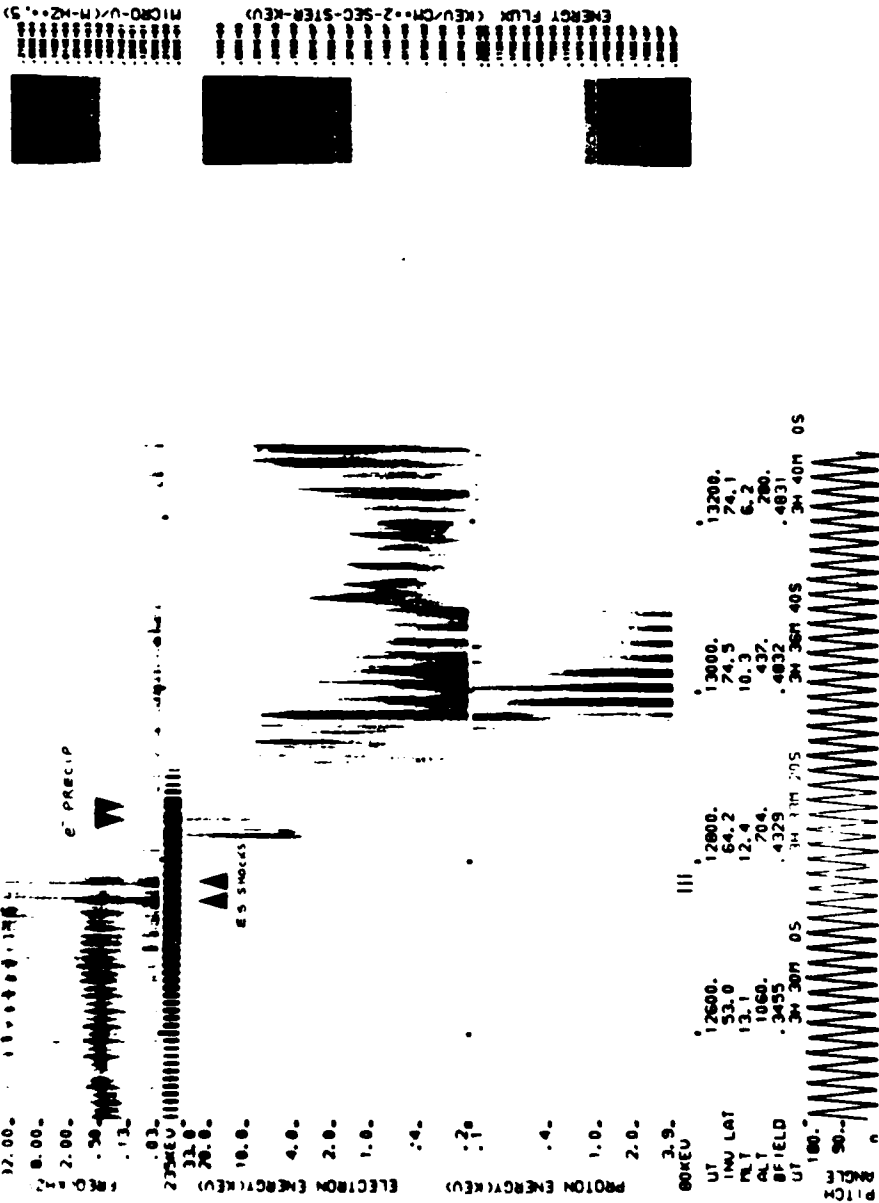


FIGURE 2.20 Spectrogram for low altitude early morning pass
on 31 December 1976.

100.00. 53-3 DAY 366 YEAR 1976 PARTICLE-WAVE SPECTROGRAM



CHAPTER 3. ATMOSPHERIC CONSEQUENCES

3.1 Introduction and Background

The earth's atmosphere below 100 km has been the object of intense investigation over the past 10-15 years. It is well known that a variety of electromagnetic and particle radiation can deposit energy in this region and thus modify the local chemistry. The remainder of this section describes the ionization and neutral chemistry with special emphasis on ozone removal. In subsequent sections we examine the three examples in Chapter 2 and their effect on ozone.

The ionosphere below 100 km, known as the "D region," has been studied by several authors with respect to ionization by solar and galactic cosmic rays (Webber, 1962), primarily in the polar cap region ($\Lambda \geq 75^\circ$): energetic electron precipitation in the auroral region, $\Lambda \simeq 70-75^\circ$ (Rees, 1963; Lerfold and Little, 1964; Bailey et al., 1970) and subauroral regions, $\Lambda \simeq 60-70^\circ$ (Bailey, 1968; Larsen, 1973; Reagan, 1977; Thorne, 1977b) and at mid latitudes, $\Lambda \leq 60^\circ$ (Potemra and Zmuda, 1970; Spjeldvik and Thorne, 1975a,b; Larsen et al., 1976). Solar UV, especially HLy α at 1216 \AA and HLy β at 1025 \AA , produces the dominant primary ion in the quiet mesosphere through ionization of NO. Also, solar UV scattered from the sunlit hemisphere has long been thought to be an important ionizing radiation source at night. Solar UV (and X-ray) radiation is strongly absorbed below about 85 km where positive ion composition is dominated by multiply hydrated ion clusters of the type $\text{H}^+(\text{H}_2\text{O})_n$, $\text{NO}^+(\text{H}_2\text{O})_n$, and $\text{H}_3\text{O}^+(\text{H}_2\text{O})_n$ (Johannesen and

Krankowski, 1972). O_2^- , the dominant negative ion below 80 km, is produced by electron attachment and interacts with CO, NO, HCO and $(H_2O)_n$ to form heavy negative ion and large water clusters.

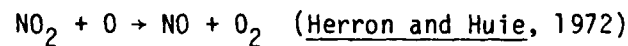
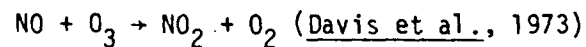
Although the precipitating particle energy is usually negligible compared to other ionizing sources below 100 km, it can be the dominant ionization source during disturbed conditions below 80 km (Thorne, 1977b). Particulate energy produces mostly primary ions O_2^+ and N_2^+ , N^+ , and O^+ of which ~94% are rapidly converted to O_2^+ through charge exchange with O_2 (Reagan et al., 1978) and ~6% to NO^+ by ion-atom interchange with N_2 and O_2 (Gunton et al., 1977). This disturbed increase in O_2^+ can thus modify the complex clustering reactions in the D-region.

A comparison of several ionizing sources is presented in Fig A.15 (Appendix A) where the overwhelming dominance of the largest solar proton event (SPE) of August 1972 is evident. Immense SPEs such as August 1972 are relatively rare and indeed REPs can frequently be the dominant ionization source below 100 km (Spjeldvik and Thorne, 1975a; Thorne, 1977b).

The production and removal of odd oxygen (O , O_3) is of paramount importance in neutral atmospheric chemistry below 100 km. Odd oxygen is produced primarily by photodissociation of molecular oxygen. The Chapman (1930) process, $O + O_3 \rightarrow 2O_2$, is now well known to account for only ~20% removal of odd oxygen. The major removal process includes catalytic loss cycles with chlorine compounds (e.g. Crutzen and Howard, 1978), and odd nitrogen and odd hydrogen species (Blake

and Lindzen, 1973). The latter two are of particular interest here since they can be significantly modified by intense particle precipitation (Thorne, 1977a,b; 1979).

The bulk of odd oxygen depletion is through the following nitrogen chemistry important particularly below 50 km (Blake and Lindzen, 1973):



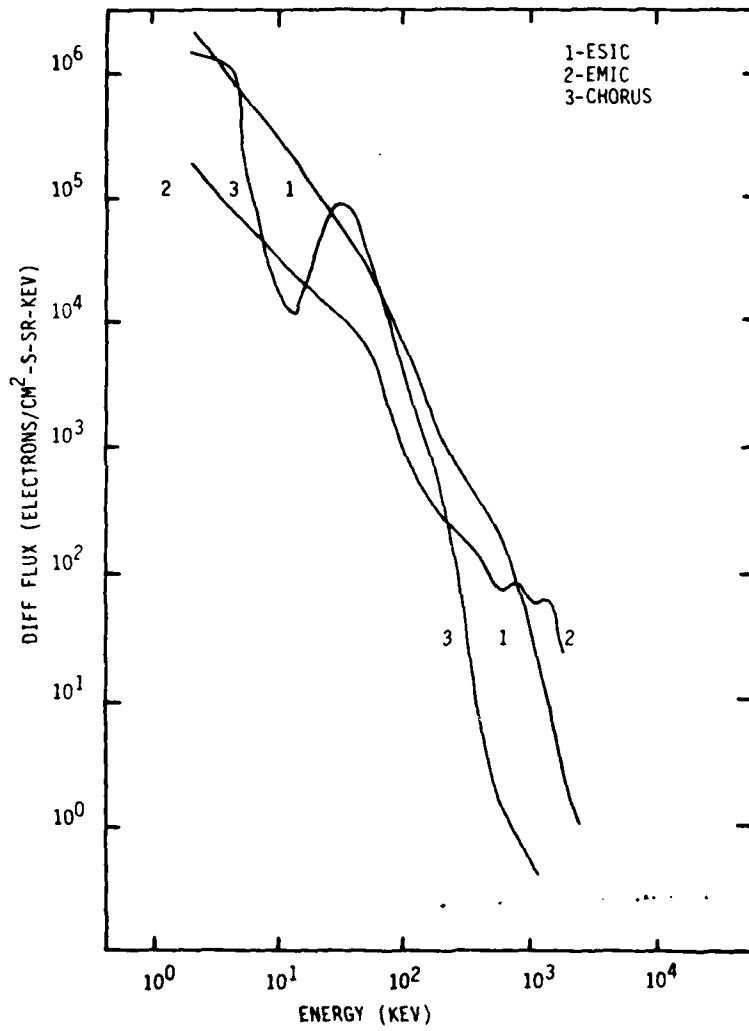
REPs can be the dominant source of NO production by dissociation of N_2 in the mesosphere and a competing source with nitrous oxide upwelling from the Earth's surface in the upper stratosphere (Thorne, 1979).

Above 40 km the odd hydrogen chemistry begins to dominate in the removal of odd oxygen (Blake and Lindzen, 1973). Here the odd hydrogen is produced mainly by dissociation of water vapor by photons and O ('D) atoms and as a by-product of O_2^+ production by energetic particle precipitation (Rowe et al., 1974). In this way, the increase in O_2^+ during REPs may also affect the mesospheric removal of odd oxygen. Since energy deposition by REPs can dominate in the mesosphere, we shall subsequently examine the removal of odd oxygen by the enhanced odd hydrogen chemistry and its ozone consequences for the three REP examples of Chapter 2.

3.2 REP Energy Spectra and Atmospheric Energy Deposition

The precipitation spectra for the three cases presented in Chapter 2 are shown in Fig 3.1. The ESIC spectrum (Case 1) has the

FIGURE 3.1 Energy spectra for cases 1,2,3 of Chapter 2.



highest flux in the 100-800 keV range. The EMIC spectrum (Case 2) dominates at energies above 1 MeV and is comparable to the previously published spectrum of Vampola (1971). The chorus event is relatively weak dominating only near 3 and 30 keV. These REP spectra were input to the UCLA (Spjeldvik, 1974) electron precipitation model (see Appendix D). The ion pair production (assuming 35 keV per pair) versus altitude is shown in Fig 3.2. The Bremsstrahlung X-ray deposition profiles were also plotted in Fig 3.2. As expected, the ESIC curve dominates in the 70 km region, and EMIC in the 50-60 km region. Whistler events are relatively non-competitive with other REP spectra but may dominate local chemistry.

The ESIC case of Chapter 2 was used to model a typical REP to calculate the total energy deposition. The latitudinal extent was approximately 200-300 km based on the invariant latitude excursions of both S3-2 and S3-3. The longitudinal extent based on local time duration was 7,000-10,000 km. The event lasted approximately 2.2 hours. The peak ion production rate occurred at ~73 km altitude with a value of $6.8 \times 10^{-3} \text{ (cm}^{-3} \text{ sec}^{-1}\text{)}$. Assuming precipitation was nearly constant and uniform throughout the events (which is not likely), a peak value of $\sim 1 \times 10^{24}$ ion pairs cm^{-1} (column) or 3.5×10^{22} keV of energy was deposited during this event.

3.3 REP Induced Hydrogen and Ozone Modifications

The REP ion production rates versus altitude were processed to produce odd hydrogen (and thus hydroxyl) according to the chemical modelling of Heaps (1978). Fig 3.3 presents this model of odd hydrogen

FIGURE 3.2 Ion pair production rate versus altitude for ESIC,
EMIC, and chorus events.

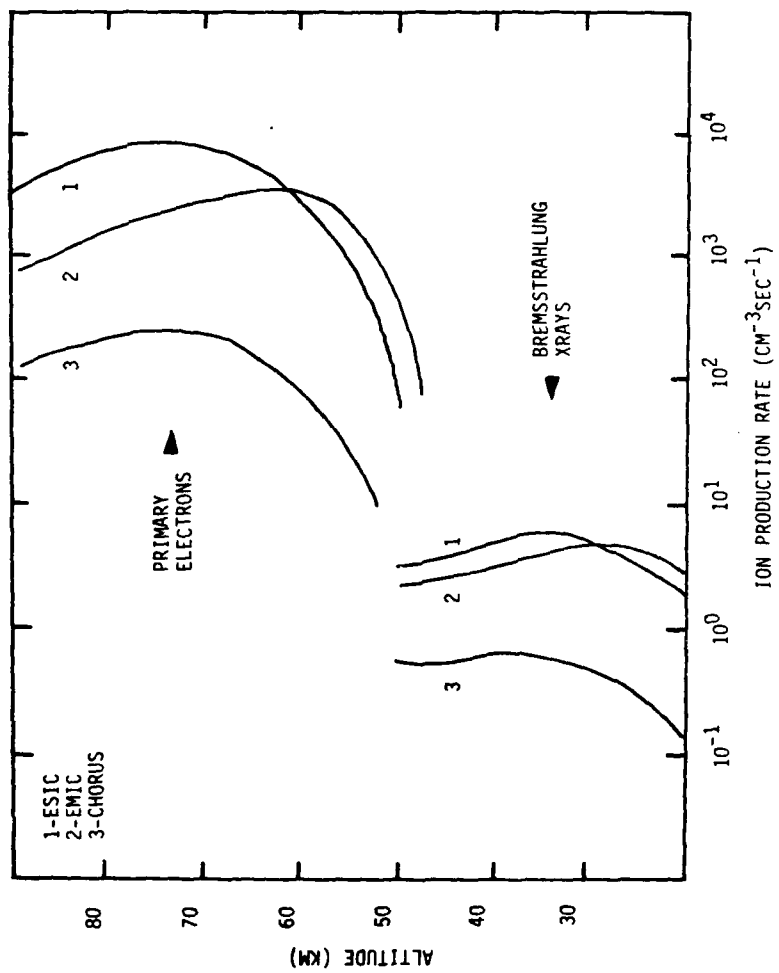
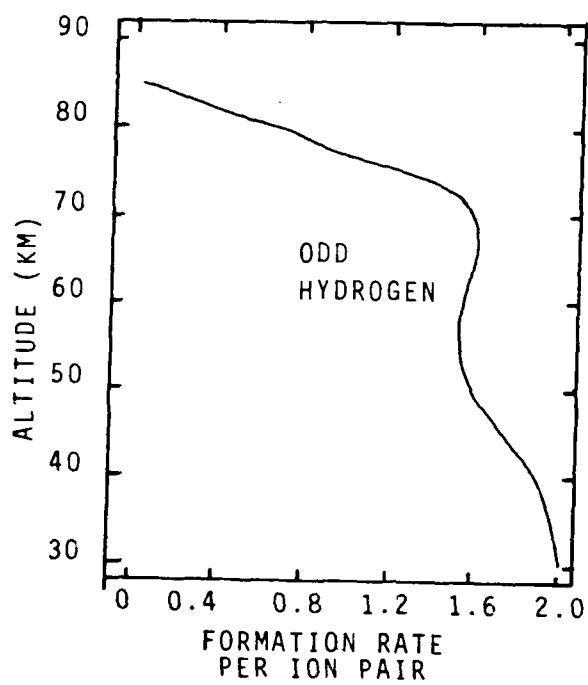


FIGURE 3.3 Altitude profile of the odd hydrogen formation rate related to ion pair production.



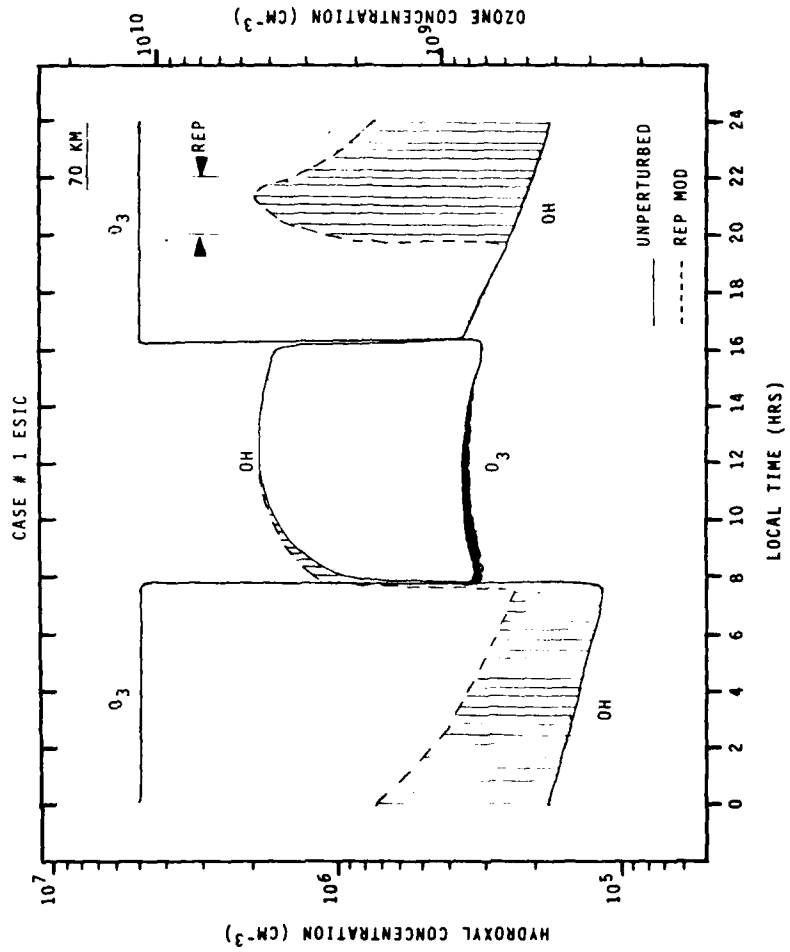
production as a function of ion pair production for disturbed conditions (higher ion production rate). The higher altitude limit for cluster ion formation is clearly indicated by the trailoff above 80 km. Since the nitrogen cycle dominates in the stratosphere, odd hydrogen is not expected to be important below 40 km.

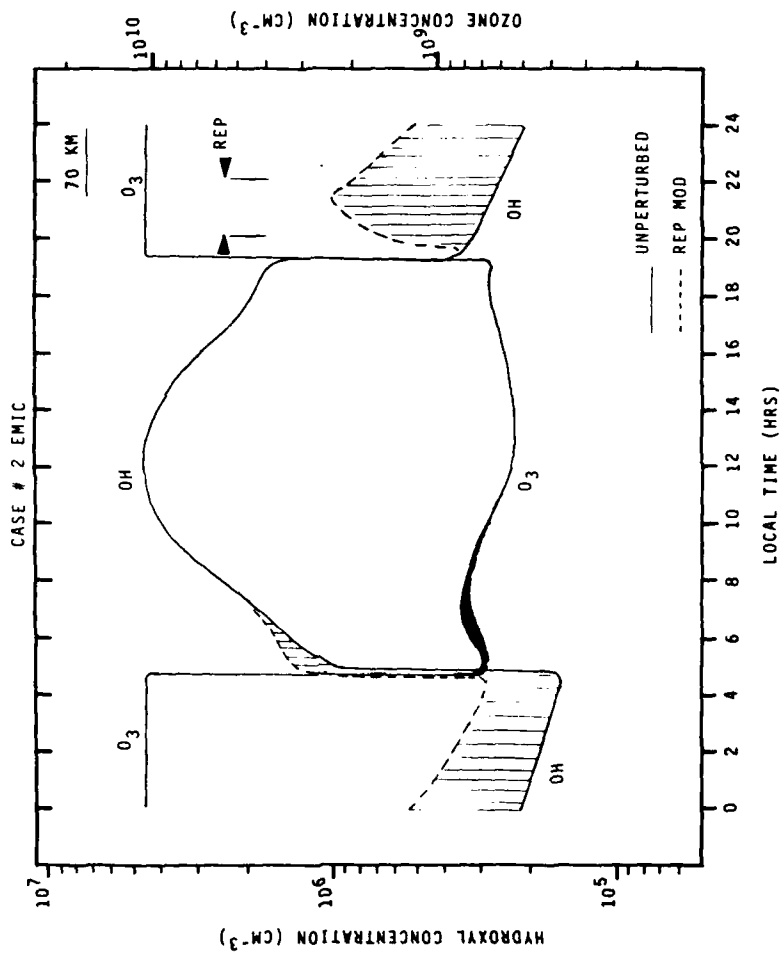
After modelling the OH increases to the REP ionization rates, a numerical simulation fashioned after Blake and Lindzen (1973) was performed to calculate the relative change in ozone concentrations. The numerical simulation differs from Blake and Lindzen's model in that two reactions (OH nighttime source reaction and Chapman's (1930) original 3 body atomic oxygen collision equation to form O_2) are added, reaction rates are updated, and solar flux and ozone concentrations vary with time of day (DeVore, 1977).

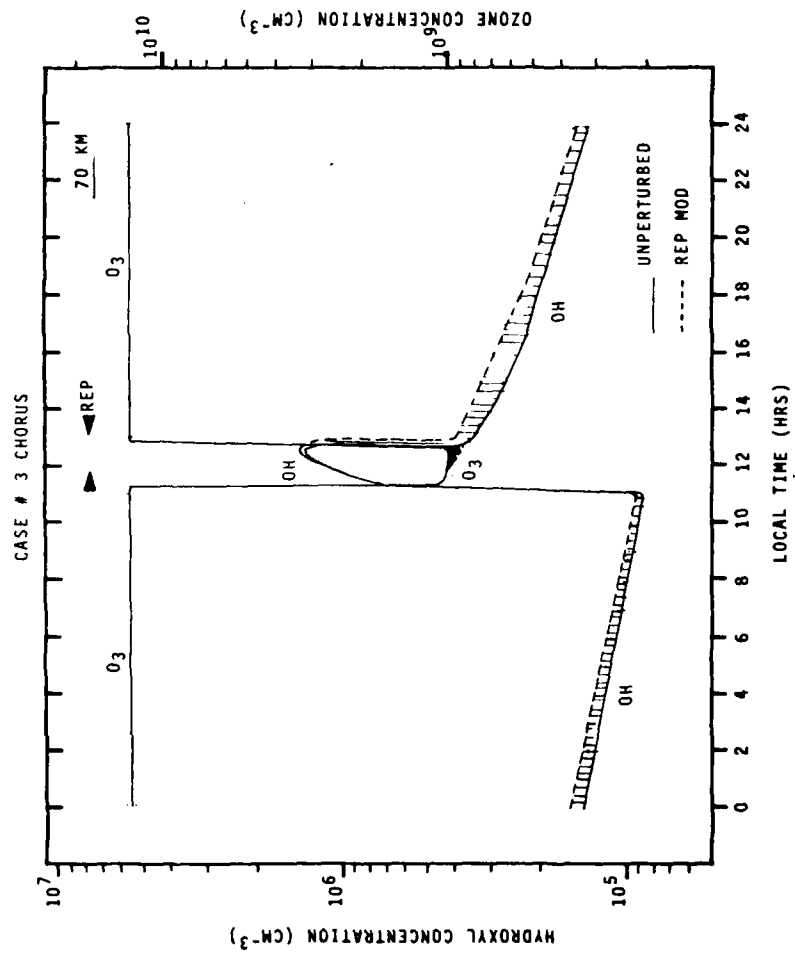
The results for Case 1 and 2 (Fig 3.4a and b) are similar since they both occur during local evening. Factors of 10 and 4 increase in OH concentrations at 70 km immediately after precipitation commenced occurred for the ESIC and EMIC cases respectively. The enhancement continued throughout the night and several hours into the day. Ozone was not affected until sunrise due to the low nighttime O concentration but showed sunrise decreases of 10-15% for both cases. These diminished ozone concentrations lasted for several hours.

The chorus event (Case 3, Fig. 3.4c) occurred near local noon with a small enhancement in OH concentration and throughout the succeeding night. The ozone concentration showed an immediate response dropping by nearly 20%. The diminution of ozone ended at local sunset.

FIGURE 3.4a,b,c Ozone and hydroxyl concentrations [cm^{-3}] versus local time (hrs) for a. ESIC, b. EMIC, and c. chorus REP events.



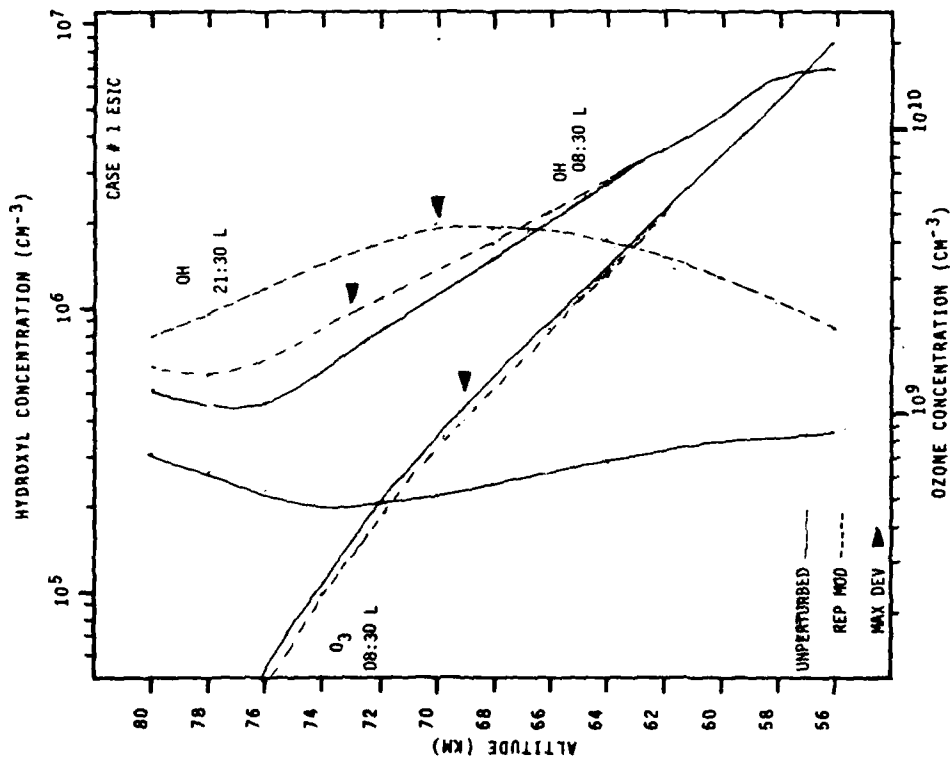


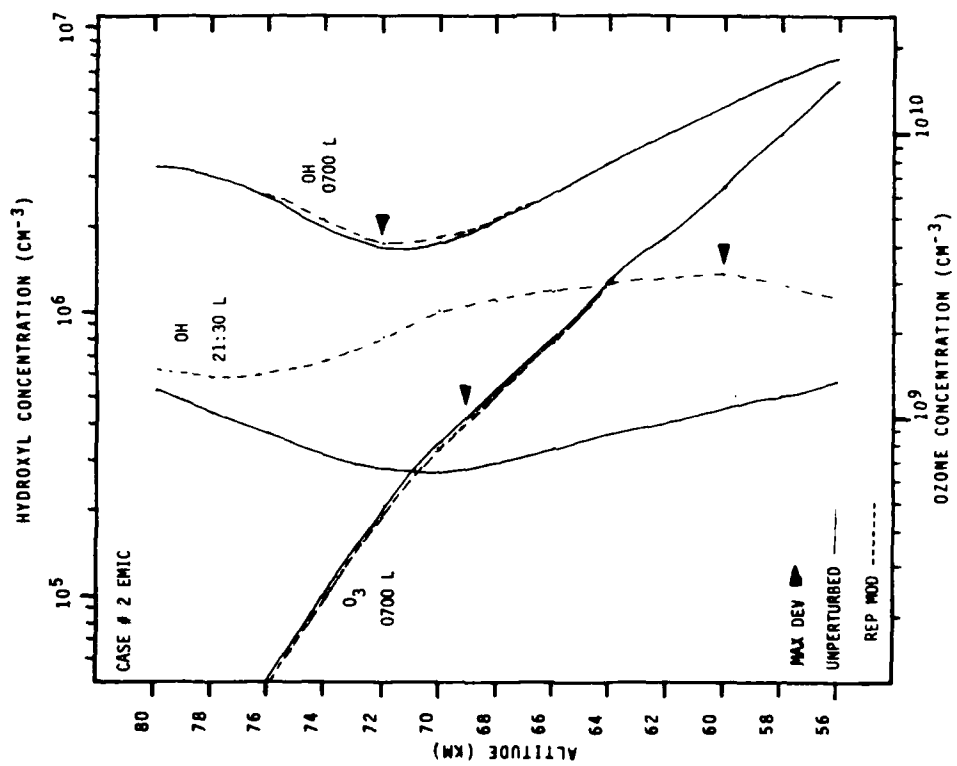


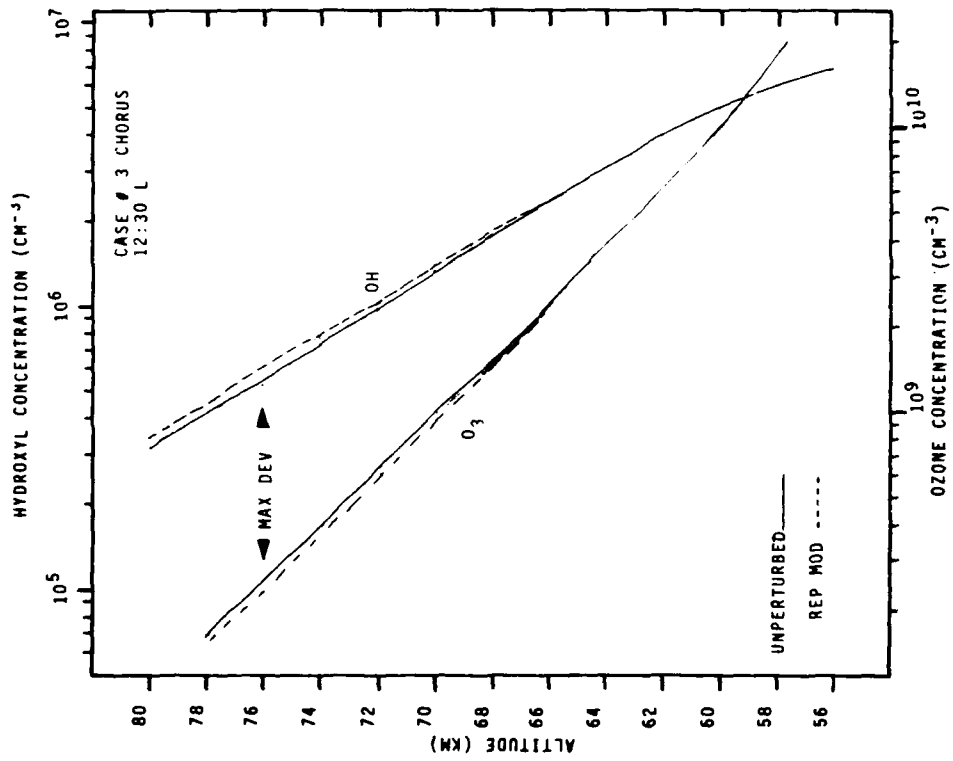
A similar simulation with an intense REP modelled after the ESIC case, this time occurring at local noon, showed only a 17% increase in OH but more than a 30% decrease in ozone.

Altitude profiles of the ozone concentrations at the local time of maximum decrease, OH concentrations at the same time and at time of maximum OH enhancement for the three cases are presented in Figs 3.5a,b,c. Maximum ozone decreases occurred at 69 km for both the ESIC and EMIC cases and at 76 km for the chorus case. Only in the chorus case, where onset of OH enhancement and ozone diminution were simultaneous, did the deviations occur at the same altitude.

FIGURE 3.5a,b,c Altitude [km] profiles of ozone concentration at time of maximum decrease, hydroxyl concentration at the same time, and hydroxyl concentration at the time of maximum enhancement for a. ESIC, b. EMIC, c. chorus REP events.







CHAPTER 4. SUMMARY, CONCLUSIONS AND OUTLOOK

I have presented here the results of two years of data analysis of fourteen months of satellite outer radiation zone coverage (Chapter 1). Theory and examples have been provided in an attempt to explain the types of REPs discovered in the analysis (Chapter 2). Finally, some atmospheric consequences were investigated (Chapter 3).

For the first time the properties of strong diffusion REPs have been studied using a relatively extensive, dense satellite data set and the following results have been obtained:

- o REPs show a seasonal variation with equinoxial maxima and solstice minima in agreement with Bailey (1968).
- o The occurrence frequency of REPs is nearly linear with general magnetic conditions again in agreement with Bailey (1968).
- o Correlation of REP producing substorms and substorm indices (AE) are weak in the limited number of cases examined in agreement with Thorne and Larson (1976) and not all substorms produce REPs in agreement with Rosenberg et al. (1972) and Thorne and Larson (1976).
- o Overall, only one out of ten REPs occur in daytime but as magnetic conditions become disturbed there is an increase in daytime REPs approaching three out of ten.
- o REPs are estimated to extend in longitude from 1000 km to 10-15,000 km in general agreement with Larsen, 1973, and Bailey, 1966.

- o 98% of all REPs occur concurrently with protons (≥ 80 keV) precipitating on strong diffusion and many features of the precipitation are similar.
- o Three distinct classes of REPs occurred: an overwhelming majority (96.5%) had strong diffusion precipitating spectra from hundreds to a thousand keV energy and all these events occurred near the outer bound of trapping simultaneously with proton precipitation; the smallest group (1.3%) had strong diffusion precipitating spectra only for electrons ≥ 1 MeV and again with concurrent proton precipitation but inward from the outer bound of trapping and probably within the plasmasphere; a slightly larger group (2.2%) had strong diffusion precipitation spectra only for electrons of a few hundred keV but with protons present with normal loss cone distributions (not precipitating on strong diffusion).

The theoretical investigation shows that REPs can be caused by radiation belt plasma interactions with plasma waves and that the electrostatic ion cyclotron (ESIC), electromagnetic ion cyclotron (EMIC) and whistler mode chorus waves can respectively account for the three distinct types of REPs discussed above.

Atmospheric consequences were investigated with the following results:

- o ESIC and EMIC REPs deposit more energy into the atmosphere than chorus REPs.
- o Immediate increases in the OH concentration up to a factor

of 10 are observed during REP occurrence and gradually return to normal within 24 hours.

- o Ozone diminution only occurs simultaneously with daytime REPs, only after sunrise with nighttime REPs, and return to normal in several hours.

The following main conclusions may be drawn from these results:

- o The ESIC wave interaction is by far the dominant cause of REPs.
- o REPs are locally important in reducing ozone concentration (10-30%).

Based on all of the above, future considerations should include the following areas of endeavor:

- o Although the concurrent proton precipitation seems to point to the anisotropic pitch angle instability at the equator rather than to a current driven instability low on the field line, results are inconclusive (Kintner et al. 1979). Resolution of the location of the instability would lead to better understanding of the trigger mechanism for REPs.
- o Relativistic electron precipitation at rates less than strong diffusion were not analyzed here due to the difficulty with distinction of the pitch angle distribution and the instrument response limits for small loss cones. Weaker events (moderate diffusion) may be locally important to ozone chemistry particularly if they occur during the daytime.

- o Though REPs locally increase the OH concentrations day or night, an investigation of long term effects on the atmosphere are important since it may be possible to dry out the mesosphere over a year's time and actually account for long term ozone buildups (Crutzen, private communication, 1980).

Appendix A - OVI-19 Data Set and Analysis

The data used for this study was collected by two magnetic spectrometers aboard Air Force Office of Aerospace Research satellite 1969-25C designated OVI-19 and provided by the instrument builder and prime investigator, A. L. Vampola, Aerospace Corp., El Segundo, Ca. The purpose of this study was to find the frequency of occurrence of REP events and determine the suitability of this type of data for future study.

The OVI-19 satellite was launched on 18 March 1969 into an elliptical orbit with apogee/perigee at 5790/467 km. Inclination was 104.8° and the period 154 min. The satellite was spin stabilized at 8.4 rpm. A three-axis magnetometer tabulated aspect information. Four hours of continuous data was collected and transmitted each tape recorder dump. Over the life of the satellite, approximately 25% data coverage was achieved.

The two spectrometers were of the 180° magnetic focusing type similar to those onboard AF/OAR satellite OV3-3 and schematically presented in Fig. A-1. (Vampola, 1969). One unit, the Low Energy Magnetic Spectrometer (LEMS), had eight electron channels from 53-444 keV with geometric factors as shown in Table A-1 and the other, the High Energy Magnetic Spectrometer (HEMS), had 16 electron detectors from 537 keV to 5.1 MeV as shown in Table A-2. The look direction of each instrument was perpendicular to the satellite spin axis (for pitch angle scanning) which was nominally perpendicular to the orbit plane. Flux counting rates were sampled once per second. As with

FIG A-1. OV3-3 180⁰ Magnetic Focussing Spectrometer (shown without baffling) from Vampola, 1969.

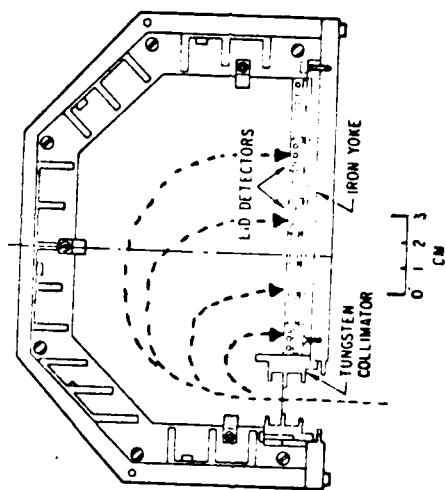


TABLE A-1. Low Energy Magnetic Spectrometer Channels, Energy, and Geometric Factors.

TABLE A-1. Low-Energy Magnetic Spectrometer

Detector	Electron Energy, kev	Geometric Energy Efficiency Factor, cm ² ster kev
1	53-14 ⁺¹⁵	0.00205
2	92-16 ⁺¹⁹	0.0833
3	139-21 ⁺²²	1.36
4	192 ± 24	2.53
5	250 ± 26	2.34
6	312 ± 27	2.16
7	376 ± 28	2.01
8	444 ± 29	1.86
9	Protons > 35 Mev, bremsstrahlung back- ground detection	

TABLE A-2. High Energy Magnetic Spectrometer Channels, Energy,
and Geometric Factors.

TABLE A-2. High-Energy Magnetic Spectrometer

Detector	Energy, Mev	Geometric Energy Factor, $\text{cm}^2 \text{ster kev}$
1	0.537- 1.14×10^3	45.7
2	0.822- 1.14×10^3	36.5
3	1.112- 1.14×10^3	30.1
4	1.409- 1.14×10^3	25.5
5	1.711- 1.14×10^3	22.1
6	2.018- 1.16×10^3	19.4
7	2.321- 1.19×10^3	17.3
8	2.632- 1.14×10^3	15.7
9	2.932- 1.14×10^3	14.3
10	3.245- 1.14×10^3	13.1
11	3.557- 1.14×10^3	12.1
12	3.863- 1.14×10^3	11.3
13	4.169- 1.14×10^3	10.5
14	4.476- 1.14×10^3	9.87
15	4.783- 1.16×10^3	9.29
16	5.091- 1.14×10^3	8.67
17	Protons >130 Mev, bremsstrahlung background detection	

the OV3-3 units, the flux determinations were limited at low counting rates by statistical fluctuations and at high counting rates by telemetry quantization (Vampola, 1969). In addition, a rectangular energy versus geometric factor response is assigned for each actual channel response for data analysis as illustrated in Fig. A-2. Instrument response to an input spectrum zero everywhere in the loss cone and isotropic at some level outside the loss cone is shown for Channel 4 of the low energy unit in Fig. A-3. An order of magnitude drop in flux response roughly occurs 7° into the loss cone.

A typical high to low altitude pass for Channel 7 of the low energy unit is shown in Fig. A-4. Inner zone (A), slot region (B), outer zone (C), and polar region (D) are distinctive features of the pass. The large variation in count rate is due to pitch angle sampling.

Fig. A-5 for an OV3-3 sample shows the outer zone at low altitude from L=3-6 with envelopes drawn for the top and bottom of the sampling scatter. The top envelope is representative of the locally trapped flux and the bottom is the flux furthest into the loss cone. When the two curves touch, as in Fig. A-5 at L=5.6, this indicates an isotropic pitch-angle distribution, as predicted for strong pitch angle diffusion. The analysis mode is then to view all of the available passes for this envelope closure. Final determination is made when flux vs pitch angle plots show isotropic fluxes filling the upward looking (downward flowing) loss cone.

In order to accomplish the task of analyzing the data set

FIG A-2. Geometric factor-energy response function assigned each channel for data analysis. Total area outside of the rectangle is about 15% but both functions contain equal areas. A slight bias towards higher energies for very steep spectra is produced (Vampola, 1969).

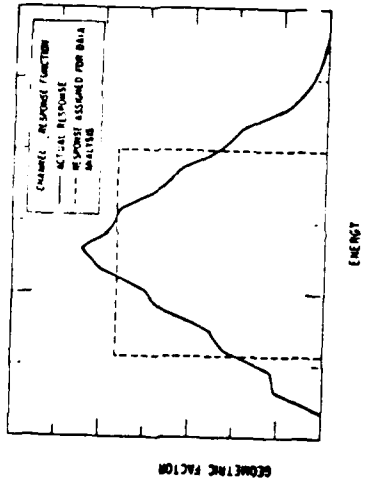


FIG A-3. Instrument response-pitch angle output.

OUT-19 LEMS CHANNEL 4 (.197 MEV) 30 DEGREE LOSS-CONE

INPUT RESPONSE: 0 FOR P.A. $\leq 30^\circ$ AND $\geq 150^\circ$
1 FOR P.A. $> 30^\circ$ AND $< 150^\circ$.

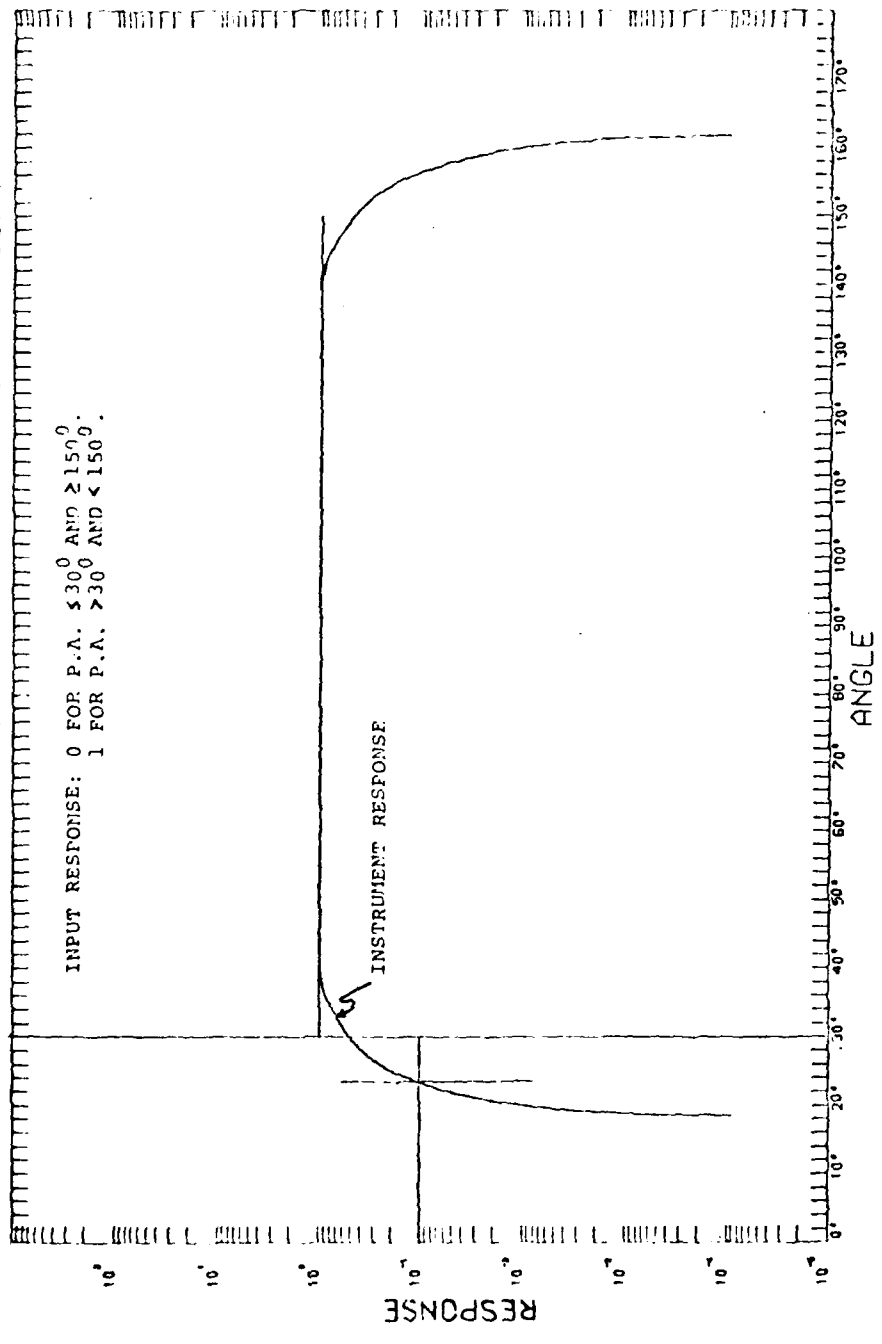


FIG A-4. Analog record of 376 ± 28 keV electron data plotted as counting rate versus time on a pseudo-logarithmic scale. Universal time, magnetic field intensity, invariant latitude, L, altitude, and MLT are listed for 200-sec intervals. Pitch-angle distribution sampling produces the large variations in counting rate in several regions of the plot (Vampola, 1971).

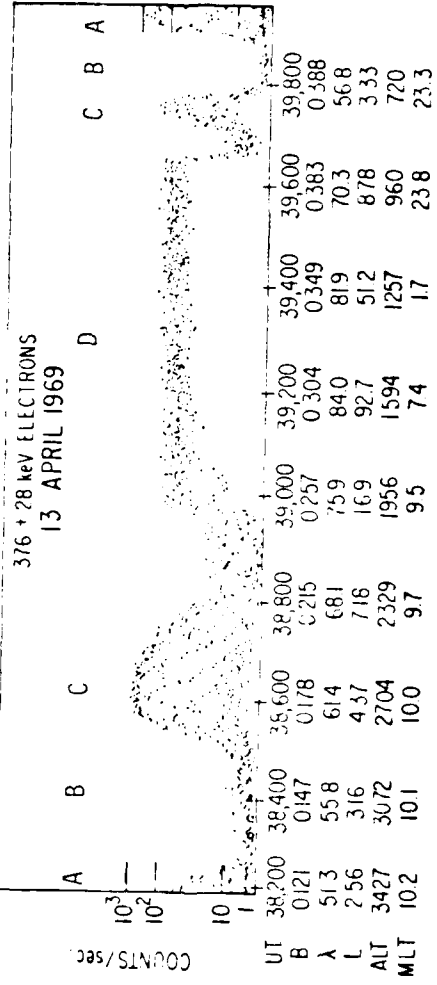


FIG A-5. Plots of instantaneous flux measurements versus L for the 712 ± 137 keV electron channel from the OV3-3 magnetic spectrometer. The scatter in the data points is due to pitch-angle sampling. Envelope curves at the maximums and minimums are shown to emphasize the data that are isotropic in pitch angle (Koons et al, 1972).

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1980 L J ANDREOLI
AFIT-CI-80-26D

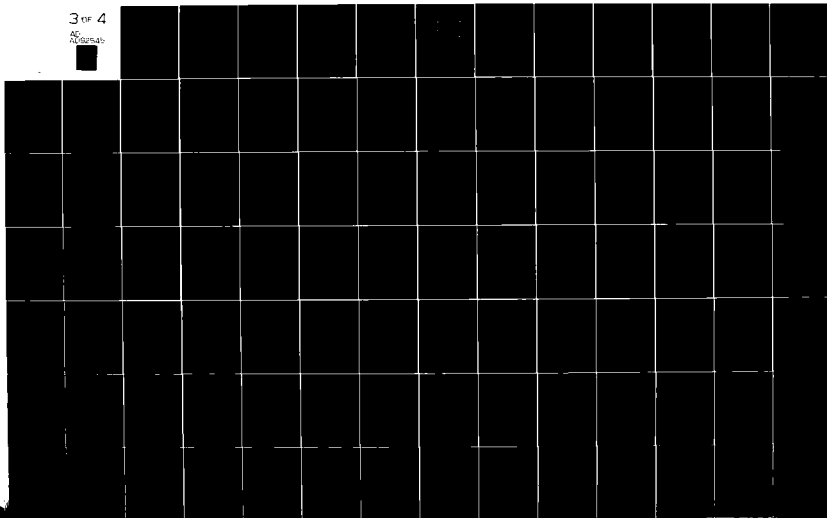
F/G 4/1

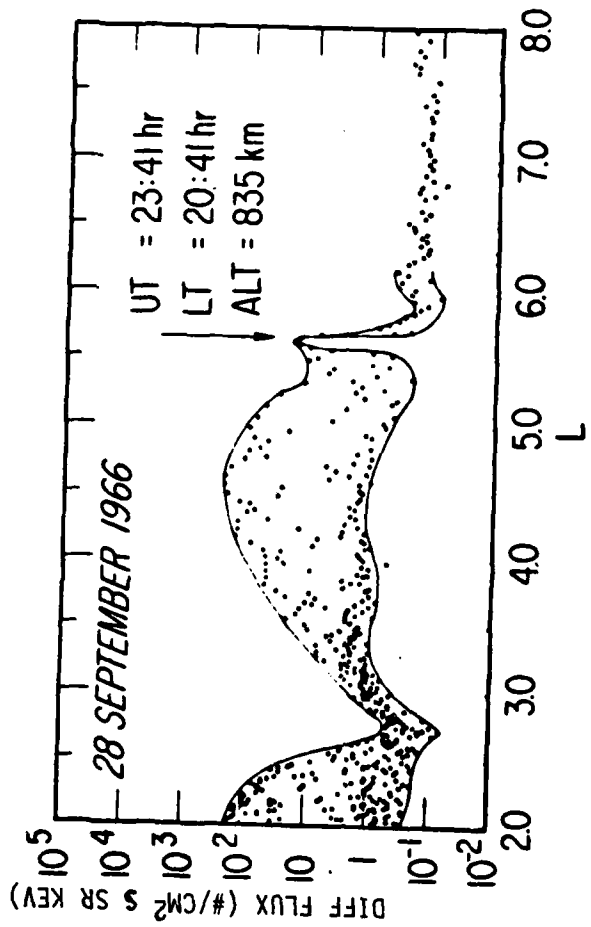
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as objectively as possible, a set of criteria was developed for REP event identification. These are illustrated in flow chart form in Fig. A-6. The first criteria was that only data in the L range 4-7 would be studied. A total of 320 passes, clustered as high as 4 passes on a given day, was extracted from a data set spanning from day 77 of 1969 to day 13 of 1970 (301 days). Because of the finite width of the field-of-view ($\sim 10^\circ$) and the fact that the satellite orbit precluded pitch angle sampling to 0° (and 180°) everywhere in the orbit, it was determined that data below 1000 km would yield loss cones large enough for a high frequency of complete loss cone sampling. This criteria eliminated 88 cases. The third criteria demanded that REP fluxes be high at higher energies ($1 \text{ MeV} \pm 500 \text{ keV}$). Seventy-two cases failed on this criteria. Next we determined that if we had data at least 10° into the loss cone, we could identify real precipitation based on the channel response previously described. This eliminated 39 cases of which 17 were less than 10° into the loss cone and 22 did not sample the loss cone at all. In some cases (19), count rate sampling and pitch angle modulation were phased in such a way as to make event evaluation impossible. These cases were labeled "instrument problems" and eliminated. Finally, 103 cases remained in which an analysis could reasonably be made. These passes are shown as passes per day versus day of the year in Fig. A-7. Of these 103 "clean cases", only three exhibited isotropic fluxes into the loss cone at relativistic energies (identified by arrows on Fig. A-7).

The REP of 14 May 1969 (day 134) is of particular interest since it exhibited the "hardest" spectrum of the three REPs (spectra

FIG A-6. Flow chart illustrating process of elimination through the
OVI-19 data set.

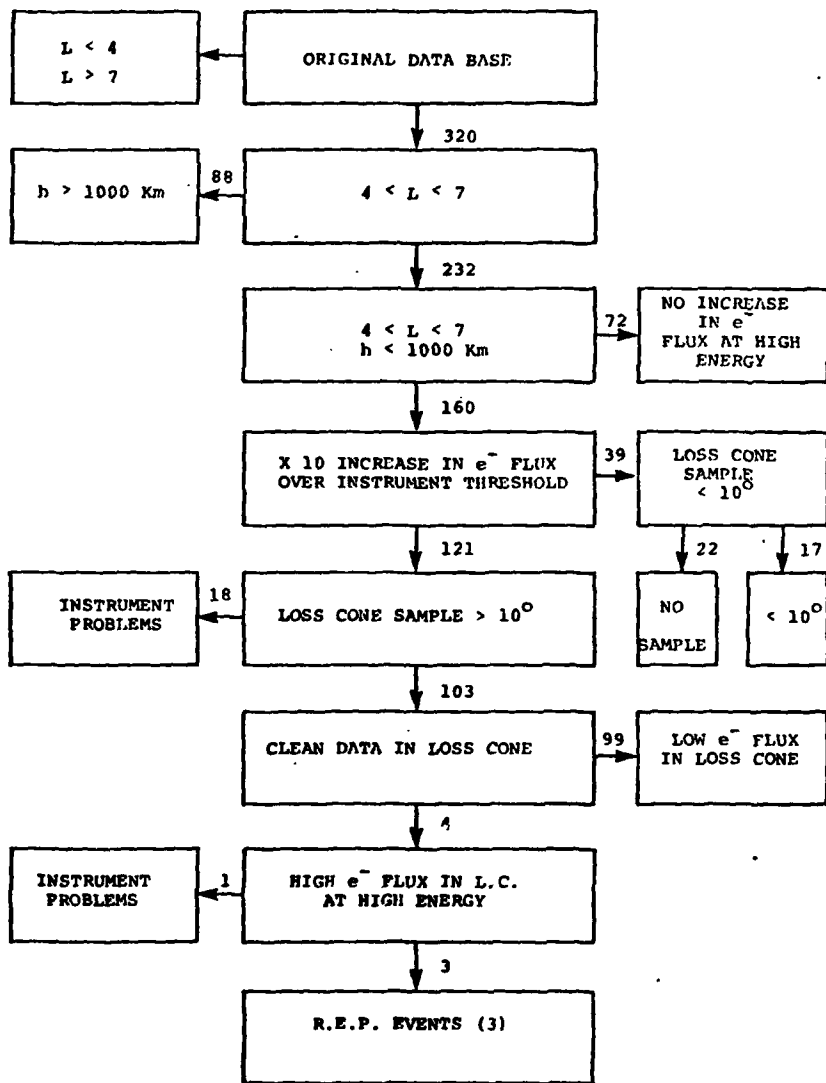
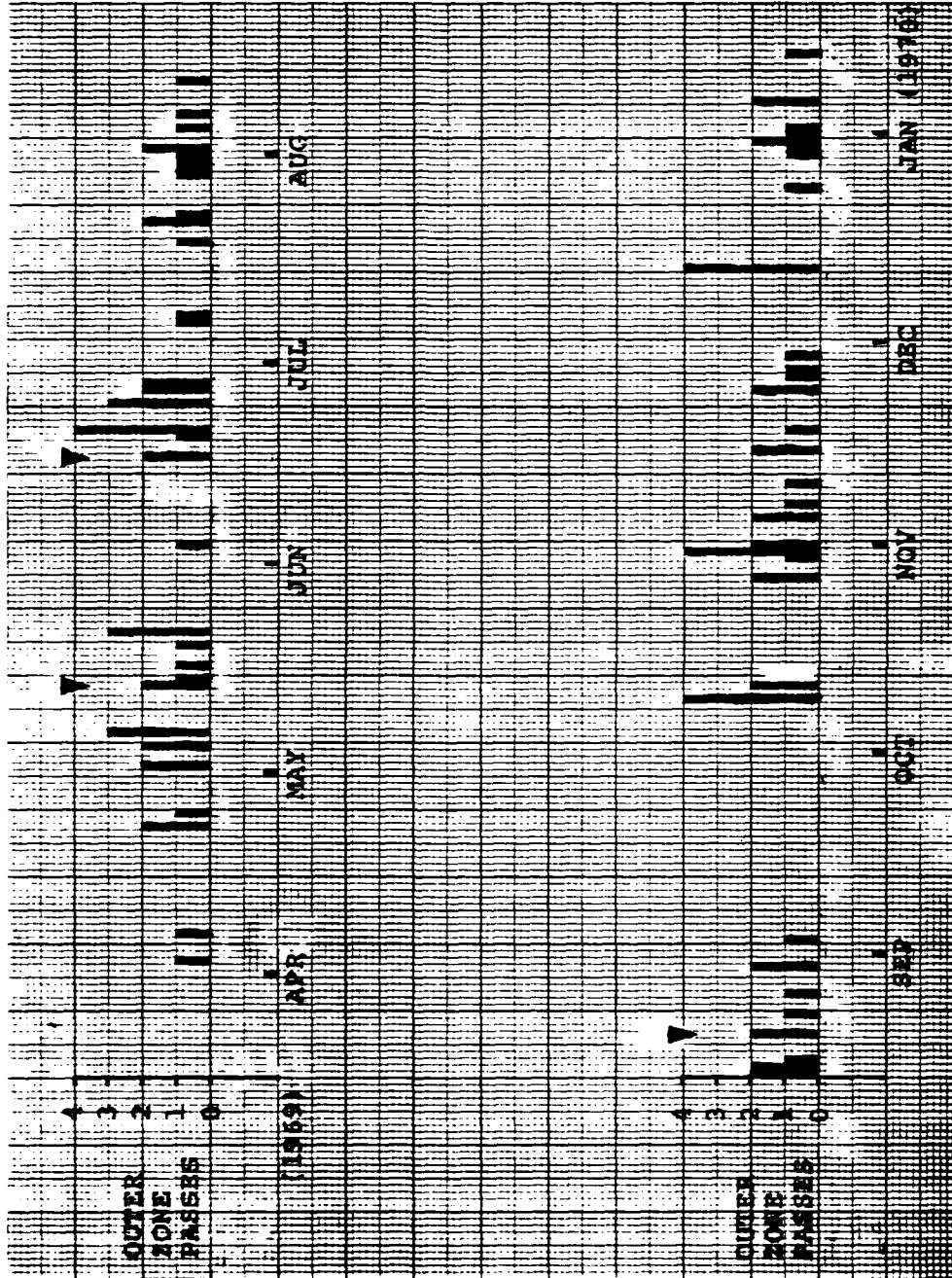


FIG A-7. Outer zone passes per day versus day of the year for the OVI-19 data set. Arrows indicate the three event days.



will be compared later) and falls in a period where Alaska forward scatter data exists. Figures A-8, a and b, show plots of differential electron fluxes versus U.T. (in seconds) for the various energy channels, LEMS Bremsstrahlung (background), and LEMS protons (>55 keV) for day 134. The higher energy electron channels exhibited low fluxes and are not shown. The outer edge of the outer zones at local morning (07:20 MLT: 79642-79659 U.T.) and local midnight (00:18 MLT; 80184-80225 U.T.) exhibit envelope closure. The local morning outer zone exhibited very high trapped fluxes (see Fig. A-9) at all energies but the loss cone ($\sim 55^\circ$) was not sampled.

The pitch angle distribution for the local midnight (80196-80225 U.T.) outer zone transit are shown in Fig. A-10 for energy channels 192, 537, 822 keV. The nominal loss cone is $\sim 66^\circ$ and the symbols "+/o" represent instrument viewing positions up/down the local magnetic field line. The plots show high fluxes in the upward viewing loss cone at several instances in all energy channels while the downward viewing loss cone generally show normal trail-off to lower values. The large variation in precipitating fluxes is due either to the transition of the satellite through regions of strong and not so strong precipitation or spatial variations in trapped flux during the 30 seconds duration of the pitch and angle plot. The plots represent approximately seven complete (0-180 $^\circ$) pitch angle distributions each and 537 keV channel has each data sample labeled with the last three digits of the U.T. in seconds. Deep in the loss cone two orders of magnitude variation occur between 210 and 218 whereas a half order of magnitude between 205 and 220 or 201 and 197 occur in the trapped distribution.

FIG A-8a. Differential electron fluxes versus UT (sec), B (Gauss),
L, LAT (deg), ALT (km), for LEMS Proton, LEMS BREM, 822,
537, 444, 376 keV Electrons for DAY 134.

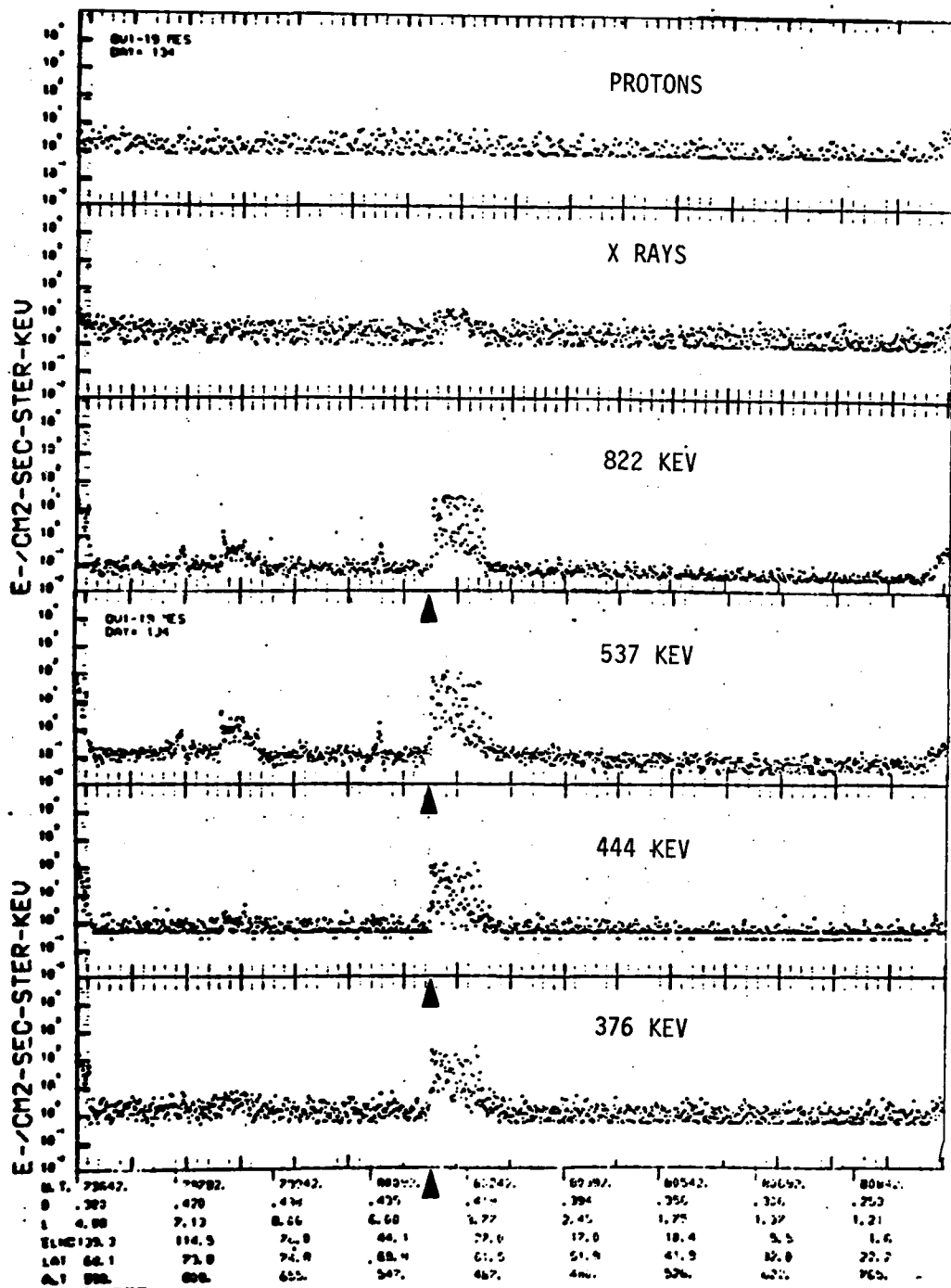


FIG A-8B. Differential electron fluxes versus UT (sec), B (Gauss),
L, LAT (deg), ALT (km), for 312, 250, 192, 139, 92,
53 keV Electrons for DAY 134.

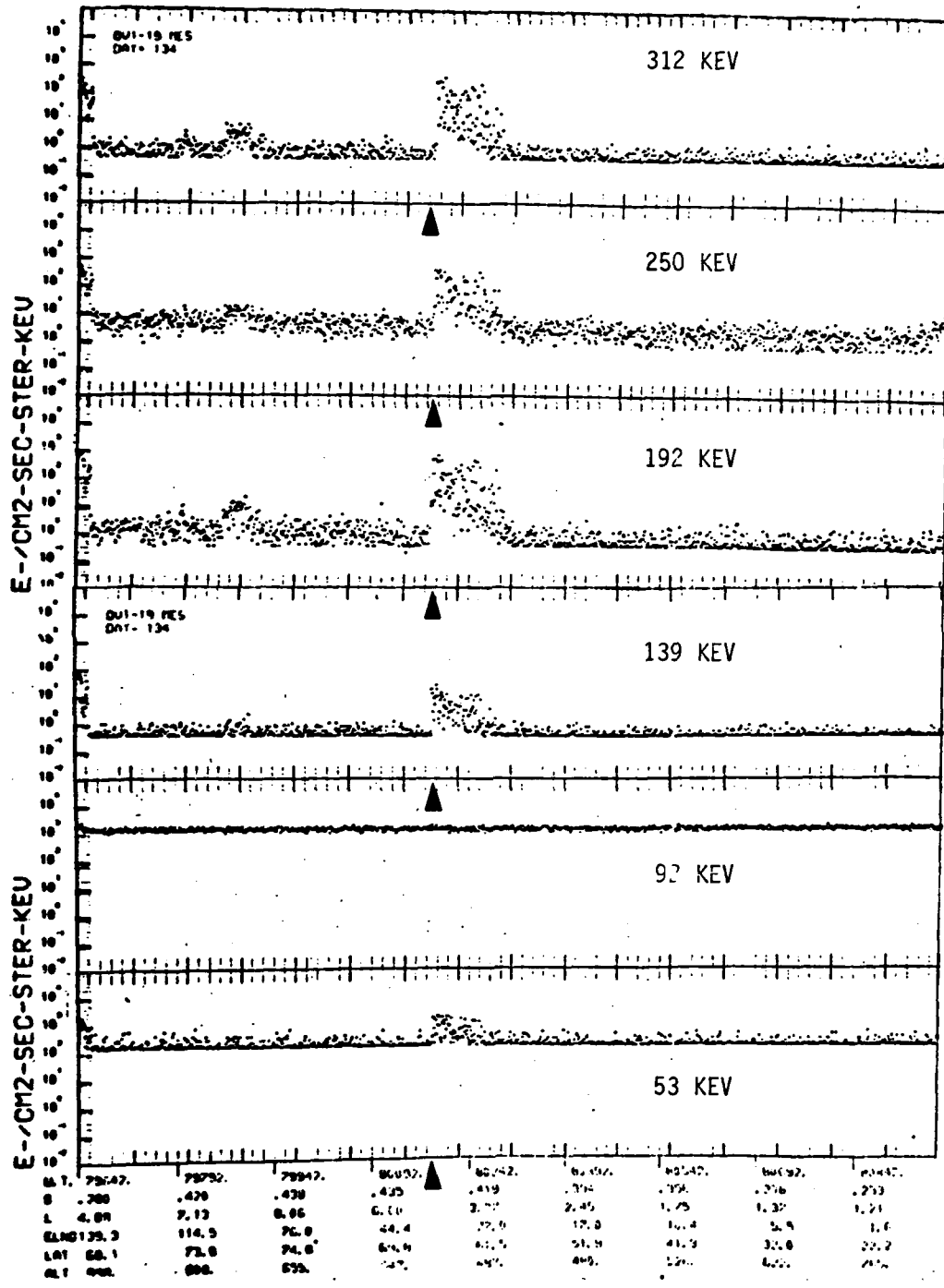


FIG A-9. Electron differential energy fluxes versus UT (hr), LT (hr),
ALT (km), L, for 192, 537, 822 keV electrons for DAY 134
(14 MAY 1969).

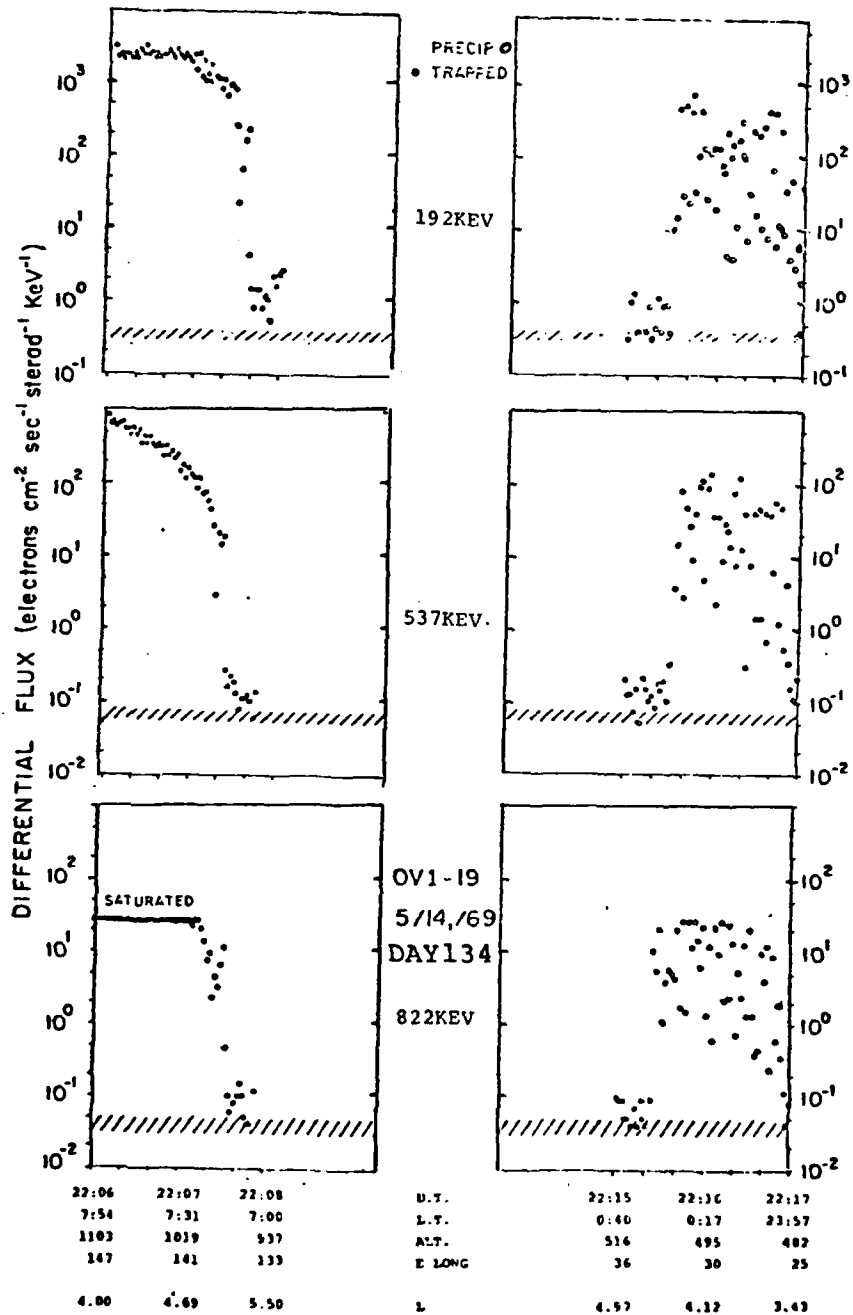
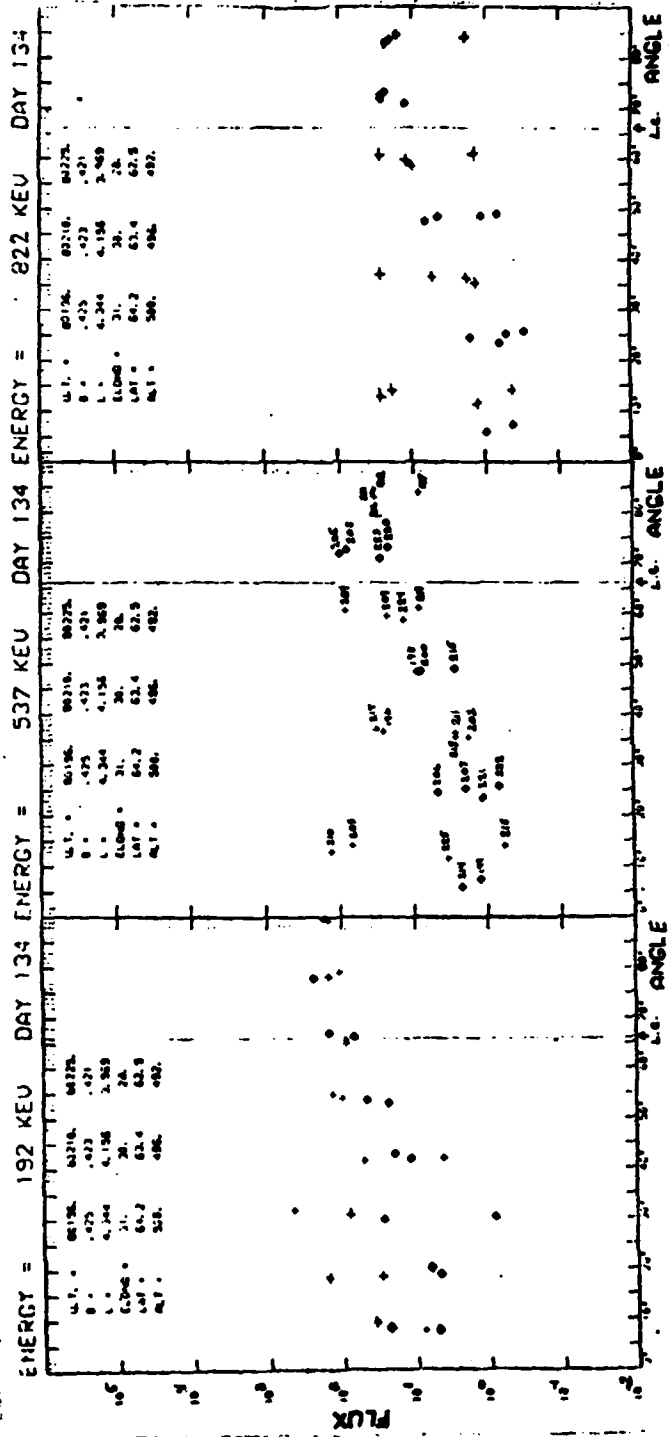


FIG A-10. Flux versus Pitch Angle for electron channels 192, 537, and 822 keV for DAY 134. Flux upper limit in channel 822 keV is 2.696×10^1 since the high count rate meter coax cable from the HEMS to the satellite data system broke during launch.



A time sequence in pitch angle versus differential flux for the same local midnight outer zone pass (day 134) for channels 444, and 537 keV are shown in Figs. A-11 a and b respectively. Data samples are subjectively connected to display the most "normal" distribution. This is difficult because of the wide spread in pitch angles ($\sim 50^\circ$) between samples and compounded by a few missing samples. The 444 keV channel is the highest energy measured by the LEMS whereas the 537 keV channel is the lowest for the HEMS. Thus, Fig. A-11 a is 165° out of phase with Figs. A-11b. Each sample is identified by the last three digits of the U.T. in seconds. "Normal" upward and downward loss cone distributions are evident at all energies after sample 213. The upper envelope (dashed line -U) connects the locally trapped flux (pitch angles $\sim 90^\circ$) and the lower (dashed line-L) connects the lowest flux into the loss cones (pitch angles $\sim 0^\circ, 180^\circ$). It is evident that the envelopes nearly close in at least a few places towards the outer edge of the outer zone though the instrument is viewing almost 2 1/2 orders of magnitude into the loss cone. Most notable is the isotropic (upward viewing loss cone) flux in Fig. A-11b at sample 188 (80188 U.T.).

The magnetic activity for this event as well as Alaska forward scatter absorption are illustrated in Fig. A-12. The Alaska forward scatter absorption on day 134 was greater than 10 db from 19:40 to 02:10 U.T. (09:40-16:10 Alaskan local time) with a peak absorption of 19 db at 20:12 U.T. The only other absorption recorded was early on days 132 and 137. Dst shows disturbed conditions

FIG A-11a. Time sequence of differential electron flux versus pitch angle for the 537 key channel on DAY 134.

537 KEV CHANNEL

DIFF FLUX (E/CM² SR S KEV)

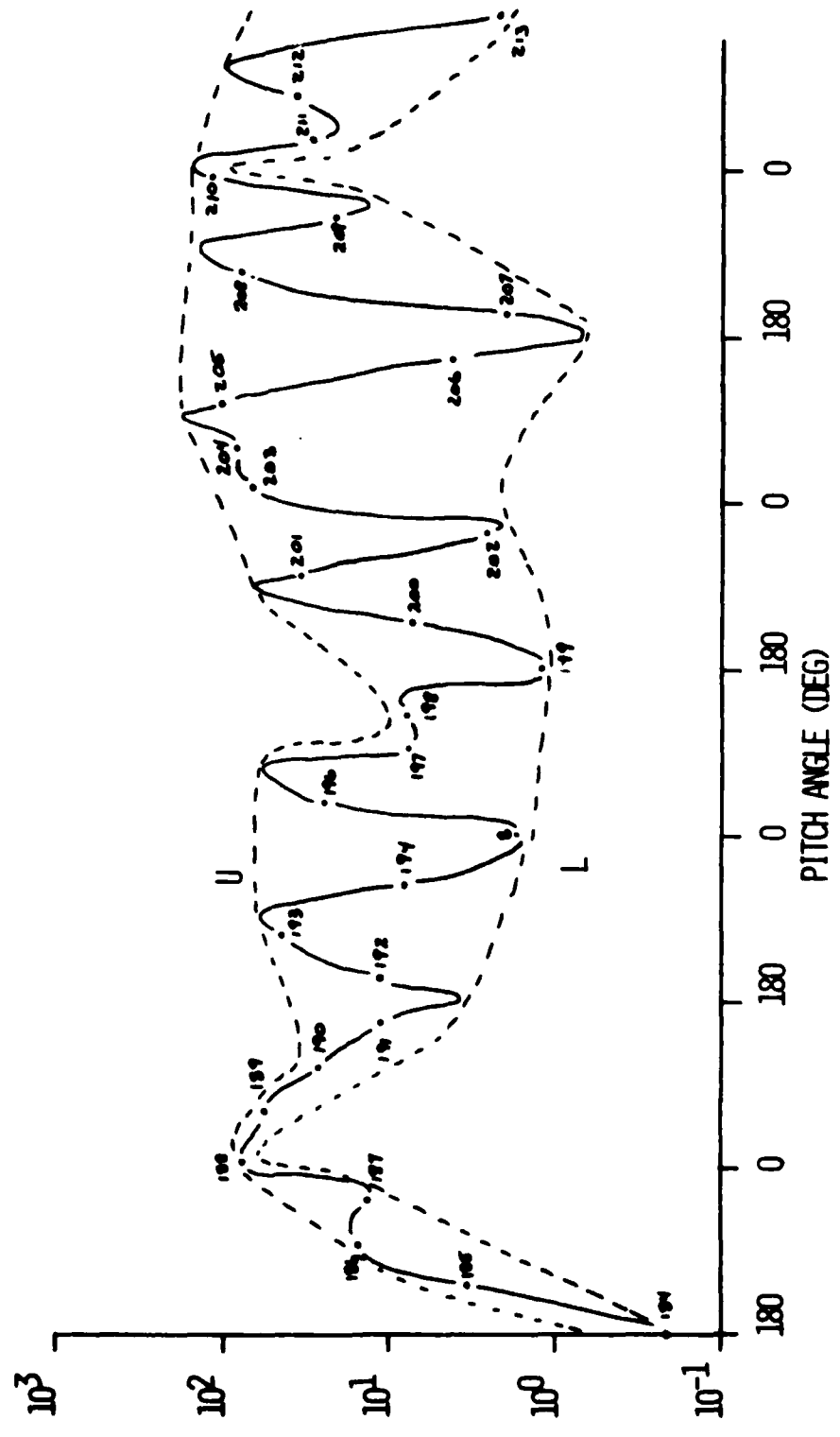


FIG A-11b. Time sequence of differential electron flux versus pitch angle for the 444 keV channel on DAY 134.

444 KEV CHANNEL

DIFF FLUX (E/CM² SR S KEV)

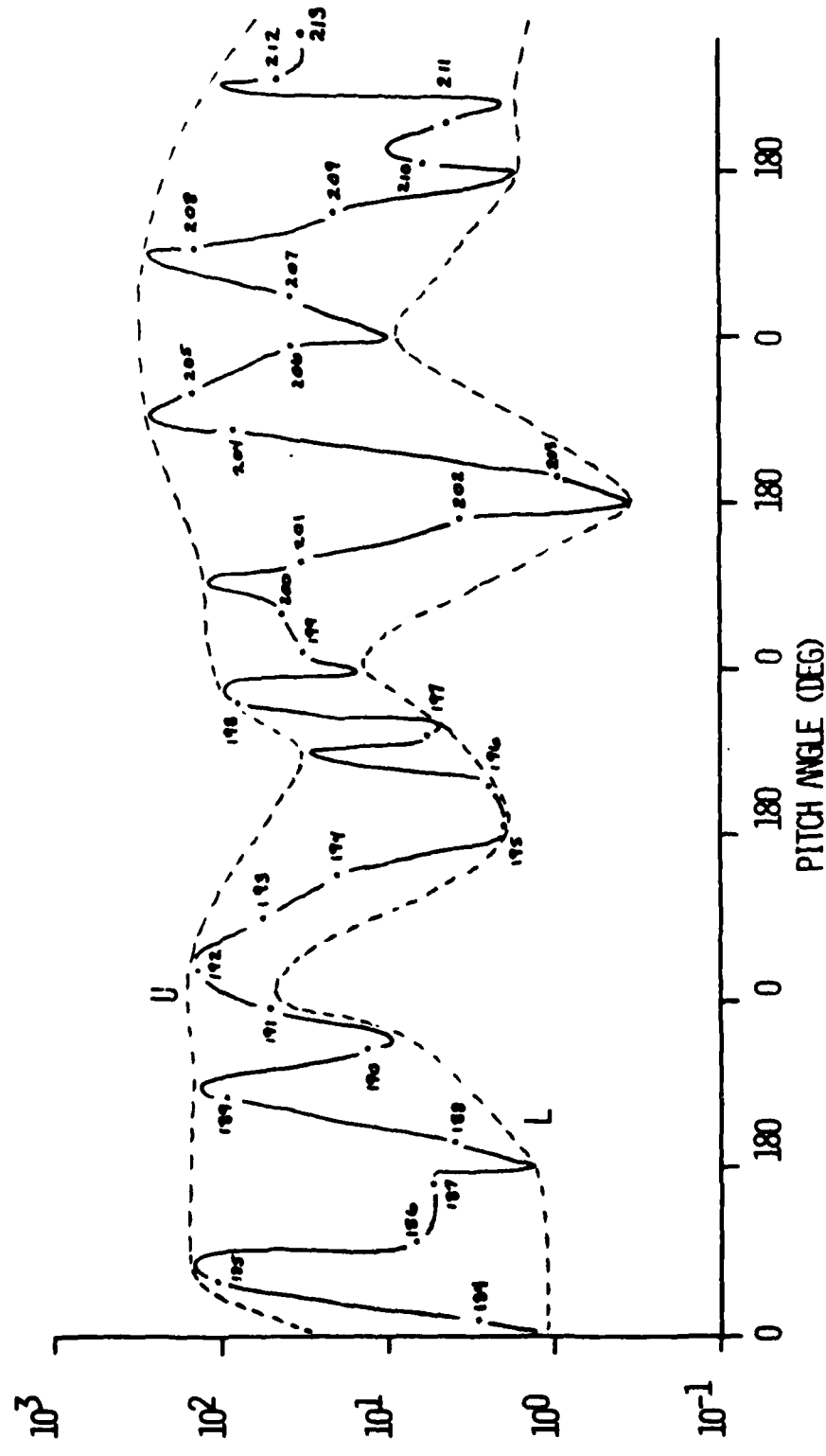
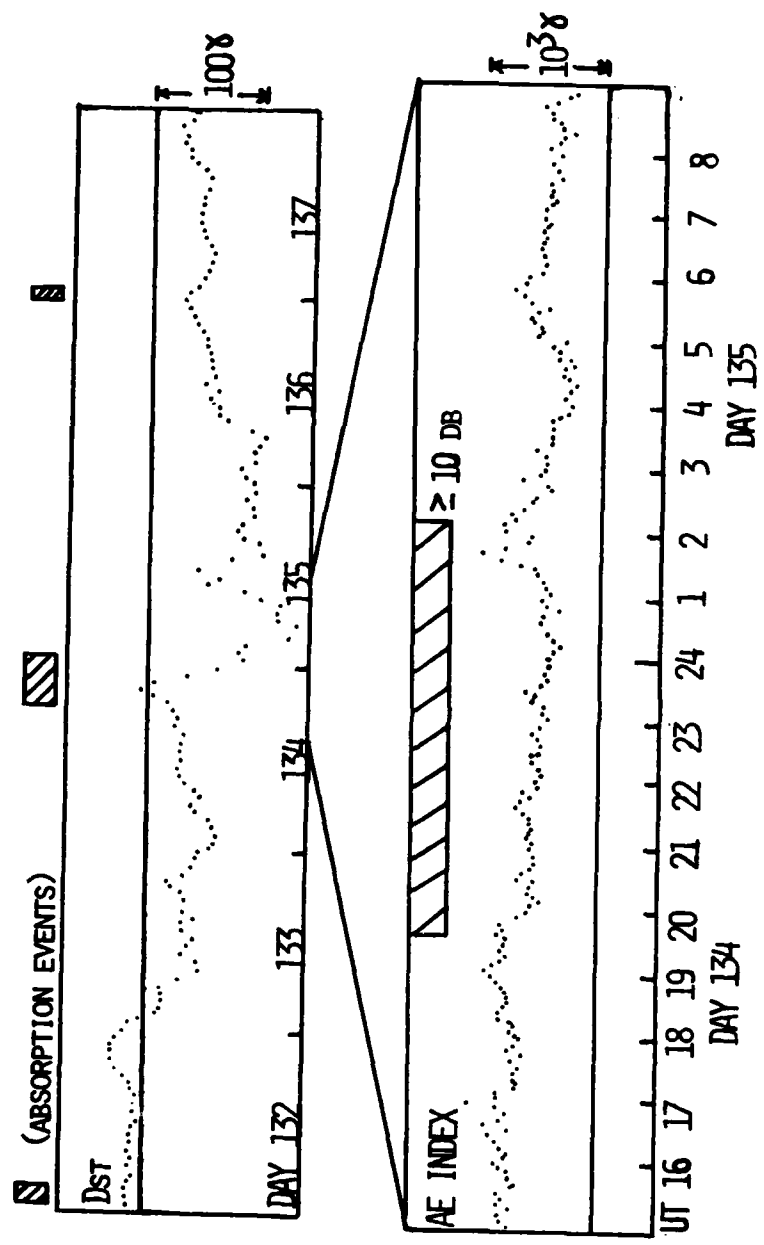


FIG A-12. Dst for days 132 through 137; AE Index for parts of the UT day for days 134 and 135; Alaskan forward scatter absorption events are represented by the hatched boxes (Bailey, Personal Communication, 1973).



from late on day 132 through recovery on day 137. A positive excursion is particularly noticeable near 22 U.T. on day 134 near the time of the satellite REP observation. This is followed by a $>100\gamma$ depression early on day 135. AE shows a 1500γ increase at 20 U.T. on day 134 two hours before the satellite observation. It is interesting to note that although day 133 and 135 are strongly disturbed no forward scatter absorption occurred above Alaska (Thorne and Larsen, 1976). A local early morning and late evening pass by OV1-19 occurred on day 135. The local early morning pass did not sample the loss cone but even though the evening pass did, no REP was observed.

All three REPs observed occurred between local midnight and local dawn. The expected frequency of occurrence of REPs based on this study is 1-13% with a 90% confidence interval. The 50% confidence level indicates a frequency of 5% (see Fig. A-13). This is consistent with 5% frequency for moderate (>8 db absorption) REPs found by Bailey (1968).

The energy spectra for the three OV1-19 REPs are compared with the OV3-3 REP measured by Vampola (1971) in Fig. A-14. It is evident that they are comparable in hardness (approach to isoenergetic). So it is not surprising that the REPs will have a similarly predicted effect on the deposition of energy (represented by ion production) in the atmosphere as shown in Fig. A-15.

The conclusions of this study are that the OVI-19 was capable of observing strong diffusion events (REPs) from a few hundred keV to 1 keV energies. More REPs would probably have been observed if

the loss cones were sampled more frequently. A similar data set with better statistics would yield more morphological information and possibly a clue to the type and relative importance of various mechanisms for causing REPs.

FIG A-13. REP frequency of occurrence for the OV1-19 Data Base.

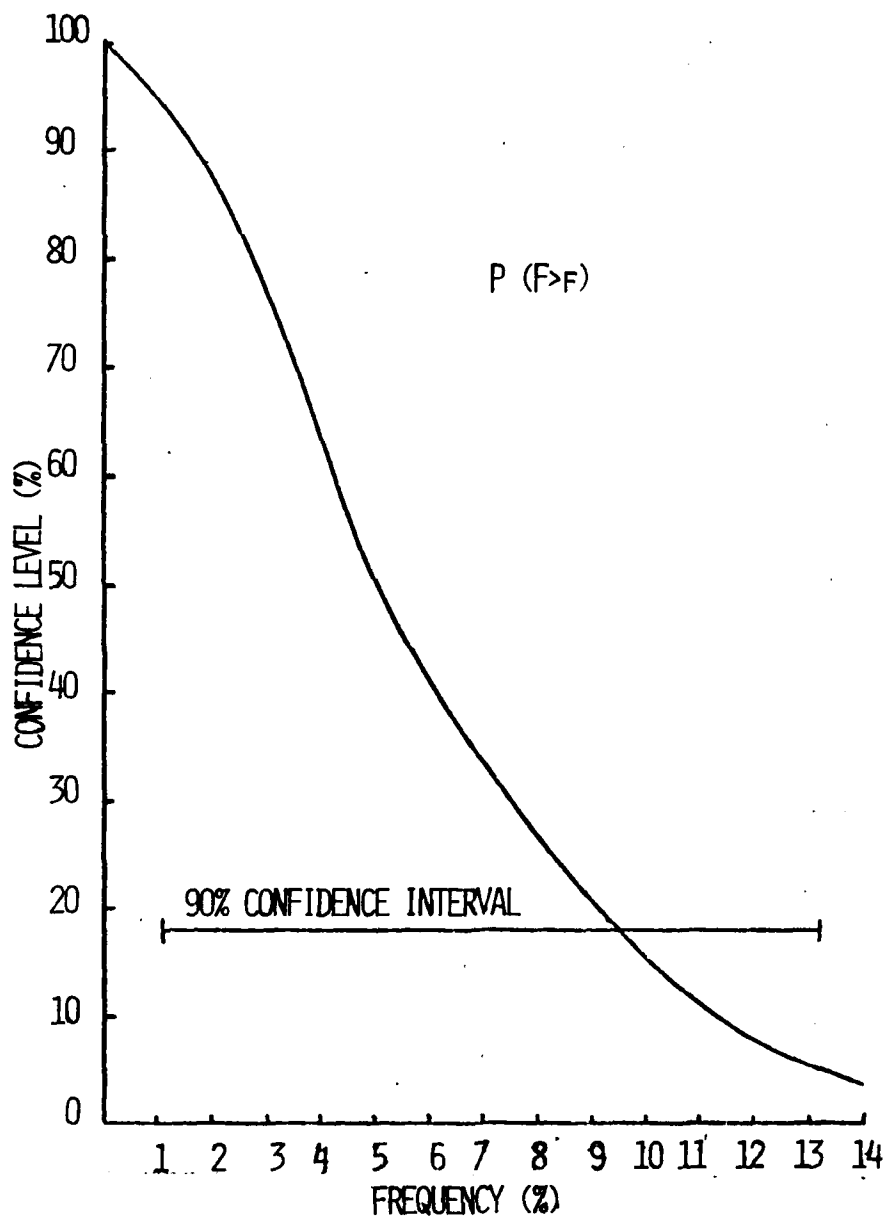


FIG A-14. REP energy spectra for the three OV1-19 events in 1969 compared with the Vampola OV3-3 spectrum of 1971.

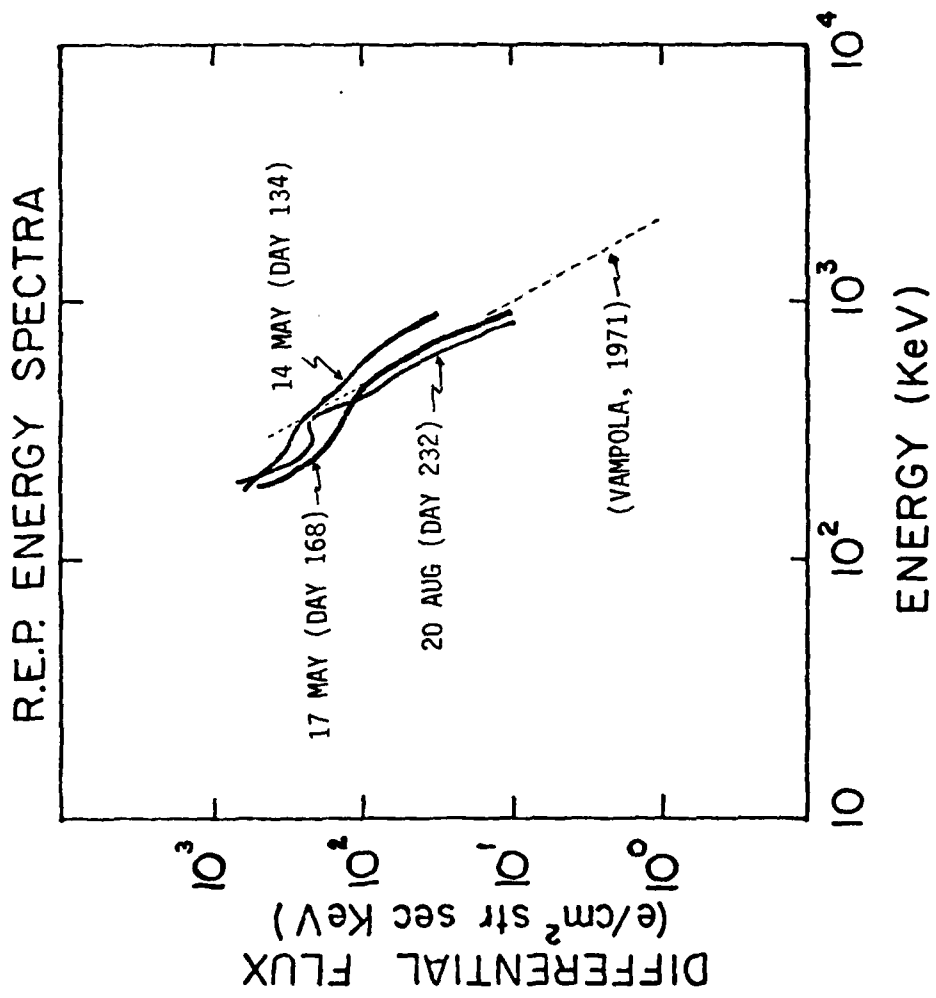
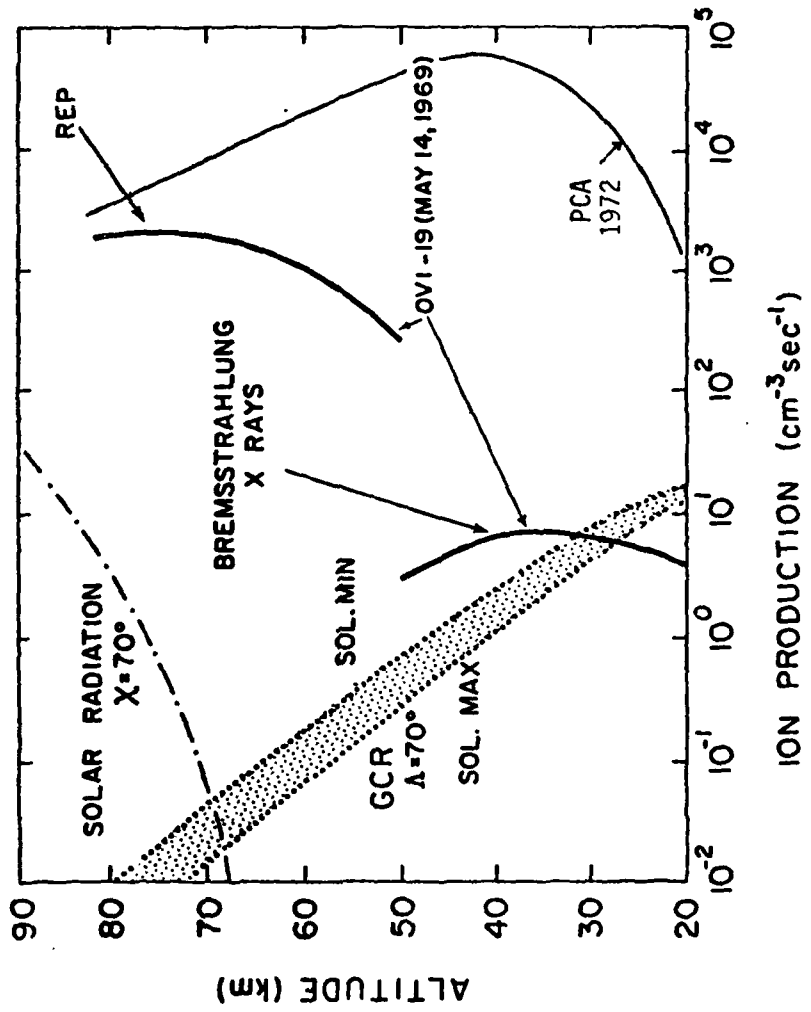


FIG A-15. Ion production rate versus altitude for the OV1-19 REP,
solar radiation, galactic cosmic rays and the solar proton
event of 1972.



Appendix B - S3-3 Data Collection

The primary data set used in this dissertation was collected by a magnetic spectrometer aboard Air Force Space Test Program satellite STP 74-2 designated S3-3 launched in the summer of 1976. The S3-3 orbit is highly elliptical (schematically shown in Fig. B-1) with apogee/perigee nominally 8000/250 km. The maximum apogee of 8035 km was reached on 3 August 1976 and continues to decrease at a rate of 0.35 km/day. Orbit inclination is 97° with a period of 180 minutes. The nodal drift rate is 0.27° /day westward, i.e., the local time of orbit nodal crossing decreases 0.049 hr/day from launch values of 10:30 (ascending node) and 22:30 (descending node) as shown in Fig. B-2, panel a. The line of apsides (major axis of orbit) was initially placed less than 90° from equator on the sun side (Fig. B-1) and slowly rotates toward the dark side at the rate of 0.94° /day from a launch value of 63° as shown in Fig. B-2, panel b. The satellite spin rate was maintained at approximately 3 rpm with the spin vector perpendicular to the orbit plane. In order to maintain the solar cells facing toward the sun (and the backside away for thermal control), the satellite was yawed 180° on days 40-48 (around 14 February) 1977.

The S3-3 scientific payloads are listed with brief descriptions in Table B-1. Of prime interest is the CRLS-217 which will be described here in detail. Data from ONR-104 and CRLS-218 are referred to in Chapter 2 and only brief details will be provided.

FIG B-1. S3-3 Orbital parameters.

APOGEE DRIFT - .94°/DAY
HIGHEST N. LATITUDE
~ 3 AUG '76

VELOCITY

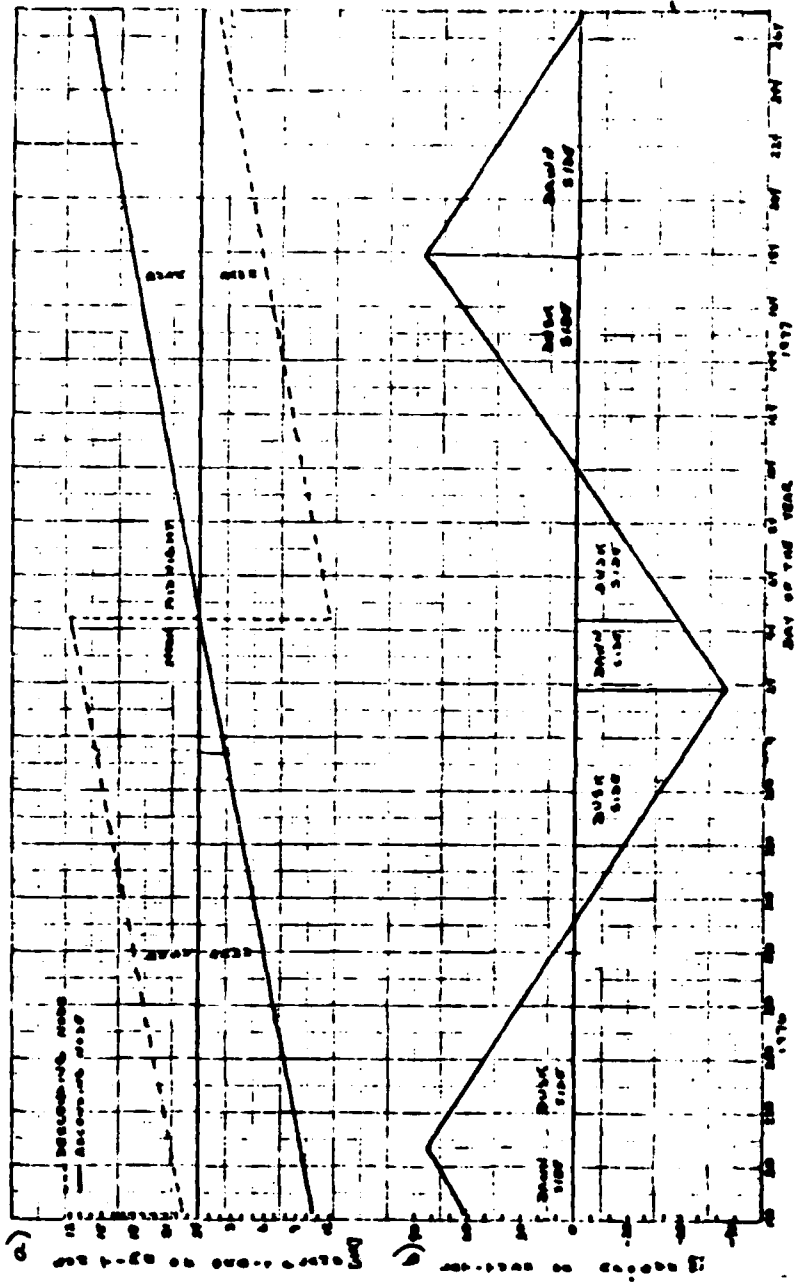


SPIN - 3 RPM
AXIS 1 TO
AVERAGE MAGNETIC
FIELD

INCLINATION - 97°
PERIOD - 180 MIN
APOGEE - 8035 KM (AUG 3, '76)

NODAL DRIFT RATE IS .27°/DAY
(I.E. LOCAL TIME OF ORBIT
NODAL CROSSING DECREASES
.049 HR/DAY)

FIG B-2. Day of the year versus a) sun time of the orbit phase
and b) latitude of apogee.



**TABLE B-1. S3-3 Scientific payloads with prime investigators
in parenthesis.**

CRL-214 Trapped Proton Monitoring (Yates) AFGL

Solid state detectors are used to measure the spatial distributions, angular distribution and energy spectrum of protons with energy between .1 and 100 MeV.

ONR-118 Low Energy Particle Spectrometer (Sharp) Lockheed

Ions and electrons with energy between 0.5 to 16 KeV and 0.07 to 24 KeV, respectively, are detected using channel multipliers. Particle energy is determined using either magnetic deflection or penetration of particles through thin foils.

ONR-104 Electric Field Measurements (Mozet) UC Berkeley

3 dipoles are used to measure the three components of the DC & AC electrostatic electric field.

CRL-215 Electric Fields and Ion Drifts (Wildman) AFGL

Plasma probes are used to detect the bulk flow of plasma and hence to infer the electric field strength. Four electrostatic analyzers are sensitive to ions in the energy range of 0.1 to 30 eV.

CRLS-217 Energetic Electron Environment (Vampola) Aerospace

A magnetic spectrometer with solid state detectors is used to measure the differential electron flux between 0.012 and 1.6 MeV, protons from 0.08 to 3 MeV, and alphas above 4 MeV.

CRLS-218 Magnetospheric Hydrogen-Electron Abundance (Fennell) Aerospace

An electrostatic analyzer is used to measure the fluxes of low energy electrons and ions at several energies; electrons from 0.17 to 8.4 KeV, ions from 0.09 + 3.9 KeV/Q.

CRLS-225 ELF/VLF Fields (Koons) Aerospace

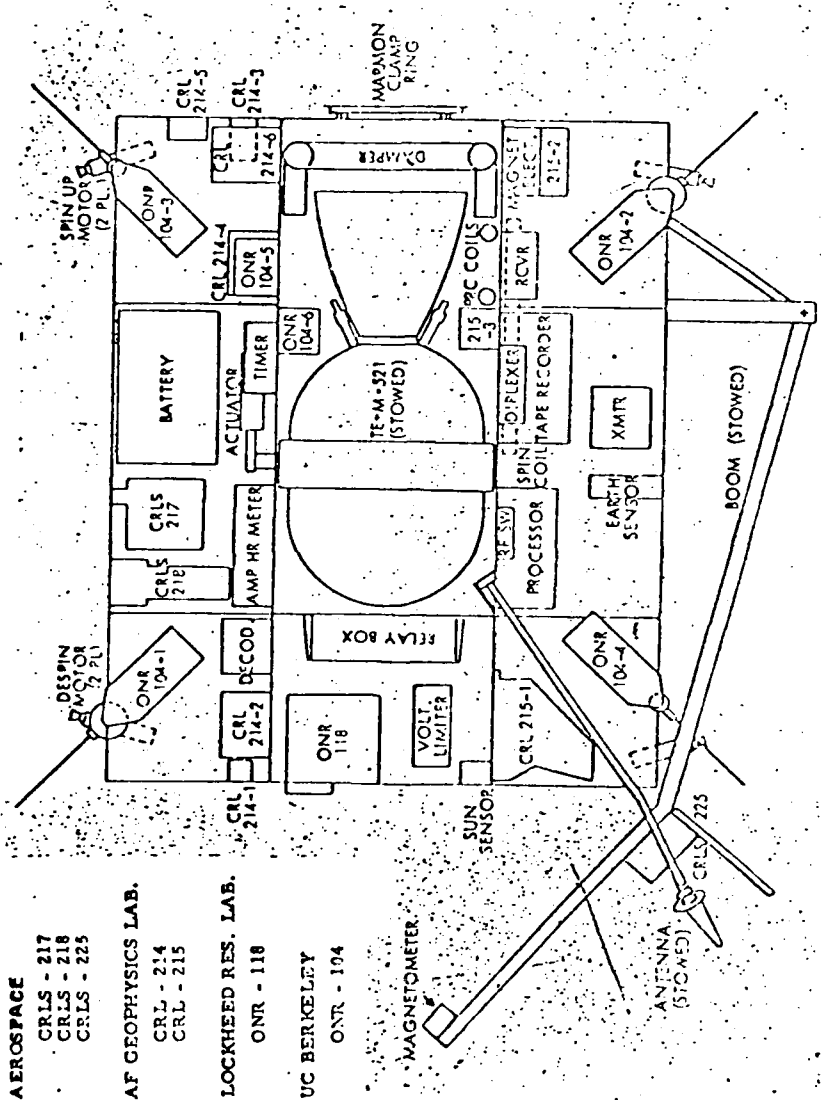
An air core magnetic loop and a dipole antenna are used to measure the intensity of magnetic and electric fields in the 0.1 to 20.0 KHz frequency range.

The satellite payload configuration is shown in Fig. B-3 with the spin vector perpendicular to the page. The CRLS-217 location is such that penetration thresholds are approximately 10 MeV for electrons and 100 MeV for protons.

CRLS-217 is a twelve differential energy channel magnetic electron spectrometer with an eight channel integral proton and alpha (1 channel) telescope. Permanent magnets (.53 and 1.64 kiloGauss) momentum analyze electrons entering the instrument aperture focusing them onto 12 silicon semiconductor detectors at the primary focus (180°). The instrument configuration is similar to the OV3-3 and OV1-19 instruments described in Appendix A except that the Low Energy Magnetic Spectrometer (LEMS) and High Energy Magnetic Spectrometer (HEMS) are built into one unit as shown in Fig. B-4. Pulseheight analysis of the energy deposits in the detectors is used to discriminate against bremsstrahlung, noise, and penetrating ion background. Also, an additional shielded detector is used to monitor the background. The electron output of each detector channel corresponds to the uni-directional electron intensity of a given energy bin (typically ± 15 keV for LEMS, ± 100 keV for HEMS) and from these outputs an energy spectrum can be obtained.

The data format is presented in Table B-2. Included are the identification code, channel, sample rate, bits/sample, and geometric energy factor for 28 CRLS-217 and 2 (21, 22) CRLS-218 measurements. Measurements 1-5 are the 5 LEMS channels while 6-12 are the HEMS, with all channel accumulators read out every 1/16 second. Note that

FIG B-3. S3-3 Payload Configuration.



- AEROSPACE**
 - CRLS - 217
 - CRLS - 218
 - CRLS - 225
- AF GEOPHYSICS LAB.**
 - CRL - 214
 - CRL - 215
- LOCKHEED RES. LAB.**
 - ONR - 118
- UC BERKELEY**
 - ONR - 104

FIG B-4. S3-3 HEMS/LEMS Layout.

TABLE B-2. S3-3 HEMS/LEMS Data Format.

MEAS.	ID CODE	CHANNEL	SAMPLES/SEC	BITS/SAMPLE	GEOMETRIC - ENERGY FACTOR*
1	ME1	E _c ~ 12 KeV	16	12**	0.0088 cm ² ster KeV
2	ME2	E _c ~ 33 KeV			0.0029
3	ME3	E _c ~ 70 KeV			0.0141
4	ME4	E _c ~ 110 KeV			0.0036
5	ME5	E _c ~ 160 KeV			0.0147
6	ME6	E _c ~ 235 KeV			6.54
7	ME7	E _c ~ 435 KeV			6.52
8	ME8	E _c ~ 655 KeV			6.12
9	ME9	E _c ~ 880 KeV			5.17
10	ME10	E _c ~ 1.11 MeV			4.57
11	ME11	E _c ~ 1.35 MeV			4.04
12	ME12	E _c ~ 1.60 MeV	16		0.012 cm ² ster
13	MPI -S	E _p > 50 KeV			
14	MP2 -S	E _p > 100 KeV	4		0.012 cm ² ster
15	MP3	3.2 MeV/E _p > 80 KeV			
16	MP4	" > E _p > 150 KeV			
17	MP5	" > E _p > 350 KeV			
18	MP6	" > E _p > 770 KeV			
19	MP7	" > E _p > 1.55 MeV			
20	MP8	E _d > 4.0 MeV - 13.2 MeV			
21	S	ESA STEP } 60x5-249			
22	L	ESA STEP }			
23	E01	E _c BKG 1	4		
24	E02	E _c BKG 2			
25	ES1	E _c ~ 12 KeV			
26	ES2	E _c ~ 70 KeV			
27	ES3	E _c ~ 235 KeV			
28	ES4	E _c ~ 880 KeV			
29	EPI	E _p > 80 KeV			
30	EPI	E _p > 4.0 MeV			

* Assumed spectrum is E⁰ with 100% efficiency.

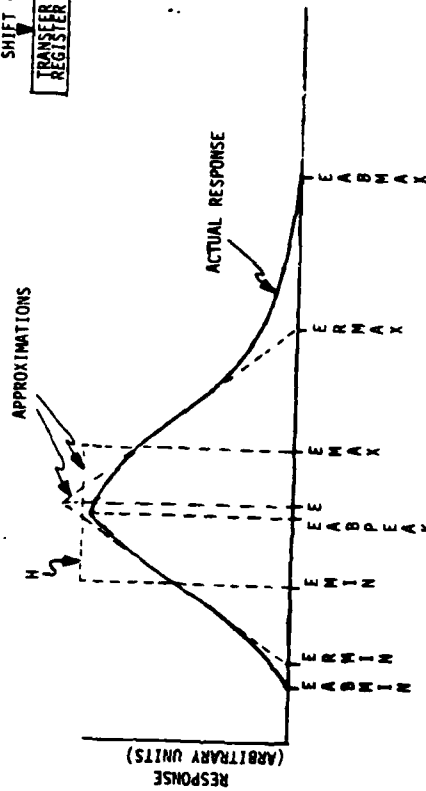
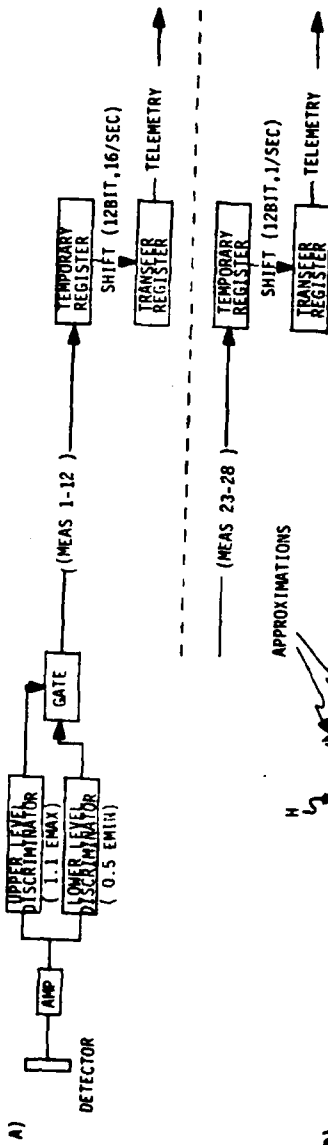
** 2¹ bit accumulator formatted into floating point 12 bit output with 0.2% resolution.

the geometric energy factor decreases with energy for the LEMS and HEMS, i.e., sensitivity decreases with increase in energy (given flux will produce less counts/sec for a reduced geometric energy factor). In addition, S3-3 HEMS channels are more than a factor of 2.5 higher in geometric energy factor than OV1-19 LEMS channels to ~450 keV (see Table A-1) but more than a factor of 5.8 lower than OV1-19 HEMS channels to ~1.1 MEV (see Table A-2). Measurements 13-19 and 20 show specs for the integral proton and alpha channels which were also sampled at a rate of 1/sec.

A functional block diagram of the instrument is shown in Fig. B-5a. The lower level discriminator is set at 50% of the minimum expected energy deposit to reject bremsstrahlung and noise and the upper level discriminator is set at an upper threshold of 110% of the maximum expected energy to reject penetrating particles (cosmic rays and energetic protons).

Each electron energy channel is calibrated to find the channel response curve from which the absolute peak (EAPEAK), minimum (EAMIN), and maximum (EAMAX) values are determined (see Fig. B-5 b). A triangle which best matches the slopes of the response curve then yields its peak (E), minimum (ERMIN), and maximum (ERMAX) values. Center values of the two triangle slopes determine EMIN and EMAX. Vertical lines are extended through EMIN and EMAX. The vertical extension of these two lines continues until the resulting rectangle area formed with line segment H equals the area under the actual response curve. This representative rectangle is then used as the

FIG B-5. a) Electron Channel component schematic; b) general response curve; c) channel values for E^0 with 100% efficiency.



c)

LENS	EMAX	EMIN	ERMAX	ERMIX	EARMAX	EARMIN	E
LENS	11.00	10.21	17.26	16.50	17.19	16.79	1.106
	11.77	10.43	17.79	16.79	17.79	16.79	1.057
	12.54	10.65	18.32	17.58	18.32	17.58	1.008
	13.31	10.87	18.85	18.37	19.06	18.37	0.959
	14.08	11.09	19.38	19.16	19.75	19.16	0.910
	14.85	11.31	19.91	19.95	20.44	19.95	0.861
	15.62	11.53	20.44	20.74	21.13	20.74	0.812
	16.39	11.75	20.97	21.53	21.82	21.53	0.763
	17.16	11.97	21.50	22.32	22.51	22.32	0.714
	17.93	12.19	22.03	23.11	23.20	23.11	0.665
	18.70	12.41	22.56	23.90	23.89	23.89	0.616
	19.47	12.63	23.09	24.69	24.58	24.58	0.567
	20.24	12.85	23.62	25.48	25.27	25.27	0.518
	21.01	13.07	24.15	26.27	25.96	25.96	0.469
	21.78	13.29	24.68	27.06	26.65	26.65	0.420
	22.55	13.51	25.21	27.85	27.34	27.34	0.371
	23.32	13.73	25.74	28.64	28.03	28.03	0.322
	24.09	13.95	26.27	29.43	28.72	28.72	0.273
	24.86	14.17	26.80	30.22	29.41	29.41	0.224
	25.63	14.39	27.33	31.01	30.10	30.10	0.175
	26.40	14.61	27.86	31.80	30.79	30.79	0.126
	27.17	14.83	28.39	32.59	31.48	31.48	0.077
	27.94	15.05	28.92	33.38	32.17	32.17	0.028

channel response for data analysis. (This is identical to the description in Appendix A and depiction in Fig. A-2 for instruments OV3-3 or OV1-19.) Values for the above parameters for an assumed spectrum of E^0 with 100% efficiency are listed in Fig. B-5 c. A typical three dimensional response is shown in Fig. B-6 for the 435 KEV channel.

The instrument loss cone response for 655 keV channel is shown in Figs. B-7 a-c as the result of an input distribution isotropic everywhere outside the loss cone and falling off about an order of magnitude per 2^0 in the loss cone for pitch angles of 8^0 , 18^0 , 28^0 , respectively. Response to a sharp cutoff can be demonstrated by a zero input distribution at the loss cone but is not represented here. Except for pitch angles less than $\sim 10^0$, an order of magnitude drop in response is about 9^0 into the loss cone. This response obviously affects the subjective selection of isotropic distributions in the loss cone and is further discussed in Appendix C.

The spin rate of ~ 3 RPM allows the instrument to complete a pitch angle plot ($0-180^0$) in approximately 10 seconds (compared to ~ 4 seconds for OV1-19). At 16 samples/sec, the instrument will collect ~ 150 samples for a pitch angle plot, nearly a factor of 40 better than the OV1-19. The S3-3 advantage is clear when the two data sets are compared for analysis.

The proton/alpha telescope consists of two silicon detectors in series as shown in Fig. B-8. Detector D1 ($100 \mu\text{m}$ thick) has a

FIG B-6. Typical three dimensional response curve for HEMS 435 keV
electron channel.

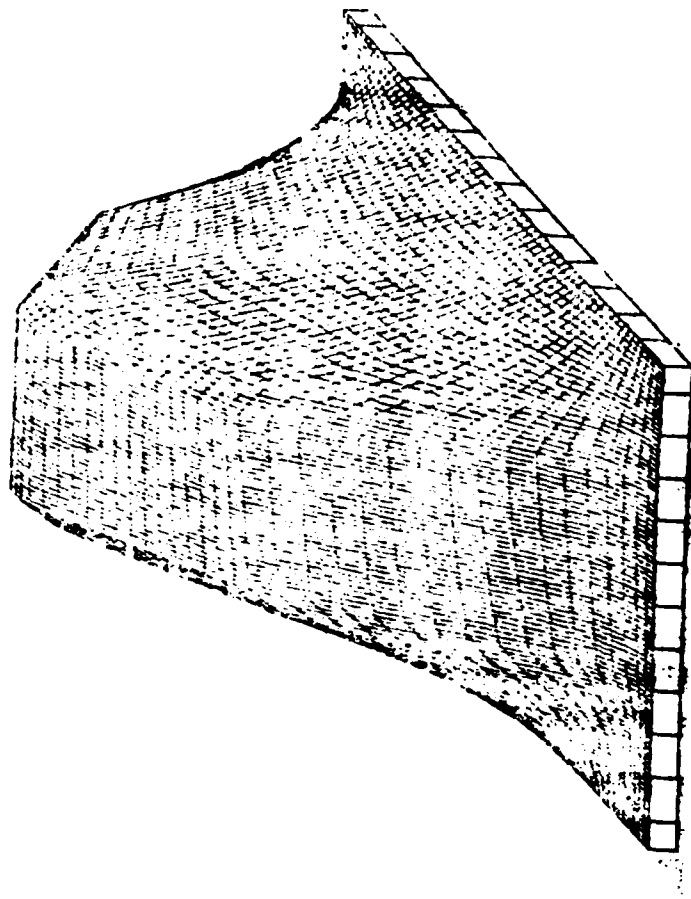
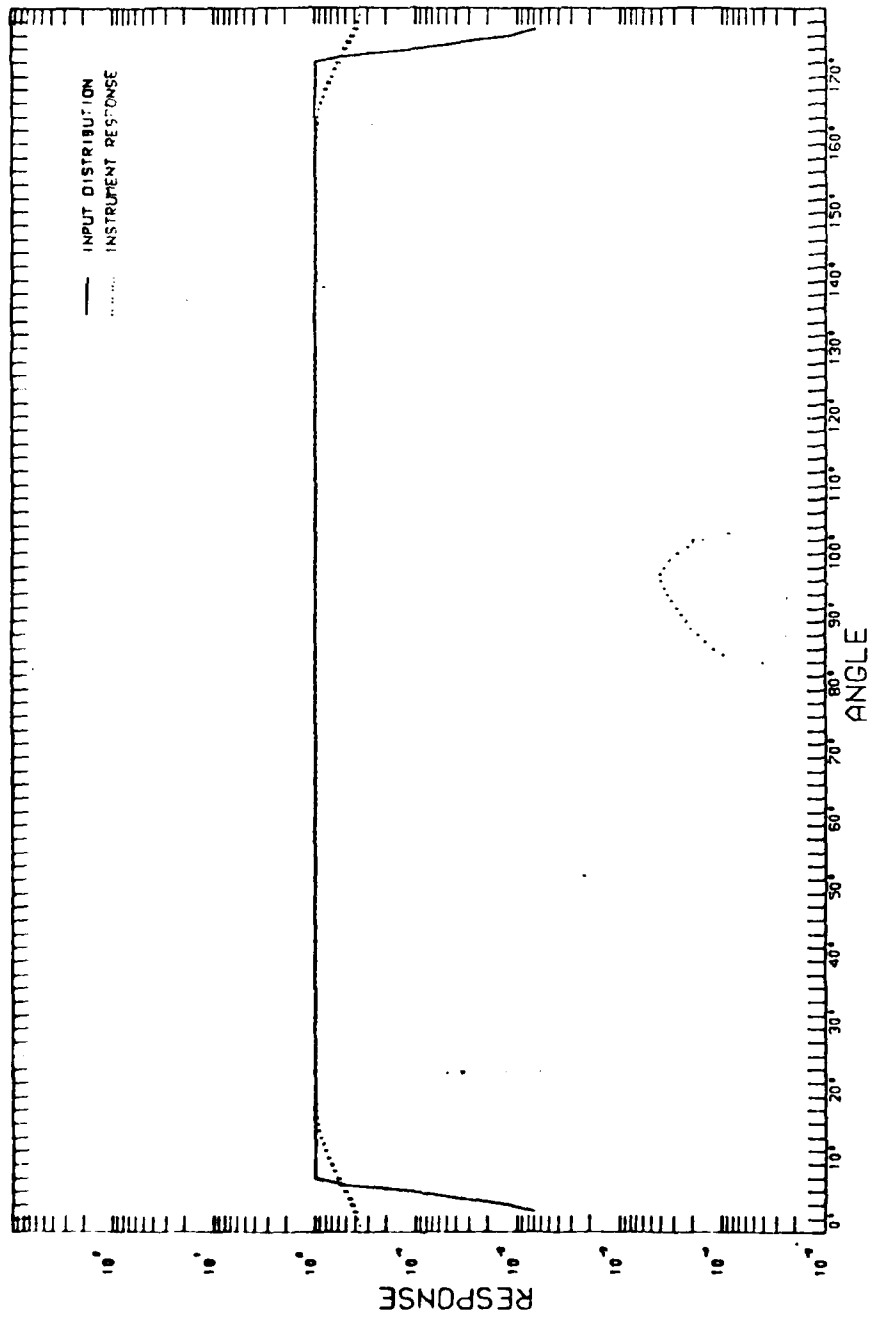
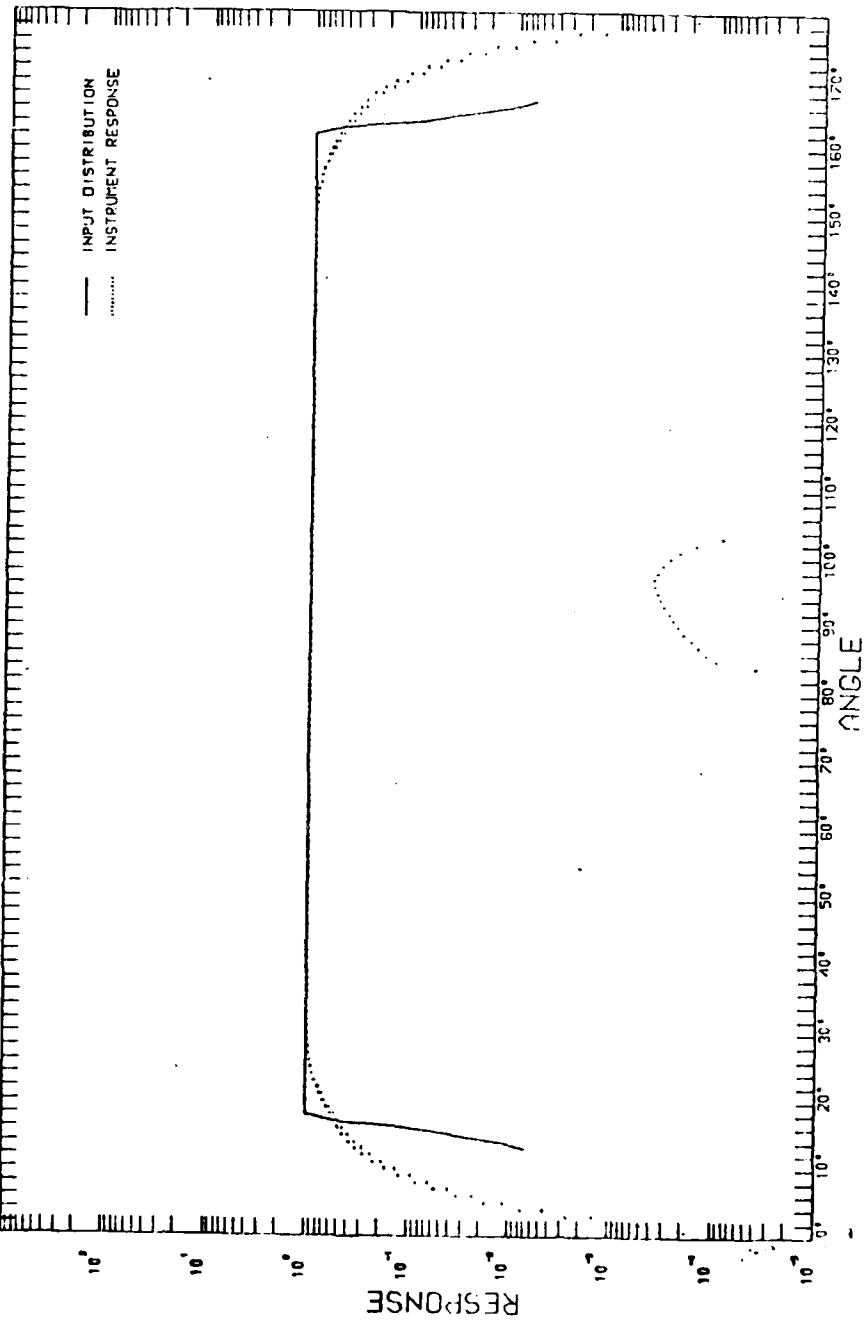


FIG B-7. Typical instrument loss cone response for the 655 keV electron channel for pitch angles 8° , 18° , and 28° .

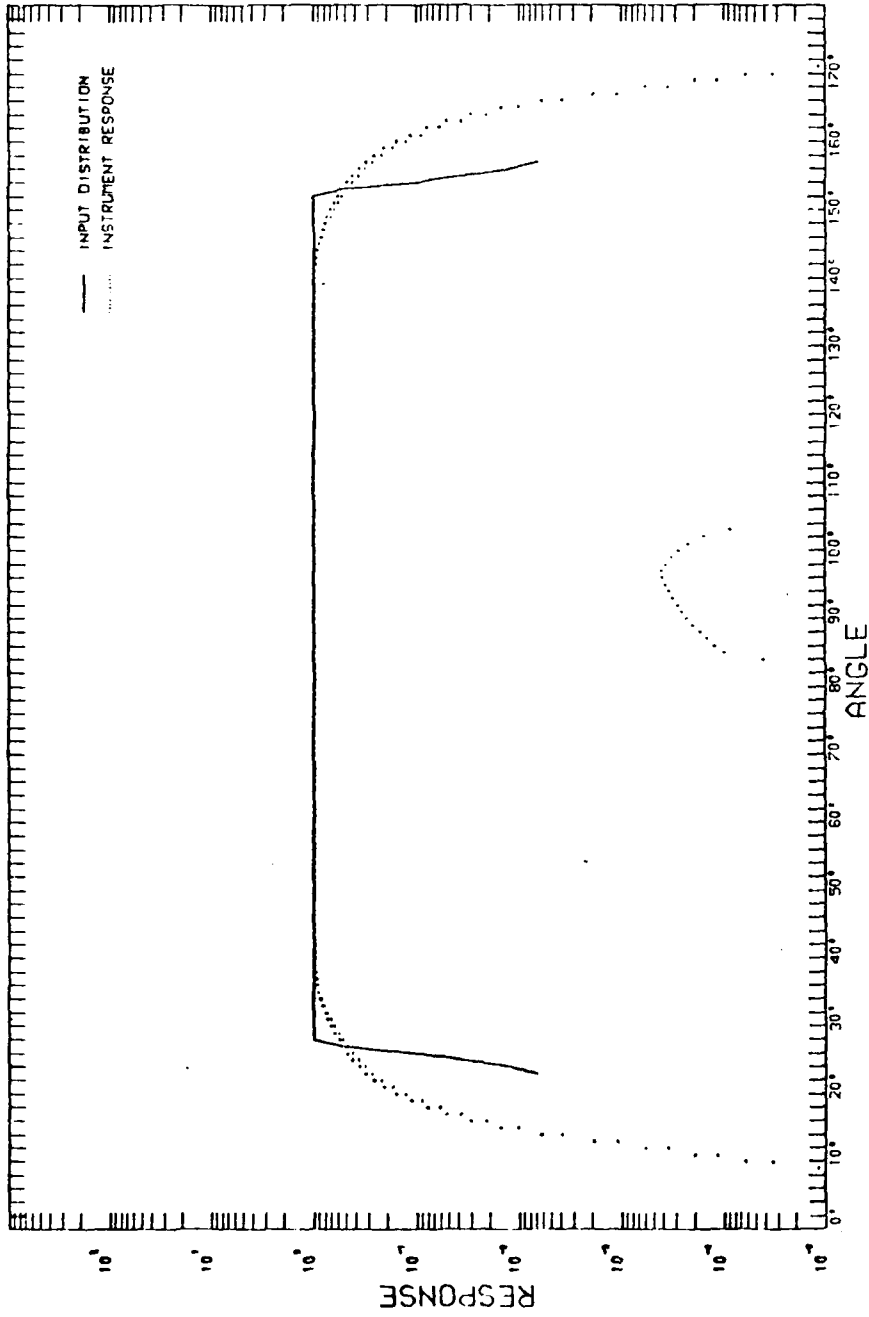
S3-3 655 KEV RESPONSES



S3-3 655 KEV RESPONSES



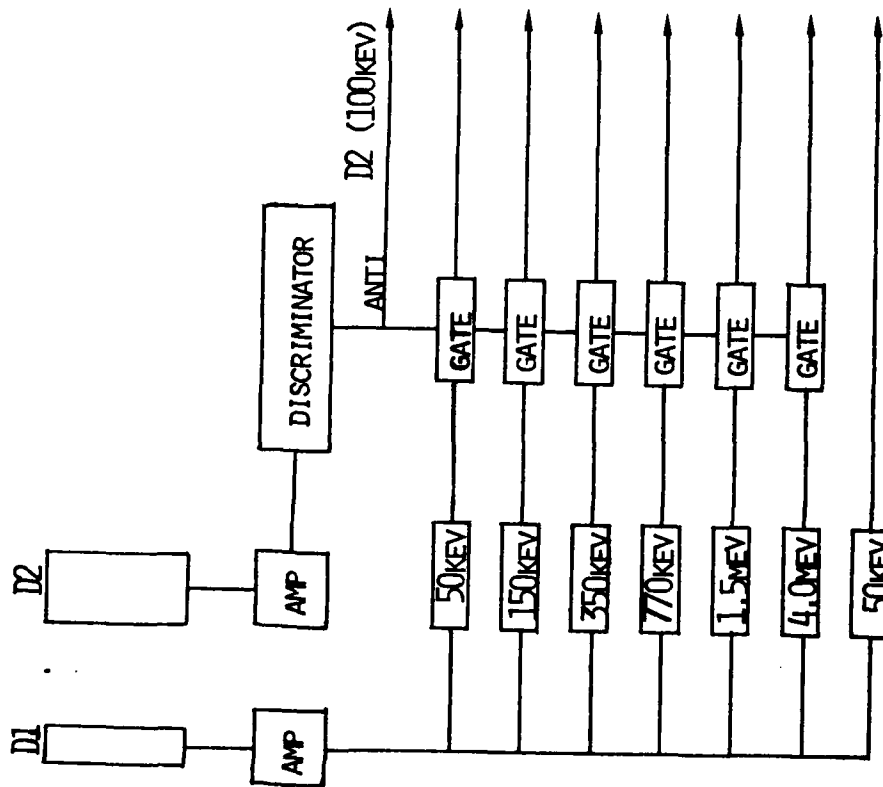
S3-3 655 KEV RESPONSES



lower threshold of 50 keV and detector D2 (200 um thick) has a lower threshold of 100 keV and is the anti to reject energetic protons and cosmic rays penetrating through D1. Particles depositing more than about 4 MeV are assumed to be alphas since $p^+ > 3.2$ MeV would pass through D1 and be antied out by D2. All channels are sampled at 4 samples/sec and $E_p > 80$ keV and $E_\alpha > 4.0$ MeV are also sampled at 1/sec.

The S3-3 wave experiment by the Berkeley group (Mozer et al., 1977) obtained the three orthogonal components of the static or fluctuating electric field by measuring the potential differences between three orthogonal pairs of separated spheres. Two pairs of spheres were separated by 37 m in the satellite spin plane and one pair by 6 m along the spin axis. One sphere on the long boom could be operated as a Langmuir probe to make relative density fluctuations. Broadband signals (.05 - 16 kHz) were transmitted in real time to ground stations for 15 min periods on about 25% of the orbits (Kintner et al., 1978). The percentage of concurrent real time wave data and stored plasma data (CRLS 217) is much less. Only filtered wave data in bands from .03-100 kHz are available for most of the stored plasma data presented here.

FIG B-8. S3-3 Proton/alpha telescope data schematic.



TO DATA BOX AND
TELEMETRY

Appendix C - REP Event Identification and Catalog Use

In order to select REP events, it was necessary to be able to identify regions of isotropic fluxes on microfiche time plots of the electron and proton fluxes indicative of strong diffusion. Fig. C-1 shows such a plot for the 12, 235, 435, 655 and 850 keV electron and 80 keV proton channels for rev 1661 on a logarithmic scale; UT is measured in sec. The microfiche displays only every fourth data point. The variation in fluxes are due to pitch angle modulation. The key feature one must recognize is the missing upward viewing loss cone indicative of a full or isotropic loss cone. This feature is readily evident throughout the 12 keV electron channel from $L = 4.0 - 8.0$ and the 80 keV proton channel from $L = 4.5 - 6.0$. The 235, 435, 655, and 850 keV channels exhibit missing loss cones at $L \sim 4.5$ with the rest of the outer zone mostly eroded. Effects such as data drops and low altitude diminished flux values similarly appear as missing loss cones so it is extremely important to follow up with examination of the actual pitch-angle plot. Fig. C-2 shows pitch-angle plots for the 235, 435, 655, and 850 keV channels of the example in Fig. C-1 for UT 61092 and 61109 secs. The magnetic field direction is zero in the southward direction at the equator so for this Southern Hemisphere example, 180° is the up-the-field-line direction. The pitch angle is approximately 20° . Also, every data point is displayed and differential flux values can be obtained by dividing by the geometric value per channel (Table B-2). The top panel shows normal loss cone distributions both up and down the field line indicative of expected

radiation belt distributions. (However, the belts are disturbed since normally quiet times produce loss cones 3 1/2 orders of magnitude down from trapped flux). The bottom panel (data toward the outer edge of the outer zone) shows the pitch angle 17 secs later. Although normal distributions are evident down the field line, viewing up the field line shows isotropic distributions in all channels. Similar plots can be obtained for the proton channels.

Since the satellite spin axis is not always directly perpendicular to the local magnetic field line and the instrument has a field of view of approximately 5° (half angle), we chose only those cases with data at least 10° into the local loss cone. A few exceptions were made when the excursion into the loss cone was greater than 6° and isotropic fluxes were measured on previous and succeeding revs. Less than 3% of the 325 events selected for further pitch-angle analysis were rejected for the 'less than 10° in the loss cone' criteria.

Apogee/perigee bias in the flux data was investigated and found to be not important except in the actual appearance of the flux versus time distributions. At apogee the loss cones are small and the counting rates are high. Here we merely insisted on the ' 10° into the loss cone' criteria for validation. At perigee the loss cones are large, the extent (in time) of the entire outer zone is narrow, and the flux counting rates of both trapped and precipitated particles is small. The learning process to identify isotropic fluxes on the microfiche plots was accomplished by a series of trial and error analyses taking about 3 months of time and 150 samples. A high

Fig. C-1. Logarithmic flux (counts/sec) versus UT (sec) for the 12, 235, 435, 655, 850 keV electron and 80 keV proton channel for rev 1661.

CRLS 217

ALT	5040.00	6237.00	6511.01	6762.00	6920.13	7105.10
LONG	113.23	111.00	110.01	108.00	106.44	105.00
LAT	04.50	04.00	03.27	02.07	01.33	01.00

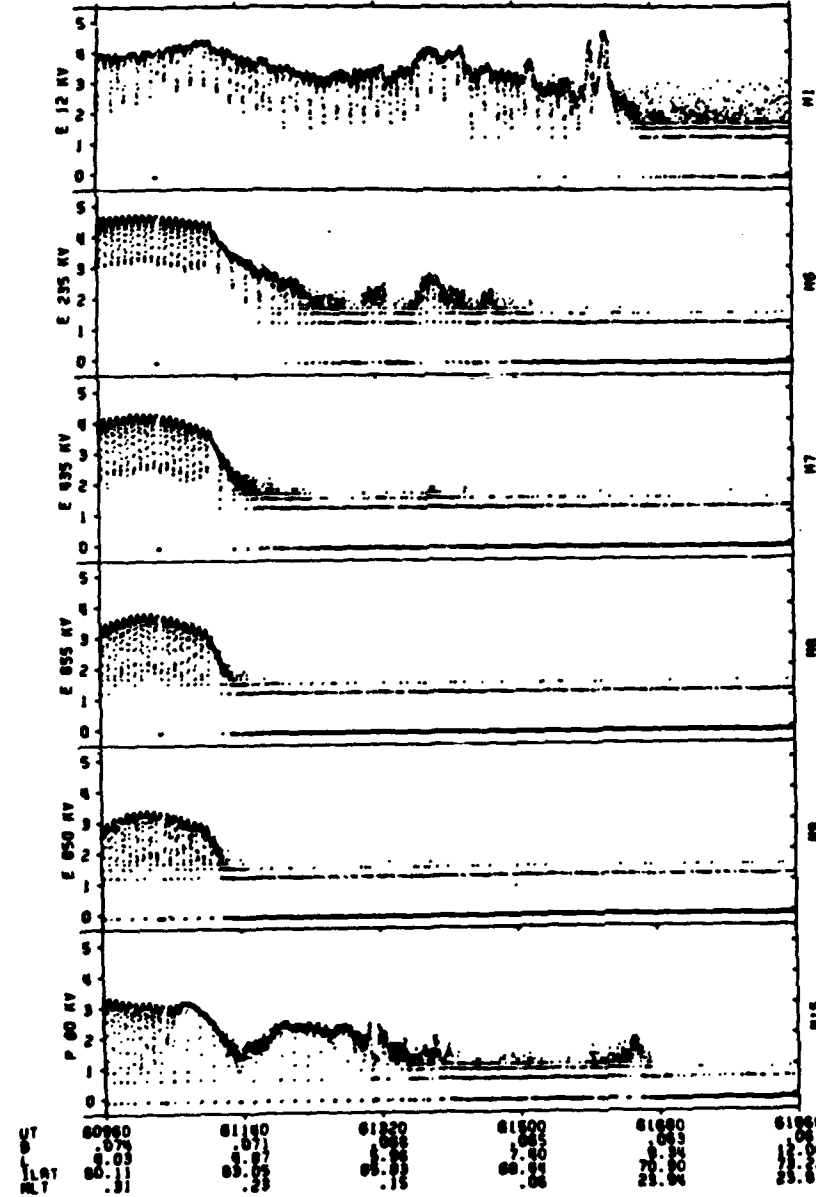


Fig. C-2. Flux (counts/sec) versus pitch angle (degrees) for rev 1661 at 61092 UT (top) and 61109 UT (bottom) for the 235, 435, 655, 880 keV channels. Loss cone is $\sim 20^\circ$ with the upward looking loss cone at 160° indicated by a vertical line.

REV 1661 30 JAN 1977 P.A. ~20°

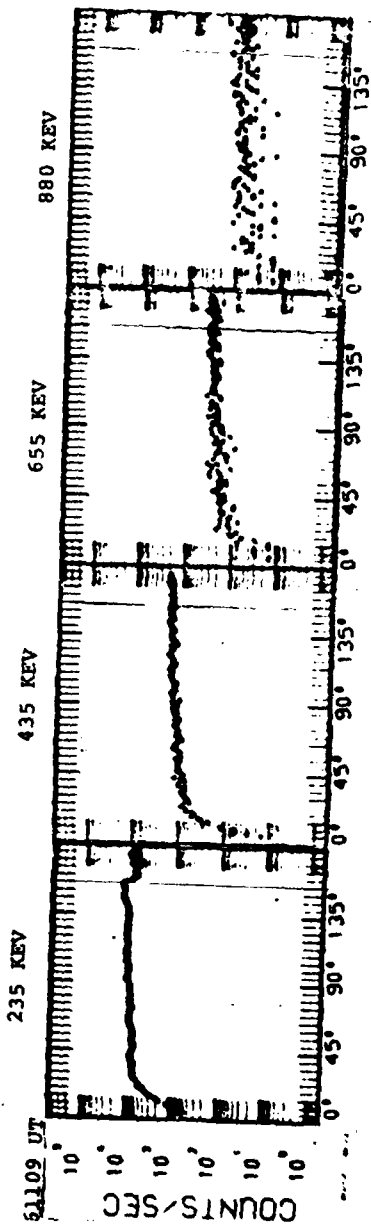
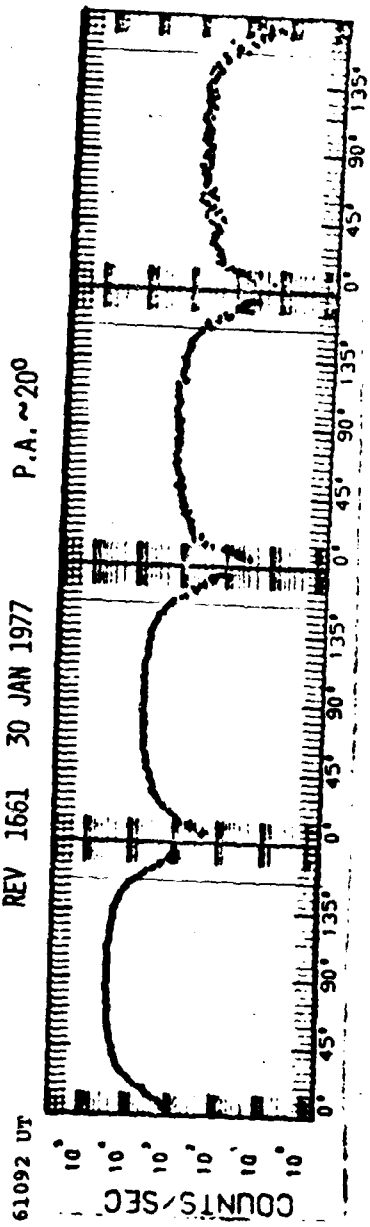


Fig. C-3. Microfiche layout with each logarithm of flux (counts/sec) as ordinate for each channel and abscissa the UT (sec) every 180 sec for 900 sec (16 samples/sec), ALT [km], LONG [deg-E], LAT [deg], B [Gauss], L [], ILAT [deg], MLT [hr]. E.G. Eight panels (15 min each) of 12, 33, 70 electrons, 8 panels 110, 160, 235 keV electrons, etc. P = Protons, A = Alphas, PD = Proton Detector, BG = Background.

16 SAMPLES/SEC								1 SAMPLE/SEC	
12 } KEV 33 } 70 }	15 MIN ▶							861 } 862 }	1 HR ▶
"	110 } KEV 160 } 235 }							"	"
"		435 } KEV 635 } 850 }						"	12 } KEV 70 } 235 }
"				1.11 } MEV 1.35 } 1.60 }				"	"
"					P17 } P18 } P80 KEV			"	850 } KEV 880 } 941.0 MEV
"						P150 } KEV P350 } P770 }		"	"
"								"	"
"								81.55 } MEV 94:00 }	"

Table C-1. Listing of individual microfiche by Readout REV NO
(Station and Number), DATE of REV (Last Digit of Yr,
Month, Day), Beginning and Ending TIME OF REV IN
HRS (Hours, Minutes, Seconds), Beginning and Ending
TIME OF REV IN SECS (Seconds), Precip Category
(1 \leq 235 keV, 2 \leq 435 keV, 3 \leq 655 keV, 4 \leq 850 keV),
COMMENTS (12 = 12 keV electron channel, etc.; 80,
150 = 80, 150 keV Proton Channel, etc; L = L value).

REV NO.	DATE	TYPE OF REV	TIME OF REV	IN MRS	TIME OF REV	IN SECS
000-0001	60711	100000	02147	00000		
MUL	0002	60712	00001	00000	00001	00000
MUL	0003	60713	00002	00000	00002	00000
MUL	0004	60714	00003	00000	00003	00000
MUL	0005	60715	00004	00000	00004	00000
MUL	0006	60716	00005	00000	00005	00000
MUL	0007	60717	00006	00000	00006	00000
MUL	0008	60718	00007	00000	00007	00000
MUL	0009	60719	00008	00000	00008	00000
MUL	0010	60720	00009	00000	00009	00000
MUL	0011	60721	00010	00000	00010	00000
MUL	0012	60722	00011	00000	00011	00000
MUL	0013	60723	00012	00000	00012	00000
MUL	0014	60724	00013	00000	00013	00000
MUL	0015	60725	00014	00000	00014	00000
MUL	0016	60726	00015	00000	00015	00000
MUL	0017	60727	00016	00000	00016	00000
MUL	0018	60728	00017	00000	00017	00000
MUL	0019	60729	00018	00000	00018	00000
MUL	0020	60730	00019	00000	00019	00000
MUL	0021	60731	00020	00000	00020	00000
MUL	0022	60732	00021	00000	00021	00000
MUL	0023	60733	00022	00000	00022	00000
MUL	0024	60734	00023	00000	00023	00000
MUL	0025	60735	00024	00000	00024	00000
MUL	0026	60736	00025	00000	00025	00000
MUL	0027	60737	00026	00000	00026	00000
MUL	0028	60738	00027	00000	00027	00000
MUL	0029	60739	00028	00000	00028	00000
MUL	0030	60740	00029	00000	00029	00000
MUL	0031	60741	00030	00000	00030	00000
MUL	0032	60742	00031	00000	00031	00000
MUL	0033	60743	00032	00000	00032	00000
MUL	0034	60744	00033	00000	00033	00000
MUL	0035	60745	00034	00000	00034	00000
MUL	0036	60746	00035	00000	00035	00000
MUL	0037	60747	00036	00000	00036	00000
MUL	0038	60748	00037	00000	00037	00000
MUL	0039	60749	00038	00000	00038	00000
MUL	0040	60750	00039	00000	00039	00000
MUL	0041	60751	00040	00000	00040	00000
MUL	0042	60752	00041	00000	00041	00000
MUL	0043	60753	00042	00000	00042	00000
MUL	0044	60754	00043	00000	00043	00000
MUL	0045	60755	00044	00000	00044	00000
MUL	0046	60756	00045	00000	00045	00000
MUL	0047	60757	00046	00000	00046	00000
MUL	0048	60758	00047	00000	00047	00000
MUL	0049	60759	00048	00000	00048	00000
MUL	0050	60760	00049	00000	00049	00000
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MUL	0052	60762	00051	00000	00051	00000
MUL	0053	60763	00052	00000	00052	00000
MUL	0054	60764	00053	00000	00053	00000
MUL	0055	60765	00054	00000	00054	00000
MUL	0056	60766	00055	00000	00055	00000
MUL	0057	60767	00056	00000	00056	00000
MUL	0058	60768	00057	00000	00057	00000
MUL	0059	60769	00058	00000	00058	00000
MUL	0060	60770	00059	00000	00059	00000
MUL	0061	60771	00060	00000	00060	00000
MUL	0062	60772	00061	00000	00061	00000
MUL	0063	60773	00062	00000	00062	00000
MUL	0064	60774	00063	00000	00063	00000
MUL	0065	60775	00064	00000	00064	00000
MUL	0066	60776	00065	00000	00065	00000
MUL	0067	60777	00066	00000	00066	00000
MUL	0068	60778	00067	00000	00067	00000
MUL	0069	60779	00068	00000	00068	00000
MUL	0070	60780	00069	00000	00069	00000
MUL	0071	60781	00070	00000	00070	00000
MUL	0072	60782	00071	00000	00071	00000
MUL	0073	60783	00072	00000	00072	00000
MUL	0074	60784	00073	00000	00073	00000
MUL	0075	60785	00074	00000	00074	00000
MUL	0076	60786	00075	00000	00075	00000
MUL	0077	60787	00076	00000	00076	00000
MUL	0078	60788	00077	00000	00077	00000
MUL	0079	60789	00078	00000	00078	00000
MUL	0080	60790	00079	00000	00079	00000
MUL	0081	60791	00080	00000	00080	00000
MUL	0082	60792	00081	00000	00081	00000
MUL	0083	60793	00082	00000	00082	00000
MUL	0084	60794	00083	00000	00083	00000
MUL	0085	60795	00084	00000	00084	00000
MUL	0086	60796	00085	00000	00085	00000
MUL	0087	60797	00086	00000	00086	00000
MUL	0088	60798	00087	00000	00087	00000
MUL	0089	60799	00088	00000	00088	00000
MUL	0090	60800	00089	00000	00089	00000
MUL	0091	60801	00090	00000	00090	00000
MUL	0092	60802	00091	00000	00091	00000
MUL	0093	60803	00092	00000	00092	00000
MUL	0094	60804	00093	00000	00093	00000
MUL	0095	60805	00094	00000	00094	00000
MUL	0096	60806	00095	00000	00095	00000
MUL	0097	60807	00096	00000	00096	00000
MUL	0098	60808	00097	00000	00097	00000
MUL	0099	60809	00098	00000	00098	00000
MUL	0100	60810	00099	00000	00099	00000

1 - only 635
2 - 235, 435
3 - 235, 435, 655
4 - 235, 435, 655, 880 on mm

COMMENTS

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98. 235° pump L-5-6
99. 235° pump L-5-6
100. 235° pump L-5-6

REV NO.	DATE	TIME OF REV	TIME OF REV
OF REV		IN HRS	IN SECS
BOS-0270-60010	01434	34200	0100
BOS-0270-60010	01434	34200	0100
BOS-0270-60012	21636	34700	0196
MIL-0277-60012	34262	01500	13643
MIL-0277-60012	34262	01500	13643
MIL-0277-60012	34725	01500	13645
POG-0282-60012	141212	153712	01132
BLA-0281-60012	153729	171046	06249
COO-0288-60013	110430	122935	39670
COO-0289-60013	110430	122935	39670
POG-0288-60013	122950	140307	04990
POG-0288-60013	122950	140307	04990
COO-0290-60014	75656	00030	00612
COO-0290-60014	75656	00030	00612
COO-0290-60014	75656	00031	00616
MIL-0295-60014	90848	105520	32920
MIL-0295-60014	90848	105520	32920
MIL-0305-60015	44905	40721	17345
MIL-0305-60015	44905	40721	17345
COO-0304-60015	40730	74741	22050
COO-0304-60015	40730	74741	22050
COO-0311-60016	14116	30600	6076
COO-0311-60016	14116	30600	6076
COO-0309-60016	30616	63940	11176
COO-0310-60016	223340	235120	01220
COO-0317-60016	235136	13200	05496
BOS-0326-60017	192537	205044	69937
COO-0323-60017	205109	222408	75069
COO-0330-60018	161745	173559	50665
COO-0329-60018	173614	191617	63376
BOS-0346-60020	195952	174452	07592
MIL-0345-60020	172500	185824	62708
COO-0358-60021	126150	141236	46324
COO-0355-60021	125204	141236	46324
COO-0353-60021	141252	155034	51172
COO-0353-60021	141252	155034	51172
POG-0362-60022	06412	110032	33053
POG-0362-60022	06412	110032	33053
COO-0360-60022	110000	124203	00162
COO-0360-60022	110000	124203	00162
POG-0360-60023	43423	74754	23783
POG-0360-60023	43423	74754	23783
COO-0368-60023	74800	02462	00400
COO-0368-60023	74815	93454	28095
COO-0371-60023	194244	204038	70964
POG-0370-60023	204054	222044	74654
MIL-0374-60024	32851	43529	12811
MIL-0373-60024	43346	62703	16426
MIL-0373-60024	43346	62703	16426
MIL-0373-60024	43346	62703	16426
MIL-0373-60024	43346	62703	16426
MIL-0373-60024	43346	62703	16426
MIL-0373-60024	43346	62703	16426
MIL-0373-60024	43346	62703	16426
COO-0377-60024	122513	133640	44713
COO-0377-60024	122513	133640	44713
COO-0376-60024	133713	152343	49033
MIL-0381-60024	121254	224701	76914
MIL-0381-60024	121254	224701	76914
COO-0379-60024	224722	2021	02042
MIL-0384-60025	01720	104220	33442
MIL-0384-60025	01720	104220	33442
POG-0383-60025	104224	121562	38844
POG-0383-60025	104251	121553	38571
BLA-0390-60025	104357	190536	05039
BLA-0390-60025	104357	190536	05039
BLA-0390-60025	181403	192536	65443
COO-0384-60025	192555	211231	69955
COO-0384-60025	192555	211231	69955
COO-0384-60026	132720	150030	48445
POG-0394-60026	132725	150030	48445
POG-0394-60026	150101	165730	04061
MIL-0398-60027	247	11343	167
POG-0397-60027	11400	30119	4440
POG-0397-60027	11400	30119	4440
IND-0401-60027	115822	131643	43102
POG-0400-60027	131702	145051	47822
POG-0400-60027	131702	145051	47822
COO-0408-60027	205455	221042	78295
MIL-0403-60027	221901	235329	80341
BOS-0408-60028	05027	101520	31827
BOS-0408-60028	05027	101520	31827
COO-0407-60028	101545	114901	36945
COO-0411-60028	124206	141350	44036
COO-0411-60028	124206	141350	44036
COO-0413-60028	174713	191159	64033
POG-0410-60028	191215	204539	69137
POG-0410-60028	191215	204539	69137
MIL-0418-60029	143906	154934	82746
COO-0417-60029	154954	173734	86994
COO-0417-60029	154954	173734	86994

80° L=8-9

330° L=11-12 (copy)
80° L=6

rapid water you drop off 10 23.1 km vs 1 km

rapid water you drop off 231 140, 80° L=6-7
note 2

Bad comparison problem

80° L=6-7

230° L=7-12, 80° L=6-8

1st dropouts

230° L=7-12, 80, 150° L=8-11

80, 150° L=4-5

330° L=5-6 235° L=7-8, 80, 150° L=5-6

80° L=5-5

80, 150° L=4-5 Bad comparison problem

80, 150° L=4-5

330° L=7-8 235° L=7-8, 80, 150° L=5-6

80, 150° L=5-6

80° L=7-10

80, 150° L=6-11

33, 235° L=7, 80, 150° L=5-8, 5-8

330° L=7-9

330° L=7-9 235° L=7-8, 80, 150° L=6-11

330° L=7-8

80, 150° L=5-7

330° L=6-5, 7-10, 235° L=7-8, 235° L=7-8, 80, 150° L=5-8

235° L=7 (copy), 80, 150° L=6-9

2, 2, 330° L=6-10, 235° L=7-8 (copy), 80, 150° L=5-7

80, 150° L=6-8

80, 150° L=6-7

2, 2, 330° L=5-8, 80, 150° L=4-5, 6

330° L=5-8

80° L=6-7

330° L=8-10, 235° L=9, 80, 150° L=6-7

12, 330° L=7-8 235° L=7-8, 80, 150° L=5-7

12, 330° L=8-9 235° L=7-8, 80, 150° L=6-7

12, 330° L=8-9 235° L=7-8, 80, 150° L=6-7

12, 330° L=8-9 235° L=7-8, 80, 150° L=6-7

12, 330° L=8-9 235° L=7-8, 80, 150° L=6-7

REV NO.	DATE	TIME OF REV	TIME OF REV
OF REV		IN HRS	IN HRS
COO 2746	78613	13221	32939
POG 2745	78613	42911	60940
MUL 2754	78614	40217	55932
POG 2754	78614	65905	83944
BJA 2761	78615	3906	23633
BOS 2766	78615	18227	171957
COO 2770	78616	1229	20937
MUL 2769	78616	36916	44985
COO 2778	78617	23451	14306
POG 2777	78617	24230	42319
COO 2784	78617	20257	222208
BJA 2798	78619	44601	64333
COO 2802	78620	41931	41657
POG 2812	78621	91525	111239
BOS 2815	78621	190349	210112
COO 2821	78622	114423	134144
BJA 2823	78622	173727	193451
BOS 2824	78622	204736	223127
POG 2825	78623	2930	22650
POG 2827	78623	52344	72107
MUL 2829	78623	111640	131415
MUL 2830	78623	143537	161054
BJA 2832	78623	220691	220404
COO 2833	78623	231958	10042
COO 2835	78624	45624	42354
IND 2837	78624	184938	124706
MUL 2839	78624	164250	184018
COO 2841	78624	194947	213655
MUL 2843	78625	42015	42641
COO 2845	78625	102228	121954
MUL 2846	78625	132916	151632
MUL 2847	78625	162052	181306
POG 2848	78625	191932	210942
MUL 2849	78625	20533	25222
MUL 2850	78626	40205	54822
COO 2852	78626	65037	85606
BOS 2853	78626	101501	115244
COO 2854	78626	130222	144921
MUL 2857	78627	3822	21536
MUL 2859	78627	33503	53219
IND 2861	78627	92007	112531
IND 2865	78627	152119	171644
IND 2869	78628	90057	105820
COO 2870	78629	83346	103110
BJA 2880	78629	182227	201949
COO 2887	78630	110028	130032
BOS 2894	78701	111902	124344
COO 2895	78701	134294	154015
MUL 2896	78701	164054	183649
BJA 2897	78701	193550	213223
BOS 2900	78702	42551	42305
COO 2901	78702	74039	91940
BOS 2902	78702	102327	121612
COO 2903	78702	131527	151244
MUL 2904	78702	161216	180920
IND 2906	78702	190832	210554
IND 2909	22222	213230	296
COO 2907	78703	10144	25901
BJA 2909	78703	48449	64041
BOS 2913	78703	192866	212544
COO 2919	78704	91226	110943
COO 2925	78705	64708	84432
COO 2928	78705	143794	163523
COO 2934	78706	81726	101445
BJA 2941	78707	45319	68042
COO 2944	22227	144152	163915
BJA 2949	78708	64121	62306
BOS 2951	78708	103057	122820
IND 2951	78708	132545	150425
BJA 2954	78708	196330	211757
COO 2957	78709	11330	31058
POG 2958	78709	50850	69731
COO 2959	78709	108827	128033
MUL 2961	78709	156419	176334
POG 2962	22228	105243	205007
MUL 2963	78710	4840	22107
MUL 2964	78710	34220	51733
POG 2966	78710	63852	83615
POG 2967	78710	95044	113250
COO 2968	78710	123928	142923
MUL 2969	78710	153805	172857
BJA 2970	78710	183509	202230
BOS 2972	78711	1013	21530
BOS 2973	78711	48142	61214
BOS 2974	78711	61350	80044
POG 2975	78711	91826	110522
COO 2976	78711	121510	140135
BOS 2977	78711	150422	169820
BJA 2978	78711	180319	199503
COO 2979	78711	209615	228258
MUL 2981	78712	24719	44110
MUL 2982	78712	142946	161809
POG 2991	78713	86833	101041
BOS 2993	78713	140626	160343
BOS 2995	78713	198922	218645
IND 2996	78713	233946	2519
BOS 2998	78714	44907	64420
POG 2999	78714	83855	94284

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REV NO.	DATE	TIME OF REV	TIME OF REV	TIME OF REV
OF REV		IN MRS	IN SECS	
COO 3000	70714	111059	123925	40259 48568
MUL 3002	70714	164540	103233	60340 64783
QUA 3003	70714	193140	212906	70300 77346
IMD 3004	70714	222812	233812	80092 13380
POG 3005	70715	124445	38212	8065 11132
BOS 3007	70715	71757	91521	26277 33321
POG 3008	70715	101445	121154	17005 43914
MUL 3009	70715	133045	150030	40445 46510
MUL 3010	70715	160741	180505	80061 86108
IMD 3012	70715	220054	235610	79254 86290
BOS 3014	70716	35350	55117	14030 21077
BOS 3015	70716	45110	64750	24670 31676
BOS 3016	70716	102103	114427	37263 42267
MUL 3017	70716	124335	144059	45815 52859
QUA 3018	70716	154407	173732	56767 63452
IMD 3020	70716	213319	233040	77599 84640
POG 3021	70717	2952	22714	1792 8034
BOS 3023	70717	62256	82021	22976 30021
BOS 3024	70717	93504	111658	34504 40618
COO 3025	70717	124920	141332	46160 51212
MUL 3026	70717	155313	171005	57195 62805
BOS 3027	70717	180913	200637	65353 72397
BOS 3028	70717	215529	230313	78929 82993
BOS 3030	70718	25057	45619	10737 17779
COO 3031	70718	60538	75255	21938 28375
COO 3032	70718	93410	104928	34450 39960
COO 3033	70718	114042	134408	42522 49560
QUA 3034	70718	144514	164234	53114 60154
POG 3035	70718	174146	193909	63704 70749
BOS 3036	70718	211643	223544	74403 81344
BOS 3038	70719	23123	42849	9083 16129
BOS 3039	70719	54735	72525	20835 26725
POG 3040	70719	82436	102157	30276 37317
COO 3041	70719	112556	131831	41156 47911
POG 3042	70719	143540	161506	52548 58506
POG 3044	70719	201052	220013	72652 79693
POG 3046	70720	20405	40120	7445 14480
BOS 3047	70720	51621	65753	10961 25073
MUL 3048	70720	83653	95422	11133 18662
BOS 3049	70720	110405	125052	13945 24252
QUA 3050	70720	143910	154722	15750 24042
BOS 3052	70720	194254	214018	70974 78018
MUL 3053	70720	234830	13548	85710 8768
BOS 3055	70721	43231	62949	14351 23509
IMD 3056	70721	73111	92620	27071 33900
MUL 3058	70721	133039	151923	40639 55116
MUL 3045	70722	94249	114012	34964 42012
BOS 3071	70723	12410	32120	4050 12000
COO 3074	70723	121122	140036	43082 50916
COO 3082	70724	114308	134032	42108 49232
MUL 3087	70725	22533	42257	8733 15777
COO 3093	70725	151022	170794	54622 61664
BOS 3096	70726	45359	65121	17639 24681
COO 3104	70727	72215	91965	24545 32505
BOS 3109	70727	101153	100912	50313 63352
QUA 3117	70728	154043	174107	54443 63667
MUL 3119	70729	31652	50521	11812 18321
COO 3121	70729	52036	72557	19716 26757
MUL 3122	70729	93020	112115	34228 40875
MUL 3123	70729	122204	141739	44524 51459
QUA 3124	70729	151644	171408	53004 62040
QUA 3125	70729	181317	201034	63597 72634
IMD 3126	70729	211141	230136	74201 82874
BOS 3128	70730	30245	45954	10945 17996
BOS 3129	70730	61124	79424	21206 28504
COO 3130	70730	90920	105253	32964 39173
COO 3131	70730	118150	134916	42710 49754
QUA 3133	70730	175046	194213	64726 70933
IMD 3134	70730	204119	223038	74479 81518
MUL 3135	70730	233751	13505	85971 8705
BOS 3137	70731	53039	72758	19030 24670
BOS 3138	70731	84408	102428	11840 17460
MUL 3139	70731	115032	132056	43112 48054
MUL 3140	70731	145440	161723	53400 60443
MUL 3144	70801	15233	34953	6753 13793
MUL 3155	70802	71707	91424	26227 33264
POG 3160	70802	215924	235622	79164 86182
POG 3164	70803	123629	143854	49349 52734
POG 3165	70804	32350	52110	12230 19270
MUL 3175	70804	100607	120325	35147 42285
POG 3179	70805	84824	102621	31704 37801
MUL 3179	70805	102640	121021	37600 43021
BOS 3181	70805	153208	165221	58920 60741
COO 3184	70806	2329	14142	1409 6102
POG 3186	70806	53713	79436	20133 27276
BOS 3187	70806	89058	103105	31850 37965
COO 3188	70806	113101	127710	41410 48440
BOS 3189	70806	143554	162358	52584 60030
COO 3191	70806	201930	221650	73170 80210
BOS 3193	70807	21235	40944	7955 14984
BOS 3194	70807	52053	70613	19235 25073
MUL 3195	70807	80523	100239	29123 36189
COO 3196	70807	111052	125907	40252 46787
QUA 3197	70807	142100	155534	51640 57394
BOS 3198	70807	165439	180189	60470 67019
MUL 3203	70808	72205	91920	26525 32064
MUL 3208	70808	220422	30	77942 30
POG 3210	70810	32816	52522	12446 19022
COO 3223	70810	101001	200716	68401 72436

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REV NO.	DATE	TIME OF REV	TIME OF REV	TIME OF REV
		IN MMS	IN MMS	IN SECS
COO 3228	70811	94850	110505	35330
MUL 3229	70811	121618	140130	44170
MUL 3230	70811	150035	165753	84035
BJA 3230	70812	154736	175140	86056
COO 3244	70813	53454	71920	20094
BOS 3249	70813	104407	202913	67447
MUL 3254	70814	102231	121944	37351
BJA 3255	70814	170608	182819	61560
BOS 3256	70814	194744	212443	71264
COO 3257	70815	12016	132028	4816
COO 3260	70815	65745	85416	25065
COO 3265	70815	104234	203945	67354
COO 3269	70816	62003	82515	23283
MUL 3279	70817	124133	140522	45493
IND 3280	70817	180053	194457	64053
COO 3281	70817	208709	225427	78429
POG 3283	70818	24950	44713	10190
COO 3285	70818	74350	94110	27830
COO 3286	70818	113902	134623	41942
BOS 3287	70818	143227	163243	82527
MUL 3289	70818	202023	225220	73703
BOS 3291	70819	22056	41814	8456
COO 3294	70819	71456	91213	26096
COO 3297	70819	194153	213910	70913
MUL 3303	70820	102330	122051	37410
COO 3310	70821	65755	85513	25075
COO 3314	70821	213940	234654	77980
BOS 3320	70822	121125	141837	44405
BJA 3324	70823	30302	50010	10902
BOS 3331	70823	204104	223820	74444
BJA 3332	70824	53006	70346	19808
BOS 3333	70824	71606	82610	26160
MUL 3334	70824	84009	112000	31209
POG 3334	70824	114049	123925	42049
BJA 3336	70824	141337	161622	51217
BOS 3340	70825	20526	40140	7526
COO 3346	70825	194211	213916	70931
BOS 3353	70825	102324	122035	37404
COO 3359	70827	65710	85426	25038
BOS 3364	70827	213839	233540	77919
MUL 3366	70828	121952	141704	44392
MUL 3376	70829	85345	105054	32025
MUL 3384	70830	82347	102101	30227
POG 3388	70830	200852	220605	72532
BOS 3391	70831	78357	98108	28437
BOS 3402	70831	92126	111834	33686
POG 3407	70832	65344	85053	24824
COO 3413	70832	213457	233201	77697
MUL 3416	70833	101619	121540	36779
MUL 3425	70834	94652	114529	35212
COO 3428	70834	203420	223139	74060
BOS 3432	70835	91805	111520	33485
BOS 3439	70835	230030	25741	80839
BOS 3442	70836	134135	153046	49298
MUL 3449	70837	42216	61920	15736
POG 3453	70837	192550	205953	68570
MUL 3458	70838	94322	114027	35002
MUL 3465	70839	91212	110922	33132
BJA 3466	70839	120621	140531	43701
IND 3469	70839	205645	225349	75405
COO 3472	70839	235341	24956	86021
MUL 3473	70810	84350	103816	31450
BJA 3475	70810	113710	133423	41830
POG 3476	70810	202535	212828	73535
BOS 3477	70810	213719	222249	77039
BOS 3478	70810	232130	1621	84099
POG 3478	70811	2959	11053	1799
BOS 3480	70811	81400	71104	18840
BOS 3481	70811	83640	109713	31000
COO 3482	70811	112000	130119	41200
MUL 3483	70811	140500	155926	52080
BJA 3484	70811	173121	185534	63001
COO 3485	70811	195433	215130	71673
COO 3486	70812	44257	63959	16477
BOS 3494	70812	192331	212034	69011
BOS 3500	70813	100604	120109	36244
MUL 3506	70814	85413	103119	30853
POG 3511	70815	211710	231412	76430
BOS 3516	70815	78100	85784	25260
MUL 3515	70815	90119	108822	32479
COO 3521	70816	53336	73440	20016
IND 3526	70816	171748	191440	62265
COO 3529	70817	28954	40284	7884
POG 3538	70818	43012	62715	16212
BOS 3542	70818	191829	210727	69029
COO 3546	70819	65430	85136	24070
MUL 3548	70819	134030	153113	49230
POG 3549	70819	161455	181134	60495
IND 3550	70819	193447	210750	78407
POG 3551	70819	221039	385	80319
POG 3553	70820	38055	80557	14335
POG 3559	70820	39029	80909	44920
COO 3554	70820	71648	83180	24200
MUL 3555	70820	101920	113717	37160
BJA 3556	70820	123640	143323	48400
BJA 3557	70820	153600	172923	60160
IND 3558	70820	182857	202553	64837
POG 3559	70820	213105	232154	77465
IND 3559	70821	4829	21795	8275
BJA 3561	70821	31687	81383	11017

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				IN SECS			
COO 3562	70921	62306	80956	21906	29396	23L76	235 L=6, 6-9 80, 10L=
POG 3563	70921	93122	110555	34282	39955	23L76	235 L=5-9 80, 10L=
GUA 3564	70921	123250	140155	48178	58815	23L76	235 L=9 80-356L76
POG 3565	70921	150058	165753	54058	61873	23L76	235 L=5 80, 10L=
IMO 3568	70921	134851	14554	88731	6384	23L76	235 L=6-70, 80, 10L=
POG 3569	70922	244521	44153	9891	18913	23L76	235 L=6-7, 8, 9 80, 10L=
MUL 3570	70922	55515	73755	21315	27875	23L76	235 L=6, 7, 8, 9 80, 10L=
MUL 3572	70922	121852	132954	44332	48594	23L76	235 L=7 80, 10L=
GUA 3573	70922	142852	168524	52132	59124	23L76	235 L=7 80, 10L=
COO 3577	70923	21253	40954	7973	14994	23L76	235 L=5, 7 80, 10L=
MUL 3578	70923	52645	70556	19405	25556	23L76	235 L=6
MUL 3579	70923	81933	100156	29973	36116	23L76	235 L=6
COO 3580	70923	111230	125757	48350	56677	23L76	235 L=6, 7 80, 10L=
GUA 3581	70923	135654	155354	50214	57234	23L76	235 L=6, 7, 8, 9 80, 10L=
BOS 3583	70923	165254	184954	67774	77794	23L76	235 L=6, 7, 8, 9 80, 10L=
COO 3585	70924	14047	33738	4047	13058	23L76	235 L=6 80, 10L=
BOS 3586	70924	50551	63328	18351	23608	23L76	235 L=6 80, 10L=
COO 3587	70924	29855	92918	29335	34158	23L76	235 L=6
COO 3588	70924	105157	122504	39117	44704	23L76	235 L=6
MUL 3589	70924	132400	152054	48240	55254	23L76	235 L=7, 80 L=6
BOS 3590	70924	161952	181642	58792	65802	23L76	235 L=6, 7, 8, 9 80, 10L=
COO 3593	70925	10713	30489	4033	11049	23L76	235 L=7, 8 80, 10L=
BOS 3595	70925	65857	85546	25137	32146	23L76	235 L=7, 8 80, 10L=
COO 3596	70925	176457	195134	38697	42694	23L76	235 L=7, 8 80, 10L=
MUL 3597	70925	132226	144726	48146	53246	23L76	235 L=7, 8 80, 10L=
GUA 3598	70925	155826	174315	57506	63795	23L76	235 L=7, 8, 9 80, 10L=
BOS 3602	70926	34115	82629	13275	19589	23L76	235 L=7, 8, 9 80, 10L=
BOS 3603	70926	62531	82217	23131	30137	23L76	235 L=7, 8, 9 80, 10L=
COO 3611	70927	55355	75040	21235	28240	23L76	235 L=7, 8, 9 80, 10L=
COO 3612	70927	92803	104910	34883	38950	23L76	235 L=7, 8, 9 80, 10L=
MUL 3613	70927	121203	134780	43923	49620	23L76	235 L=7, 8, 9 80, 10L=
BOS 3614	70927	145321	165512	53612	60252	23L76	235 L=7, 8, 9 80, 10L=

aliability in identifying REPs was obtained.

Fig. C-3 shows how each microfiche is layed out and Table C-1 is a listing of the individual microfiche by Rev Number, Date, UT of Rev in hrs and secs, and notes on electron and proton precipitation events. Approximately 1212 microfiche were examined three times by one individual (the author) and once by another graduate student (T. Kelley) for comparison. A list of 358 outer zone transits were identified for further discrimination by pitch-angle. 313 outer zone REP revs were confirmed by the pitch-angle analysis and are listed in Table C-2. The remaining case was eliminated by the loss-cone criteria previously described, bad data drop outs, low flux values, and spike events. This table also subjectively groups the revs into 173 'events' by combining consecutive revs with strong precipitation. Since each revs last 3 hrs and there was no AE data set available to determine substorm duration this grouping maybe off as much as 25% in total event number.

Table C-2. Revs with strong precipitation \geq 235 keV electrons by EVENT (one or more consecutive revs with strong precipitation), DATE (last digit of year, month, day), Readout REV (first letter of station, number), LAT (North, degrees), East LONG (degrees), L-VAL, MLT (hrs), $j(e/cm^2 s str keV)$ = differential flux value of precipitating electrons from 235 keV to highest energy channel at which strong precipitation occurred, $J(p/cm^2 s str)$ = integral flux of strong diffusion precipitation of protons per energy channel if any, 80 p⁺ L-VAL = L value of 80 keV proton precipitation if any, strong precip 12, 33, L-VAL = check if 12, 33 keV electron channels had strong precip and L value of 12 keV channel, UT (sec) = time at which 235 keV electron precip occurred.

EVE N	DATE	REV	LAT	LONG	LVAL	MLT	E55	E45	J [°/hr. obs.]			J [°/hr. obs.]			SOPT	STAGUS PROG		(1/2)
									850	1.1	1.35	1.6	80	150		850	1.2	
1	60714	B49	68.1	293.7	6.17	21.2	306+1			1.6	1.66+4.66+3	6.5-6	✓	6-12	16.2			
2	60809	C252	62.0	598	6.4	20.4	765+1			3.33+4.833+3	5-6.5	✓		6.4	15.8			
3	60824	M375	65.1	528	7.0	10.3	506+2			8.33+3	9-10	✓		6-8.5	8.6			
4	60824	M381	47.6	914.2	7.1	20.1	163+1			2.5+4.833+3	6.7-6.6	✓			22.5			
5	60825	C386	56.8	554.8	7.8	20.2	306+1			3.5+5.833+4	5.3-7.5	✓		6-2.5	19.5			
6	60827	M398	82.2	12.0	7.9	21.8	306+1			1.66+4.85+3	5-10	✓		5-10	1.2			
7	60827	M398	42.5	274.7	6.2	19.5	1.53+2.307+1			8.33+4.85+4	5-7	✓						
8	60827	C405	44.3	319.9	5.7	20.1	46+2	353+1		4.5+4.833+3	4-8	✓		6	22.1			
9	60828	P410	63.1	321.9	9.5	9.7	6.12+1			6.66+3	8-10	✓		8-11	11.2			
10	60906	B477	77.9	305.3	5.3	22.0	1.0+3	1.98+2	8.66+1	7.7+0	4.95+0	2.5+4.833+3	✓	6-12	20.1			
11	60914	C547	65.4	206.5	6.4	22.5	1.53+2	1.2+0		7.5+3	1.3+3	5-6	✓	5-7	6.4			
12	60918	B575	45.0	157.3	4.3	20.3	2.65+1	1.8+1		1.1+5.75+4	8.3+3	4-7	✓	4-7	8.9			
13	60919	I584	50.4	508.1	5.4	8.3	9.2+2			8.5+3	6.7+2	7-9	✓	5-6	10.9			
14	60920	M589	7.5	95.9	7.0	8.5	1.1+3	2.8+1		8.5+3	9-12	9-12	✓	6-10	1.7			
15	60920	I592	68.0	2.2	7.1	9.5	7.65+1			2.5+3	10-20	10-20	✓	6-9	7.7			
16	60920	B595	71.3	219.8	7.5	7.8	1.4+3	2.3+1		1.5+3	5-6	5-6	✓	5-20	14.3			
17	60921	C599	61.4	55.5	6.0	3.7	3.06+1	6.1+0		7.5+4.10+4	5-6	5-6	✓	4-5-20	2.2			
18	60921	B599	77.0	285.0	6.1	22.2	3.06+2	3.08+1		1.0+5.16+4	5-6	5-6	✓	5-8	2.3			
19	60921	B598	65.0	50.5	6.0	9.1	1.2+3			9.3+4	3.3+4	5-6	✓	6-11	4.5			
20	60921	B602	47.5	118.6	6.0	19.4	4.6+1						✓	6-8	11.4			
21	60922	I608	74.8	110.8	10.5	8.7	3.06+1						✓	1-2-17	1.3			

DATE	REV	LAT	LONG	(g)	L-VAL	MLT	d [1/2.5, 4.5, 10.0]			T [1/2.5, 4.5, 10.0]			STRAPE PERIOD		L-VAL	UT		
							435	455	500	11	1.35	1.6	80	150			350	12
17	60829	C 619	-51.2	123.1	8.0	19.6	1.2E2					2.5+5	6.6+4	2.5+3	✓	✓	6-10	11.1
18	60926	C 644	-76.8	290.2	6.7	21.5	3.5+1					3.3+4	4.2+3		✓	✓	5-6	1.5
19	60929	C 665	-57.0	182.0	6.9	21.1	1.5+2					2.3+2			✓	✓	6-8	7.01
20	61002	I 639	53.4	288.0	7.2	7.4	7.6+2	3.0M1				2.3+3			✓	✓	5-12	11.8
21	61003	I 697	-75.8	14.1	7.2	2.0	4.6+1	1.5+1				7.5+3			✓	✓	6-8	3.2
	61003	I 697	-73.5	280.1	6.0	21.8	1.5+2	2.0+1	7.0+0			4.2+4	6.7+3		✓	✓	6-7	3.5
22	61016	I 803	-76.2	264.6	6.7	5.0	6.1+2					1.0+4	5.8+3		✓	✓	5-8	9.8
23	61017	C 808	-58.0	62.0	6.2	4.0	4.6+1	1.4+1				2.3+3	1.6+3		✓	✓	4-8	6
	61017	C 808	-73.0	280.0	6.0	19.6	4.6+2	4.6+1				2.8+4	8.3+3		✓	✓	2-5	1.9
24	61021	C 845	-69.3	283.8	7.3	16.7	4.6+1					7.5+3			✓	✓	6-6.5	1.1
25	61028	C 901	-56.0	253.7	6.5	5.2	1.1+2								✓	✓	6.5	12.9
26	61031	C 924	-75.3	255.7	6.0	4.7	7.7+1					2.3+4	5.0+4		✓	✓	4-7	9.9
27	61110	I 004	59.5	272.7	6.4	4.4	4.6+1					2.5+3			✓	✓	6-9	13.2
28	61111	I 1011	-65.3	255.5	6.4	1.6	9.1+1					1.6+4	7.5+3		✓	✓		1.6
	61111	I 1011	-69.4	238.0	7.3	19.6	4.6+1					2.5+3			✓	✓	6-7	1.9
29	61112	I 1025	-50.5	252.2	5.5	3.0	3.0+1					2.5+4	7.5+3		✓	✓	5	22.1
30	61114	C 1034	-64.3	35.2	7.1	1.5	2.0+1					4.2+4	1.7+4		✓	✓	6-9	6.9
	61114	C 1034	71.9	39.0	7.5	7.0	1.4+2								✓	✓	6-10	3.3
31	61201	C 1174	-58.9	178.0	7.2	4.5	1.8+1					2.5+3			✓	✓	6	14.8
32	61204	C 1202	70.5	188.5	8.7	12.4	5.0+1					2.3+3			✓	✓	7-12	1.5
33	61209	C 1234	-74.0	273.8	7.0	2.3	4.6+1	4.6+0				6.7+3	2.5+3		✓	✓	5-9	7.1

E V E N T	DATE	REV	LAT	LONG	I-VAL	MLT	u [km ² /m ² hr]			S [km ² /hr]			BOPT I-VAL	STAGE	Metric	3 (HRS) UT		
							485	485	1.6	80	150	350					12	33
34	6/12/11	I1250	632	22.9	7.0	0.5	485	485	1.1	1.35	1.6	80	150	350	✓	33	1-VAL	6
35	6/12/12	M1266	573	272.5	8.8	14.0	655	850	1.1			1254	33+3		✓		5-9	6
36	6/21/9	P1317	604	4.7	51.6	0.6						67+3	8.3+2		✓		6-13	20.2
37	6/12/21	M1338	48.5	109.7	7.0	1.82						42+4	1.7+4		✓		5-6	1.6
38	6/12/22	G1340	679	322.5	5.6	0.6						83+3	1.7+3		✓		7	18.7
39	6/12/22	D1342	64.9	236.6	5.8	2.9	1.2+1	1.6+1	7.5+0			2.5+4	5.8+3		✓		5-6	9.8
40	6/12/22	E1345	44.8	114.5	5.4	2.0						5.0+3			✓		5-6	9.8
41	6/12/23	I1350	66.2	388.2	6.1	2.8						83+3			✓		4-6	15.5
42	6/12/25	I1369	50.3	166.0	68.3	2.9						1.1+4			✓		5	18.5
43	6/12/27	G1380	67.8	338.1	6.5	0.1						1.7+3			✓		6	9.5
44	6/12/27	M1386	48.5	128.1	7.3	2.0						2.5+3			✓		5-8	14.8
45	6/12/29	I1398	58.8	307.0	5.8	1.2						3.3+3	9.2+2		✓		5-8	20.7
46	6/12/29	G1404	55.5	42.0	6.8	0.1						8.3+3			✓		23.4	2.6
47	6/12/30	I1406	71.5	277.6	7.3	0.6						2.5+3	1.7+4	3.3+3	✓		6-7	17.2
48	6/12/31	G1413	73.0	130.0	6.0	12.1						1.7+4			✓		6-7	5.0
49	6/12/31	G1420	58.8	202.0	7.0	2.5						9.2+3	2.5+3		✓		4-7	8.0
	6/12/31	C1444	63.0	221.0	6.5	2.6						7.5+4	1.7+4		✓		7	22.7
												9.2+3	2.5+3		✓		6	4.9
												8.3+3	8.3+3		✓		6-8	7.8
												2.5+3			✓		7	9.7
															✓		7	10.4

F V E N T	DATE	REV	LAT	LONG (E)	L-VAL	MIT	d [m/s ² max]				1.6	S (m/s ²)			STR	L-VAL	4RS UT
							285	435	655	850		1.1	1.35	80			
50	70103	01408	57.8	12.0	5.8	23.8	1.5+2	1.5+1		1.35	1.7+4	3.3+3	✓	6	6	4	
51	70104	01450	-44.0	156.0	5.2	2.2	1.5+2	1.2+2	2.3+1	7.7+0	4.2+4	8.3+3	✓	6.2	>5	15.1	
52	70105	I MISS	64.5	236.2	6.5	2.1	1.5+1				1.7+4	2.5+3		6	9.5		
53	70108	01485	-49.2	77.7	6.6	0.3	1.5+2	3.1+1			8.3+3	1.7+3		6.5	20.0		
54	70115	01535	71.0	98.5	6.5	11.8	1.5+2				8.3+3		✓	>6	>5	5.5	
55	70116	I 1517	-71.0	308.0	7.8	23.2	3.0+1				6.7+3		✓	6-7	8	9.4	
56	70117	MISS	-44.0	107.1	6.0	0.6	1.4+1				4.2+3		✓	6	6	17.6	
57	70118	M 1560	62.3	270.5	5.0	0.5	3.0+1				2.5+4	8.3+3	✓	5	5.55	5.7	
58	70123	01600	69.2	277.8	7.1	23.7	4.6+1				2.5+3		✓	7	4.5	4.5	
59	70124	01608	69.2	293.0	7.0	23.6	1.5+1				1.7+3		✓	>6	>6	4.2	
60	70130	01656	63.0	275.5	5.0	0.04	4.6+2	7.7+1	1.5+1		2.5+4	6.7+3	✓	>5	>5	5.5	
	70130	M 1658	56.8	235.8	5.0	1.06	6.1+1	1.5+1	4.9+0		1.7+4	2.5+3	✓	4.55	>4.5	8.4	
	70130	M 1658	-49.2	195.8	5.0	0.8	3.0+2	4.6+1	1.6+1		1.7+4	3.3+3	✓	4.55	>5	11.2	
	70130	01659	-45.4	154.1	4.6	0.8	1.4+2	1.5+1			7.5+4	8.3+3	✓	4.55	>4.5	14.1	
	70130	01661	37.9	112.4	4.85	0.2	1.5+3	3.1+2	5.4+1	4.5+0	8.3+4	2.5+4	✓	4.5-9	4-9	17.0	
	70130	01661	-41.8	64.5	4.5	23.6	6.1+2	7.7+1	1.6+1	4.1+0	4.2+4	1.7+4	✓	4-6	4-6	20.0	
	70131	01664	26.2	323.2	5.9	23.2	1.5+2	4.6+1	8.2+0	5.3+0	7.5+4	2.5+4	✓	4-5	4-5	23.0	
	70131	01664	61.3	280.4	4.95	23.9	4.6+1				7.5+4	1.7+4	✓	5-6	5-6	2.9	
	70131	01665	62.2	248.3	4.95	23.9	3.0+1				NO DATA		✓	4.8	4.8	5.2	
61	70131	I 1665	-49.7	199	5.1	1.0	1.7+1				3.3+4	2.5+3	✓	5	5	10.9	
	70131	MISS	44.1										✓	5.2-5.5	6	16.8	
62	70131	I 1665	41.6	113.8	5.7	0.1	4.6+2	6.1+1			8.2+3	7.5+3	✓	5.5-6.0	6	17.8	
	70131	M 1670	46.4	67.0	6.2	23.2	3.0+1				8.3+3	4.2+3	✓				
63	70202	I 1681	50.7	205.3	4.85	0.9	3.0+1				7.5+4	1.7+4	✓	4.5	4.2-5	10.4	
	70202	I 1684	-47.4	162.5	5.6	0.1	2.0+1				2.5+4	2.5+3	✓	6	13.9		

DATE	REV	LAT	LONG	L-VAL	MIT	T ₀ (sec)			T ₁ (sec)			T ₂ (sec)			12	13	14	15
						235	435	635	850	1.1	1.35	1.6	80	150				
70202	I 1874	40.5	121.2	49.75	8.9	6.1+1	7.6+0					4.2+4	8.3+3	5-9	✓	✓	5-15	16.2
70202	H 1486	19.2	235, 235	15.5	23.5	11.5	11.5											
70203	O 1478	1.4	235, 235	16.5	11.5	11.5	11.5											
70203	M 620	7.9																
70203	I 1692	46.9	160.2	5.9	0.8	2.2+1						6.7+4		6	✓	6-7	13.0	
70205	O 1702	48.3	172.5	6.2	0.8	1.5+2	2.0+1					1.7+4	6.7+3	6.5-6	✓	6	12.5	
70208	H 1734	75.2	142.1	57.10	1.4	1.5+2	3.0+1					1.6+4	8.3+3	5-10	✓	4-10	14.6	
70208	H 1734	41.0	97.5	5.7	23.4	1.5+2	2.0+1					8.3+4	1.7+4	4-10	✓	4-10	17.5	
70209	I 1741	73.5	50.1	6.8	11.2	3.0+2						2.5+3		4-10	✓	4-10	7.3	
70222	O 1749	68.4	57.3	6	1.5	4.6+1						8.3+3		6	✓	4-8	21.8	
70225	I 1864	71.2	280.4	6.7	0.2	1.2+2	1.5+1					2.5+4	7.5+3	6-8	✓	7-6	5.6	
70225	I 1820	87.8	203.3	6.8	23.3	1.5+2	3.0+1					1.7+4	5.0+4	7	✓	7	11.5	
70225	I 1870	41.9	156.2	7.6	23.7	3.0+1						8.3+3	4.2+3	6-7	✓	6-8	12.7	
70302	O 1709	64.9	219.1	7.8	22.8	1.5+1						6.7+3	8.3+2	6	✓	5-10	9.9	
70307	O 1747	66.9	270.8	6.2	22.4	1.4+1						1.7+3		5.7	✓	6.5	4.0	
70307	H 1748	57.0	232.0	6.0	23.1	1.5+1						3.3+3		5.5	✓	5-7	6.8	
70307	H 1748	56.0	233.5	5.8	22.7	1.2+0						DATA	DATA				8.3	
70309	H 1764	56.6	243.5	48.5	23.1	4.6+2	2.0+1					4.2+4	8.3+3	4.8-5	✓	4.6-5	6.1	
70309	O 1765	71.7	2.4	7-7.6	10.7	1.5+1						8.3+3		8	✓	7-8	12.7	
70309	P 1763	41.9	75.9	5.8	21.8	1.5+2	1.5+1					6.7+4	1.7+4	4-9	✓	4-10	17.7	
70309	C 1962	73.0	273.8	6.0	13.3	1.5+3						8.3+3		7-14	✓	6-7	18.6	
70309	P 1970	51.3	244	5.7	21.3	6.6+2	7.7+1	6.5+0				5.0+4	2.5+3	4.8-7	✓	5.6	20.8	
70310	O 1772	3.0	176.4	35.2	9.9	7.2+1						7.5+3		12	✓	7-10	9.4	
70310	H 1775	65.2	41.5	9.3	22.6	2.7+3	4.6+1					5.8+4	2.5+3	6-8	✓	6-8	14.5	
70310	I 1775	43.6	121.6	6.8	20.1	1.5+1						3.5+4		5.5-4.5	✓	6.6	20.6	
70310	I 1977	56.6	25.1	7.0	20.1	1.5+1									✓			
70311	O 1800	68.4	285.5	62.6	21.8	4.6+2	3.5+1					8.3+3	6.7+3	5.5-7.0	✓	6-8	2.7	
70311	O 1986	62.0	340.8	53.5	21.0	7.6+2	1.6+2	8.2+1	3.5+1	7.7+0		8.3+4	2.5+4	4.9-7.0	✓	5.8	23.3	

EVENT	DATE	REV	LAT	LONG	L-VAL	MLT	235	435	δ (°/hr)		1.6	J (°/hr)				20°/hr	12	33	L-VAL	MRS UT
									655	850		80	100	350	350					
76	70312	01987	2-21	29	28	22.8	35+1		1.1	1.35		8.3+4	1.6+4	5-6.5	✓	✓	5-6.5	6	5-6.5	5.3
	70312	H1988	46.6	172.8	5.9	23.2	46+2					8.3+3	4.2+3	5.5-6.5	✓	✓	5.5-6.5	6	5.5-6.5	10.9
	70312	I1991	41.4	86.5	5.8	21.9	46+1					8.3+4	2.5+4	5-6	✓	✓	5-6	6	5-8	16.8
	70312	01994	55.5	347.6	47.5	21.3	15+3		1.1+1	7.4		5.8+4	8.5+3	4.2+3	✓	✓	4.2+3	6	>5	22.9
	70313	H1988	65.2	252.9	6.4	22.8	14+1					5.8+4	8.5+3	5-6	✓	✓	5-6	6	6-7	5.0
	70313	P1989	10.7	11.3	285	1-4	day and night								✓	✓	5-6.3	6	6-8	22.6
77	70313	02004	58.2	352.0	55.4	20.7	30+2	20+1				1.7+4	7.5+3	5-6.3	✓	✓	5-6.3	6	6	13.2
78	70314	I2008	41.8	139.2	6.0	22.7	12+2	15+1				1.7+4	2.5+3	5-6	✓	✓	5-6	6	6	7.2
79	70315	H2014	60.4	221.2	6.2	23.1	3.5+1	9.2+0				8.3+3	1.7+3	6	✓	✓	6	6	4.0	
80	70316	P2020	65.3	265.6	5.9	22.2	3.0+1	1.4+1				1.7+4	1.7+3	6	✓	✓	6	6	4.0	
81	70316	P2029	53.9	7.7	5.3	21.0	3.3+2	2.5+2	3.9+0			5.8+4	1.7+4	4.8-6	✓	✓	4.8-6	6	5.4-6	21.6
82	70317	02034	48.3	151	6-6.2	20.5	3.0+1					1.7+4	3.3+3	6	✓	✓	6	6	21.2	
	70318	02034	47.5	267.1	6	21.6	3.0+1					4.7+4	2.5+3	5-6	✓	✓	5-6	6	5.5-6.5	3.4
	70318	02034	51.5	195.5	54-5.9	23.0	4.6+1	9.2+1	2.1+1			1.7+4	2.5+3	5	✓	✓	5	6	5.5-7	9.1
83	70319	02051	53.7	20.7	5.7	20.7	3.0+2					8.3+3	2.5+3	5.5	✓	✓	5.5	6	20.6	
84	70320	02052	65.4	282.9	5.2-5.5	21.7	1.5+2	2.0+1				8.3+3	2.5+3	5.5	✓	✓	5.5	6	5.5	2.7
85	70320	02058	10.5	27.0	5.5	20.8	4.6+1					1.7+4	6.7+4	5-5.5	✓	✓	5-5.5	6	5.5	20.2
	70321	H2025	10.0	32.4	5-6	20.9	1.4+2					1.7+5	7.5+3	4-7	✓	✓	4-7	6	5-7	19.8
86	70321	02067	47.0	222.9	6.4	22.1	4.6+1					7.5+3		6	✓	✓	6	6	6.4	9.2
87	70322	P2071	60.1	220.9	6.4	22.1	4.6+1					2.5+3		6	✓	✓	6	6	6.4	9.2
88	70323	02077	46.3	230.9	6	22.3	9.2+1					2.5+3		6	✓	✓	6	6	4.6	7.6
	70323	P2079	64.0	207.0	5.3-7.3	23	4.6+1					2.5+3	8.3+3	5.9	✓	✓	5.9	6	5-9	7.6
	70323	P1019	51.8	235	day and night	MLT: 9.5, 10.3						8.3+3	8.3+3		✓	✓		6	5-9	7.6
89	70323	I2082	47.7	88.1	6	21.6	1.5+2					4.3+3	4.2+3	5.5-6	✓	✓	5.5-6	6	6	16.1

T W R E	DATE	REV	LAT	LONG	L-VAL	ALT	j [km ² sec ⁻²]				1.6	j [P/sec ²]				350	30,000	12	33	L-VAL	MCD UT
							435	465	850	1.1		1.35	80	150	370						
90	70324	C2084	61.2	304.9	45-47	21.2	7.6+2	1.5+2	3.3+1	850	1.1	1.35	1.6	6.7+4	1.7+4	456	✓	✓	4.5	1.3	
	70324	H2085	61.3	287.1	51.5	21.9	1.5+2							8.3+3	1.7+3	5.5	✓	✓	5.5	4.3	
	70324	E2086	61.4	215.3	58.64	22.8	3.0+1							2.5+3		5.9	✓	✓	5.5	7.2	
	70314	P2087	-49.1	178.7	6.0	22.8	4.6+1							5.0+3		6			6	10.0	
91	70325	E2089	52.1	181.6	6.5	22.8	1.7+1							5.0+3		6			6	10.0	
92	70326	P2104	72.3	146.7	57.6	22.3	6.1+1							2.5+4	1.7+3	5.5+5	✓	✓	6	12.1	
	70326	P2106	72.4	99.8	6-10	21.5	4.6+1							8.3+3	2.5+3	5.5	✓	✓	6	15.1	
	70326	P2106	-45.8	56.1	57-6.1	20.8	1.5+2							8.3+3	2.5+3	5.5	✓	✓	6	18.1	
93	70328	P2119	52.9	237.3	6	21.9	1.5+1							2.3+3	2.5+3	6			6	7.1	
94	70329	E2126	-67.7	282.2	7.2	22.9	1.5+2							8.3+3	7.5+4	6-7			6	5.5	
95	70331	H2144	58.6	210.1	6.2	21.7	1.2+2							8.3+3	2.5+3	6			6	9.0	
96	70401	H2182	157.2	215.3	6.5	21.7	1.5+1							2.5+3		6.5		✓	6.5	8.6	
	70401	H2182	149.7	170.7	8.3	22.5	2.0+1							2.5+3		6			6	10.1	
97	70404	P2177	44.1	143.2	5.9	21.9	6.1+1	3.0+1						5.0+4	8.3+3	6		✓	6	12.0	
98	70405	C2182	-71.4	260.4	6.2	21.4	2.0+2	4.6+1	8.2+0					2.5+4	8.3+3	6		✓	6	3.0	
	70406	C2189	78.7	313.4	5.6	20.8	7.6+1	4.6+0						8.3+4	8.5+4	5-6		✓	5.5	1.2	
	70406	P2186	71.6	265.1	6	21.2	6.1+2	4.6+1						4.2+4	1.1+4	5.5-6		✓	6	2.6	
	70406	C2194	63.7	198.3	6.7	21.6	9.2+1							2.5+4	4.2+3	5.5-6		✓	6	9.7	
100	70406	E2197	76.2	325.5	4.9-4.9	20.2	2.0+2	2.0+1						5.0+5	1.7+5	1.0+4	✓	✓	4.5-5.5	23.2	
	70407	E2200	63.9	278.9	4.3-5	21.2	1.2+2	6.1+1	3.3+1	1.8+1	5.8+0			1.7+5	8.3+4	3+3	✓	✓	4.5-8	2.2	
	70407	H2201	-42.4	158.4	5.1	21.8	3.0+2	2.0+1	3.3+1	3.5+0				8.3+4	1.5+4	1.7+3		✓	5.1	10.8	
	70407	E2202	39.7	115.2	5.1	21.3	4.6+1	8.2+0						8.3+4	2.5+4	1.7+3		✓	6	13.7	
	70407	E2202	-37.0	4.0	12.5	11.5	4.6+1							2.5+3	8.3+1	9-16		✓	8-13	14.3	
101	70408	C2204	-70.8	276.0	6.0	20.9	2.0+2	1.5+1						2.5+5	2.5+4	6.7+3			6	1.9	

E V E N T	DATE	REV	LAT	LONG	L-VAL	MLT	d (0/1000)		D (P km)		12	33	L-VAL	UT
							850	11	1.35	1.6				
102	70408	02011	40.8	75.2	4.9	20.8	655	850	1.1	1.35	1.6	33	1-VAL	16.3
103	70409	02288	48.0	105.5	5.7	22.1	1.2+2	0.6+1	1.1+1	3.9+0	✓	✓	6	10.1
104	70410	02256	54.3	162.2	7.6	22.2	6.1+1				✓	✓	6	9.8
70410	02288	46.5	82.9	6.0	20.5	1.5+2	2.7+0				✓	✓	6	15.6
105	70411	02230	68.5	249.9	5.7	21.5	1.5+2	2.1+1			✓	✓	5-7	3.7
106	70419	02297	58.5	20.6	14-65	21.7	4.6+3				✓	✓	3.5-12	6.4
70419	02297	7.6	10.8	6.7	9.8	1.5+2					✓	✓	6-7	7.3
70419	02288	47.1	172.3	4.8	21.7	4.6+2	4.6+1				✓	✓	4-5	9.4
70419	02297	54.9	169.5	4.4	21.4	3.0+2	4.6+1				✓	✓	4-4.5	10.7
70419	02300	63.4	130.1	4.8-5.2	22.0	1.5+3	6.8+1				✓	✓	5-5.5	13.6
70419	02301	54.7	35.5	5.7	19.5	7.7+1					✓	✓	5-5.5	18.2
70419	02302	53.9	35.5	5.7	19.5	7.7+1					✓	✓	5-5.5	18.2
70419	02303	53.9	35.5	5.7	19.5	7.7+1					✓	✓	5-5.5	18.2
70420	02304	42.1	301.1	5.4	22.1	7.6+2	1.5+1				✓	✓	6	22.5
70420	02305	45.4	238.6	5.7-6.4	21.2	1.1+3	7.7+1	1.1+1			✓	✓	5-6	1.6
70420	02306	43.3	212.4	6	21.9	6.1+1	1.2+2	1.5+1			✓	✓	6-6.5	4.5
70420	02307	57.1	174.8	5-6.6	21.7	9.2+2	7.7+0				✓	✓	6	6.1
70420	02307	47.8	132.8	7.1	21.0	7.7+1	6.1+1	7.8+0			✓	✓	6-7	8.9
70420	02307	64.0	42.2	5.5	22.6	1.5+2	4.6+1	1.6+1			✓	✓	7	11.8
70421	02311	20.0												16.1
107	70421	02318	70.7	354.1	5.95	20.6	2.0+2	3.5+1	1.6+1	3.9+0	4.5+0	✓	✓	2.8
108	70422	02323	72.7	256.3	5.8	20.8	3.0+1					✓	✓	2.4
109	70423	02333	45.2	105.5	6.0	20.3	5.1+1	1.6+1				✓	✓	13.6
110	70424	02347	48.4	159.1	5.8	21.3	7.7+1	1.5+1			✓	✓	12	9.8
70425	02348	61.3	160.9	5.6	21.3	6.1+2	9.2+1	1.1+1			✓	✓	5-7	11.2
111	70428	02381	52.4	359.4	5.4	22.3	2.0+2	1.2+1			✓	✓	5-5.5	21.5
112	70430	02384	43.2	265.9	6.1	20.9	4.6+1					✓	✓	5-6

STATION	DATE	REV	LAT	LONG	L-VAL	MLT	d (P/Sec - 4 km)			E (P/Sec - 4 km)			80°	12	33	L-VAL	UT											
							655	850	1.1	1.35	1.6	80						150	350									
115	70502	P2404	56.9	275.5	4.6	21.0	1.5+3	1.5+2	435	235	435	1.5+2	1.5+1	1.8+1	655	850	1.1	1.35	1.6	80	150	350	1.7+3	4.5+12	4.5-5	2.9		
	(70502)	P400	(6.0)				20.15+2					6.1+1	1.5+1												4-16	4-1	8.5	
116	70502	P2404	51.7	192.4	4.6	20.6	6.1+1	1.5+1				6.1+1	1.5+1												4-16	4-1	8.5	
117	70502	H2407	22.7	66.8	6.6	22.4	7.4+3	1.4+2				7.4+3	1.4+2												5.5-6	4.5-6	17.2	
	70502	C2408	45.9	324.9	4.1	19.1	3.1+2	7.0+1				3.1+2	7.0+1												5.5-6	4.5-6	17.2	
	70502	C2408	52.6	169.4	6	11.3	1.5+3	4.6+1				1.5+3	4.6+1												6	6	22.1	
	(70502)	H2410	(2.0)																							4.5-6	4.5-6	17.2
118	70500	P2111	48.7	239.5	57.6	20.5	1.5+2					1.5+2													4.5-6	4.5-6	17.2	
	70500	C2411	41.8	297.8	6	21.4	4.6+2	1.5+1				4.6+2	1.5+1												5.5-6	6-6.5	1.1	
119	70510	B2473	57.2	18.4	6	21.9	6.1+1					6.1+1													5.3-6	5.3-6	19.9	
	70511	P2404	(6.0)																							4.5-5	5-5.5	7.8
	70511	C2477	49.0	196.2	4.5	20.0	7.0+3	4.0+2				7.0+3	4.0+2												4.5-5	4.5-5	18.5	
	70511	B2477	45.0	179.0	6.3	21.0	1.5+1	2.0+1				1.5+1	2.0+1												4.6-6	4.5-4	16.6	
	70511	H2400	55.1	66.9	4.6	21.4	2.0+1					2.0+1													3.7-9	4.5-4	16.6	
120	70512	H2482	39.9	285.8	5.7	20.8	2.7+1	9.2+0				2.7+1	9.2+0												5-6	6	1.6	
	70512	H2482	40.0	246.0	48.5	20.1	6.1+1	1.1+1				6.1+1	1.1+1												4.5-6	4.5	4.6	
	70512	H2482	38.8	210.9	4.2	20.2	7.0+2	3.0+2				7.0+2	3.0+2												4.5-6	4.5	4.6	
	70512	H2482	49.0	201.8	4.3+4	20.1	1.1+2	3.1+1				1.1+2	3.1+1												4.5	4.7	7.4	
	70512	C2487	59.4	119.9	5.1	21.0	1.2+2					1.2+2													4.5-5	4.5-5	19.1	
	70512	H2488	61.7	72.3	5.9	21.6	1.5+1					1.5+1													5.5-5	5.5-5	16.1	
121	70515	C354	53.8	360.0	6	22.1	1.5+2	1.5+1				1.5+2	1.5+1												5-2-6	6	20.8	
122	70516	H2523	41.1	314.5	53.5	21.1	1.1+3	1.5+2				1.1+3	1.5+2												5.4-4.5	5-5.5	23.5	
123	70517	S2558	64.4	108.0	6.4	21.1	1.5+2	1.5+1				1.5+2	1.5+1												5.6-6.2	6-5	19.9	
124	70518	B2533	51.6	237.1	7.4-7.7	19.8	1.5+1					1.5+1													5.5-8	7.5-8	4.8	
125	70518	C2533	52.5	332	7.7	21.6	6.1+1					6.1+1													6-7	6-7	22.5	
126	70519	B2546	63.9	303	9.0	21.9	3.5+1					3.5+1													6-7	6-7	22.5	
																										6.6	6.6	18.9

S Y E T	DATE	REV	LAT	LONG	L-VAL	MLT	J [°/ω, Δ km]			J (P, Δ km)			800° L-VAL	12	97	L-VAL	10 (RES) UT
							850	11	1.35	1.6	90	110					
127	70520	G2855	50.5	347.2	6.5	21.5	1.5+1	6.5	1.35	1.6	2.5+1	6.7+1	5.5+6	✓	✓	6.5	21.7
128	70524	B2587	55.9	8.6	6.8	21.5	4.6+1	8.3+3			8.3+3	1.7+3	6-6.5				19.9
129	70610	B2722	45.0	234.8	5.3	18.7	7.4+1 4.6+1 1.5+2	8.3+3	1.2+1	NO PA	8.3+3	1.7+3	5.3+6	✓	✓	9.5	30.4
130	70702	G2905	53.1	333.6	7.0	17.5	7.7+0	8.3+3			8.3+3	1.1+3	5.5-7	✓	✓	5-10	4.1
131	70710	H2963	62.4	306.8	56-78	5.7	3.5+2	1.1+1			4.2+3	1.1+3	6-9	✓	✓	5-5	13.7
132	70710	C2968	63.9	74.4	7.0	19.0	1.5+1	8.3+1			8.3+1		5.8-7	✓	✓	5-5	19.7
133	70710	G2970	53.2	338.3	6.7	19.2	2.0+2	7.7+0			8.3+3	2.5+3	5-5.7	✓	✓	4.9-14	7.0
134	70716	A3015	53.4	324.2	5.5	5.5	5.0+3	3.5+4			3.5+4	2.5+3	5-15	✓	✓	4.7-6	12.8
135	70716	H3017	77.6	239	5.8	3.9	3.5+1	6.7+4			6.7+4	8.3+3	5	✓	✓	6-12	13.8
136	70718	G3034	49.0	200.2	7.2	3.2	1.5+1	2.5+3			2.5+3		5.8-7.5	✓	✓	6.2	15.0
137	70720	H3033	57.8	76.6	7.3	4.08	3.1+1	8.3+2			8.3+2		6.2	✓	✓	6.2	18.0
138	70722	H3065	46.1	277.2	6.9	4.3	2.0+1	2.5+3			2.5+3		7	✓	✓	5.5-8	0.8
139	70723	B3074	61.4	49.6	7.0	3.6	6.1+1	4.2+3			4.2+3		5.6-7	✓	✓	5-7	10.0
140	70729	C3124	66.0	5.0	4.5	4.3	3.1+0	8.3+3			8.3+3	2.5+3	7.3	✓	✓	7	2.8
141	70805	H3172	78.5	300.1	5.4	7	1.2+3	1.5+1			1.7+3	8.3+2	7.5-10	✓	✓	4.5	5.8
142	70806	C3184	63.0	44.5	7.8	4.5	1.2+2	1.7+3			1.7+3	8.3+2	7.8	✓	✓	4.7	11.2
70806	C3184	43.5	325.0	6.7-8	4.3	1.5+2	1.5+1	2.5+3			2.5+3	8.3+2	7.8	✓	✓	7.8	4.8
70806	C3184	41.8	285.6	6	3.6	1.1+2		6.7+4			6.7+4	8.3+3	5.7-8	✓	✓	5-10	6.1
70806	C3184	46.5	238.7	6.7	2.6	4.6+2	3.1+1	5.8+1			5.8+1	1.7+3	6-7	✓	✓	6-7	8.9
70806	B3184	48.9	186.0	7	2.1	3.1+1	1.5+1	1.7+4			1.7+4	4.5+3	6-7	✓	✓	6-7.5	11.9

B Y E N T	DATE	REF	LAT	LONG	L-VAL	MIT	j [°/sec. x sec.]			T [°/sec. x sec.]			10	33	L-VAL	
							6.5	8.5	1.1	1.5	1.6	80				
M3	70807	B3196	330	8270	7-10	4.4	1.5+1			2.5+3	6.5-7	✓	✓	6-11	5-7	
	70807	H3194	440	2944	6-7	3.6	1.5+2			1.7+4	6-7	✓	✓	6-7	8-V	
	70807	C3196	418	2988	6-5	2.8	1.5+2			8.3+4	5.5-7	✓	✓	11.4	14.6	
	70807	C3197	589	1930	7	2.1	1.2+2			4.2+4	5.9-7	✓	✓	6-8	17.6	
	70807	B3198	647	1443	7-9	2.5	1.5+2			1.7+4	6-8	✓	✓	6-8	8.0	
M4	70808	B3202	484	2972	6-8	3.7	1.2+2			8.3+3	6-6.5	✓	✓	5.5-7	4.2	
	70810	P3218	495	3510	6	4.5	1.5+2	1.2+1		4.2+3	5.5	✓	✓	5.5-8	19.1	
M6	70810	P3223	633	1207	5.7-8	3.0	1.5+2	1.2+1		2.5+4	5.5-8	✓	✓	5.5-8	11.0	
	70814	H3154	410	2479	6-9	2.7	1.4+2	1.5+1		1.7+4	5.3-6.0	✓	✓	7	17.2	
M8	70814	B3254	629	1467	7	2.2	1.1+1			2.8+4	6-7	✓	✓	6	7.6	
	70815	C3260	478	2980	6-9	3.5	1.1+2	9.2+0		8.3+3	5.5-6	✓	✓	6	9.1	
M9	70818	C3185	493	2142	6-7	2.9	1.6+1			8.3+3	6-6.6	✓	✓	7	12.1	
	70818	C3186	470	2210	5-8	2.5	1.5+1			1.7+3	6-7	✓	✓	5.5-8	15.4	
	70818	B3285	430	2211	6-7	2.6	1.6+1			1.7+4	6-8	✓	✓	6-8	21.2	
	70818	H3294	643	836	6-5	3.4	1.5+1			8.3+3	6-6.7	✓	✓	6.5	7.8	
150	70819	C3204	478	210	8	4.3	1.5+1			4.2+3	8	✓	✓	6	8.6	
	70819	C3194	425	2809	6-3	3.0	1.6+1			8.3+3	6-6.4	✓	✓	6	20.7	
	70819	C3214	666	879	6.5-8	1.2	1.4+1	4.6+0		8.3+3	5.5-6	✓	✓	6.5-8	11.2	
151	70820	H3300	510	339	1.7-8	2.0	1.4+2	1.5+1		2.5+3	6.5	✓	✓	7.5	11.1	
	70822	B3353	435	2380	5.5	2.1	1.6+1	1.3+1		4.3+3	5.5-5	✓	✓	4.5-7	13.6	
153	70820	H3360	562	2010	6.3	1.4	1.5+1			8.3+3	6	✓	✓	7	9.6	
	70823	H3370	472	2175	7	2.1	1.5+1			1.7+3	5.5-7.8	✓	✓	6.3	9.1	
155	70830	H3354	440	2651	6.8	2.4	1.2+0			1.7+3	6.4	✓	✓	6.8	22.5	
	70832	C3413	603	56	5.5	2.8	1.2+0			1.7+4	6	✓	✓	6		

S U B T	DATE	REV	LAT	LONG	L-VAL	MLT	Q35	V35	G55	j (e ₀ to e ₁ h ₀)			J (P ₀ to P ₁)			L-VAL	HCS OT
										850	1.1	1.35	1.6	80	150		
157	70709	I3469	64.3	57.8	6.6	2.4	1.5+1							6	6-6.5	21.9	
	70709	C3472	60.8	343.5	6.9	3.3	3.1+1							6.9	5.5-6	9.4	
158	70910	H3473	4.5	253.5	5.8-6	1.8	4.6+1							6	6	21.4	
	70910	B3475	12.0	65.9	5.9	2.3	1.5+2	4.6+0						6	6	21.4	
159	70910	P3476	63.2	65.9	5.9	2.3	1.5+2	4.6+0						6	6	21.4	
160	70711	B3478	41.7	262.2	5.5	1.95	1.1+2	6.1+0						5-5.5	5.5	8.8	
	70711	C3482	41.6	211.4	5.6-6	1.1	1.4+2	1.2+1						5-6	5.5	11.9	
	70914	H3483	57.7	163.9	5	1.4	1.4+2	3.1+1						4.5-5	4.5-5	14.9	
	70711	C3484	67.5	110.5	5.1-12	1.7	4.6+1	6.1+0						4.5-18	5-18	18.0	
161	70712	B3486	4.3	79.3	6	2	1.5+1							6	6	20.1	
162	70713	B3500	46.9	229.9	5.0	1.3	2.2+2	3.1+2	1.6+2	7.1+1				4.8-5	4.7-5.4	10.8	
	70918	B3500	67.8	21.7	4.6-7	1.2	2.2+2							7	7	11.9	
163	70715	H3515	48.4	243.7	6.1	1.3	2.6+1	1.5+1						5.5-15	5-11	9.8	
	70915	H3515	55.5	38.6	5.9-7	11.5	2.2+2							5.5-15	6-10	10.9	
164	70918	P3528	48.2	207.8	6.9	2.6	3.1+1							6	6.9	5.2	
165	70719	C3544	46.0	274.3	6.6	1.8	1.5+1							6.6	6.6	7.6	
166	70919	I3550	74.8	92.0	10-10.6	1.2	7.7+0							6.2-10	6.2-9	12.6	
167	70920	P3553	43.8	326.8	4.5-5	2.7	6.1+2	6.1+1	8.2+0					7.3-5	4	4.1	
	70920	P3553	66.2	330.2	4.5-5	2.8	3.1+2	3.1+1	4.9+0					4-6	4-6	5.5	
	70920	C3554	69.9	290.5	5.1	4.2	2.0+2	1.4+1						5	5	8.4	
	70920	H3555	78.0	10.0	5-6	4.9	1.5+3							6	4-10	11.8	
	70920	B3556	63.8	182.7	6.5-8.5	0.1	3.1+1							4.2-5	7	13.1	
	70920	C3557	51.5	188.1	5.2	4.4	4.6+2							4.2-5	4.5-5.5	14.4	
	70920	C3557	57.1	142.1	4.3-8.6	1.3	6.1+2	4.6+1	6.5+0					4.5-14	4.1-4.8	16.0	
	70920	C3557	45.3	141.1	5.5-6	3.4	1.2+3							5-6	5-6	17.5	
	70720	I3559	46.6	97.7	5.4	1.7	3.1+1							4.5-11	5-12	22.0	
	70920	P3559	72.5	31.9	7.4	1.9	1.5+1							6.9	5-6	22.0	
	70920	P3559	53.7	56.6	6.9	1.9	1.5+1							5.5-11	5-6	23.2	

DATE	REV	LAT	LONG	L VAL	MLT	j [°/min. x sec]				j [°/min. x sec]				1/2	80% L VAL	1/2	3/3	2-VAL	MKS UT
						435	460	65	850	1.1	1.35	1.6	80						
168	70921	035614	52.5	330.055-12	26	1.512	4640											5-16	5.7
	70921	035614	52.9	323.1 5.4	26	6.141	3.111											5-7	3.8
	70921	035614	53.5	289.8 6.2	4.1	6.141												5-1-6	6.5
	70921	035614	54.1	516.8 6.284	4.2	12+2												5-8	7.8
	70921	035614	54.7	295.0 5.975	5.5	7.7+3												5-8	7.8
	70921	035614	55.4	273.6 9.7	9.4	1.5+2												5-8	7.8
	70921	035614	56.1	146.4 5.1	1.0	1.1+2	3.1+1											5-8	7.8
	70921	035614	56.8	86 6-20	3.0	2.0+2												5-8	7.8
	70922	035614	57.5	336.1 6.3	2.6	6.1+3	1.5+2	7.1+0										5-8	7.8
	70922	035614	58.2	182.7 7.0	14.6	1.5+2												5-8	7.8
	70922	035614	58.9	143.5 4.8	14.5	1.5+3												5-8	7.8
	70922	035614	59.6	297.7 5.4	2.1	1.5+2	1.5+1											5-8	7.8
	70922	035614	60.3	311.0 6.3	3.5	1.5+1												5-8	7.8
	70922	035614	61.0	156.5 4.8	9.7	3.5+2	6.1+1	1.6+1										5-8	7.8
	70922	035614	61.7	285.0 6.8-7	11.8	1.4+2												5-8	7.8
	70922	035614	62.4	168.0 1.5	5.6	2.0+1												5-8	7.8
169	70923	035804	61.7	207.1 5.7	.6	1.4+2	6.1+1	1.6+1										5-8	7.8
	70923	035804	62.4	234.6 8.9	6.8	2.0+1												5-8	7.8
	70923	035804	63.1	160.1 5.3+3	.6	3.1+2	4.6+1	1.8+0										5-8	7.8
	70923	035804	63.8	164.1 5.1	3.7	1.1+4												5-8	7.8
	70923	035804	64.5	112.1 6.5	1.0	4.6+1												5-8	7.8
	70923	035804	65.2	118.7 5.8	2.8	6.1+1	1.2+1											5-8	7.8
	70923	035804	65.9	387.9 5.9	2.0	3.1+1												5-8	7.8
170	70924	035804	66.6	165.6 6.7-7	.2	2.0+1												5-8	7.8
	70924	035804	67.3	105 6-12	.5	6.1+1												5-8	7.8
	70924	035804	68.0	355.7 6.2	2.5	6.1+1	7.7+0											5-8	7.8
	70924	035804	68.7	125.7 6.9	.6	4.6+1												5-8	7.8
	70924	035804	69.4	310.7 7.9	3.4	4.6+1												5-8	7.8

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Appendix D - Computation of Electron and X-ray Energy Deposition Profiles

The precipitation of energetic electrons into the atmosphere is simulated by a finite difference numerical solution to the multi-dimensional Fokker-Planck diffusion equation developed by Walt et al. (1968). The technique was slightly modified to handle other than isotropic pitch angle distributions by Spjeldvik (1974) but this modification was not used here since the distribution over the downward hemisphere at 300 km is isotropic for all cases considered. The conversion of the Walt et al. (1968) program to the UCLA computer by Spjeldvik (1974), however, was used. Spjeldvik (1974) contains the computer listing.

The Walt program solves the following final version of the Fokker-Planck equation as derived by Chappell (1968):

$$\cos \alpha \frac{\partial}{\partial s} \left(\frac{f}{r} \right) = -\frac{\partial}{\partial \alpha} \frac{\sin \alpha}{R_s L} \frac{\partial}{\partial \alpha} \left(\frac{f}{r} \right) + \frac{(E+1)^2 2\pi}{E^2 (E+2)^2} \left(\frac{q^2}{m_e c^2} \right)^2 \sum_i m_i Z_i^2 \ln \eta_i^{-1} \frac{1}{\sin \alpha} \frac{\partial}{\partial \alpha} \left[\sin \alpha \frac{\partial}{\partial \alpha} \left(\frac{f}{r} \right) \right] + \frac{2(E+1)}{E^{1/2} (E+2)^{1/2}} \frac{\partial}{\partial E} \left[\frac{(E+1) 2\pi}{E^{1/2} (E+2)^{1/2}} \left(\frac{q^2}{m_e c^2} \right)^2 \sum_i m_i Z_i \ln \sigma_i \left(\frac{f}{r} \right) \right],$$

where Table D-1 defines the symbolism. Notice that the stochastic averages $\langle (\Delta E)^2 \rangle$, $\langle (\Delta \alpha \Delta E) \rangle$, and $\langle (\Delta E \Delta r) \rangle$ (small) and higher order terms have been eliminated. Steady state ($\frac{\partial f}{\partial t} = 0$) is assumed since precipitation is steady on times scales much longer than atmospheric transit times for precipitating electrons. Also, radial

Table D-1 Nomenclature (Spjeldvik, 1974)

f = electron distribution function averaged over cyclotron phase
 t = time
 c = speed of light
 R_E = radius of the Earth
 L = L-value, field line equatorial distance in units of R_E
 E = electron energy in rest mass units
 α = local pitch angle
 α_0 = equatorial pitch angle
 r = electron gyro radius
 s = length along the field line
 q = electron charge
 m_e = electron mass
 e = the natural logarithm base, 2.7182...
 n_i = number density of atmospheric species i
 Z_i = atom number of atmospheric species i
 l_i = average excitation potential of atmospheric species i
 η_i = minimum scattering angle corresponding to maximum impact parameter (for thermal plasma, the Debye-length; for the neutral atmosphere, the atomic radius)
 $\sigma_i = (e (E+2)/l_i)^{1/2}$
 $\langle \rangle$ denotes stochastic ensemble average.

diffusion is ignored ($\frac{f}{r} = \text{constant}$).

The program outputs the amount of average energy deposited per unit volume at each altitude below the incident altitude. Also, assuming an average of 35 eV is required to produce each ion pair, the rate of ion pair production is output versus altitude.

The complete altitude profile of energy deposition by bremsstrahlung is computed from the Berger, Seltzer, Maeda (1974) Monte Carlo simulation method. This work is essentially an extension of the earlier (Berger, Seltzer, Maeda, 1970; Berger and Seltzer, 1972) Monte Carlo simulation methods by this same group. Basically, a computation is made of 10,000 electron trajectories for each incident electron energy from 2 keV to 2 MeV for electrons injected into the atmosphere at 300 km altitude. The trajectories are histories of individual primary and secondary electrons as they coulomb scatter (elastic by nuclei collisions, inelastic by orbital electron collisions) off neutral constituents and deposit energy at various altitudes. In addition, the production of bremsstrahlung quanta in these interactions and their Compton scattering and consequent photoelectron absorption of the bremsstrahlung quanta in the atmosphere is followed. The Berger, Seltzer, Maeda (1974) results of energy deposition as a function of atmospheric depth (mass thickness) by bremsstrahlung from monoenergetic incident electron beams for wide-area precipitation flux isotropic over the downward hemisphere is given in Table D-2 in the form of energy deposited per unit mass thickness (ev/g cm^{-2}) per incident electron energy. From this

TABLE D-2. Energy deposition function for Bremsstrahlung, for the case of uniform wide-area precipitation of an electron flux isotropic over the downward hemisphere. The quantity given is $A_{BR} (Z_m/T_0)$, in units of $\text{cm}^2\text{gm}^{-1}$. Number in parenthesis indicates powers of ten. (Berger, Seltzer, Maeda, 1974)

$\frac{z}{T_e}$ ($\text{g cm}^{-2} \text{ kV}^{-1}$)	2000	1000	500	200	100	50	20	10	5	2
	$A_{\text{br}}(z)/T_e \text{ cm}^2/\text{g}$									
2.0(-6)	6.9(-4)	1.0(-3)	2.2(-3)	5.0(-3)	1.1(-2)	7.3(-3)	3.4(-3)	2.0(-3)	7.2(-4)	8.6(-6)
4.0	7.6	1.1	2.4	5.3	1.7	2.2	1.7	1.7	2.7	1.4
8.0	8.4	1.2	2.5	5.0(-3)	7.0(-3)	1.1	1.7	2.0(-3)	8.8(-4)	7.2(-4)
2.0(-5)	9.3	1.3	2.6	5.0	1.7	2.2	1.7	2.0(-3)	8.8(-4)	8.6(-6)
4.0	9.9	1.4	2.6	5.0	1.7	2.2	1.7	2.0(-3)	8.8(-4)	8.6(-6)
8.0	1.1(-3)	1.4	2.6	5.0	1.7	2.2	1.7	2.0(-3)	8.8(-4)	8.6(-6)
2.0(-4)	1.1	1.2	1.0	1.8	1.7	2.2	1.7	2.0(-3)	8.8(-4)	8.6(-6)
4.0	6.8(-4)	5.5(-4)	5.2(-4)	6.5(-4)	8.4(-4)	1.1	1.7	2.0(-3)	8.8(-4)	8.6(-6)
8.0	4.5	3.2	2.8	3.3	4.4	1.1	1.7	2.0(-3)	8.8(-4)	8.6(-6)
2.0(-3)	2.9	1.0	1.5	1.5	1.9	2.3	2.0	2.0(-3)	8.8(-4)	8.6(-6)
4.0	1.0	1.3	0.9(-5)	8.8(-6)	1.0	1.1	1.3	2.0(-3)	8.8(-4)	8.6(-6)
8.0	9.2(-5)	8.3(-5)	0.5	5.2	5.4(-5)	5.4(-5)	6.7(-5)	2.8	3.3(-5)	5.1(-7)
2.0(-2)	1.7	2.8	2.8	2.4	2.3	2.0	1.6	3.0(-6)	5.2(-6)	5.1(-7)
4.0	1.3(-6)	5.5(-6)	0.5(-6)	1.1	1.1	2.0	1.6	3.0(-6)	5.2(-6)	5.1(-7)
8.0	2.4(-7)	2.4(-7)	1.5	3.8(-6)	3.0(-6)	3.3(-6)	3.3(-6)	3.9(-7)	3.9(-7)	3.9(-7)
2.0(-1)	6.0(-9)	6.0(-9)	6.0(-9)	1.7(-7)	5.4(-7)	3.1(-7)	3.0(-8)	3.9(-8)	3.9(-8)	3.9(-8)
4.0	5.0(-8)	5.0(-8)	5.0(-8)	5.0(-8)	2.4(-8)	2.8(-8)	2.8(-8)	2.8(-8)	2.8(-8)	2.8(-8)

starting point, we proceeded to work towards the final product, a profile of ion pair production by bremsstrahlung from an incident electron energy spectrum.

The CIRA (1965) reference atmosphere, Fig. D-1, was used to scale atmospheric depth to a representative altitude. The same reference atmosphere was used to compute the atmospheric density (g/cm^3) versus altitude (km). At each altitude the density was multiplied by the energy deposition function $A_{\text{BR}}[\text{keV/g cm}^{-2}]$, to yield the stopping ratio, $\text{BR}[\text{keV/cm, column}]$, normalized to an incident electron of a given incident energy (T_0). The resulting column stopping ratios, $\text{Br}(\text{altitude})$, are listed in Table D-3 and graphically represented in Fig. D-2. Due to the irregularity of the curves and diminishing contribution compared to electron production above 50 km in Fig. D-2, only values below 50 km were considered. For every 5 km of altitude between 15 and 50 km (inclusive), a linear (log-log) interpolation of incident electron energy (T_0) versus column stopping ratio of Bremsstrahlung photons was constructed. For example, at 15 km altitude, $\text{Br} = 10^x$ where $x = 1.018342561 \log T_0(\text{keV}) - 11.85638142$ for initial energies T_0 between 200-500 keV.

A computer program using the linear (log-log) interpolation curves as described above was written to calculate the production of ion pairs versus altitude for an input electron differential energy spectrum isotropic over downward hemisphere at 300 km by solving the equation

$$\text{Prod}_{\text{IP}}(\text{alt}) = \frac{1}{\text{EMI}} \iint \text{Br} \cdot \text{Dif}(E) \cdot d\Omega dE$$

FIG D-1. CIRA Reference Atmosphere of depth versus altitude (1965).

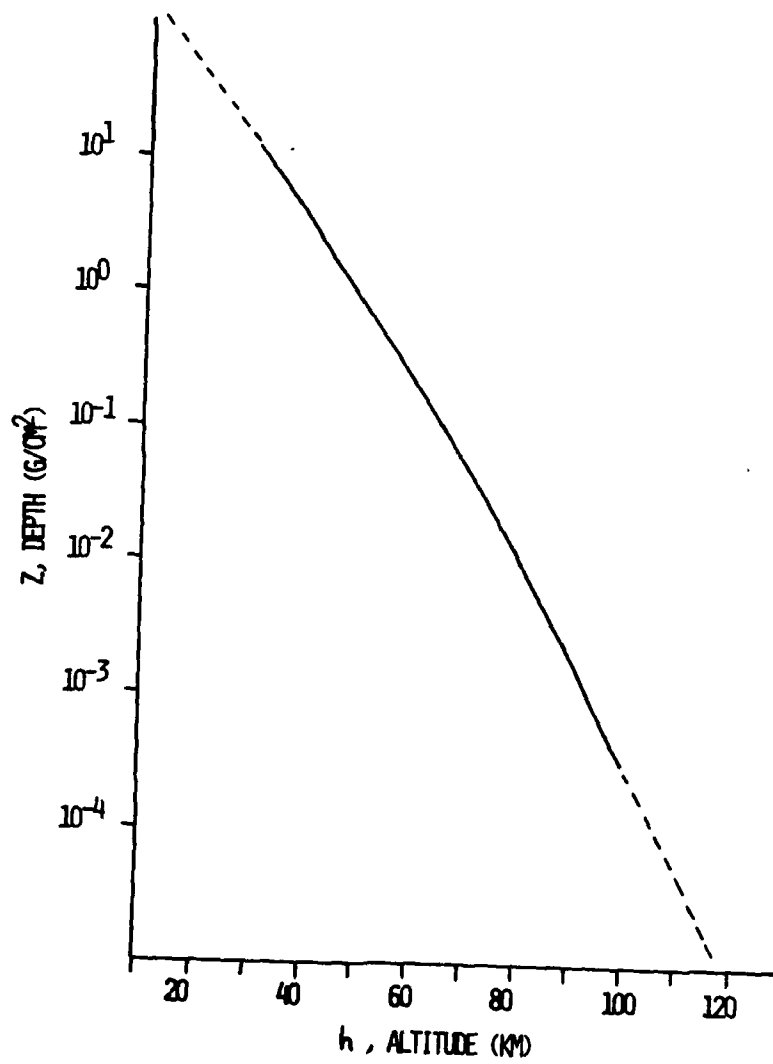


TABLE D-3. B_r (keV/cm, column), stopping ratio, versus altitude for given incident electron energies.

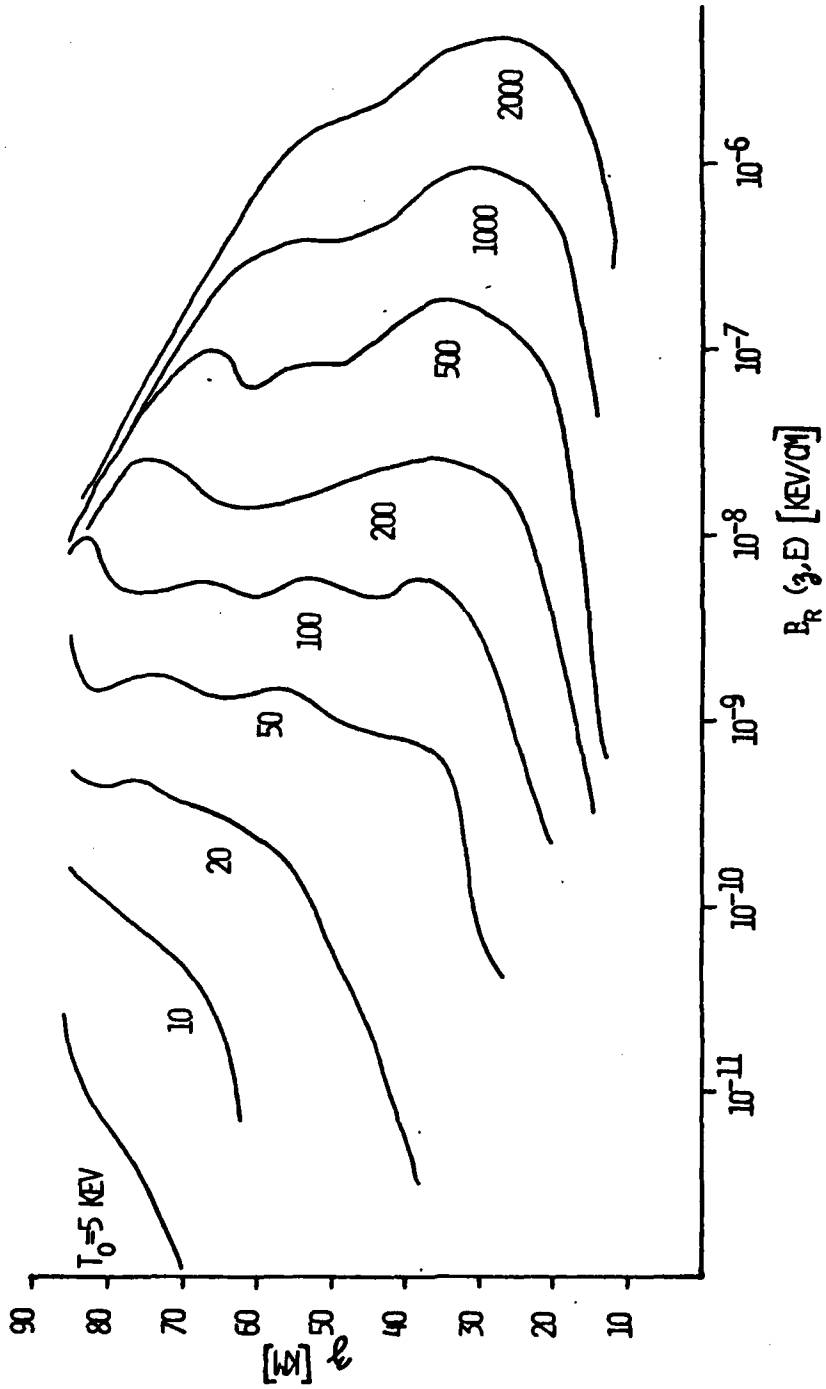
T₀ (KEV)

	2500	1000	500	200	100	50	20	10	5	Z
86	8.9 (-9)	6.3 (-9)	6.9 (-9)	6.3 (-9)	6.9 (-9)	2.3 (-9)	4.8 (-10)	1.2 (-10)	2.3 (-11)	1.1 (-13)
82	1.9 (-8)	1.4 (-8)	1.5 (-8)	1.4 (-8)	1.0 (-8)	1.4 (-9)	4.3 (-10)	1.1 (-10)	8.9 (-12)	1.3 (-14)
77	4.2 (-8)	3.0 (-8)	3.1 (-8)	2.5 (-8)	4.3 (-9)	1.4 (-9)	4.0 (-10)	6.8 (-10)	4.1 (-12)	
72	1.2 (-7)	8.1 (-8)	7.8 (-8)	2.2 (-8)	5.2 (-9)	1.8 (-9)	3.7 (-10)	5.8 (-11)	1.6 (-12)	
67	2.5 (-7)	1.8 (-7)	1.0 (-7)	1.6 (-8)	5.5 (-9)	1.4 (-9)	3.3 (-10)	3.5 (-10)	6.9 (-11)	
62	5.4 (-7)	2.9 (-7)	6.3 (-8)	1.6 (-8)	4.6 (-9)	1.3 (-9)	2.8 (-10)	7.3 (-12)		
56	1.1 (-6)	2.8 (-7)	7.2 (-8)	1.6 (-8)	5.2 (-9)	1.4 (-9)	1.5 (-10)	2.0 (-12)		
50	1.4 (-6)	2.3 (-7)	7.8 (-8)	1.8 (-8)	5.6 (-9)	1.0 (-9)	6.9 (-11)	4.1 (-13)		
45	1.8 (-6)	3.9 (-7)	1.0 (-7)	2.1 (-8)	4.7 (-9)	8.5 (-10)	2.3 (-11)			
38	3.2 (-6)	7.1 (-7)	1.8 (-7)	2.6 (-8)	6.0 (-9)	7.1 (-10)	3.2 (-12)			
33	1.4 (-6)	9.5 (-7)	1.6 (-7)	2.5 (-8)	4.5 (-9)	1.8 (-10)				
27	5.2 (-6)	7.9 (-7)	1.3 (-7)	2.1 (-8)	1.5 (-9)	4.0 (-11)				
20	3.2 (-6)	5.8 (-7)	7.1 (-8)	3.2 (-9)	2.3 (-10)					
14	6.2 (-7)	6.2 (-8)	7.7 (-10)	3.0 (-10)						

B_r (KEV/CM, COLUIN)

(M) 4

FIG D-2. $B_r(z,E)$ versus altitude.



where EMI = average minimum ionizing energy taken here to be 35 eV.

Dif (E) = differential electron flux at 300 km

[#/cm² sec str keV]

dΩ = solid angle ≈ 2π here

dE = 1 keV intervals

Prod (alt) = ion pair production rate at each altitude

[#/cm³ sec]

Also output is the total energy (keV) deposited at each altitude per (cm³ str sec). The computer listing follows:

TABLE D-4. Bremsstrahlung deposition computer listing.

```

C *****
C THIS PROGRAM CALCULATES THE ALTITUDE VS ENERGY DEPOSITION PROFILE FOR A
C GIVEN PRECIPITATING ELECTRON SPECTRUM. THE ENERGY DEPOSITION FUNCTION,
C BR, IS OBTAINED FROM THE BERGER, SELTZER, AND MANDA (JATP, 1978, 36, 591)
C FULL-SCALE CARDO CALCULATIONS FOR UNIFORM WIDE AREA PRECIPITATION OF AN
C ELECTRON FLU ISOTROPIC OVER THE DOWNWARD HEMI SPHERE. ATMO DEPTH VS ALT
C IS OBTAINED FROM CIRAS, 1965 (MEAN STRUCTURE).
C *****
C DIMENSION E(2070), DIF(2000)
C INPUT - NUMBER OF DATA SETS .
C READ (5,2) NDATA
C 2 FORMAT (I2)
C INPUT - DIFFERENTIAL ELECTRON FLUX (#/LL TSE) 1.0, 20, 50, 100, 200, 500.
C 10 FORMAT (8E10.5)
C 1 READ (5,10) DIF(1), DIF(11), DIF(41), DIF( 91), DIF(191), DIF(491)
C 10 FORMAT (8E10.5)
C ENERGY, E, FROM 10 TO 2000 KEV
C DD 1) J=1,1991
C 11 E(J)= 10.* (J-1)
C DIFFERENTIAL FLUX SPECTRUM IS GENERATED BY LINEAR FIT (DN-LOG-LOG BLDT)
C BETWEEN INPUT VALUES FROM ABOVE.
C 00 30 I=1,1991
C IF (I.GT.931.AND.I.LT.1981) GO TO 61
C IF (I.GT.491.AND.I.LT. 991) GO TO 59
C IF (I.GT.191.AND.I.LT. 491) GO TO 57
C IF (I.GT. 31.AND.I.LT. 191) GO TO 55
C IF (I.GT. 11.AND.I.LT.  31) GO TO 53
C IF (I.GT.  1.AND.I.LT.  11) GO TO 51
C IF (I.GT.  1.AND.I.LT.  11) GO TO 47
C GO TO 30
C 47 IF (DIF(I).NE.DIF(11)) GO TO 48
C DIF(I)=DIF(11)
C GO TO 80
C 48 YARG=ALOG10((DIF(11))/DIF(I))
C YARG=ALOG10(E*(DIF(11)/DIF(I)))
C Y=ALOG10(DIF(I))
C X=ALOG10(E(I))
C GO TO 99
C 91 IF (DIF(11).NE.DIF(111)) GO TO 88
C DIF(11)=DIF(111)
C GO TO 50
C 88 YARG=ALOG10((DIF(111))/DIF(11))
C YARG=ALOG10(E*(DIF(111)/DIF(11)))

```

```

Y=ALC310(DIF(11))
X=ALC310(E(11))
GO TO 92
93 IF (DIF(41).NE.DIF(91)) GO TO 54
DIF(11)=DIF(41)
GO TO 80
94 YARG=ALOG10((DIF(91))/(DIF(41)))
XARG=ALOG10((E(91))/(E(41)))
Y=ALC310(DIF(41))
X=ALC310(E(41))
GO TO 99
95 IF (DIF(91).NE.DIF(191)) GO TO 56
DIF(11)=DIF(91)
GO TO 80
96 YARG=ALOG10((DIF(191))/(DIF(91)))
XARG=ALOG10((E(191))/(E(91)))
Y=ALC310(DIF(91))
X=ALC310(E(91))
GO TO 99
97 IF (DIF(191).NE.DIF(491)) GO TO 58
DIF(11)=DIF(191)
GO TO 80
98 YARG=ALOG10((DIF(491))/(DIF(191)))
XARG=ALOG10((E(491))/(E(191)))
Y=ALC310(DIF(191))
X=ALC310(E(191))
GO TO 99
99 IF (DIF(491).NE.DIF(991)) GO TO 60
DIF(11)=DIF(491)
GO TO 80
100 YARG=ALOG10((DIF(991))/(DIF(491)))
XARG=ALOG10((E(991))/(E(491)))
Y=ALC310(DIF(491))
X=ALC310(E(491))
GO TO 99
101 IF (DIF(991).NE.DIF(1991)) GO TO 62
DIF(11)=DIF(991)
GO TO 80
102 YARG=ALOG10((DIF(1991))/(DIF(991)))
XARG=ALOG10((E(1991))/(E(991)))
Y=ALC310(DIF(991))
X=ALC310(E(991))
GO TO 99
103 YARG=ALOG10((YARG/XARG)**(Z-X))
DIF(11)=10**((YARG/XARG)**(Z-X))
99 CONTINUE
C
PERFORM INTEGRAL OF BR-DIP*DE AT EACH 8 KM ALTITUDE FROM 15 TO 80 KM.
WRITE(6,13)
13 FORMATT(1,10,=MULTIPLY--DE*ANGLE--152--ARCOS--1.51--ANGLE--01--04
14 ALLST:01,4(1/ALLT))
15
100 IF (Z.50.15.) GO TO 200
15 11-50-20.) GO TO 200

```

```

108 IF (Z.EQ.25.) GO TO 208
109 IF (Z.EQ.30.) GO TO 188
110 IF (Z.EQ.35.) GO TO 165
111 IF (Z.EQ.40.) GO TO 145
112 IF (Z.EQ.45.) GO TO 125
113 TOTAL=0.0
114 DO 73 I=1,1391
115 IF (I.GT.91.) GO TO 110
116 X=2.736669403*ALOG10(E(I))-13.72349432
117 GO TO 120
118 IF (I.GT.491.) GO TO 115
119 Y=1.534787611*ALOG10(E(I))-11.51973073
120 GO TO 120
121 Z=3.333366123*ALOG10(E(I))-13.74010285
122 BR=10.
123 XRAY=)IF(I)BR
124 TOTAL=TOTAL+XRAY
125 CONTINUE
126 GO TO 900
127 TOTAL=2.0
128 DO 71 I=1,1991
129 IF (I.GT.41.) GO TO 130
130 Y=1.923061797*ALOG10(E(I))-15.73033048
131 GO TO 140
132 IF (I.GT.191.) GO TO 135
133 X=2.323664191*ALOG10(E(I))-13.02100080
134 GO TO 140
135 Y=1.93720917*ALOG10(E(I))-12.1317653
136 BR=1.
137 XRAY=)E(I)BR
138 TOTAL=TOTAL+XRAY
139 CONTINUE
140 GO TO 900
141 TOTAL=0.0
142 DO 72 I=1,1991
143 IF (I.GT.41.) GO TO 180
144 Y=5.331193138*ALOG10(E(I))-18.13209481
145 GO TO 160
146 IF (I.GT.91.) GO TO 155
147 X=2.97532307*ALOG10(E(I))-14.0683287
148 GO TO 160
149 X=2.72102793*ALOG10(E(I))-13.30446827
150 BR=10.
151 XRAY=)IF(I)BR
152 TOTAL=TOTAL+XRAY
153 CONTINUE
154 GO TO 900
155 TOTAL=0.0
156 DO 73 I=1,1931
157 IF (I.GT.91.) GO TO 170
158 Y=3.086467478*ALOG10(E(I))-16.08185425
159 GO TO 180
160 IF (I.GT.131.) GO TO 175
161 X=3.174702312*ALOG10(E(I))-13.06232314

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178 X= 2.126634192 * ALOG10(E(I)) - 12.64791195
180 RM=10
XRAY=OIF(I)*BR
TOTAL=TOTAL+XRAY
73 CONTINUE
GO TO 900
185 TOTAL=0.0
DO 74 I=1,1991
IF (I.GT.91.) GO TO 187
X= 9.78727076 * ALOG10(E(I)) - 18.09799954
GO TO 200
187 IF (I.GT.191.) GO TO 190
X= 2.96048826 * ALOG10(E(I)) - 14.44437471
GO TO 200
190 IF (I.GT.491.) GO TO 195
X= 2.09319926 * ALOG10(E(I)) - 12.25577687
GO TO 200
195 X= 2.510322759 * ALOG10(E(I)) - 13.60796849
200 RM=10
XRAY=OIF(I)*BR
TOTAL=TOTAL+XRAY
74 CONTINUE
GO TO 900
205 TOTAL=0.0
DO 75 I=1,1991
IF (I.GT.191.) GO TO 210
X= 3.91528047 * ALOG10(E(I)) - 16.57187825
GO TO 220
210 IF (I.GT.491.) GO TO 215
X= 2.110124616 * ALOG10(E(I)) - 12.64594502
GO TO 220
215 X= 2.683161107 * ALOG10(E(I)) - 14.19255332
220 RM=10
XRAY=OIF(I)*BR
TOTAL=TOTAL+XRAY
75 CONTINUE
GO TO 900
225 TOTAL=0.0
DO 76 I=1,1991
IF (I.GT.491.) GO TO 230
X= 3.358323442 * ALOG10(E(I)) - 16.78530729
GO TO 235
230 X= 2.75125017 * ALOG10(E(I)) - 14.8773525
235 RM=10
XRAY=OIF(I)*BR
TOTAL=TOTAL+XRAY
76 CONTINUE
GO TO 900
240 TOTAL=0.0
DO 77 I=1,1991
IF (I.GT.491.) GO TO 245
X= 1.618342361 * ALOG10(E(I)) - 11.88638142
GO TO 255

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OUT PUT EXAMPLE

ALTITUDE (KM)	FLUY (E/LLST)	PROD (1/LLT)
50.0	.628195E-01	11.28
48.0	.463822E-01	8.33
40.0	.412F86E-01	7.41
35.0	.179990E-01	3.23
30.0	.85377E-02	1.53
26.0	.012803E-01	0.16
20.0	.170397E-03	0.00
15.0	.043316E-06	0.00

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