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**The John A. Blume Earthquake Engineering Center**

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**SEISMIC HAZARD MAPPING  
FOR GUATEMALA**

by

**Anne S. Kiremidjian**

**Haresh C. Shah**

**Lester Lubetkin**



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## PREFACE

This is the first report of the study on Seismic Risk for Guatemala. This phase of the study includes the collection and description of geological and seismological environment and the use of that information in developing iso-acceleration maps of Guatemala. This phase of the study gives results on future probable seismic hazard for Guatemala which can be utilized by engineers in formulating the seismic codes, by planners in future developments and by the insurance companies in evaluating the seismic risks and rates. Various state-of-the-art models are used to develop this hazard information.

Phase II of the project will involve the following topics.

- A thorough study of the 1973 UBC, the 1976 UBC and the ATC-3 work currently being completed. The above codes will be evaluated with the seismicity of Guatemala in mind. Such an evaluation will permit the engineers and planners in Guatemala to appreciate and understand the current code levels and their relationships with future seismic loading demands.
- A detailed discussion on the purpose and effects of earthquake codes will be presented. This will include the history of earthquake loading criteria, the relation of design loads and quality of structures, the objectives and qualities of workable seismic codes and the role of design detailing and design forces in providing a safe economic construction.

- Introduction to the Proposed Seismic Design Provisions.  
This will take into account the current codes, their advantages, their shortcomings and the available solutions to eliminate or reduce the shortcomings. This part of the study will explain in detail the concept of acceptable risk and the associated loading levels from the hazard maps of Guatemala. The study will also develop the shape and levels of various design spectra. These spectra will include the site characteristics information.
- A detailed description of the type of structural systems, their effects on the design level and a step-by-step design procedure will be developed.
- Based on the spectral approach of seismic design, a simplified equivalent static load method, similar to the 1976 UBC method of design, will be developed. It should be emphasized here that a workable code should have the following four ingredients.
  - 1) Simplicity.
  - 2) Rationality.
  - 3) Freedom to use responsible ingenuity for special structures.
  - 4) Reward and encouragement for using dynamic analysis when merited by the complexity of a given structure.In developing the methodology in part II of this study, the above four ingredients will be kept in mind.
- Finally, a detailed comparison with the proposed methodology and the 1976 UBC will be made. This will be done with the seismic environment of Guatemala in mind.

It is estimated at this time that the results of phase II study will be submitted before the end of September 1977.



## TABLE OF CONTENTS

|  | <u>Page</u> |
|--|-------------|
| ACKNOWLEDGMENTS . . . . .  | ii          |
| PREFACE . . . . .  | iii         |
| Chapter I. INTRODUCTION . . . . .  | 1           |
| I-1 Introduction . . . . .   | 1           |
| I-2 Some Basic Concepts Concerning Hazard and Risk . . . . .                 | 3           |
| Chapter II. GEOLOGIC SETTING OF GUATEMALA . . . . .                          | 7           |
| II-1 Introduction . . . . .  | 7           |
| II-2 Nuclear Central America . . . . .                                       | 9           |
| II-3 Volcanic Activity . . . . .   | 17          |
| II-4 Faulting . . . . .  | 18          |
| Chapter III. SEISMIC DATA BASE . . . . .                                     | 29          |
| III-1 Introduction . . . . .   | 29          |
| III-2 Data Analysis . . . . .  | 31          |
| III-3 Source Location and Seismicity . . . . .                               | 35          |
| III-4 Confidence Levels . . . . .  | 54          |
| III-5 Conclusion . . . . .   | 56          |
| Chapter IV. PROBABILISTIC SEISMIC LOADING . . . . .                          | 57          |
| IV-1 Introduction . . . . .  | 57          |
| IV-2 Iso-Acceleration Maps for Guatemala . . . . .                           | 58          |
| Chapter V. SEISMIC RISK ZONING . . . . .                                     | 75          |
| V-1 Concept of Return Period and Acceleration Zone<br>Graphs (AZG) . . . . . | 75          |
| V-2 Seismic Risk Zoning . . . . .  | 79          |
| V-3 Concluding Remarks on Seismic Hazard Maps . . . . .                      | 94          |
| Chapter VI. SUMMARY, CONCLUSION AND FURTHER RESEARCH . . . . .               | 97          |
| VI-1 Summary of Work on Seismic Hazard Mapping . . . . .                     | 97          |
| VI-2 Conclusion from the Present Study . . . . .                             | 98          |
| VI-3 Further Research . . . . .  | 100         |
| REFERENCES . . . . .   | 103         |
| Appendix A POISSON MODEL . . . . .   | 107         |
| Appendix B BAYESIAN HAZARD MODEL . . . . .                                   | 127         |
| Appendix C THE FEBRUARY 4, 1976 GUATEMALAN EARTHQUAKE . . . . .              | 153         |
| Appendix D EARTHQUAKE DATA . . . . .   | 163         |
| Appendix E COMPUTER PROGRAMS . . . . .                                       | 187         |



ERRATA

Seismic Hazard Mapping  
for Guatemala  
(Report No. 26)

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## CHAPTER I

### INTRODUCTION

#### Scope

In this chapter the general background under which this study was initiated is presented. The concepts of seismic hazard and risk are described. The main objectives of the current study are presented.

.....

#### I-1 Introduction

On February 4, 1976 at 3:01 A.M. (local Guatemalan time), central Guatemala experienced a major earthquake. The surface-wave magnitude ( $M_g$ ) of this earthquake was 7.5 and the body wave magnitude ( $M_b$ ) was 5.8. The hypocentral location was 15.32°N latitude and 89.08°W longitude with a focal depth of about 5 kilometers. Various aftershocks followed the main event for days and weeks. It is estimated that millions of dollars of damage and thousands of lives were lost. The resulting economic hardships and the disruption in the way of life can not be fully estimated.

In the chaos of rescue, public care and debris removal which are the usual results of such earthquakes, it is very difficult for public officials and engineers to be concerned about why some buildings survive and others collapse. The building efforts which follow such events bring up many questions. Are existing design requirements adequate? What are the future seismic hazards? What should be the acceptable level of seismic risk? How should the information on future seismic hazards and acceptable risk be translated into a rational, simple and acceptable design methodology? Should similar land uses be permitted in the future for areas which suffered major damage? These and numerous other questions become especially relevant

after a significant damaging event. The loss of life and damage brought about by the earthquake-induced landslides in Guatemala City demonstrated the need for proper land use planning.

The decision processes which lead to the answers for the numerous questions posed in the previous paragraph are a complex mixture of socio-economic constraints, political expediency and the general engineering knowhow. In times when no significant earthquake events have taken place, the decision making processes go on at a slow rate, while the decisions immediately after the damaging event are often based on expediency and, at times, on incomplete and irrational analysis. The attention is only focused on collapsed or heavily damaged structures and the reaction is that these failures should never be allowed to happen. The public, through their representative officials, demand doubling or tripling of the existing design levels -- with resulting higher costs and delayed construction.

Under this barrage of emergency actions, public pressure and all sorts of "consulting advice" from experts from all over the world, it is remarkable that the engineers, planners and decision makers in Guatemala decided to initiate a systematic study of seismic hazard and risk analysis with long range perspective. Such a study has two basic goals.

- 1) To develop a seismic hazard map for Guatemala, based on all the available seismological and geological information.
- 2) Based on the information of future seismic hazards in various parts of the country and based on the concept of acceptable level of risk, develop a rational and simple design methodology for lateral load resistance.

This report is the result of achieving the first goal mentioned above. The study was conducted at Stanford University and supported by the following organizations in Guatemala.

1. Cámara Guatemalteca de la Construcción
2. Universidad de San Carlos de Guatemala
3. Banco Nacional de la Vivienda
4. Instituto Nacional de Electrificación
5. Banco de Guatemala

In a second report (to be published in the future) the probabilistic hazard information will be used to develop a design methodology which if implemented could help in reducing the future seismic risk to an acceptable level. A detailed comparison between the suggested method and the current and proposed California codes will also be presented.

The results of seismic hazard evaluation presented in this report should be used as a base for planning and decision making in Guatemala. Such a study can also be utilized by land use planners, investors and insurance companies. In general, the results of this study (Phase I) can provide professionals in Guatemala with tools and procedures for risk analysis and planning.

## I-2 Some Basic Concepts Concerning Hazard and Risk

In order to convey the importance of seismic hazard and risk analysis to the reader, some basic notions are presented in this section. In the earthquake engineering literature, there is in general, ambiguity regarding two words. They are: Hazard and Risk. Seismic hazard is regarded by many to be synonymous with seismic risk. Earthquake engineers and planners use these two words loosely and interchangeably in their work. There is some danger in this ambiguity since these two words within the context of earthquake engineering have different meanings.

Seismic hazard is defined as "expected occurrence of a future adverse seismic event".

Seismic risk is defined as "expected consequences of a future seismic event".

Consequences may be life loss, injury, economic loss, function loss and damage. Expected hazard and expected risk have an implication of future uncertainty. Hence, it is not surprising that principles of probabilistic forecasting and decision making are essential in any seismic hazard or seismic risk analysis.

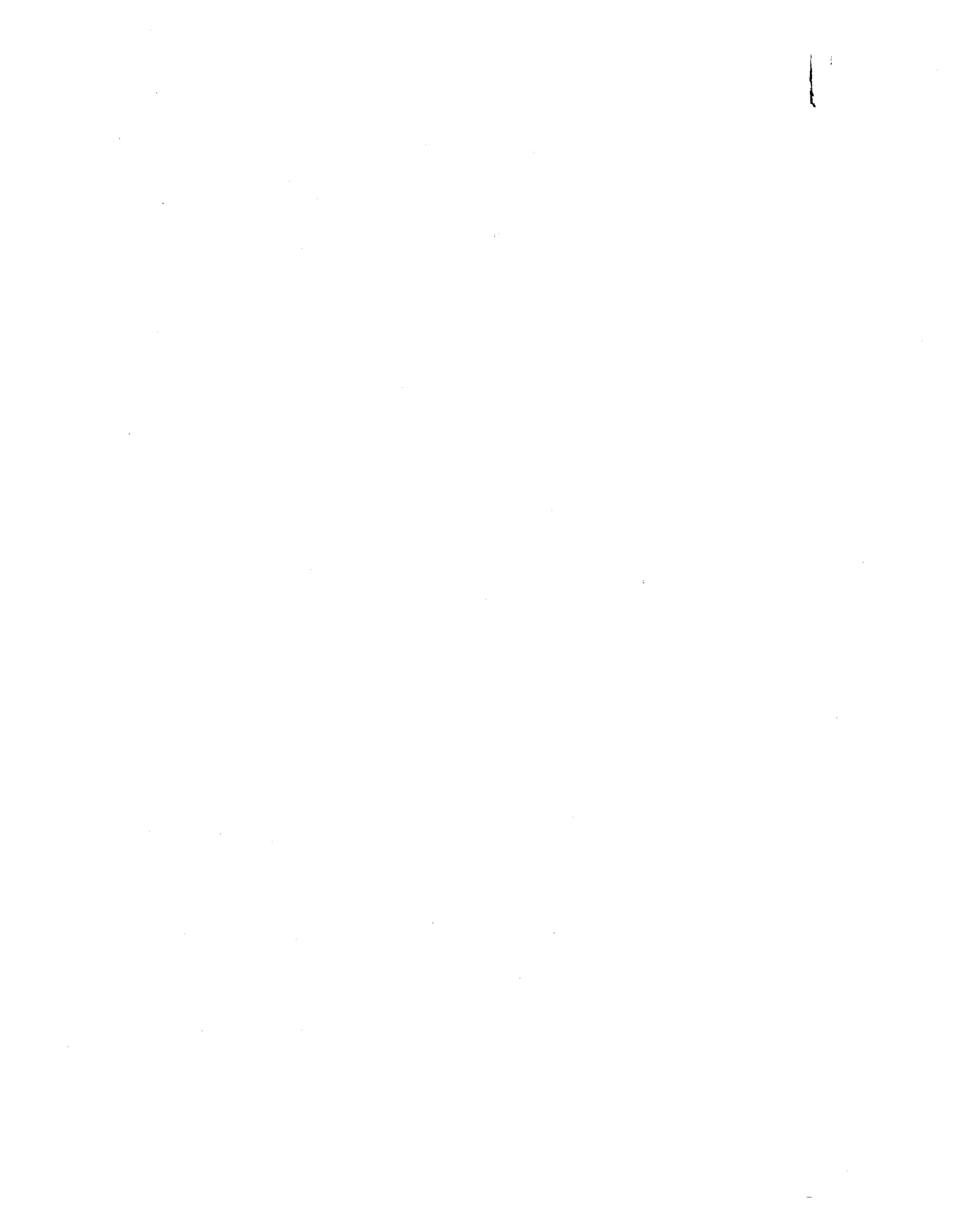
In a recent report to the United States Congress (1972) by the U.S. Executive Office of the President, Office of Emergency Preparedness, the following two recommendations were made.

- 1) The development of seismic hazard maps is an essential first step in hazard reduction and preparedness planning.
- 2) The greatest potential for reducing the loss of life and property from earthquakes lies in restructuring the use of land in high risk areas and in imposing appropriate structural engineering and materials standards both upon new and existing buildings.

As can be seen from above, it is essential that a seismic hazard map be prepared for Guatemala as a first step.

Finally, it should be kept in mind that the work and results presented here depend on the available data base and information. The reliability of results are at best as good as the reliability of the data on which the results are based. It is very easy to attack and criticize any work from the point of view of data reliability. However, it is very difficult to obtain long-range reliable data. The best available information through various organizations and researchers have been used in this study. The forecasts and predictions are based on those data. However, if in the future more reliable data are available, the model can easily accommodate the inclusion of new information and update the results. At this time, the

authors of this report feel that the results presented here represent the "best available" estimates of the future forecasts.



## CHAPTER II

### GEOLOGIC SETTING OF GUATEMALA

#### Scope

In this chapter the general geologic and seismologic setting of Guatemala and surrounding regions is presented. A description of the major plates and plate boundaries is given along with their relationship to seismically active zones. Guatemala has been divided into five geologic and structural regions which are individually discussed. The volcanic activity is described as are major faults and structural trends. This chapter concludes with some suggestions relating to seismic zoning.

#### II-1 Introduction

Guatemala is located within the circumpacific "Ring of Fire", a zone of intense seismic and volcanic activity. The modern plate tectonics model explains this high level of activity as being associated with the motions of crustal plates and the interactions at their boundaries. Most of Central America lies on the western portion of the Caribbean Plate. However, Guatemala is situated athwart the boundary between the Caribbean Plate to the south and the North American Plate to the north (see Figure 2-1). This boundary is marked by the generally east-west trending Cayman (or Bartlett) trough, a deep submarine fault valley over 4 km. deep. To the west, the landward extension of this trough coincides with the Motagua and Polochic fault zones. The motion along this boundary is predominantly left lateral (opposite side moving to the left) movement (Molnar and Sykes, 1969).

West of the Caribbean Plate lies the Cocos Plate, the boundary between the two being the Middle America Trench. The Cocos Plate is moving generally northeast and being thrust under the Caribbean Plate (Molnar and Sykes, 1969; Jordan, 1975). The Middle America Trench is due to this underthrusting, or subduction, of the Cocos Plate beneath the Caribbean Plate. Other

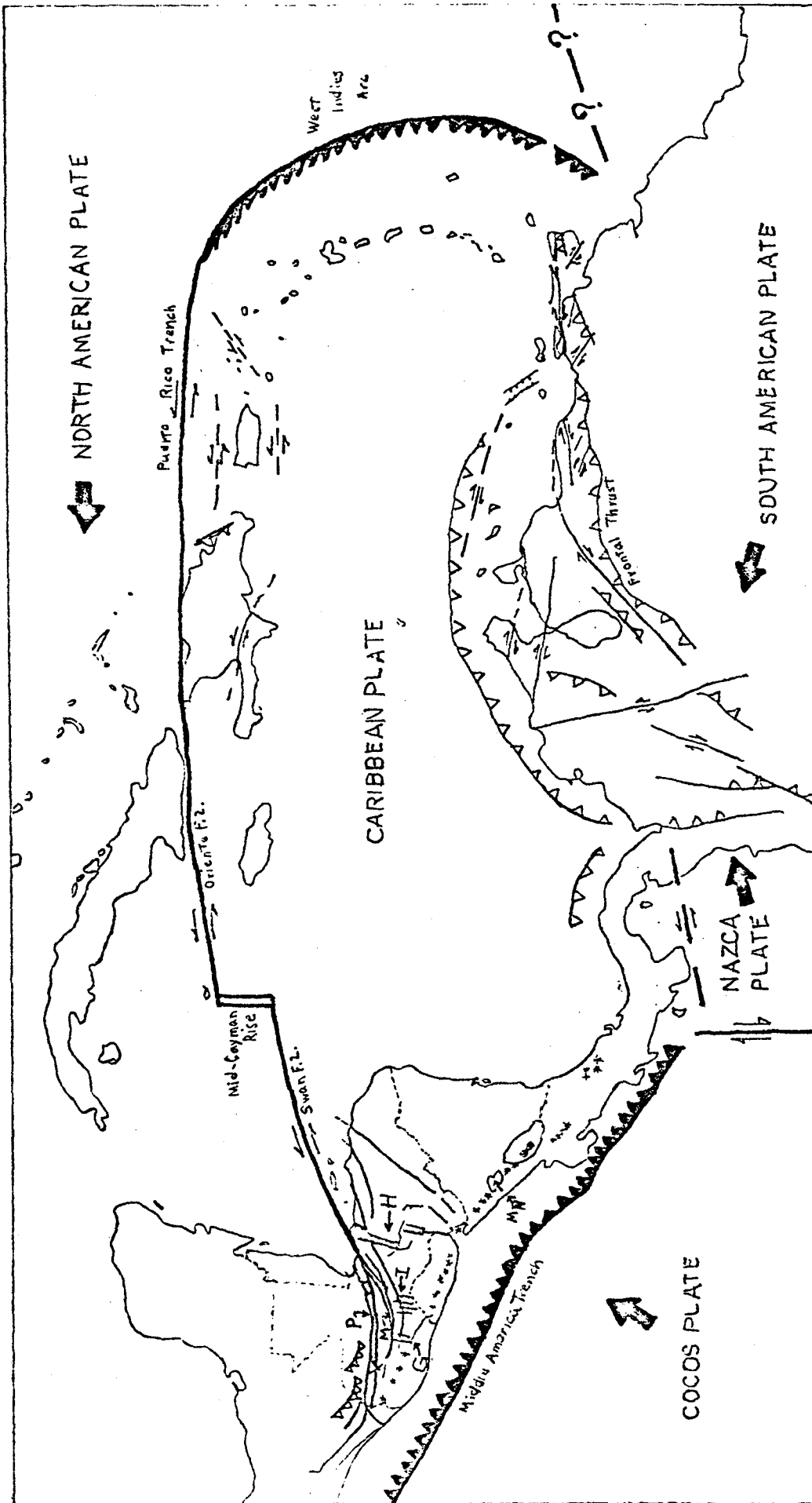


Figure 2-1. Structure and boundaries of the Caribbean Plate.  
 P--Polochoic fault zone, M--Motagua fault zone,  
 G--Guatemala City graben, I--Ipala graben, H--Honduras  
 depression, MN--Managua, Nicaragua  
 (From Billeau, 1976, modified from Jordan, 1975)



features normally associated with a "subducting" plate boundary are also observed along the western and central portions of Central America, including a northwesterly, arcuate trending chain of andesitic stratovolcanoes, a band of shallow to intermediate depth earthquake foci, and an oceanic trench as previously mentioned.

The Middle America Trench is a four to five kilometer deep depression located approximately 100 kilometers west of Central America, extending from southern Mexico southward to Costa Rica. Along the landward margin of the trench is a band of intense shallow seismic activity, with earthquake foci increasing in depth to the northeast, defining a northeast dipping slab-like zone. This zone of friction is commonly known as the Benioff Zone.

A chain of Quaternary and active volcanoes parallels the trench and is located about 100 to 200 kilometers northeast of the trench and directly above the earthquakes with focal depths of about 100 to 200 kilometers (intermediate depth foci).

Guatemala is located due east of the junction of the Cocos, Caribbean and North American Plates. This feature is known as a triple junction. The tectonics associated with triple junctions is often quite complex and confusing.

## II-2 Nuclear Central America:

Schuchert (1935) names "the ancient folded and faulted mountain land of Central America" Nuclear Central America. This portion of Northern Central America includes Chiapas of Mexico, Guatemala, Belice, Honduras, El Salvador and much of northern Nicaragua. Guatemala itself can be further subdivided into five "morphotectonic" units, as described by Dengo and Bohnenberger (1969). These morphotectonic units were "established on the basis of their internal constitution and external relief" (Guzman and DeCserna, 1961), and are a combination of the physiographic and tectonic

features of the various subdivisions. The five subdivisions are the Pacific Coastal Plain, the Pacific Volcanic Chain, the Volcanic Ranges and Plateaus, the Sierras of Northern Central America and the Yucatan Platform (see Figure 2-2). Other investigators, such as Bonis (1967), Schuchert (1935), and Brineman and Vinson (1961), have used various different divisions and classification systems. However, the classification system used herein was selected because it appeared to be the most complete and comprehensive.

#### Pacific Coastal Plain:

The Pacific Coastal Plain is a narrow coastal plain about 50 kilometers wide, located along the Pacific side of Northern Central America, extending from El Salvador northward into southern Mexico. This coastal plain is composed of a thick accumulation of sands, gravels, pumice and bouldary laharic deposits (mudflows of volcanoclastic materials), most of which were derived from the adjacent volcanic chain and highlands. The boundary between this unit and the Pacific Volcanic Chain is thought to be a series of major faults which have been buried under alluvial deposits (Agos, 1958), based on the interpretation of airborne magnetometer data.

#### Pacific Volcanic Chain:

A chain of active and Quaternary andesitic volcanic cones extends from southernmost Mexico, across Guatemala into El Salvador, paralleling the Pacific Coast with a generally northwest trend.

This precipitous volcanic chain stretches longitudinally across much of Central America, and can be divided into several short segments (see Figure 2-3) based on prominent volcanic lineaments, seismicity and quaternary structures (including faulting) (Carr, 1976; Carr, Stoiber and Drake, 1974). The boundaries separating these different segments are often zones of transcurrent faulting, and commonly appear to be tectonically very active.

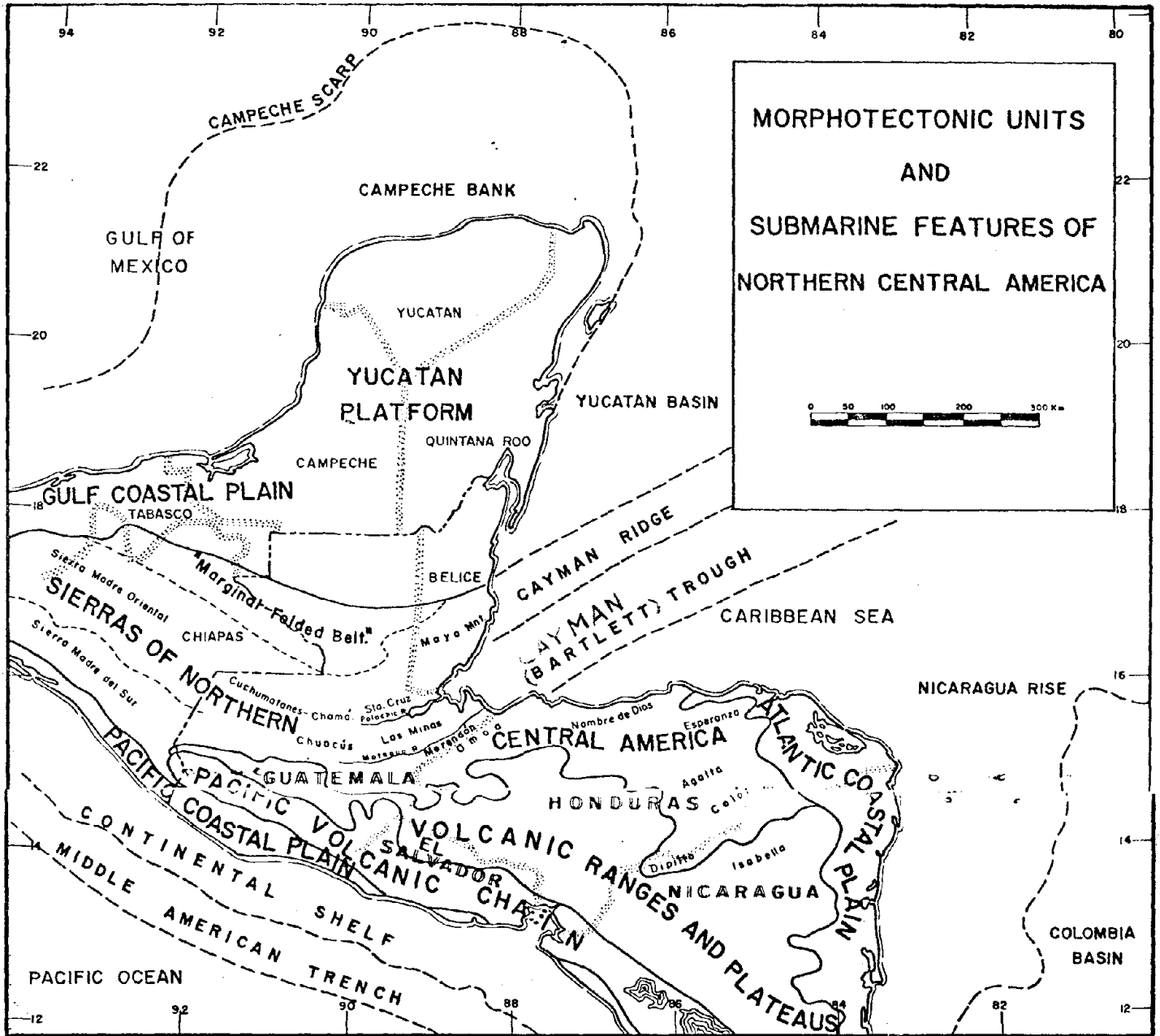
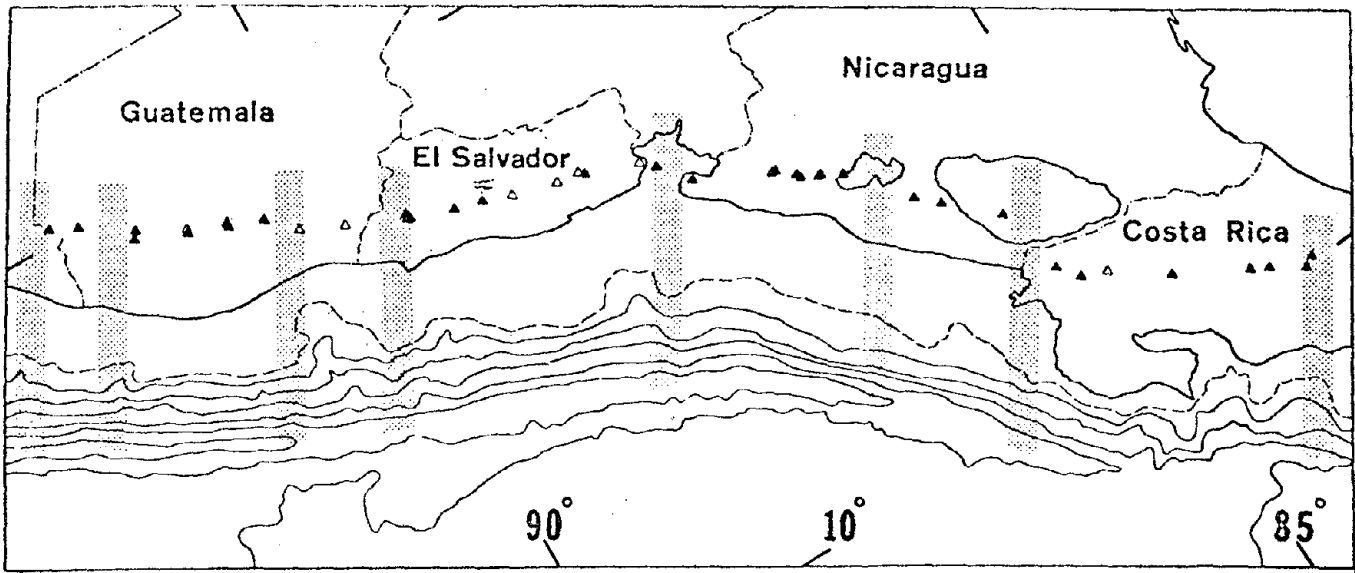
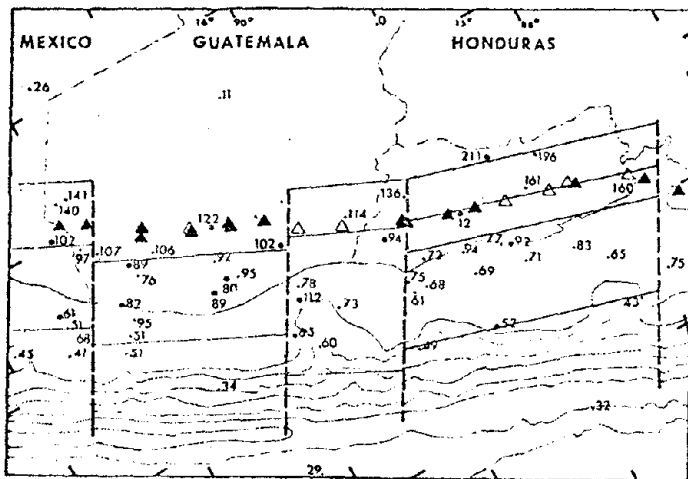


Figure 2-2. Map showing Morphotectonic units and Submarine features of Northern Central America  
(From Dengo & Bohnenberger, 1969)



Volcanic segments of Central America. Solid triangles are volcanoes with historic eruptions (from Mooser et al. 1958), with the addition of Arenal Volcano in Costa Rica. Open triangles are volcanoes with solfatara activity (from Mooser et al., 1958), with the addition of Moyuta Volcano in southeastern Guatemala and the deletion of Zuñil. Bathymetry is from Fisher (1961): the dashed contour is the 100-fathom contour; the next contour is the 500-fathom interval. Stippled areas represent boundaries between adjacent segments.

(From Carr, Stoiber & Drake, 1974)



Seismicity of northern Central America. Focal depths in kilometers are written next to epicenters (dots). Circled dots are epicenters of 15 earthquakes used in JHD. Dashed lines are proposed discontinuities in the inclined seismic zone. Straight lines are isobaths of the deep seismic zone. Contour interval is 50 km and contour nearest the Middle America Trench is the 50-km contour. Unlabeled country is El Salvador. Solid triangles are volcanoes with historic eruptions, and open triangles are volcanoes with solfatara activity (Stoiber and Carr, 1973). Bathymetry is from Fisher (1961). Contour nearest the coast is the 100-fm contour. Next contour is the 500-fm contour, and subsequent contours have a 500-fm interval.

(From Carr, 1976)

Figure 2-3

### Volcanic Ranges and Plateaus:

The Volcanic Ranges and Plateaus, or Volcanic Highlands as described by Williams (1960), form the southern part of Northern Central America, extending across southern Guatemala from the Pacific Volcanic Chain eastward into western Honduras and northern El Salvador (see Figure 2-2). The Tertiary (Miocene to Pliocene, or 25 to one million years old) rocks of the highlands are partly lava flows, ranging from basalt to rhyolite, but are more commonly rhyolitic and dacitic pumice deposits with interbedded fluviatile (river deposited) tuffaceous sediments and a few lenses of diatomite (a deposit formed from silica rich microscopic organisms) (Williams, 1960). Most of these deposits are products of fissure eruptions, rather than of large composite cones, which developed later during Quaternary times. The volcanic deposits overlies a rugged surface cut in mid-Cretaceous plutonic rocks (diorites, granodiorites and granites), early Cretaceous limestones and older metamorphic rocks (serpentinites, mica schists and minor phyllitic shales, marbles, quartzites, and calc-silicate contact rocks).

During and after late Pliocene time, these volcanic deposits were faulted and folded, the trends being generally north-south in the north eastern portion and more northeast-southwest to east-west throughout the rest of the highlands.

Quaternary volcanism consists mainly or wholly of basaltic volcanoes and abundant basaltic cinder cones, occurring as parasitic cones on the flanks of the volcanoes, and as independent features often aligned along faults (Williams, McBirney and Dengo, 1964). Quaternary pumice is relatively scarce, compared to that discharged by the volcanic chain to the west, and to those deposits produced during the period of Tertiary volcanism. The Quaternary volcanoes of this region are widely scattered

along fissure systems, often lying on north-south trending faults (Williams, McBirney and Dengo, 1964).

#### Sierras of Northern Central America:

An arcuate series of subparallel high mountain ranges, convex southward, extends across central Guatemala from Chiapas, Mexico to the Caribbean Sea (see Figure 2-2). The northern ranges of this series appear to be structurally continuous with the submarine topographic high, known as the Cayman Ridge (Dengo and Bohnenberger, 1969). Also the submarine Cayman (Bartlett) trough is continuous structurally with some of the longitudinal valleys which separate several of the mountain ranges, namely the long, narrow fault controlled Polochic and Motagua Valleys.

These mountain ranges can be subdivided, as suggested by Dengo and Bohnenberger (1969) into a northern and a southern group of mountain ranges. The northern group includes the northwest trending Sierra Cuchumatanes, the east-west trending Sierra de Chama and the generally east-northeast trending Sierra de Santa Cruz Range. The Maya Mountains of Belice are also included in this northern group of mountains. The southern group of ranges includes the Sierra de Chuacus, which trends nearly east-west, the east-northeast trending Sierra de las Minas and Sierra de Merendon Ranges and the El Mico Mountains. The Sierra de Santa Cruz and Sierra de las Minas ranges are separated by the Polochic Valley, whereas the Sierra de las Minas and the Sierra de Chuacus are separated from the Sierra de Merendon by the Motagua Valley.

The northern group of mountain ranges are composed primarily of Permian (270 to 220 million years old) and Cretaceous (135 to 70 million years old) limestone which now forms a "series of parallel, tightly folded ranges thrust faulted toward the north and modified by later normal

faults" (Dengo and Bohnenberger, 1969) (see Figure 2-4). These thrust faults trend subparallel to the arcuate trend of the mountain ranges and dip to the south. The normal faults are vertical or steeply dipping and often trend approximately north-south, truncating and offsetting the reverse (thrust) faults (Williams, 1960).

Late Cretaceous and Early Tertiary marine clastic sediments of the Sepur formation are preserved in the synclinal troughs and fault controlled valleys of the northern ranges (Bonis, 1967). This formation consists mainly of shales, calcarenites (calcium carbonate rich sandstone) and conglomerates. Pre-Mesozoic sedimentary rocks (sandstone, shale, conglomerate and phyllite of the Santa Rosa Group) are found in the Maya Mountains, and Sierra de los Cuchumatanes, as are Paleozoic metamorphic rocks and Jurassic-early Cretaceous continental and marine redbeds (mainly sandstone and conglomerate) of the Todos Santos formation (Bonis, 1967). Ultramafic rocks, predominantly serpentinites, also occur in the northern ranges, making up much of the Sierra de Santa Cruz (Dengo and Bohnenberger, 1969).

To the north the rocks are less intensely folded and thrust faulted, forming an intermediate zone between the main fold belt of the high sierras to the south and the nearly flat lying rocks of the Yucatan Platform to the north. This series of low ranges and intervening lowlands makes up the marginal folded belt (see Figure 2-2).

The southern group of mountain ranges consists primarily of pre-Pennsylvanian (before about 300 million years ago) schists, gneisses, amphibolites and marbles of the Chuacas Series (McBirney, 1963), granitic batholiths and ultramafic rocks, mainly serpentinites. Paleozoic and Mesozoic sedimentary rocks (carbonates, shales, sandstones, conglomerates, and phyllites) occur only locally within these ranges. Within the Motagua

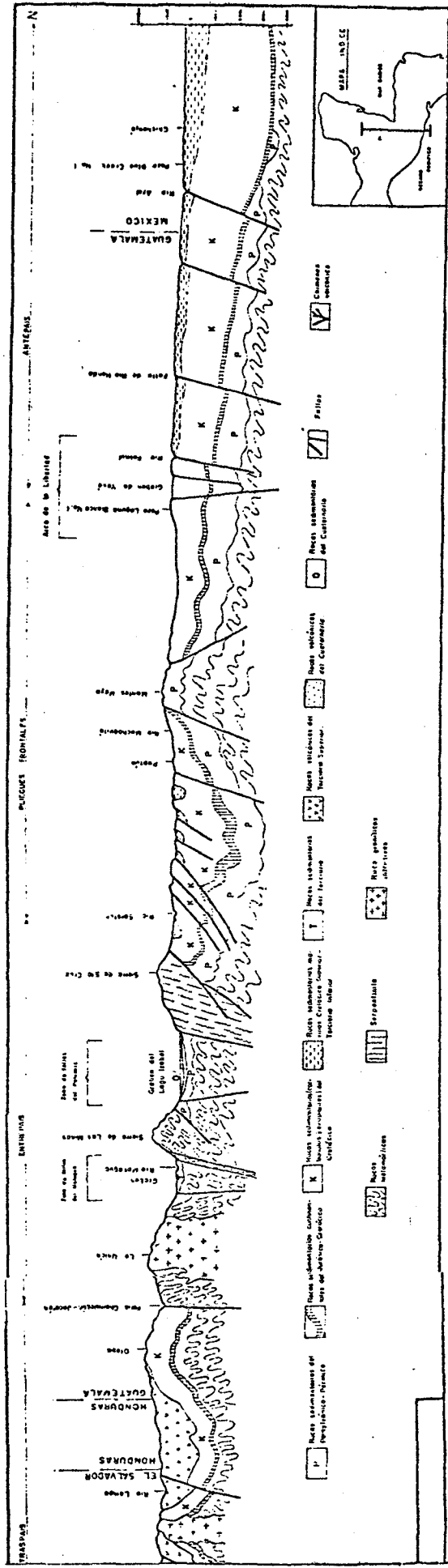


Figure 2-4. Generalized geologic cross section through Guatemala, showing the different surface and subsurface rock units and the major structures. (From Dengo, 1968)

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Valley, and bounded by the Motagua fault system are Cretaceous to Eocene (135 to 40 million years old) and redbeds of the Subinal formation. Also within the Motagua Valley, and other intervening valleys, are small areas of Quaternary pumice deposits.

The faulting in the southern group of mountain ranges is less dominated by east-west trending, south dipping thrust faults, but rather by a few major left lateral strike-slip faults, namely the Cuilco-Chixoy-Polochic fault, the Motagua fault and the Jocotan-Chamelecon fault to the south.

#### Yucatan Platform:

The Yucatan Platform, or Peten Lowland (Bonis, 1967) extends across northern Guatemala from Tabasco, Mexico into northern Belize (see Figure 2-2). The southern limit of the Yucatan Platform is marked by the La Libertad Arch, (Vinson and Brineman, 1961). Nearly flat lying Paleozoic and Eocene marine sediments cover much of the northern portion of the area, overlying Cretaceous carbonate and evaporite rocks, which crop out extensively in the southern part of the Yucatan Platform.

An extensive karst topography has developed on the carbonate rocks, including underground drainages and large caverns. Collapse of these caverns has produced earthquakes which commonly have shallow epicenters and fairly local felt areas (Lomnitz, 1974).

#### II-3 Volcanic Activity:

Quaternary volcanic activity is restricted to the southern part of the country, occurring along the Pacific Volcanic Chain and in the Volcanic Ranges and Plateaus region, east of the main chain. The types of Quaternary deposits include andesitic and basaltic flows, and pumice deposits laid down by torrential floods which removed airborne pumice falls from the

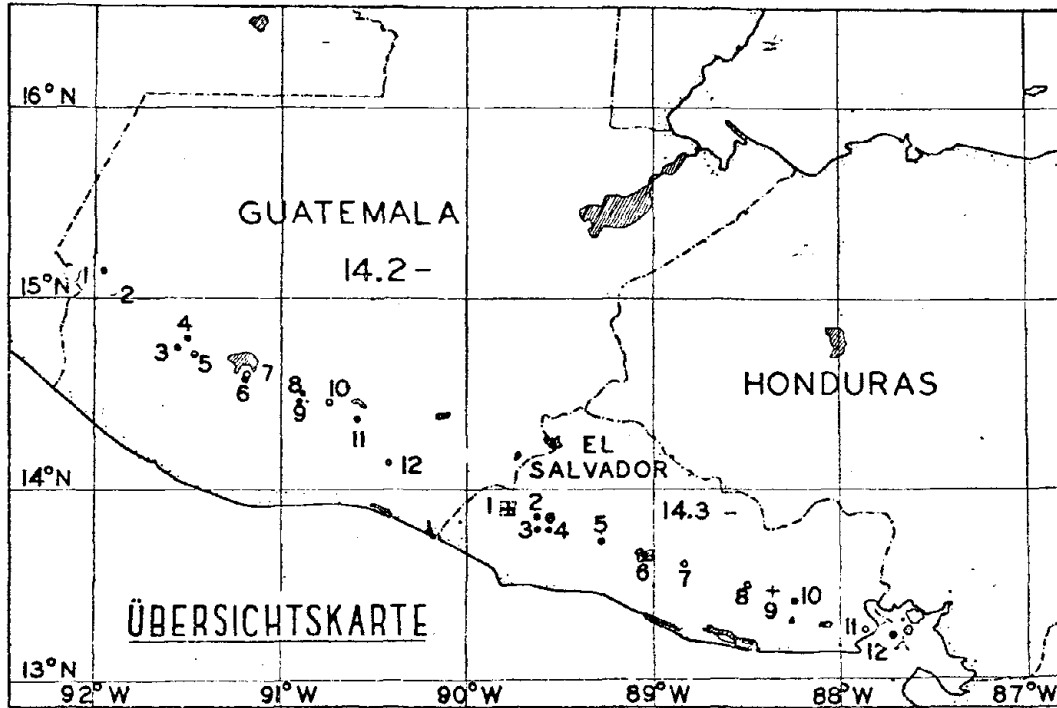
surrounding hills, and less commonly the product of glowing avalanches.

The principal group of volcanoes in Guatemala (see Figure 2-5) lie in a nearly straight line for 122 kilometers from Volcano Siete Orejas to Volcano Pacaya, at which point the line splits with one strand to the north linking the twin volcanoes of Lake Ayarza, the other strand, offset several kilometers to the south, links volcano Tecuamburo and Volcano Moyota. To the north, near the Guatemala-Mexico border, lie Volcan Tacana and Volcano Tajumulco, which are offset to the north of the main chain. Near the Guatemala-El Salvador border, the cones are more widely scattered and irregularly distributed.

Activity within the past few centuries has been recorded from several volcanoes. Cerro Quemado, south of Quetzaltenango, erupted most recently in 1785 with fumarolic activity noted by recent investigators. In 1853, Volcano Atitlan, south of Lake Atitlan, erupted. This volcano had had prior periods of explosive activity, however, no recent eruptions have been recorded. Volcano Pacaya, located south of Guatemala City has been intermittently active over the last several hundred years. The two most active volcanoes in Guatemala are Volcano Fuego, nearly 40 kilometers southwest of Guatemala City, and Santiaguito on the southern flank of Volcano Santa Maria, about eight kilometers south of Quetzaltenango. Both of these volcanoes have continued to show signs of activity up to the present. The basaltic cinder cones and related flows of eastern Guatemala are thought to be no more than a few thousand years old (Williams, 1960).

#### II-4 Faulting:

Guatemala can be subdivided into three major structural regions, each characterized by different types, orientations and intensities of faulting and folding. These regions generally coincide with some of the



Index map showing the sites of the volcanic centres of Guatemala and El Salvador.

- Figure 2-5.
- |                  |                 |
|------------------|-----------------|
| 1. Tacaná        | 7. Toliman      |
| 2. Tajumulco     | 8. Acatenango   |
| 3. Santa Maria   | 9. Fuego        |
| 4. Cerro Quemado | 10. Agua        |
| 5. Zuñil         | 11. Pacaya      |
| 6. Atitlán       | 12. Tecuamburro |

major morphotectonic subdivisions of Dengo and Bohnenberger (1969), and include: the south dipping thrust faults and normal faults of the northern group of mountain ranges and the Yucatan Platform; the east-west longitudinal strike-slip faults of the southern group of mountain ranges; and the normal and transcurrent fault zones of the Volcanic Ranges and Plateaus and Pacific Volcanic Chain. This transcurrent fault zone includes N30-45°E trending left lateral strike-slip faults, a conjugate set of N45-65°W trending normal and right lateral strike-slip faults, plus a set of normal (extensional) faults trending generally north-south. These various regions, and their fault zones, may be the result of different stress systems operating at different times, or one stress system of very large dimensions, portions of which have been active at different times (Dengo and Bohnenberger, 1969).

The northern system of thrust faults and subsequent normal faults are thought to be of late Cretaceous and Tertiary age. The thrust faulting began during late Cretaceous times and climaxed in early Eocene time. Uplift of the Yucatan Platform and adjacent regions took place during late Eocene and Oligocene times (50 to 25 million years ago). The normal faulting is thought to have occurred during this period of uplift. Present day earthquakes in this region are more commonly due to limestone cavern collapse, as previously mentioned, rather than of tectonic origin (Lomnitz, 1974).

The arcuate (convex southward) generally east-west trending longitudinal faults of the southern mountain group make up the second subdivision. This includes from north to south, the Cuilco-Chixoy-Polochic fault zone, the Motagua fault zone and the Jocotan-Chamelecon fault zone (see Figure 2-6). The Polochic fault zone extends from the Caribbean Sea to west of the Guatemala-Mexico border, stretching nearly 400

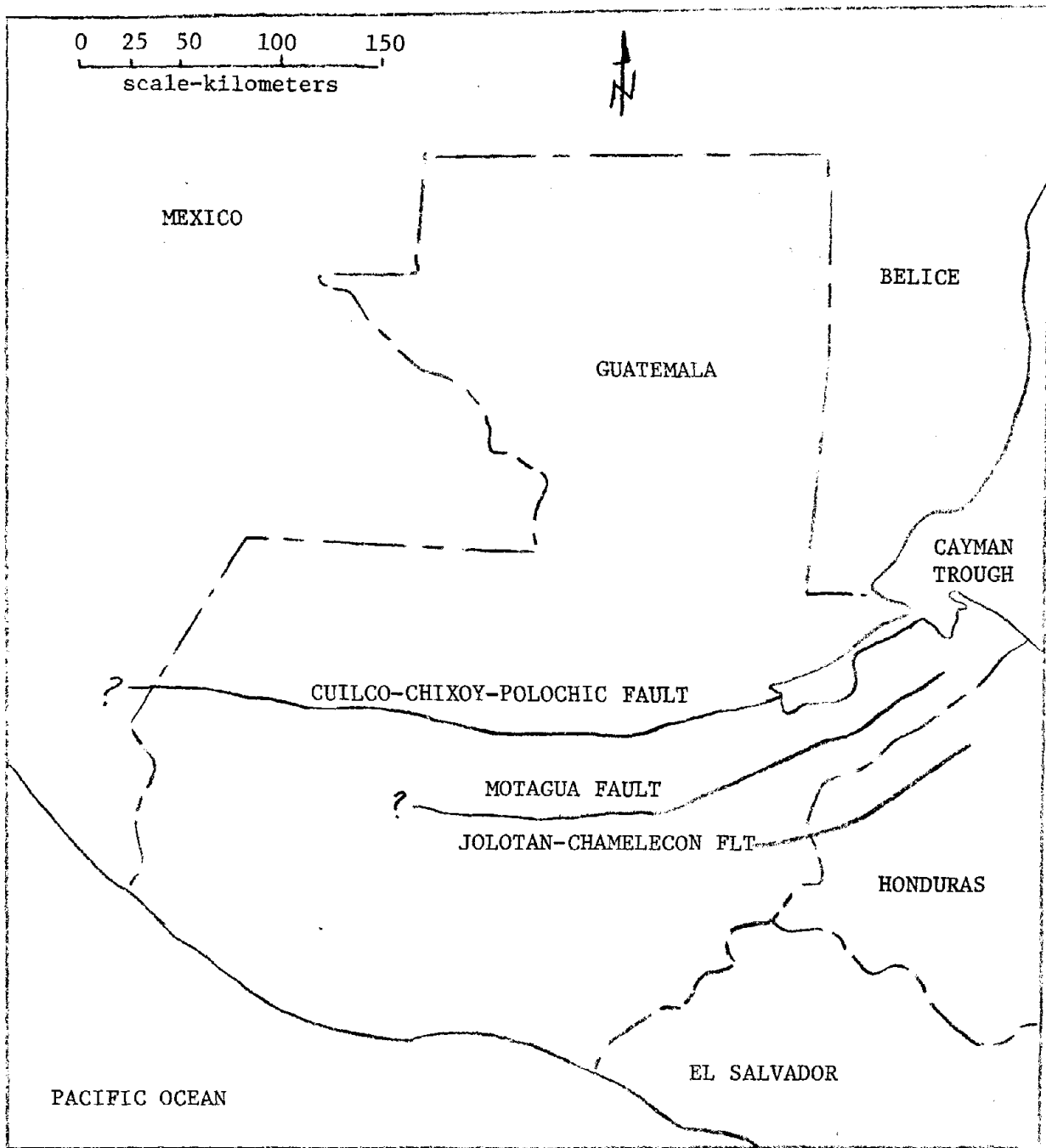


Figure 2-6. Map of Guatemala, showing the locations of the major longitudinal fault zones.  
 (Modified from Anderson and others, 1973)

kilometers, whereas the Motagua fault zone extends from near the Caribbean Sea to the west nearly 300 kilometers. The western extent of the Motagua fault is uncertain because Quaternary volcanic deposits of the Pacific volcanic belt conceal the fault trace (Anderson, 1973; Dengo and Bohnenberger, 1969). The Jocotan-Chamelecon fault zone trends generally northeast-southwest, subparallel to the eastern half of the Motagua fault zone, and extends across northern Honduras into eastern Guatemala, a distance of 125 kilometers. The Motagua and Polochic faults are the landward projection of the Cayman trough, a deep submarine trough which marks the northwest boundary of the Caribbean Plate, as mentioned previously (see Figure 2-1).

Until recently, movements along the Cuilco-Chixoy-Polochic, Motagua and Jocotan-Chamelecon fault zones were thought to be in the vertical direction, with only minor left lateral displacement, based on geologic field evidence in Guatemala. However, many investigators are now postulating large magnitude left lateral strike-slip faulting along these fault zones, suggested by the linearity and width of the fault zones, offset streams, geomorphic features such as beheaded valleys, shutter ridges and chains of scarplets (Kupfer and Godoy, 1967) and a series of inverted open S-shaped folds in Late Tertiary sediments deposited in the Polochic fault zone, which show drag features close to the faults (Dengo and Bohnenberger, 1969). Left lateral transcurrent faulting is further supported by the modern plate tectonics model for this area, earthquake focal mechanism studies (Molnar and Sykes, 1969), and by the displacements observed along the Motagua fault zone following the 1976 Guatemala earthquake (Plafker & Bonilla, 1976).

Recent activity along these major faults is indicated by the abundance of fault scarps and offsets of recent geomorphic features, as well

as the occurrence of earthquakes with epicenters along the faults. A more complete discussion of the 1976 rupture along the Motagua fault is given in Appendix C.

Quaternary faulting in the third region includes three major fault trends; transcurrent N30-45°E faults, right lateral and normal N45-65°W faults, and generally north-south extensional faults (Carr, 1976). These faults are most readily observed in the Volcanic Ranges and Plateaus region of Figure 2.

The two major left lateral, northeast trending faults of this region, as described by Carr (1976), are the Palin Shear Zone and the Rio Paz Shear Zone, about 75 kilometers to the southeast of the Palin Shear Zone. Carr suggests a relationship between these left lateral faults and the segmentation of the volcanic chain to discontinuities in the inclined seismic zone, or Benioff Zone (Carr, 1976).

The northwest trending faults may be a conjugate system to the northeast trending faults. Carr (1976) suggests that they often have right lateral motion, however, many also show a component of normal displacement (vertical movement) (Williams, McBirney and Dengo, 1964). One example of these faults is the Median Trough in El Salvador. The Jalapatagua fault of southeastern Guatemala, southeast of Guatemala City, exhibits right lateral strike-slip motion. Northwest trending faults have also been noted to be associated with, and in close proximity to, the volcanic chain of the Pacific.

The third principal fault orientation of this region is in the north-south direction. These faults are primarily normal faults, producing north-south trending grabens (down-dropped blocks), such as the Guatemala City and Ipala grabens. These faults also tend to control the distribution of Quaternary volcanics. These north-south trending faults

become much more dominant in the northern and northeastern portion of the Volcanic Ranges and Plateaus unit, whereas to the south, all three fault trends occur together.

All three fault trends of this southern region show signs of recent activity, based on geomorphic evidence and the cross cutting relationships of the various faults. Recent activity along the north-south faults is demonstrated by the occurrence of north-south trending scarps offsetting Quaternary pumice deposits and fresh cinder cones (Dengo and Bohnenberger, 1969). The Mixco fault (see Figure 2-7), west of Guatemala City, exhibited dip-slip movement (Plafker & Bonilla, 1976) in association with the February 4, 1976 Guatemala earthquake, further supporting that the north-south trend is still active (Appendix C).

Seismic zoning, based on a thorough understanding of the nature of faulting, the tectonic history and present day regime, and other critical factors, is essential for safe growth in a seismically active region, such as Guatemala. A few important points relating to seismic zoning follow:

1. Numerous active and potentially active faults are observable in Guatemala, as described above. Recognition of these active faults is not as difficult in Guatemala as in other countries, because of the abundance of Quaternary volcanic deposits (for dating the faulting), and the existence and preservation of geomorphic features such as fault scarps, fault valleys, etc., as well as other factors. Not all of the active faults ruptured during the 1976 Guatemala earthquake. Therefore, in seismic zoning, all of the active faults should be considered capable of rupturing during future earthquakes, except where specific studies suggest otherwise.



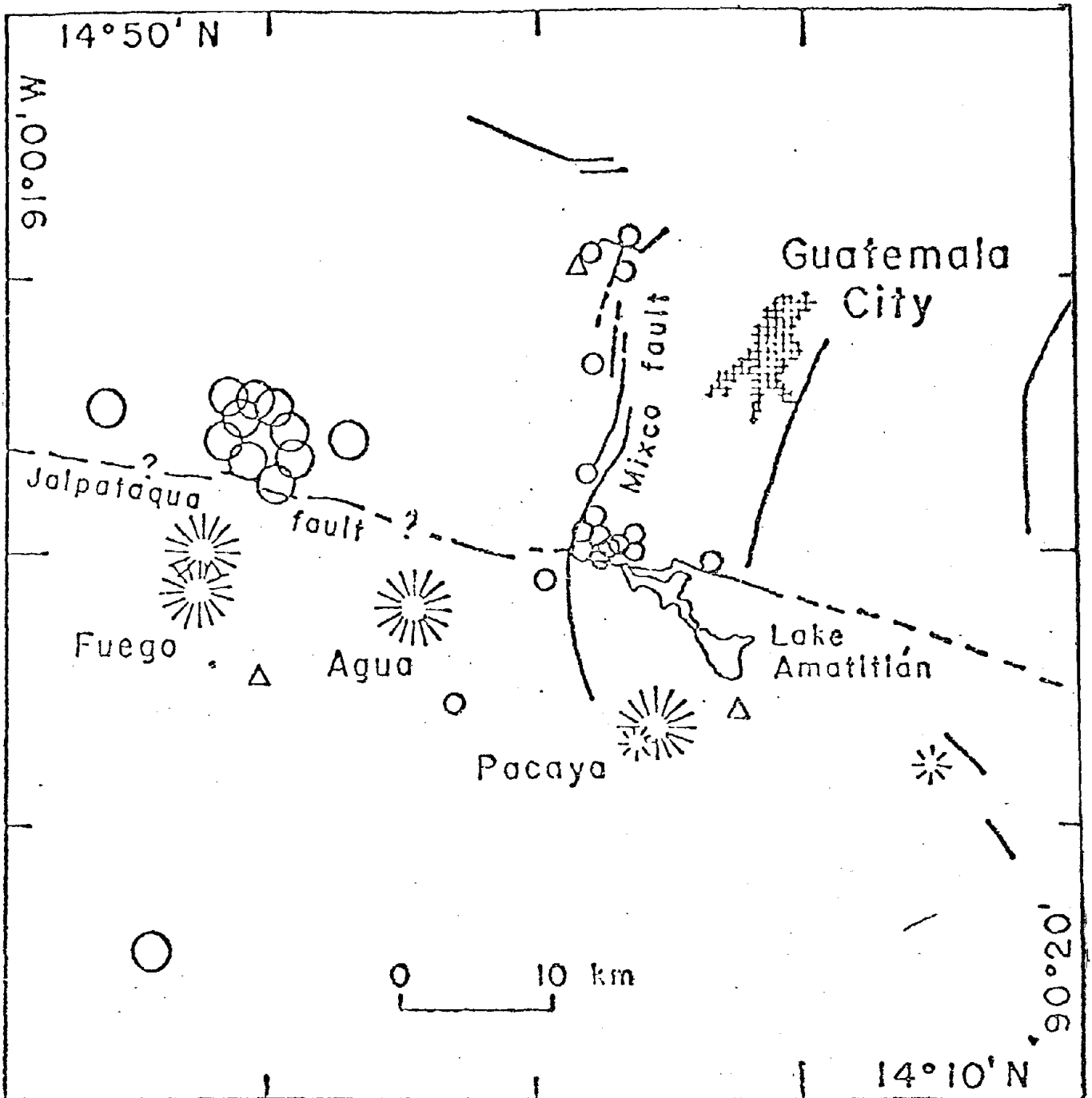


Figure 2-7. Location of Mixco Fault. Open circles are epicenter locations from 1973-1974. The size of the circles reflects the estimated error in the calculated location.

2. It should be noted, as was observed during the 1976 Guatemala earthquake, that several apparently unrelated faults can rupture and/or show other signs of activity during a single event. Therefore, recognition of more than the dominant orientation of active faulting is necessary for safe seismic zoning.

3. The definition of an "active" fault is not yet agreed upon. The number of years since the most recent movement along the fault, before it can be considered "inactive", is still unresolved. Many geologists use 40,000 years since the last movement as the criterion of "active". The valid argument has been made that even 10,000 years or less is economically impractical when considering structures with an anticipated lifetime of less than 100 years. A definition developed by Nichols and Buchanan-Banks (1974) is that "faults are regarded as active and of concern to land-use planning when there is evidence that they have moved during historic time or, through geologic evidence, there is a significant likelihood that they will move during the projected use of a particular structure or piece of land". An understanding of the recurrence interval and nature of faulting is necessary in selecting a minimum age since the most recent movement. However, it is our opinion that strategically important and large public structures should not be located over defined faults, unless the most conservative (40,000 years) definition of "active" is used.

4. The faults that are designated as "active" should be considered as zones, rather than a line of movement. It has been observed that certain faults (such as the San Andreas fault in California) tend to rupture repeatedly along nearly the same line (Nichols and Buchanan-Banks, 1974). However, in some regions many recent fault scarps are located

within wider fault zones; while in alluvium, the fault rupture could occur outward from the previously "recognized" break. The width of this "zone" depends on the nature of the faulting, the orientation of the fault plane, and the materials being faulted, along with several other factors. Once the fault zone can be reliably delineated, it is our opinion that at least the vital structures, such as hospitals, police and fire stations, etc., should be located at least 30 to 45 meters from these fault zones.

5. The type of material on which the structure rests, and the degree of saturation of the material are extremely important.

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## CHAPTER III

### SEISMIC DATA BASE

#### Scope

This chapter describes the available seismological data for the country. Its characteristics and limitations are discussed. Based on the fault locations, eighteen line sources are modeled, their seismicity determined and the confidence on the recurrence relationship presented.

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#### III-1 Introduction

A difficult but essential task for the current study is the acquisition of past earthquake data. In general it is necessary for the data to contain information on the severity of the earthquake, the epicentral location, the focal depth to the hypocenter, and the time of occurrence.

The most commonly used parameters representing the severity of the seismic events are the Modified Mercalli intensity, the Richter magnitude and the peak ground acceleration. The Modified Mercalli intensity is a subjective scale of the damage at a site. Forecasting of MM intensities can be useful in determining the future damage potential in a region, however they cannot be related to structural response characteristics for analysis and design purposes. Thus, in the present study MM intensity data will be used only when it can be translated into either a Richter magnitude or a peak ground acceleration.

The Richter magnitude is related to the overall energy release of a seismic event. It is important to note that the Richter magnitude conveys

information only at the epicenter (or hypocenter). The parameter popular among earthquake engineers which represents the ground motion at a site is the peak ground acceleration. It is frequently used for analysis and design of structures, consequently it will be employed in the present report as the parameter for determining the "seismic loads" at a site. Recently other load measuring parameters such as root-mean-square of ground motion, stress drop, and seismic moment have become popular among seismologists and engineers. However these parameters have been determined only for a few past earthquake events and thus cannot be used in the present study.

In addition to the above specified characteristics of the data base, it is desirable to have strong motion accelerograph records from major earthquakes and also information on how the energy (or the peak ground acceleration) attenuates from the source to the site. Only a few accelerograms of very low acceleration amplitudes are available for the country and hence are hardly sufficient for determining the attenuation characteristics in Guatemala.

Before carrying the discussion further into the analysis and use of the available data certain observations should be made regarding the type, the amount and the reliability of the data base used in the current work.

A shortcoming common to all the data sources is that the frequency of recorded earthquakes increases with time, (i.e. the number of seismic events recorded increases with each year). This nonhomogeneity in the data base is due not because there is necessarily an increase in the seismic activity in the region, but because of the improvement of instruments, measurements and measurement analysis techniques resulting in a bias towards recent years. This non-uniformity in the data can be observed to be true for all historical-time-dependent records and is a "fact of

life." Thus one cannot get away from it.

The other problems associated with the data base deal specifically with the variability of the epicentral location, the Richter magnitude and the focal depth. Variations in epicentral location are dependent on the seismological station reporting the earthquakes. For earlier seismic events only the general region of the earthquake occurrence is reported where on occasions it encircles an area of 100 km radius. Furthermore, for many of these events the focal depth is not determined due to lack of seismograph records close to the epicenter. In cases where a focal depth is reported, its reliability can also be questioned. Similarly, the value of Richter magnitude is not always recorded and when recorded it contains certain error. The variability of the data and its effect on the modeling has been extensively studied (Kiremidjian and Shah, 1975) previously. An assessment of the uncertainties of the data associated with the present study will be made at the end of this report.

### III-2 Data Analysis

Past earthquake data was collected for a region spanning from 12.5°N to 19.5°N and from 86.5°W to 93.5°W. These boundaries are chosen so that they are at least one degree (longitude or latitude) outside of the border of Guatemala. The two main sources of information that are considered for data acquisition are:

- The NEIC-NOAA data file containing events from January 1900 to August 1973 and referred to hereafter as Data File 1;
- The Preliminary Data Epicenters of the NEIC-USGS covering the period from September 1973 to February 1976, and referred to hereafter as Data File 2.

In addition, the events reported by the Bulletin of the Seismological Society of America and the Observatorio Meteorologica Nacional of

Guatemala were used to check the completeness of the data. Information on the major shock of the February 4, 1976 earthquake and its aftershocks was obtained from the United States Geological Survey (Page ed., 1976) and records of the Observatorio Meteorologico Nacional. In spite of the complementarity of the different data sources, a large number of events remain insufficiently documented in order to be used as such in the analysis. Thus in cases where basic information such as epicentral location or magnitude is missing, the event is disregarded. Records with a Richter magnitude smaller than 3.0 are also not considered.

Both Date Files 1 and 2 are very similar thus they will be discussed together. When complete, the information for a given seismic event contained in these files includes the time of occurrence, the epicentral location (degrees longitude and latitude), the depth to the hypocenter (km), and the magnitude. The magnitude is given in terms of one of the following:

- (1) CGS  $M_b$  average (body wave magnitude)
- (2) CGS  $M_s$  average (surface wave magnitude)
- (3) Richter magnitude  $M$ .

For the study at hand, it is necessary that all the data be expressed in terms of a single magnitude parameter. The Richter magnitude is chosen for that purpose and whenever only  $M_b$  values are reported these values are converted to Richter magnitude. It is known that for a given part of the world, the Richter magnitude and CGS  $M_b$  are linearly related such that

$$M = a + b M_b \qquad 3-1$$

In order to determine the coefficients  $a$  and  $b$ , a regression analysis was performed for all earthquakes of which  $M$  and  $M_b$  were known using the



total data of Central America. The Richter magnitude was then obtained by substituting the value of  $M_b$  in Equation 3-1.

The focal depth given in these files is expressed either in kilometers or by a letter symbol N (0 - 33 km).

From Data File 1, 979 events were reported and from Data File 2, 174 events were obtained, thus constituting a total sample of 1153 past earthquake records containing complete information. The epicenters and the magnitudes of these events are shown on Chart 1. Table 3-1 lists these records as a function of focal depths. A listing of all the data appears in Appendix D.

The general seismic pattern of Guatemala is studied from Charts 1 to 6. The country is divided into the following regions of tectonic activity:

(1) The Benioff Zone dipping northeast under the Guatemalan coast. This zone is marked by a numerous earthquakes covering the entire range of magnitudes. It extends several hundred kilometers beneath the earth's surface. Some of the deepest earthquakes recorded along the Benioff Zone go as far as 299 km projecting around the central portion of the country. The Benioff Zone in this part of the Central Americas has been observed (Carr, 1976) to start out with an extended shallow section and then dip at about  $40^\circ$  angle under the coast of the country. This observation is confirmed by looking at Charts 2 to 6. The range of shallow epicenters extends from the Middle American trench almost to the coastline forming a rather diffused zone.

(2) The Motagua and Polochic fault zones are associated with and extend into the Cayman (Bartlett) trough. In this region of the country primarily shallow earthquakes are observed (5 - 33 km) and the magnitudes are found up to 7.5 on the Richter scale. However, due to the lengths of these

TABLE 3-1

Earthquake Data Sorted According to Depth of Hypocenter

| <u>Number of Earthquakes</u> | <u>Depth Range (km)</u> |
|------------------------------|-------------------------|
| 81                           | 0 - 9                   |
| 37                           | 10 - 19                 |
| 34                           | 20 - 29                 |
| 298                          | 30 - 39                 |
| 53                           | 40 - 49                 |
| 65                           | 50 - 59                 |
| 116                          | 60 - 69                 |
| 87                           | 70 - 79                 |
| 72                           | 80 - 89                 |
| 60                           | 90 - 99                 |
| 63                           | 100 - 109               |
| 39                           | 110 - 119               |
| 27                           | 120 - 129               |
| 18                           | 130 - 139               |
| 13                           | 140 - 149               |
| 18                           | 150 - 159               |
| 11                           | 160 - 169               |
| 8                            | 170 - 179               |
| 6                            | 180 - 189               |
| 10                           | 190 - 199               |
| 20                           | 200 - 219               |
| 7                            | 220 - 239               |
| 5                            | 240 - 259               |
| 4                            | 260 - 279               |
| 1                            | 280 - 299               |

fault zones a maximum magnitude of 8.5 is believed to be possible on either of them. This information is directly incorporated in the seismic modeling of the country (see Table 3-2). These faults are of particular importance because of their capability for large magnitude earthquakes, their shallowness and their proximity to highly populated areas. A smaller, but still important fault that can be linked with these zones is the Jocotan. Its characteristics were already discussed in Chapter II.

(3) The line of volcanoes from northwest to southeast (Cordillera de los Marrabios) represents sources of future seismic activities. Volcanic eruptions are seldom by themselves sources of seismic activity, and in the past various earthquakes have been recorded preceding volcanic eruptions. For this reason earthquakes "associated" with volcanic activity were treated in the model (Chapter II) as shallow tectonic earthquakes.

(4) Several shallow and short faults have been found in the vicinity of Guatemala City of which the Mixco has been observed to have the most recent movement. Their geologic characteristics differ from the major fault zones considered (see Chapter II). Only a few earthquakes have been associated with the Mixco fault, but their severity is of particular concern because of its nearness to Guatemala City.

(5) The Yucatan Platform is observed to have a very low seismicity. In this region no earthquakes of tectonic origin have been recorded.

### III-3 Source Location and Seismicity

Based on the above observations, the total number of events was divided into 18 seismic line sources. Table 3-2 identifies the sources, the number of events and the focal depth range for each source. Charts 2 to 6 show the epicenters of past earthquakes associated with each source. Most of the sources are composed of several line segments in

TABLE 3-2

Seismic Sources For Guatemala

| SOURCE   | NUMBER OF EVENTS | NAME OF SOURCE     | DEPTH (km) |
|----------|------------------|--------------------|------------|
| S1       | 56               | Motagua            | 33         |
| S2       | 16               | Polochic           | 40         |
| S3       | 31               | Line of Volcanoes  | 30         |
| S4       | 67               | Pacific Coastline  | 30         |
| S5       | 19               | Pacific Coastline  | 34         |
| S6       | 5                | Near Mexico Border | 30         |
| S7       | 7                | Near Mexico Border | 35         |
| S8       | 16               | Mixco              | 30         |
| S9       | 17               | "Mexico"           | 60         |
| S10      | 10               | "Mexico"           | 75         |
| S11      | 7                | Jocotan            | 30         |
| S12 - B1 | 244              | Benioff 1          | 30         |
| S13 - B2 | 87               | Benioff 2          | 50         |
| S14 - B3 | 192              | Benioff 3          | 70         |
| S15 - B4 | 196              | Benioff 4          | 90         |
| S16 - B5 | 111              | Benioff 5          | 130        |
| S17 - B6 | 35               | Benioff 6          | 180        |
| S18 - B7 | 37               | Benioff 7          | 225        |

order that they fit the curvature of a given fault zone. Appendix D gives a listing of the earthquakes included in each source. The depth of each source was computed as an average hypocentral depth of all events included in the sources. Earthquakes with no or limited information were not included in this averaging process. However, they were considered in determining the location and the seismicity of the sources.

The recurrence relationship for each individual source was obtained by fitting a regression line through the data for that source and obtaining

$$\ln N(M) = \alpha + \beta M \quad 3-2$$

where  $N(M)$  = Number of events above magnitude  $M$

$M$  = Richter magnitude

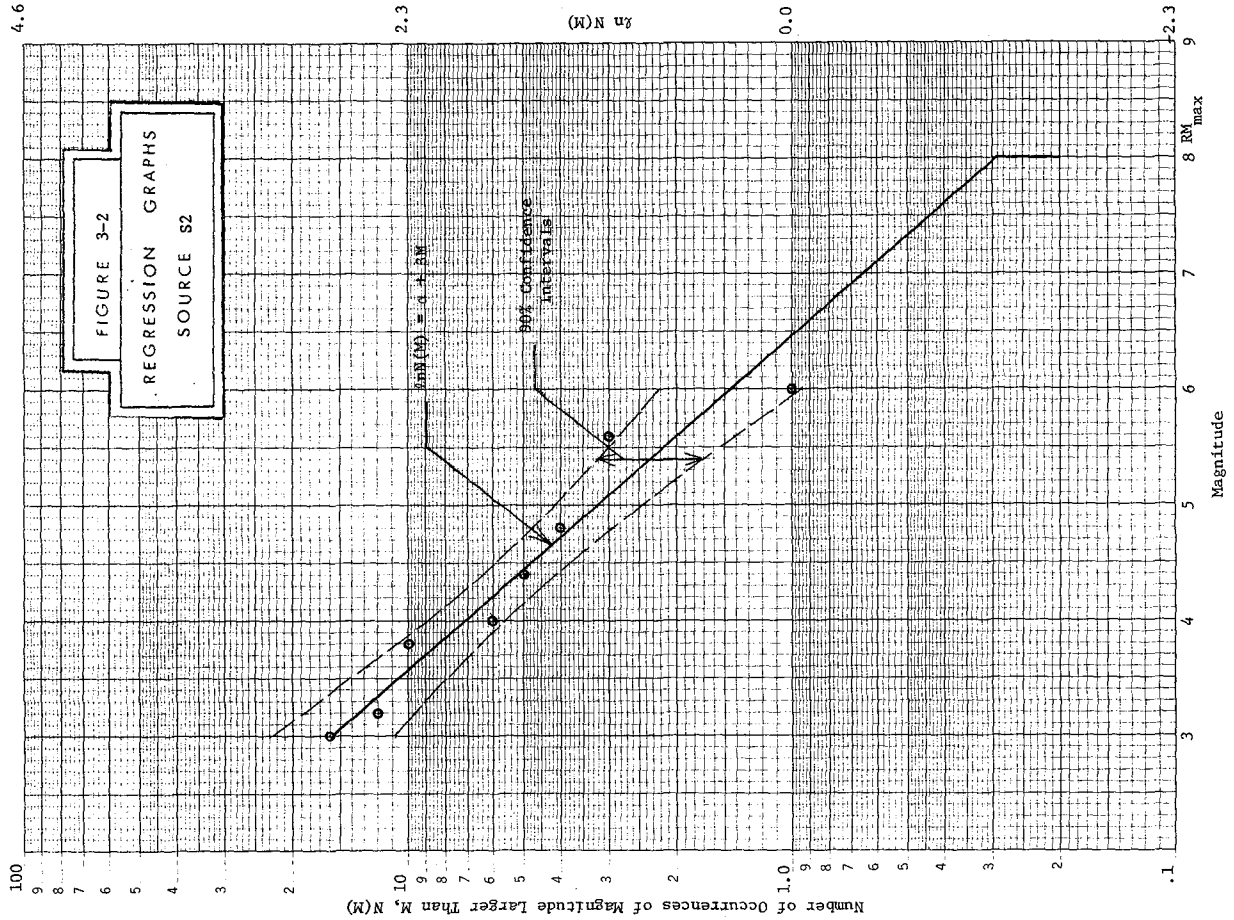
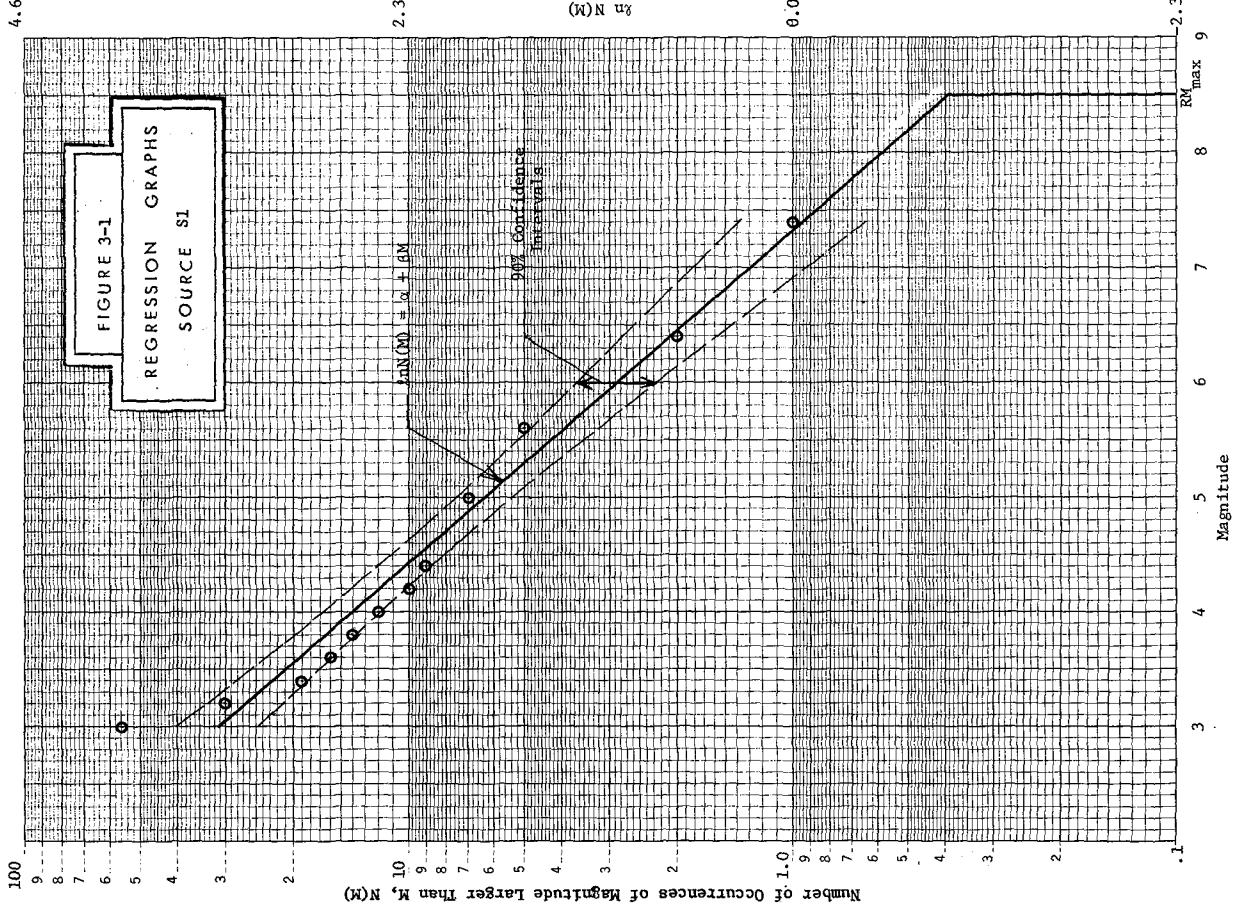
$\alpha$  and  $\beta$  are regression constants.

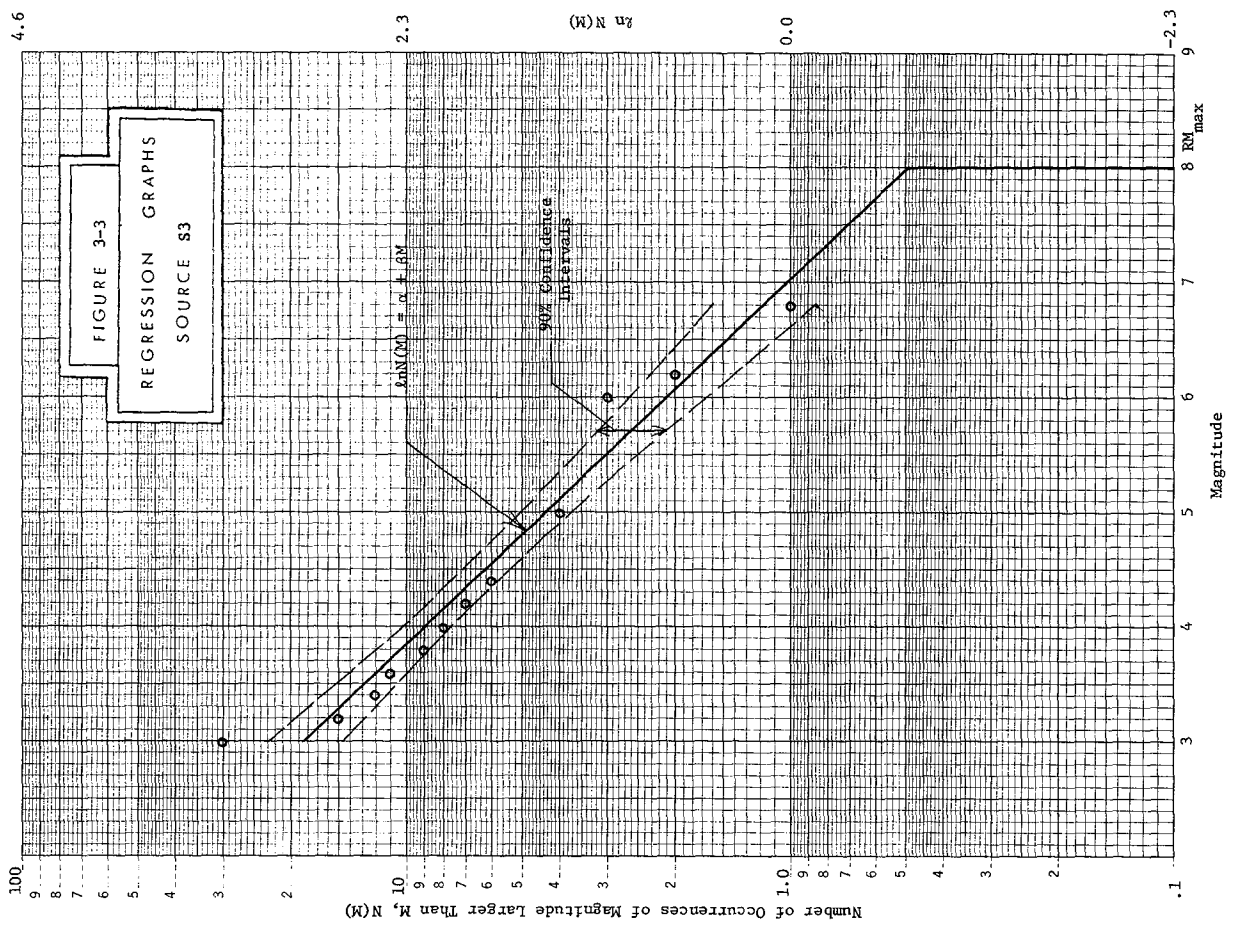
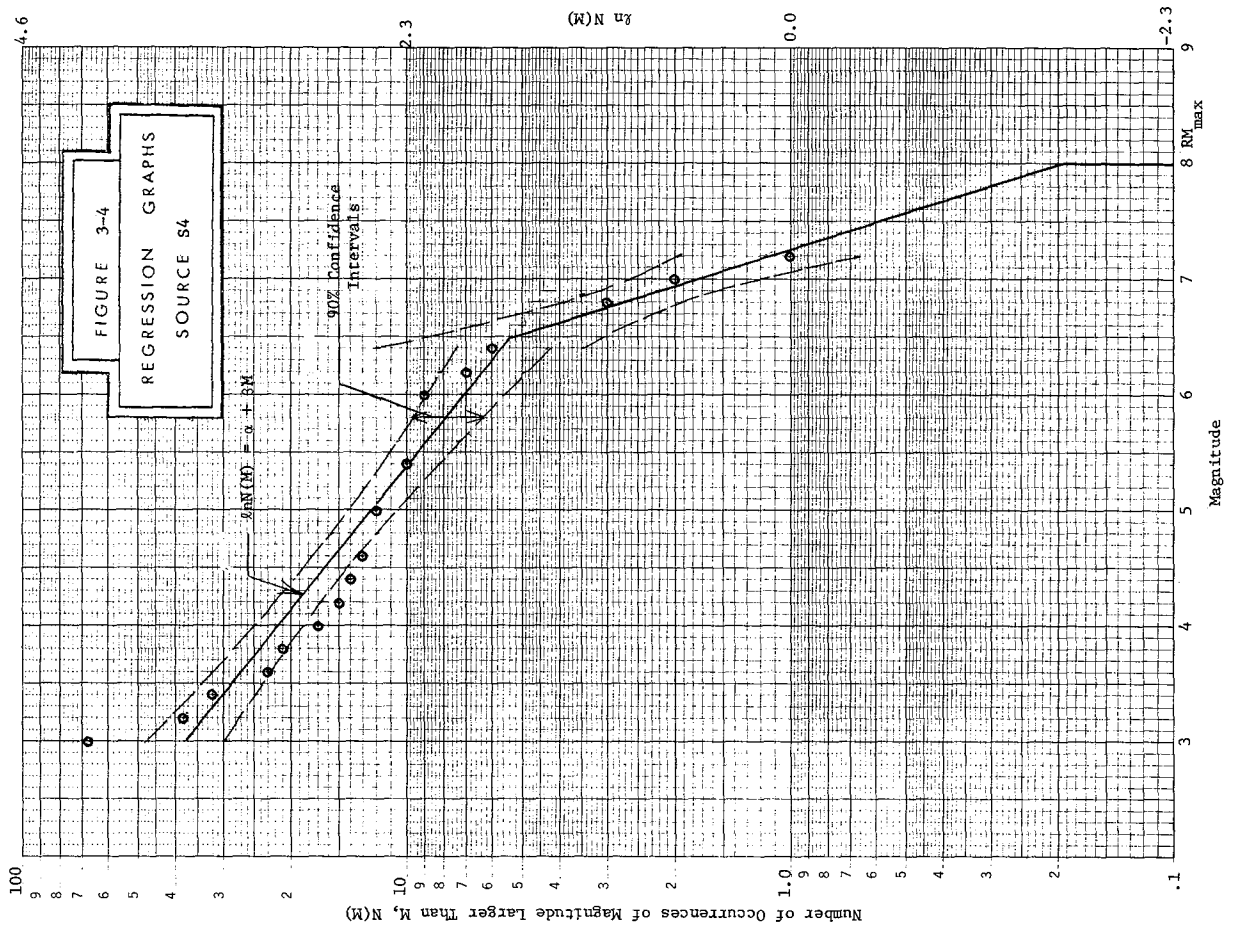
$\alpha$  is a measure of the number of events above magnitude 0 for a given source and  $\beta$  is a measure of the seismic severity for a given source. The larger the negative value of  $\beta$ , the smaller the seismic severity. For many sources, a single regression line gave erroneous results because the interpolation of the line beyond the range of data indicated unreasonably high magnitude occurrences. For such cases, two regression lines were fitted to the data, and a geologically consistent upper magnitude value was used for cutoff. (See Figures 3-1 through 3-18.) Table 3-3 gives a summary of  $\alpha'$  and  $\beta$  values for each source and the magnitude cutoff point for each source.

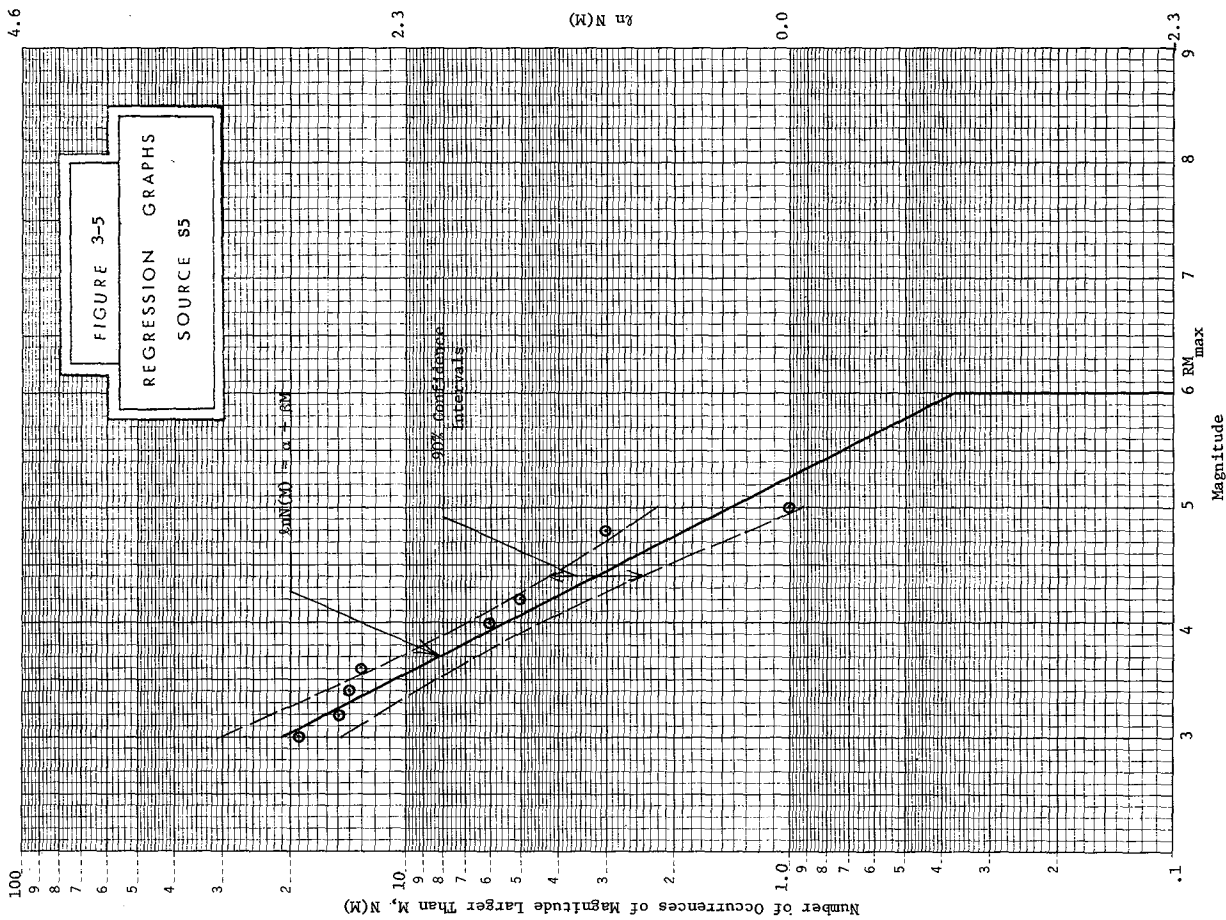
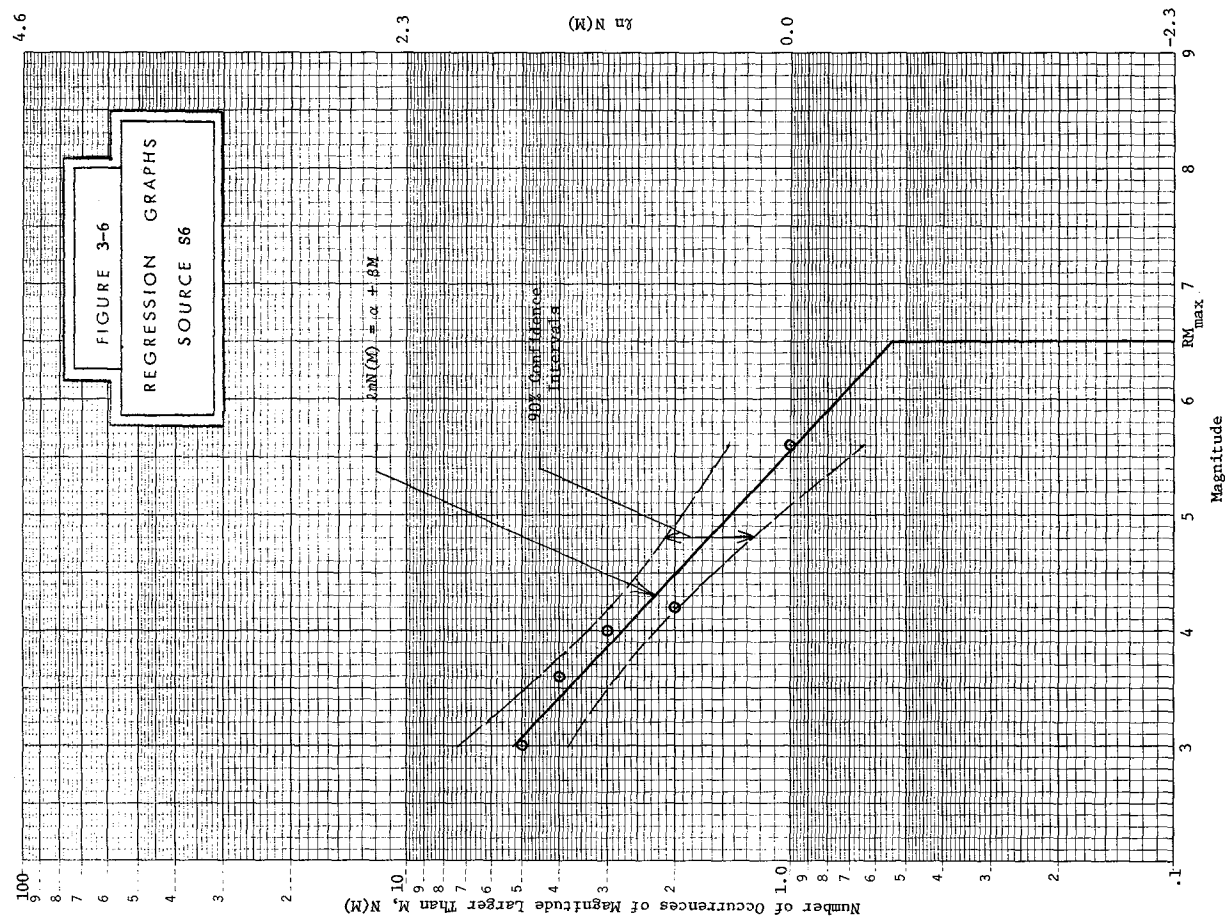
$$\text{Let } N'(M) = \frac{N(M)}{AT} \quad \text{for an area source} \quad 3-3a$$

$$= \frac{N(M)}{LT} \quad \text{for a line source} \quad 3-3b$$

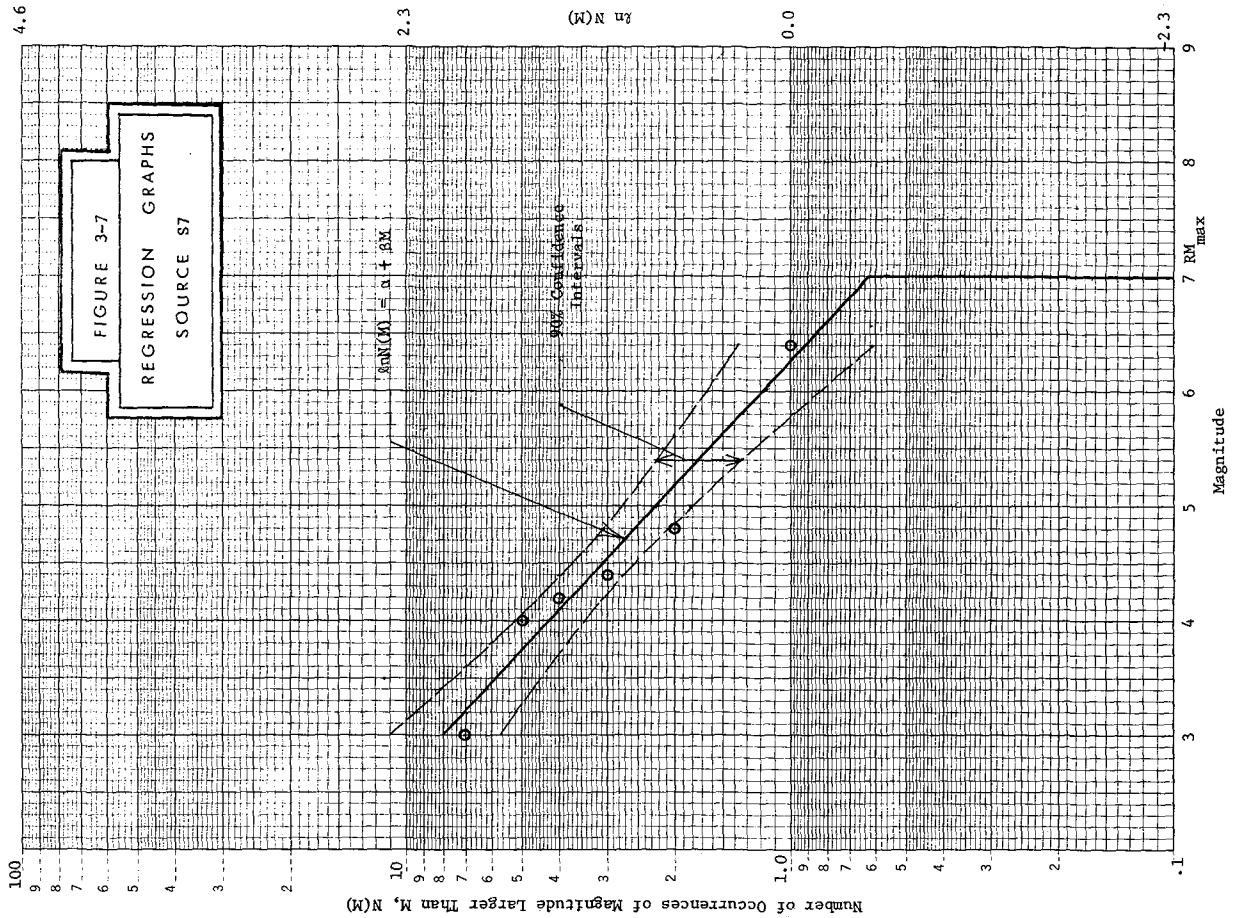
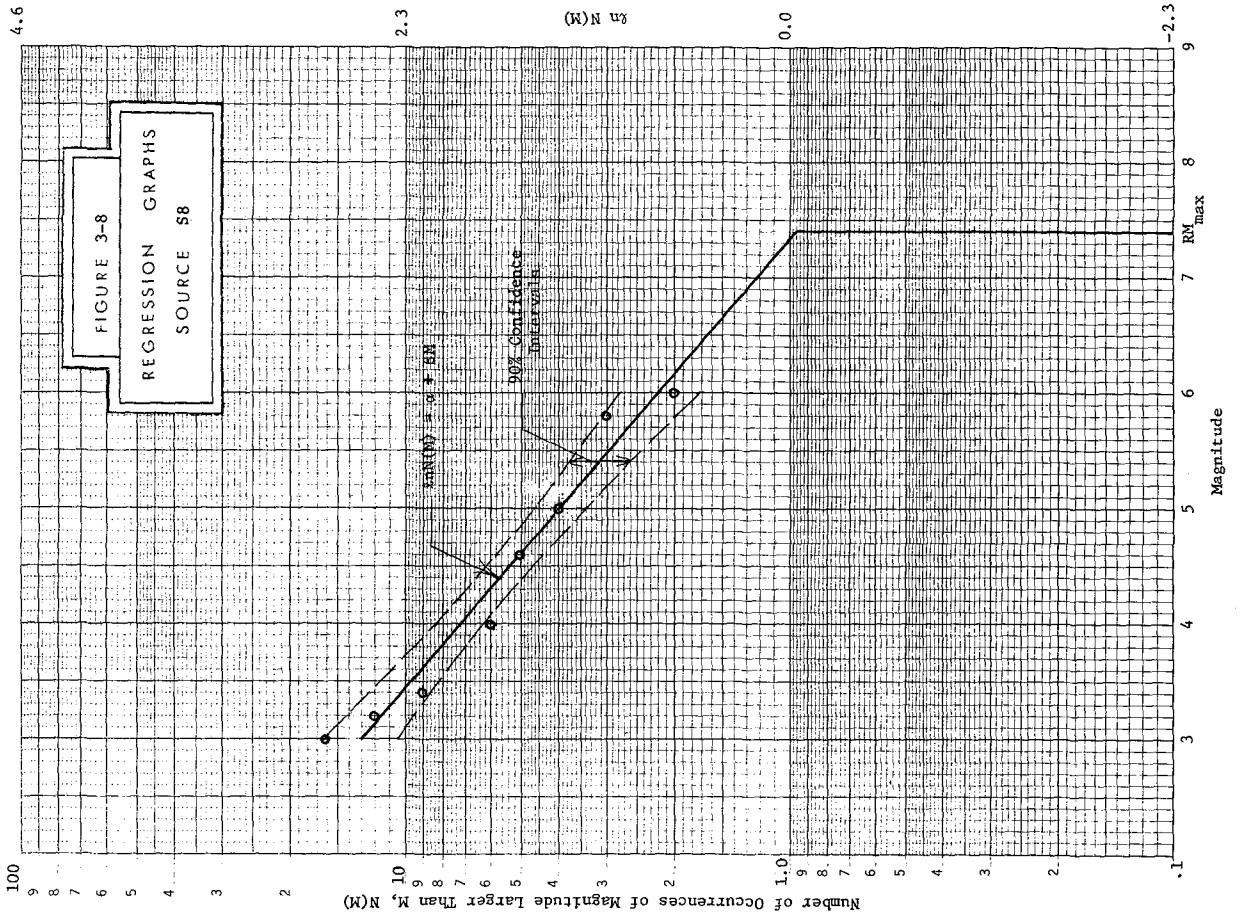
where  $L$  = Length of the line source

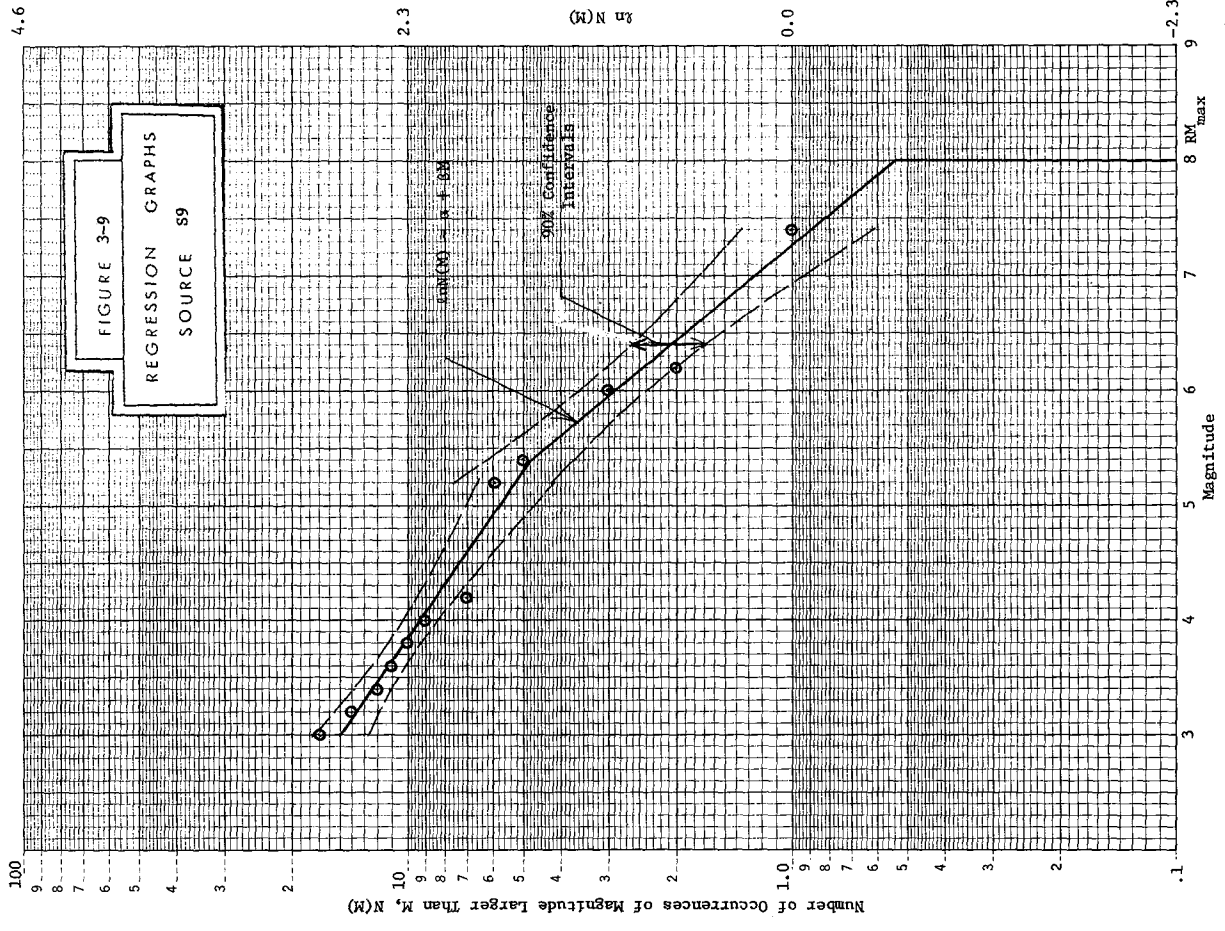
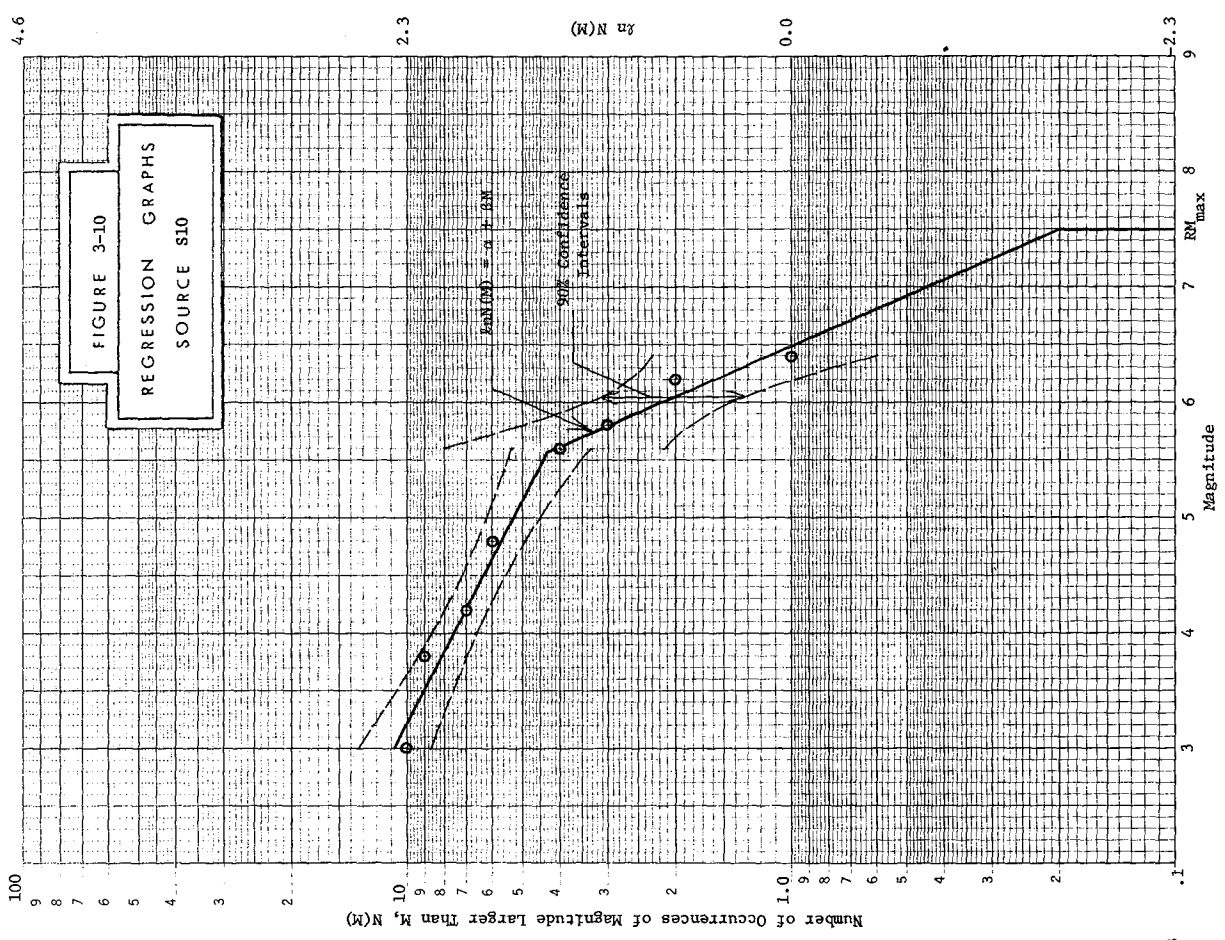


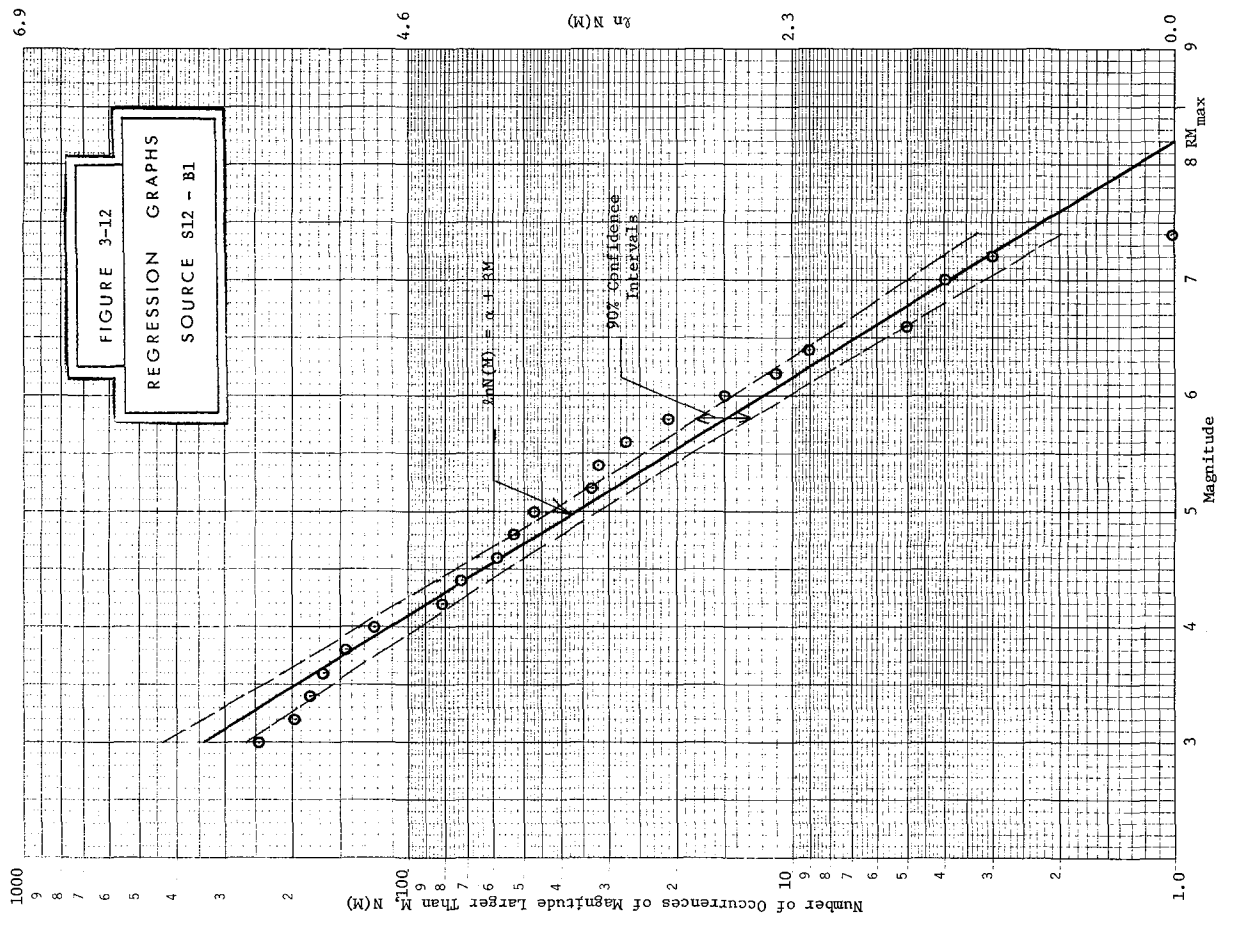
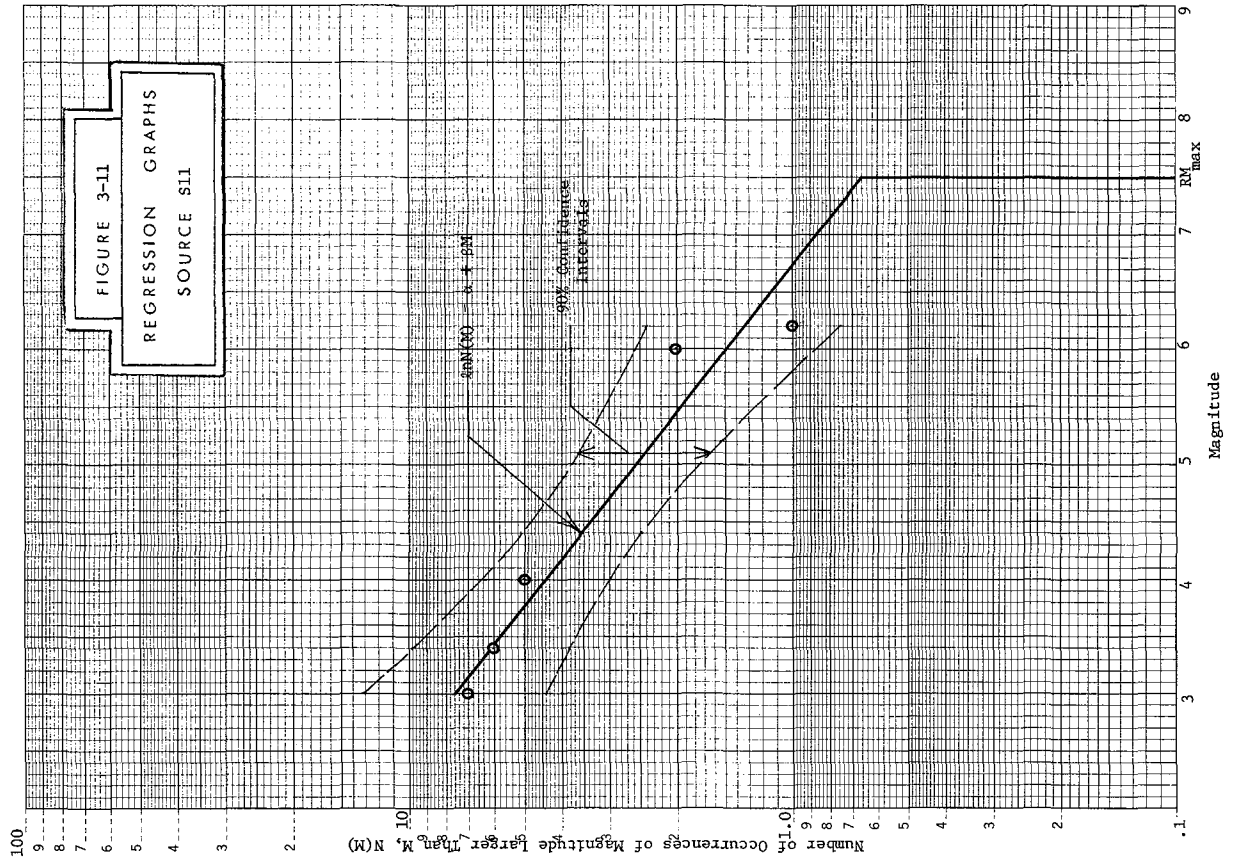


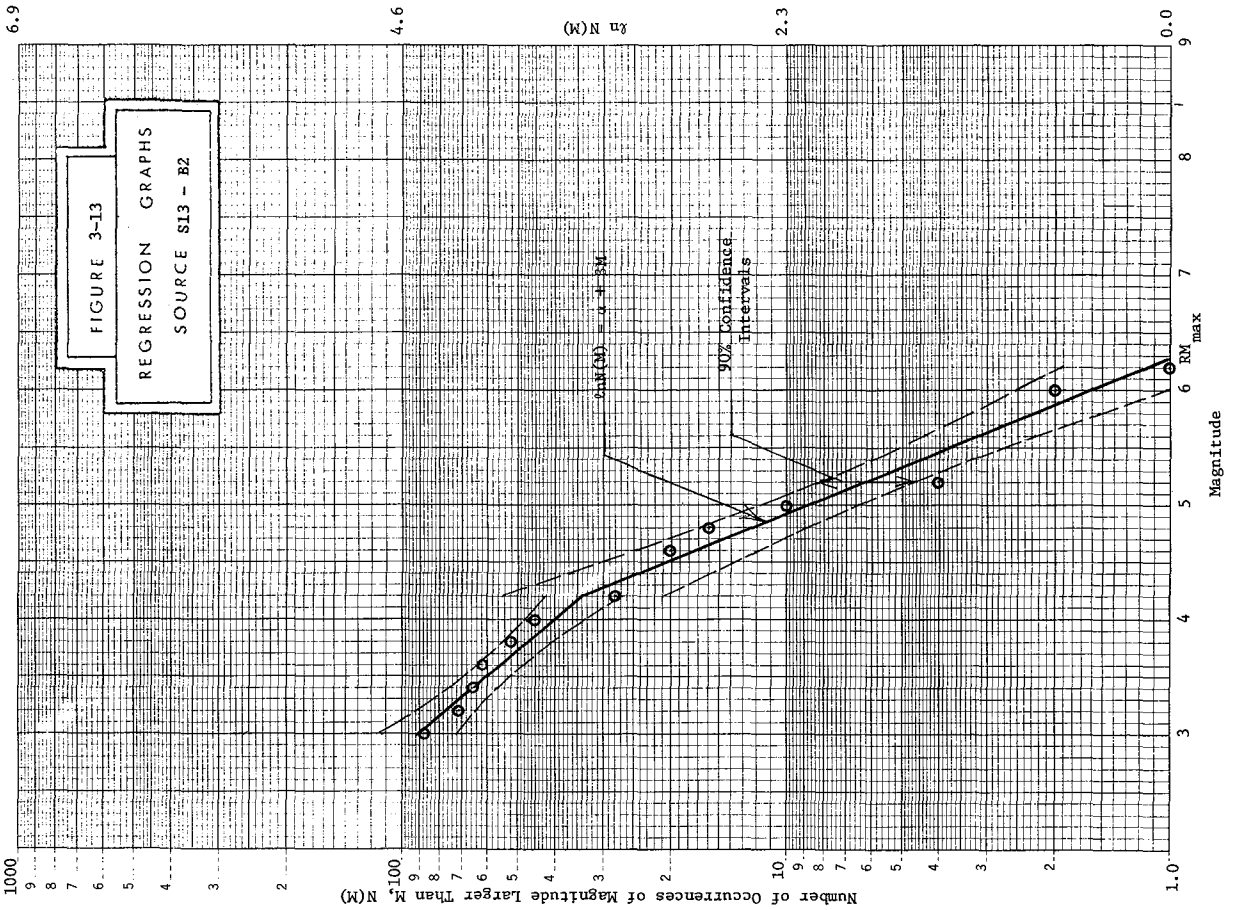
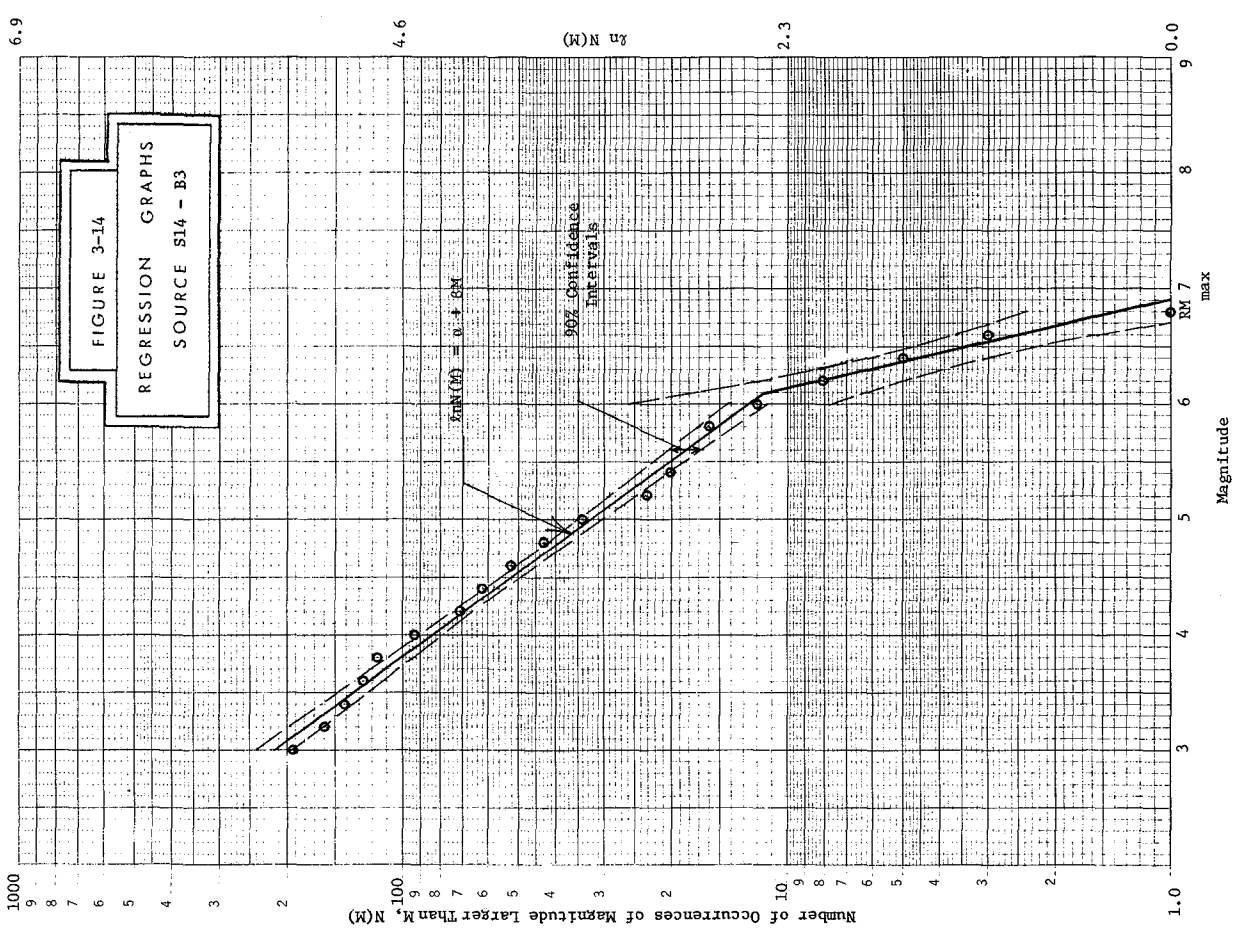


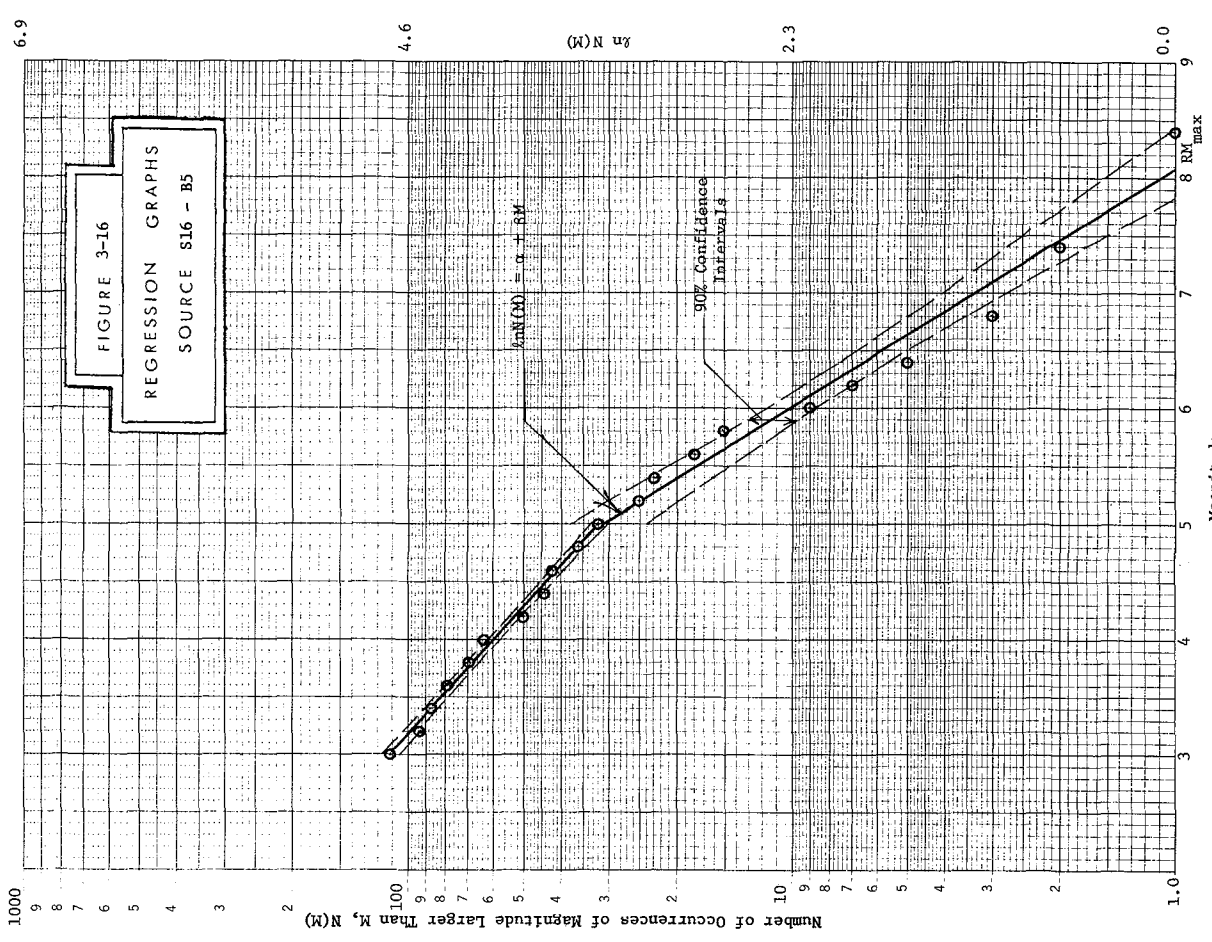
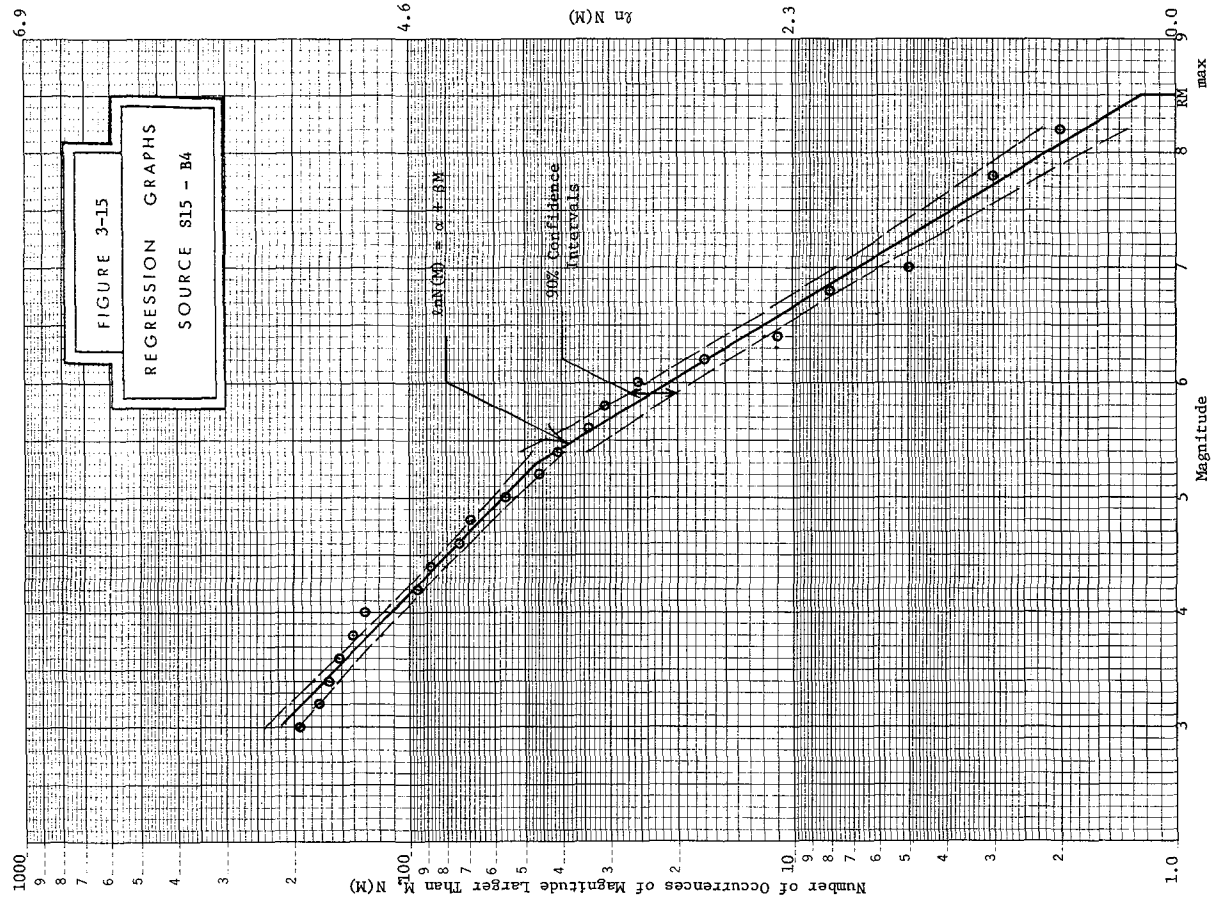












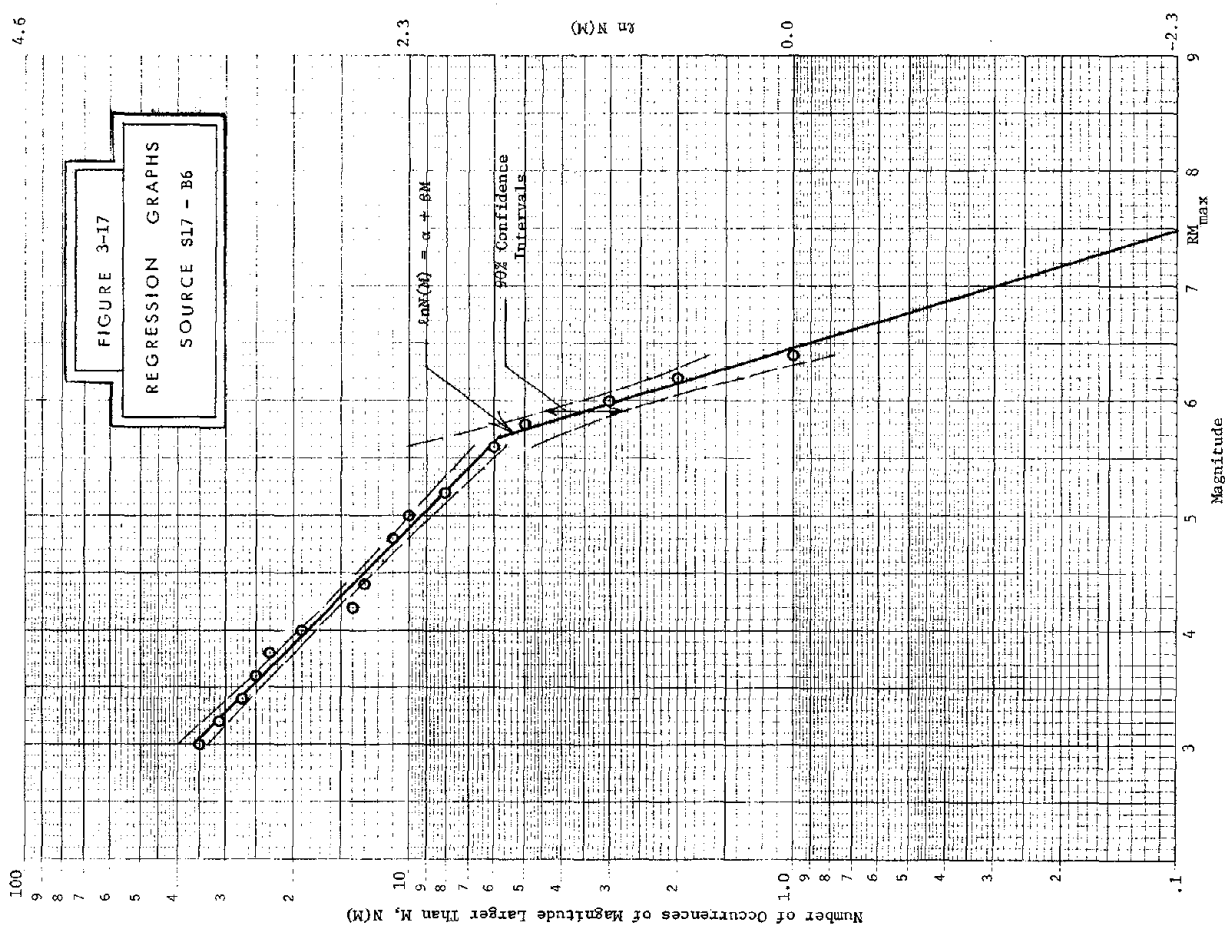
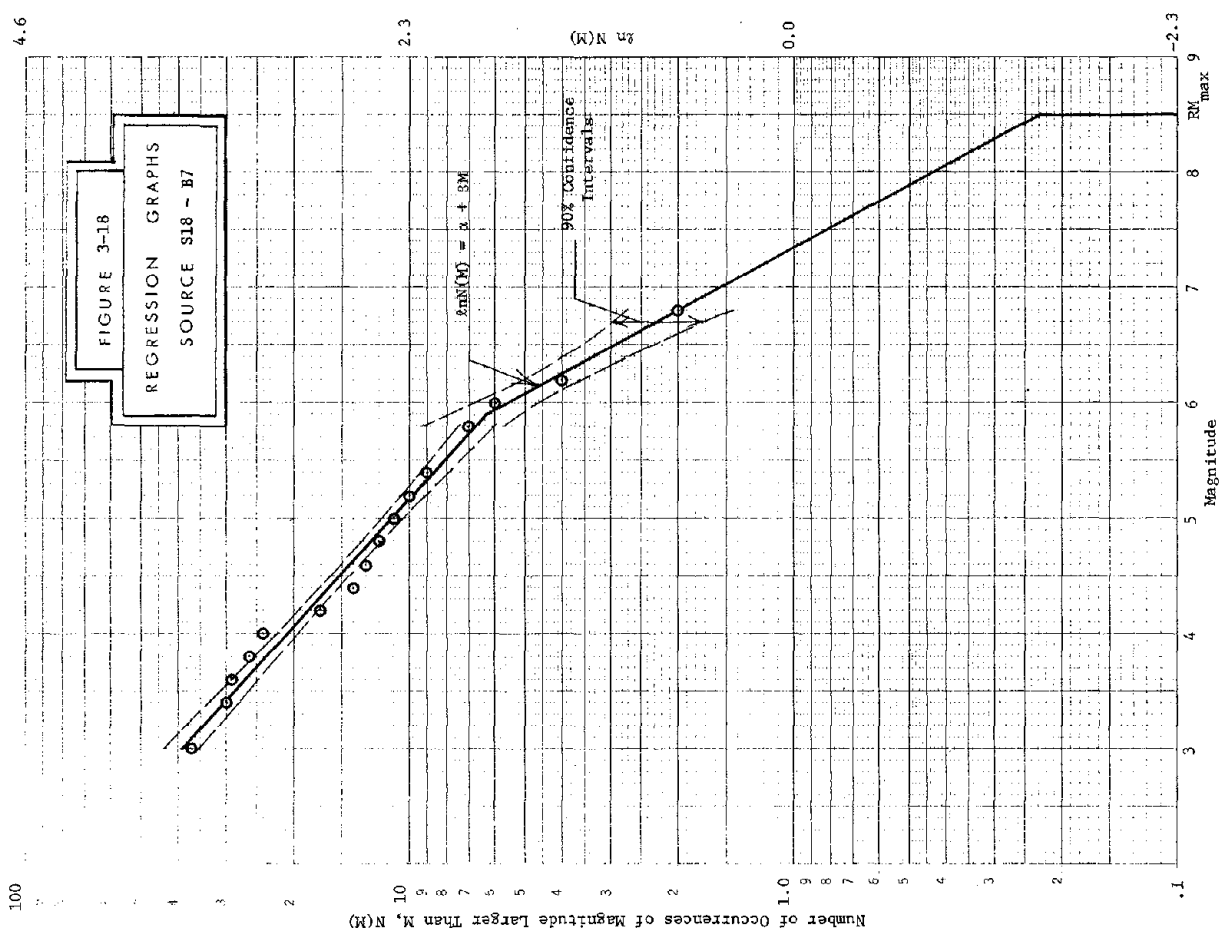


TABLE 3-3

## Line Source Information

| Source   | $\alpha_1'$ | $\beta_1$ | $\alpha_2'$ | $\beta_2$ | $RM_{\max}$<br>(recorded) | $RM_{\max}$<br>(cutoff) | Length<br>(degrees) | RMBK |
|----------|-------------|-----------|-------------|-----------|---------------------------|-------------------------|---------------------|------|
| S1       | -0.0632     | -0.7955   |             |           | 7.5                       | 8.5                     | 4.84                |      |
| S2       | -0.4323     | -0.7961   |             |           | 6.0                       | 8.0                     | 3.52                |      |
| S3       | -0.2486     | -0.7211   |             |           | 6.8                       | 8.0                     | 2.72                |      |
| S4       | -0.0023     | -0.5579   | 10.4009     | -2.1618   | 7.3                       | 8.0                     | 2.68                | 6.49 |
| S5       | 2.1556      | -1.3430   |             |           | 5.0                       | 6.0                     | 1.84                |      |
| S6       | -0.4873     | -0.6460   |             |           | 5.7                       | 6.5                     | 0.79                |      |
| S7       | -0.0187     | -0.6363   |             |           | 6.5                       | 7.0                     | 0.73                |      |
| S8       | 1.2284      | -0.5930   |             |           | 6.1                       | 7.4                     | 0.30                |      |
| S9       | -0.8131     | -0.4753   | 1.0776      | -0.8249   | 7.5                       | 8.0                     | 1.87                | 5.41 |
| S10      | -1.7146     | -0.3591   | 5.1519      | -1.5889   | 6.5                       | 7.5                     | 2.36                | 5.58 |
| S11      | -0.7026     | -0.5397   |             |           | 6.3                       | 7.5                     | 1.12                |      |
| S12 - B1 | 2.9900      | -1.1107   |             |           | 7.4                       | 8.5                     | 6.34                |      |
| S13 - B2 | 0.6477      | -0.8098   | 4.3405      | -1.6915   | 6.3                       | 8.5                     | 7.15                | 4.19 |
| S14 - B3 | 1.8929      | -0.9447   | 14.1711     | -2.9741   | 6.8                       | 8.5                     | 7.39                | 6.05 |
| S15 - B4 | 1.0108      | -0.6659   | 3.4927      | -1.1340   | 8.3                       | 8.6                     | 7.72                | 5.30 |
| S16 - B5 | 0.1986      | -0.6222   | 2.5188      | -1.0979   | 8.4                       | 8.6                     | 7.83                | 4.88 |
| S17 - B6 | -0.7732     | -0.6795   | 8.1302      | -2.2485   | 6.5                       | 8.5                     | 8.02                | 5.67 |
| S18 - B7 | -0.8487     | -0.6290   | 3.0325      | -1.2867   | 6.9                       | 8.5                     | 8.02                | 5.90 |

A = Area of the area source

T = Time for which the data was obtained

$N'(M)$  = Normalized mean number of events above magnitude M for unit-time (1 year) and unit-area or unit length.

Then

$$\ln N'(M) = \alpha' + \beta M \quad 3-4$$

where  $\alpha' = \alpha - \ln(AT)$  for area source

$= \alpha - \ln(LT)$  for line source.

Table 3-3 shows values of  $\alpha'$ ,  $\beta$  and the upper cutoff magnitude as described previously. The table gives the values of  $\alpha'$  and  $\beta$  normalized with respect to degrees of latitude and longitude and develop the forecasting model. RMBK is the Richter Magnitude where the slope of the recurrence lines change.

#### III-4 Confidence Levels

In addition to the  $\alpha'$  and  $\beta$  parameters, the 90% confidence intervals are computed for all regression lines. These intervals are measures of the accuracy of the regression fit to the data. The 90% confidence intervals indicate that there is 0.90 probability that the true value of  $\ln N(M)$  lies within these bounds (see Benjamin and Cornell, 1970, for further detail). For the cases of double regression lines, two separate confidence intervals are obtained for each segment. The discontinuity at the intersection of the two regression lines is to be expected since the least-squares fit for each segment is performed using its corresponding set of data independent of each other. The confidence lines for all the sources appear close to the regression line showing high confidence in the recurrence relationships. The confidence bounds are particularly good for sources S12 - B1 to S18 - B7, and are worse for S10 and S11. In general the intervals are narrow whenever there are many data points.



TABLE 3-4

## Statistics of Regression Lines

| Source   | $\sigma_1$ | $\rho_1$ | $V_1$ | $\sigma_2$ | $\rho_2$ | $V_2$ |
|----------|------------|----------|-------|------------|----------|-------|
| S1       | 0.245      | 0.955    | .109  |            |          |       |
| S2       | 0.253      | 0.930    | .150  |            |          |       |
| S3       | 0.221      | 0.949    | .120  |            |          |       |
| S4       | 0.235      | 0.887    | .084  | 0.159      | 0.970    | .177  |
| S5       | 0.276      | 0.927    | .148  |            |          |       |
| S6       | 0.141      | 0.963    | .147  |            |          |       |
| S7       | 0.163      | 0.956    | .145  |            |          |       |
| S8       | 0.153      | 0.960    | .087  |            |          |       |
| S9       | 0.111      | 0.912    | .048  | 0.140      | 0.972    | .135  |
| S10      | 0.089      | 0.955    | .046  | 0.193      | 0.932    | .242  |
| S11      | 0.254      | 0.930    | .210  |            |          |       |
| S12 - B1 | 0.292      | 0.965    | .085  |            |          |       |
| S13 - B2 | 0.128      | 0.900    | .032  | 0.287      | 0.957    | .149  |
| S14 - B3 | 0.100      | 0.988    | .025  | 0.244      | 0.952    | .168  |
| S15 - B4 | 0.081      | 0.978    | .018  | 0.166      | 0.979    | .067  |
| S16 - B5 | 0.041      | 0.991    | .010  | 0.213      | 0.968    | .102  |
| S17 - B6 | 0.073      | 0.985    | .026  | 0.154      | 0.966    | .148  |
| S18 - B7 | 0.086      | 0.976    | .031  | 0.078      | 0.987    | .053  |

The constant standard deviation  $\pm \sigma$  and the coefficient of variation  $V$  of  $\ln N(M)$ , which are indicators of the variability of the estimated regression line, are also determined for each source. The values of  $\pm \sigma$  and  $V$  are given in Table 3-4. Most of the sources have low coefficients of variation  $V$ . It should be noted that in the model used for the probabilistic acceleration computations (Section III-3) a maximum of four sources have an affect on any given point in the worst cases. For such points the largest coefficient of variation due to the input data is found to be 14.5% assuming the contribution from all four sources are independent and the following relationship is true:

$$V_y^2 = V_{x1}^2 + V_{x2}^2 + V_{x3}^2 + \dots + V_{xn}^2 \quad 3-5$$

Most points will have a value for  $V$  much lower than 14.5%.

The correlation coefficients,  $\rho$ , shown in Table 3-4 express the degree of linear dependence of  $\ln N(M)$  to  $M$ . For almost all sources  $\rho$  is greater than 0.9, a value of 1.0 indicating perfect linear correlation, thus denoting relatively high correlation between the variables.

### III-5 Conclusion

Some of the limitations of the data were already discussed throughout the chapter. The variability of the regression equation was determined for each source and an upper bound uncertainty in the final probabilistic loads due to them were also obtained. Other sources of error in the modeling, such as fault locations, focal depths, and cutoff magnitudes, will be discussed through subsequent chapters.

It should be emphasized that as additional data becomes available to give more reliable information on epicentral locations, the methodology presented in this research project will be able to modify the results accordingly.

## CHAPTER IV

### PROBABILISTIC SEISMIC LOADING

#### Scope

In this chapter, the Poisson model used to analyse the available data is discussed briefly. The available data presented in Chapter III is used to obtain iso-acceleration maps of Guatemala. In addition, the cumulative probability distributions on peak ground acceleration for eight cities in Guatemala are also presented.

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#### IV-1 Introduction

In Chapter III, the data base was discussed and the limitations of the available information and the approximations made in using the seismic data for Guatemala were reviewed. The recurrence relationships associated with all the seismic sources for the region were also presented. These recurrence relationships give the mean number of events of magnitude greater than  $M$  due to a seismic source and a time period. The normalized recurrence relationships are obtained when the mean number of events above a specified magnitude  $M$  is normalized with respect to the time period of the historic data and the length of the source in the case of line sources or the area of the source in the cases of area sources.

To determine the seismic risk in Guatemala, it is necessary to forecast earthquake events in this region. The three statistical models for forecasting that have been used previously are:

- (1) Poisson model
- (2) Markov model
- (3) Bayesian model

In the present study, the Poisson model used by Shah et al. (1975) is adapted because of its simplicity, its widespread use in literature, and because the results it gives are very similar to results arising from more complex models such as the Markov Chain model. The development of the Poisson model is presented in Appendix A for completeness and further reference. In addition the Bayesian model of Mortgat (1976) is also used to develop seismic hazard maps for Guatemala for peak ground acceleration as well as for time duration of an earthquake. A review of the Bayesian model and the resulting hazard maps from its application to Guatemala are discussed in Appendix B.

#### IV-2 Iso-Acceleration Maps for Guatemala

Considering the Poisson model, the seismic sources and their corresponding frequency relationships were already discussed in the previous chapter. The maximum Richter magnitude was assigned to each source by considering the largest earthquake event associated with each source, the length of the source and the fault system to which it may belong. The values of the largest Richter magnitude recorded ( $RM_{\max}$  recorded) and the cutoff Richter magnitude ( $RM_{\max}$  cutoff) for each source is listed in Table 3-3. The attenuation relationship (see Equation A-9) used in the modeling is Esteva's (1973) equation given below:

$$a = \frac{5000 \exp(0.8M)}{(R_h + 40)^2} \quad 4-1$$

where  $a$  = Peak Ground Acceleration in  $\text{cm}/\text{sec}^2$

$M$  = Richter magnitude

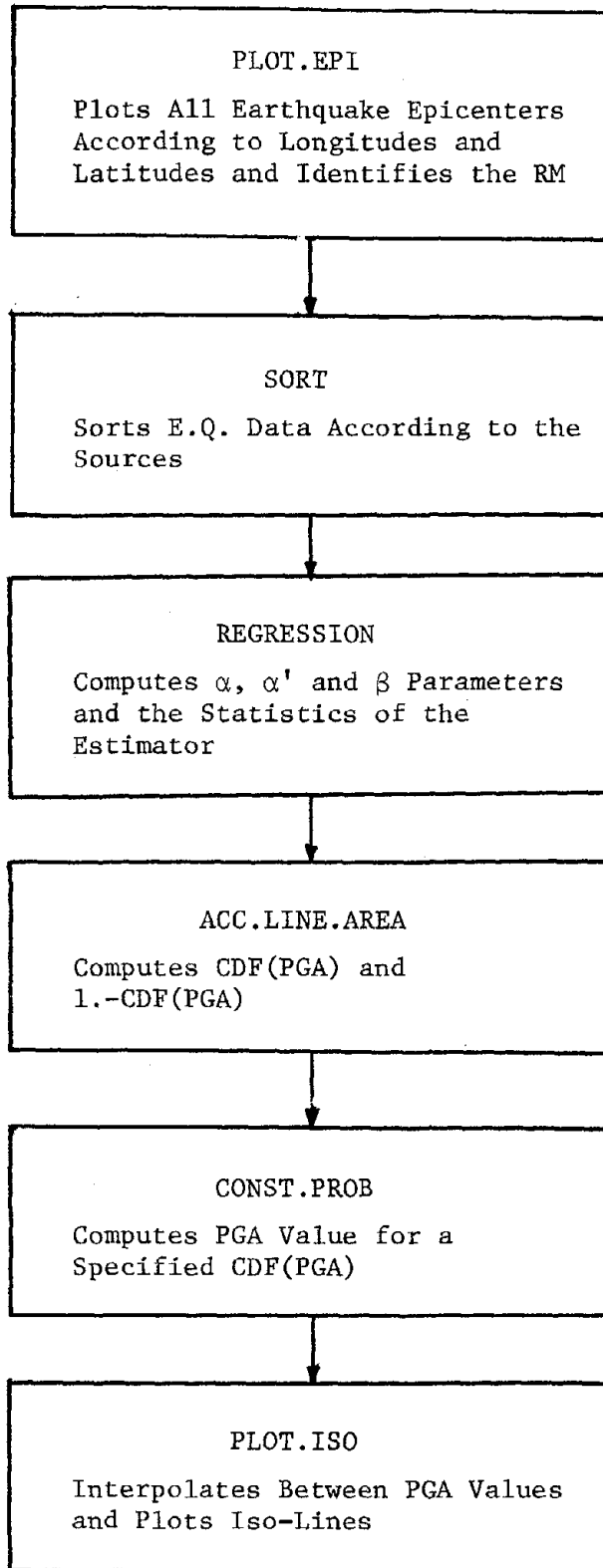
$R_h$  = Hypocentral distance in km.

Since no strong motion records from major earthquake events or high amplitude accelerograph records are available for the Guatemalan region, it was

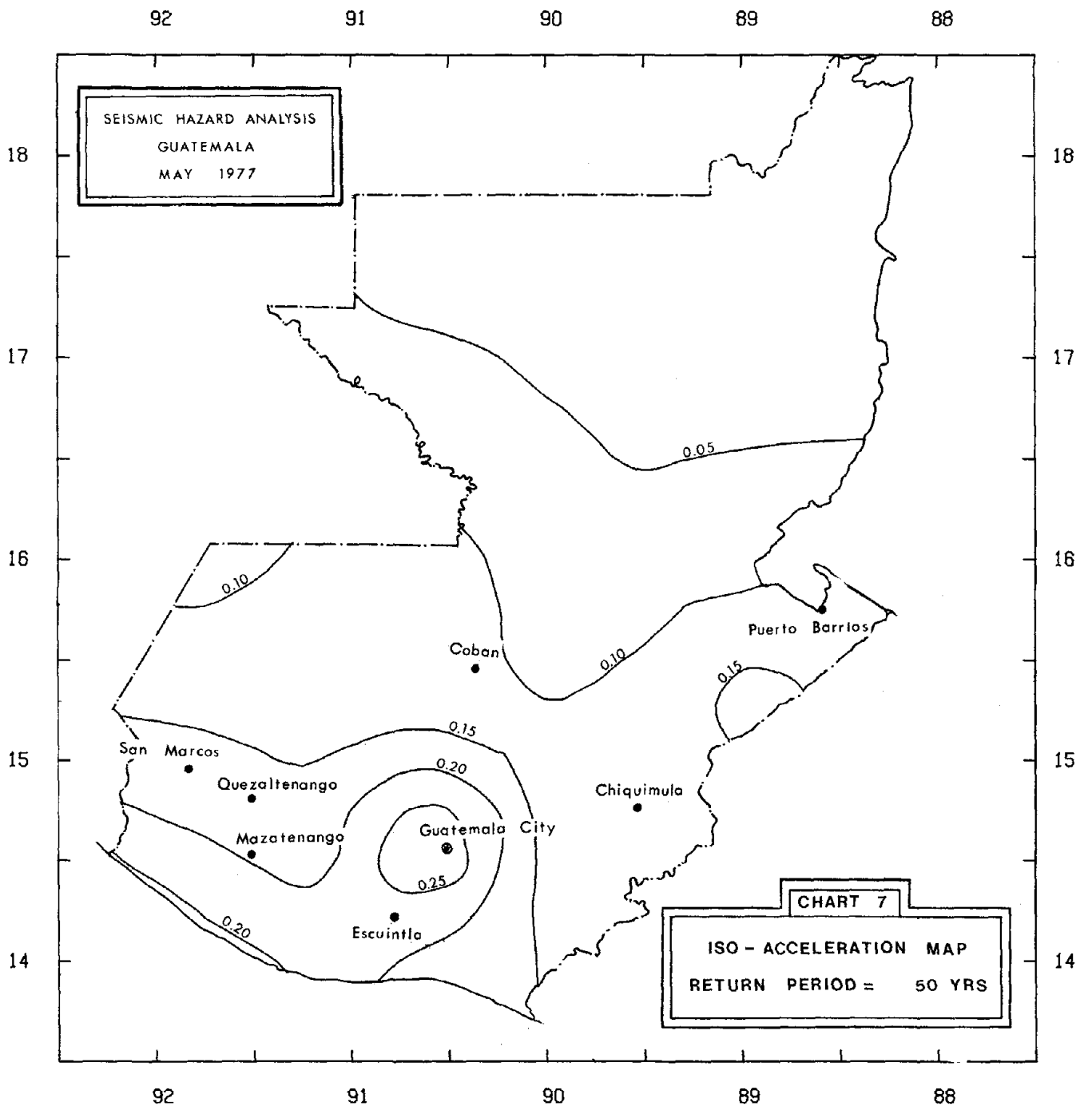
not possible to adjust Esteva's equation or to obtain an attenuation relationship specific to Guatemala. The general feeling among Guatemalan geologists and seismologists is that the ground motion attenuates at a slower rate than in California, for example, however the peak ground acceleration values corresponding to each Richter magnitude are presumed to be lower. It is hoped that with the better instrumentation and additional strong motion recording in the future, these hypotheses can be scientifically substantiated, and consequently can be used for developing a new attenuation relationship. Such changes, it should be noted, can be easily incorporated in the seismic hazard model.

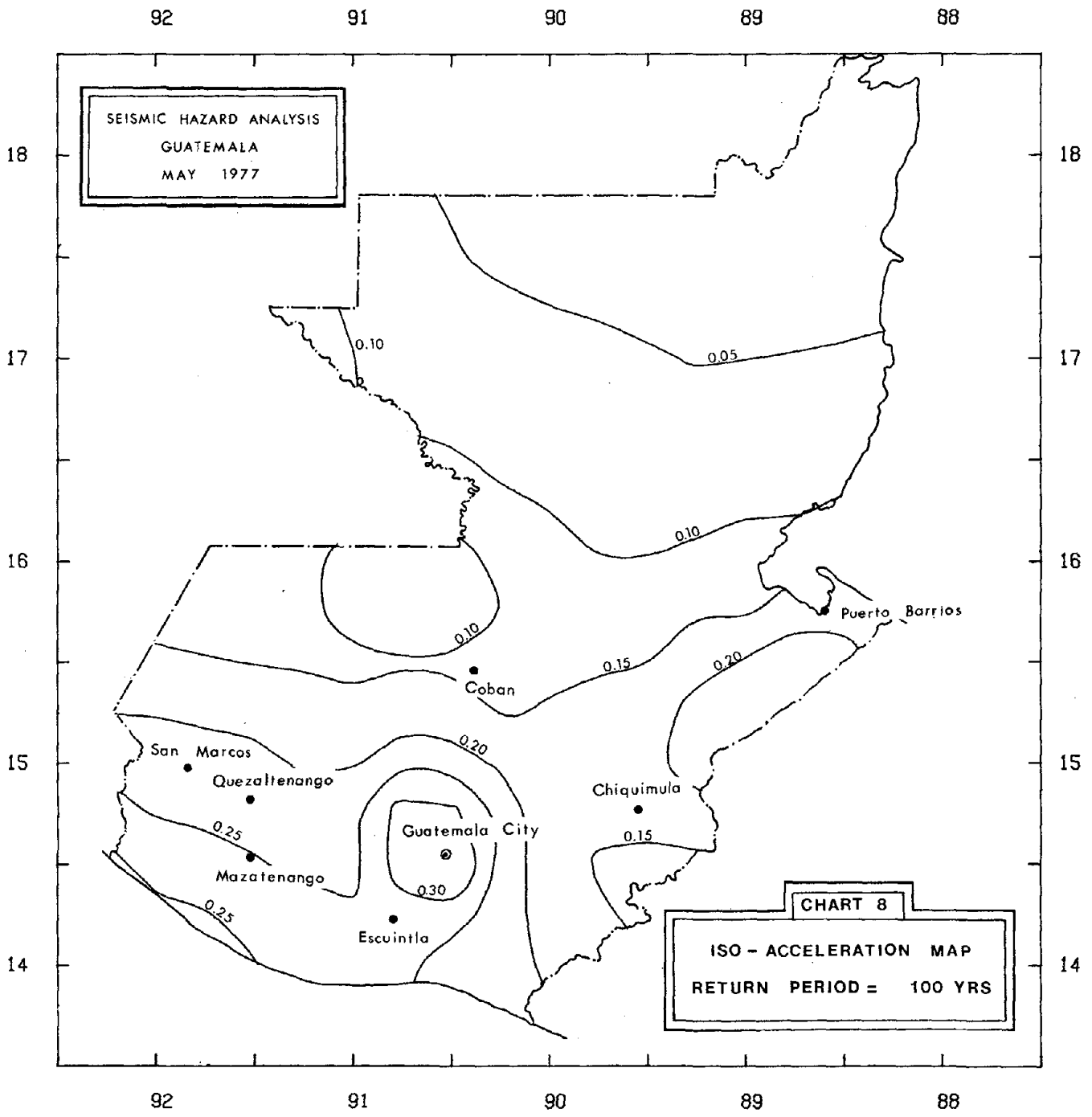
All of the above information along with the source location and focal depth, are put together in Equation A-34 (see Appendix A) to compute the cumulative probability distributions on peak ground acceleration. The entire country of Guatemala is divided into a grid of  $1/4^\circ$  longitude by  $1/4^\circ$  latitude spacings. The CDF of peak ground acceleration at each node on the grid is evaluated using the computer program ACC.LINE.AREA (see Appendix E). Then for a specified probability of exceedence, in other words for a specified hazard level, the peak ground acceleration values computed from the CDF's at each node are obtained using the computer procedure CONST.PROB. The iso-acceleration lines, which are lines of equal acceleration, are drawn by interpolation between nodal values of PGA. The program PLOT.ISO used in obtaining the hazard map contains both an interpolation routine and a graphical plotting routine. The flow chart below shows the major steps in the development of the map and the computer programs used at each step.

Iso-acceleration maps also referred to as seismic hazard maps are obtained for return period of events of 50, 100, 500, and 1000 years. Charts 7 to 10 show the iso-acceleration maps for these return periods.

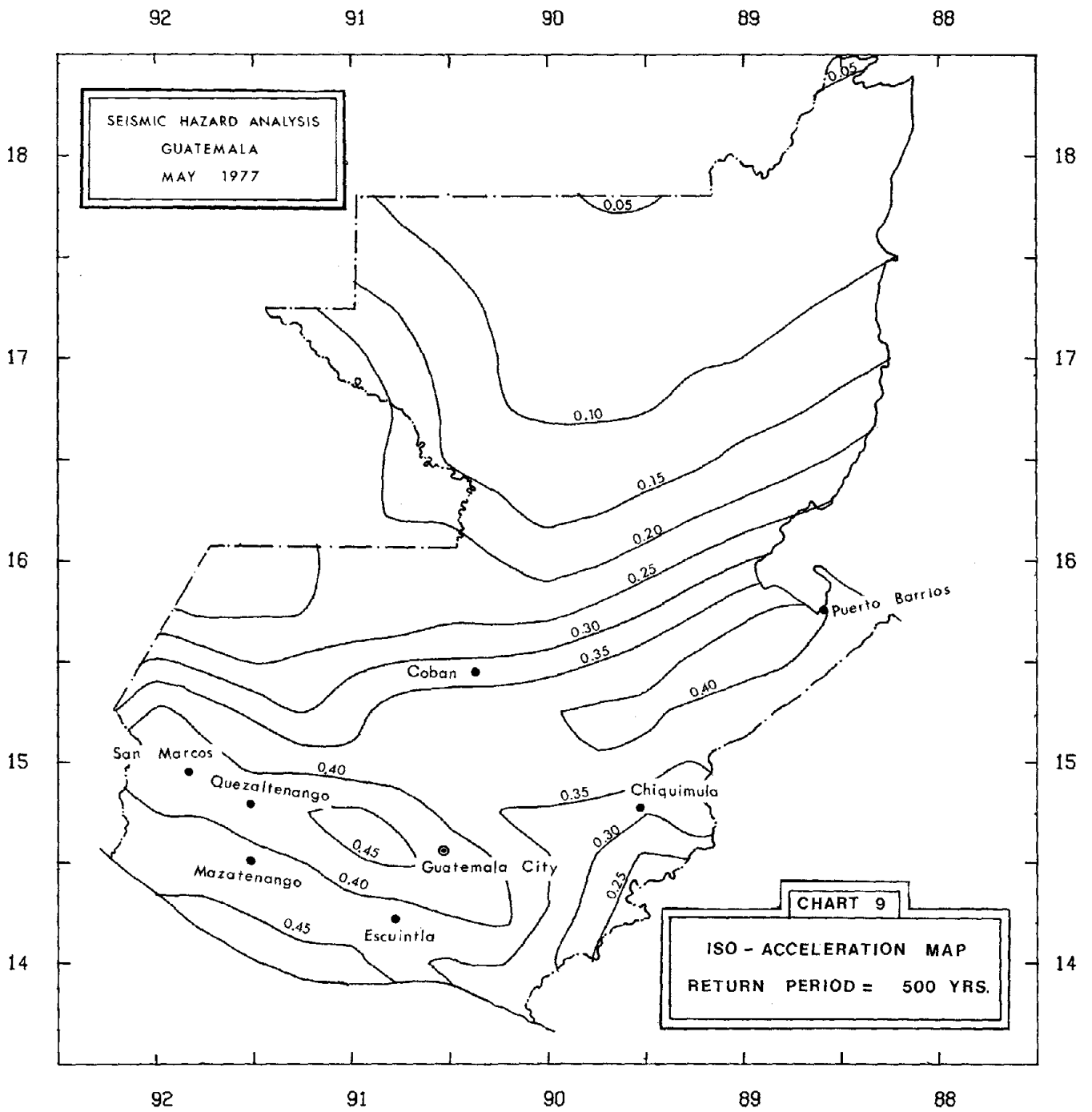


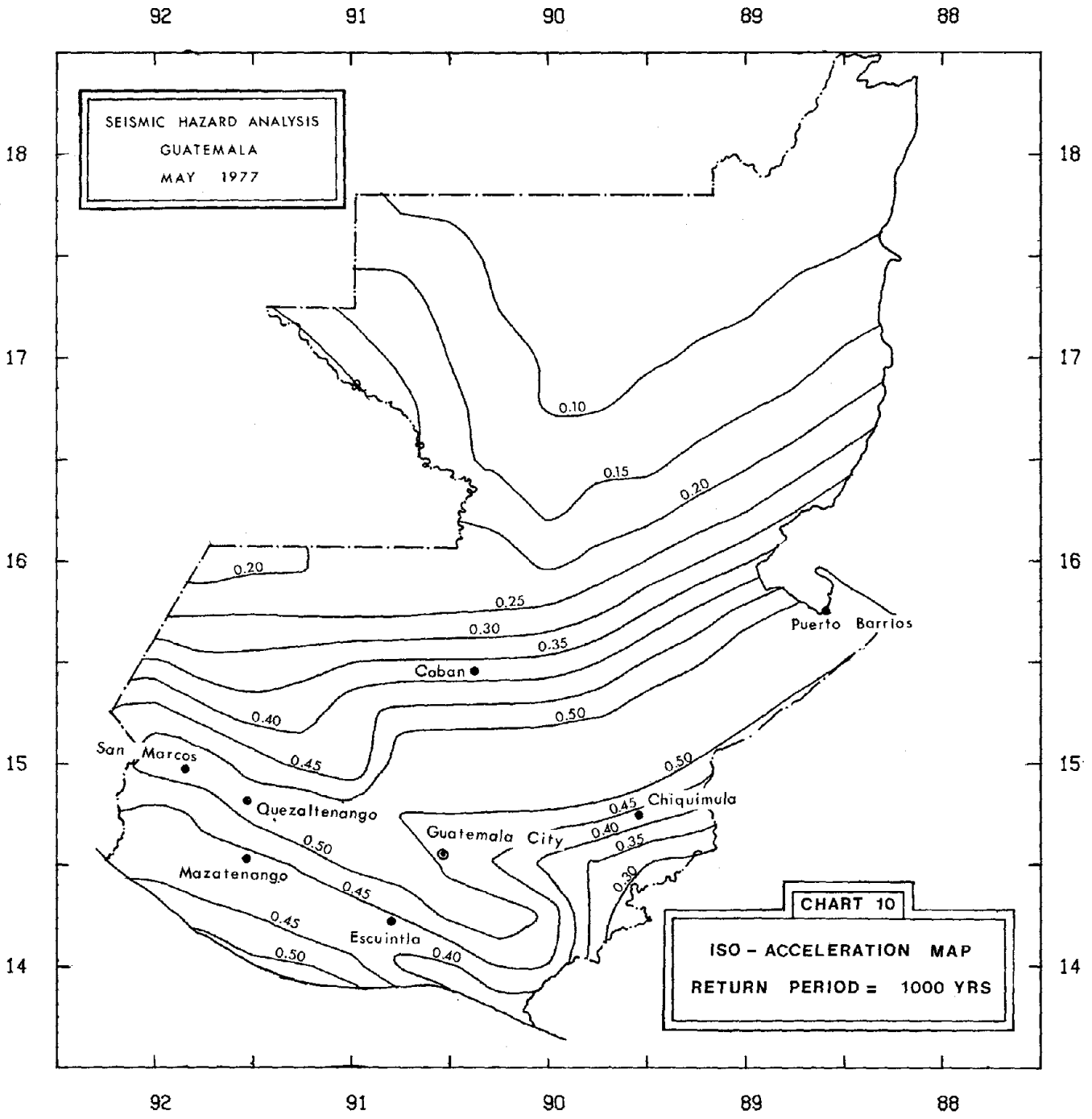
A listing of all programs is given in Appendix E.











The concept of return period and economic life in conjunction with exceed-  
ence probabilities will be discussed in Chapter V. Detailed methodology  
describing the use of these maps for structural design will be presented  
in the Part II report of the present study.

From these maps a low seismicity range is observed in the northern  
part of Guatemala. The high seismicity region in the south-southwestern  
part of Guatemala reaches a peak at the intersection of the Motagua fault,  
the Benioff zone and the Mixco fault. These high and low regions are in  
compliance with both the geologic as well as the seismologic data for the  
country.

In addition to the seismic hazard maps for the whole country, the  
following cities were studied in detail:

1. Guatemala City
2. Quezaltenango
3. Mazatenango
4. Chiquimula
5. Puerto Barrios
6. Escuintla
7. San Marcos
8. Copan

Figures 4-1 through 4-8 show the cumulative distribution function of peak  
ground acceleration for each of the cities. The results are presented  
for 50 years of future exposure time. The seismicity of Guatemala City  
appears to be the highest and for Coban it is the lowest.

The implications of these probability values and the corresponding  
acceleration values will be discussed in Chapter V.

When the cumulative distribution plots for different cities are com-  
pared, one can see the relative seismicity in terms of peak ground accel-  
eration for each city. In conclusion, it can be said that one method of

representing seismic risk is by means of cumulative distribution function plots of Figures 4-1 through 4-8.

The engineering interpretation of these results will be presented in the Part II report of the present study. It should also be pointed out that the iso-acceleration maps and any zoning based on such maps only represent macro characteristics. The macrozoning of the country should be modified with site-specific micro characteristics to microzone a given region. In that case, the local geotechnical and geological features (such as those discussed in Chapter II) should be incorporated together with the macro characteristics presented in this chapter.

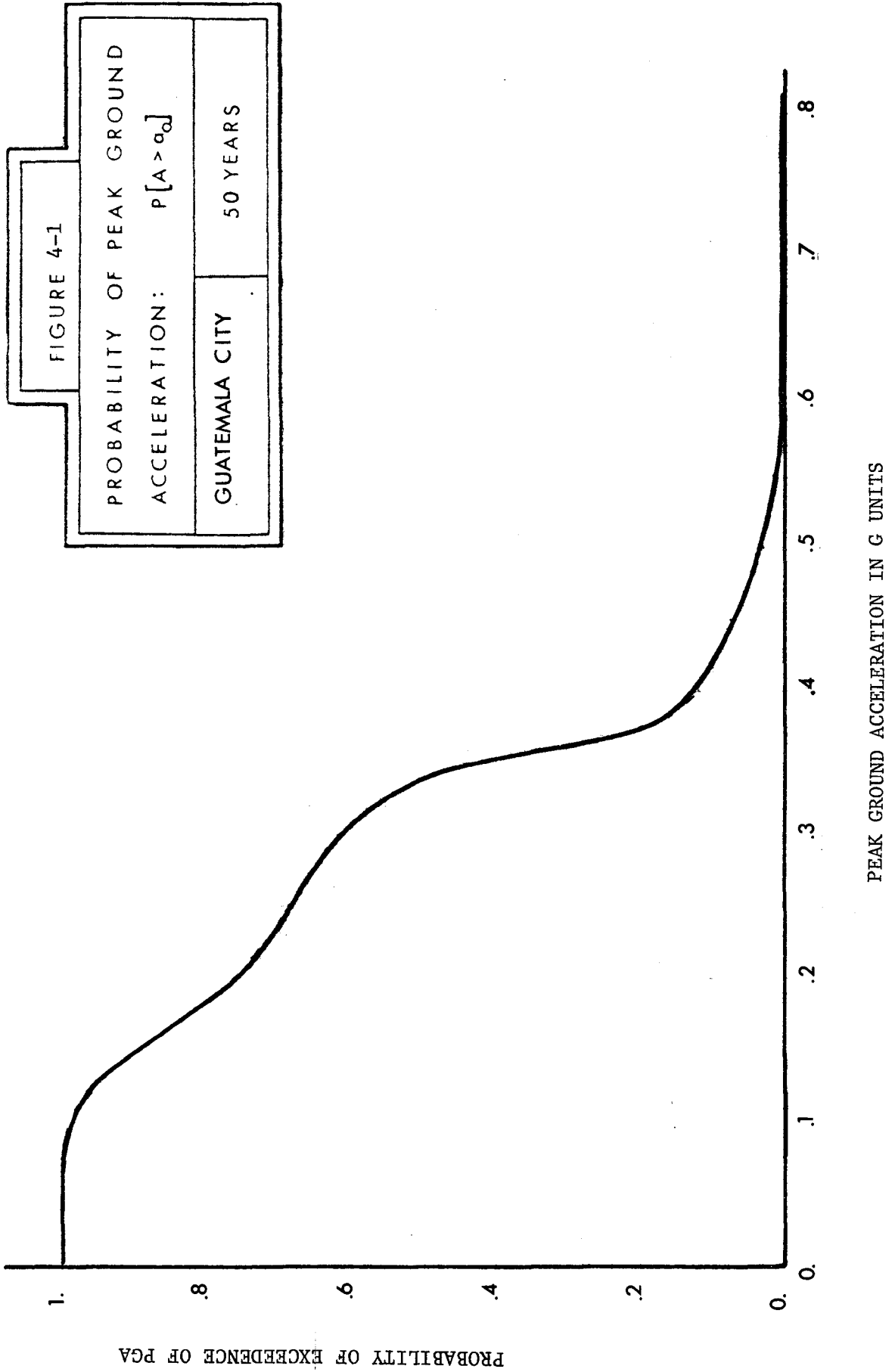


FIGURE 4-1

PROBABILITY OF PEAK GROUND

ACCELERATION:  $P[A > a_d]$

GUATEMALA CITY

50 YEARS

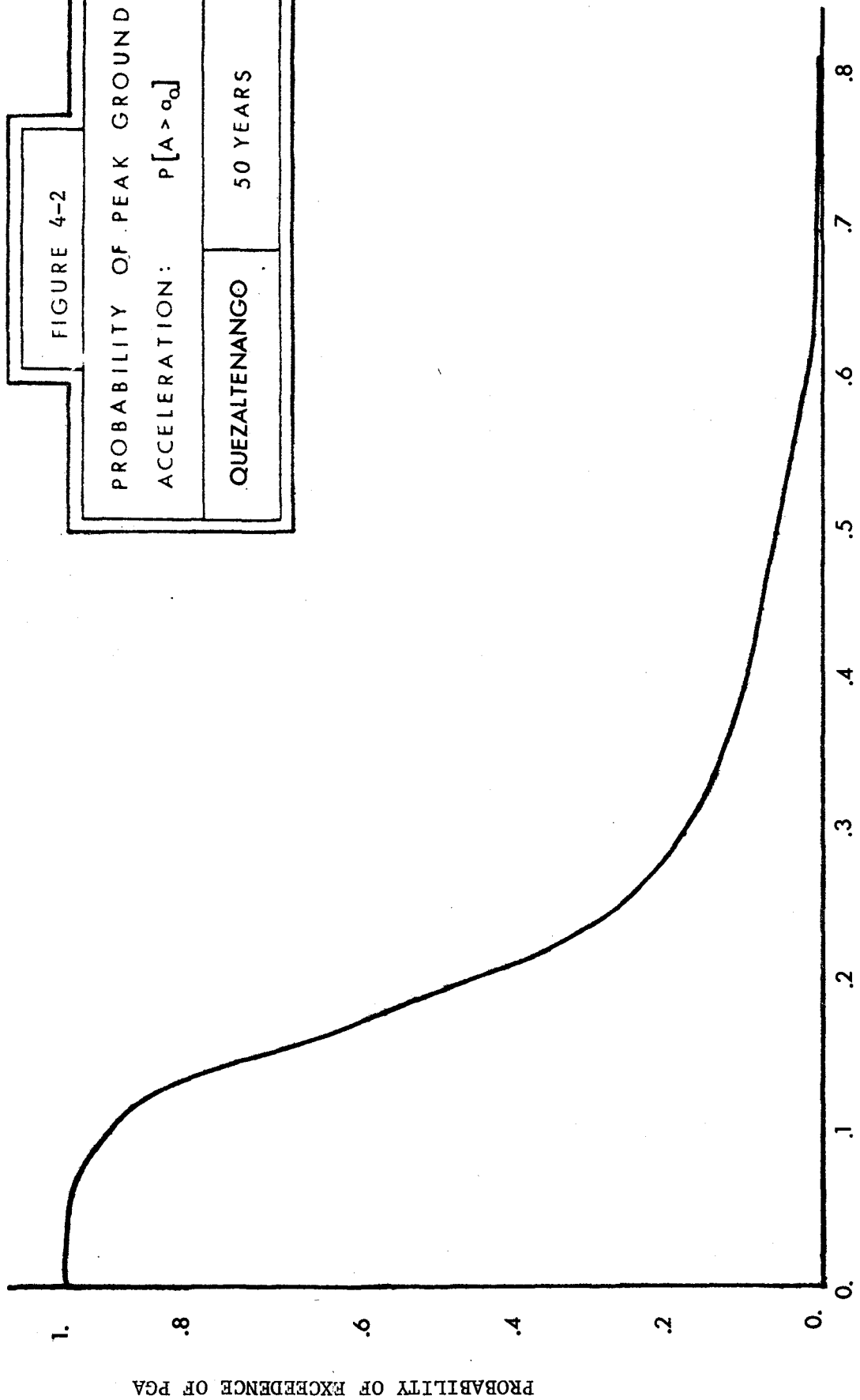


FIGURE 4-2

PROBABILITY OF PEAK GROUND  
ACCELERATION:  $P[A > a_0]$   
QUEZALTENANGO 50 YEARS

PROBABILITY OF EXCEEDENCE OF PGA

PEAK GROUND ACCELERATION IN G UNITS

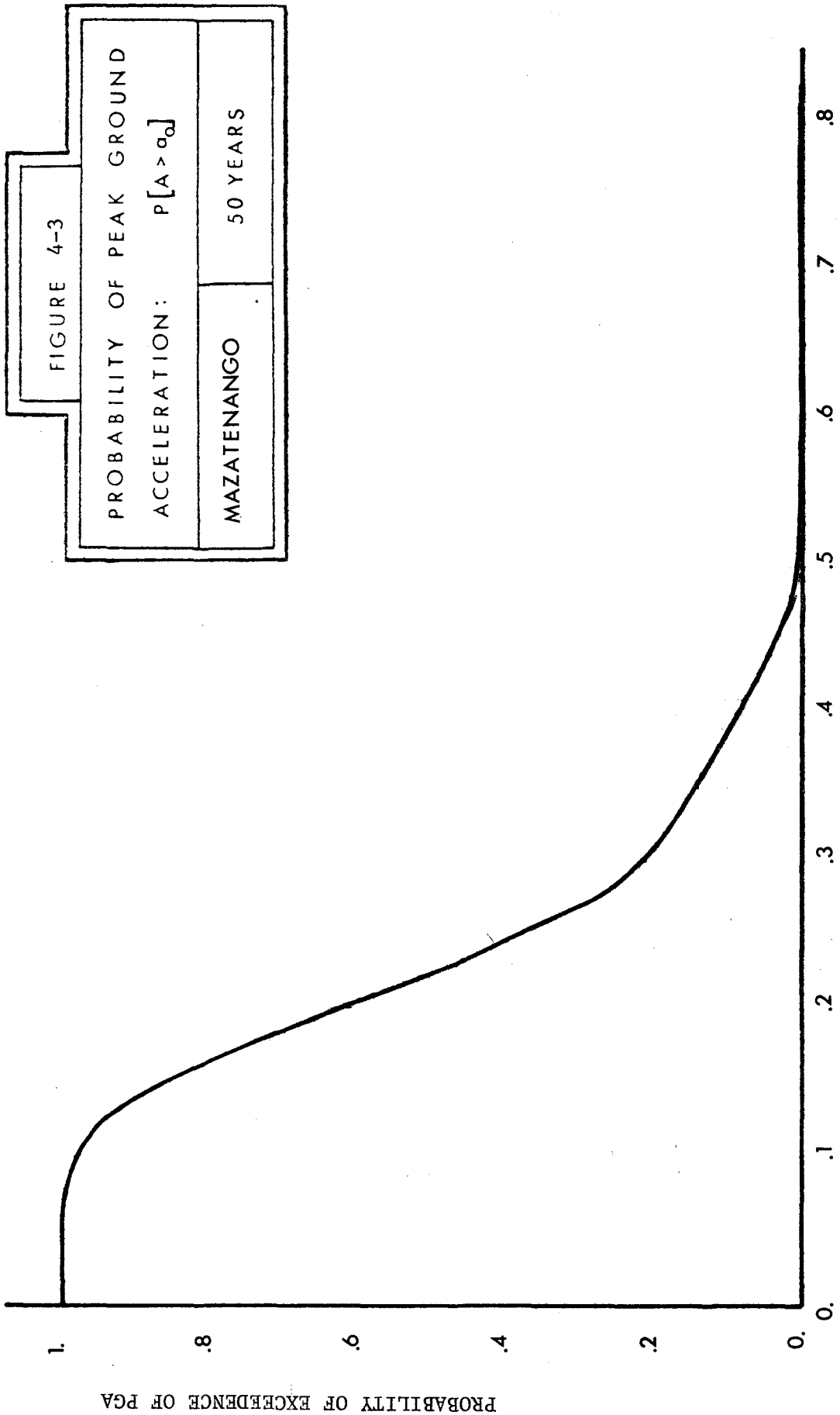


FIGURE 4-3

PROBABILITY OF PEAK GROUND  
ACCELERATION:  $P[A > a_0]$   
MAZATENANGO 50 YEARS

PROBABILITY OF EXCEEDENCE OF PGA

PEAK GROUND ACCELERATION IN G UNITS

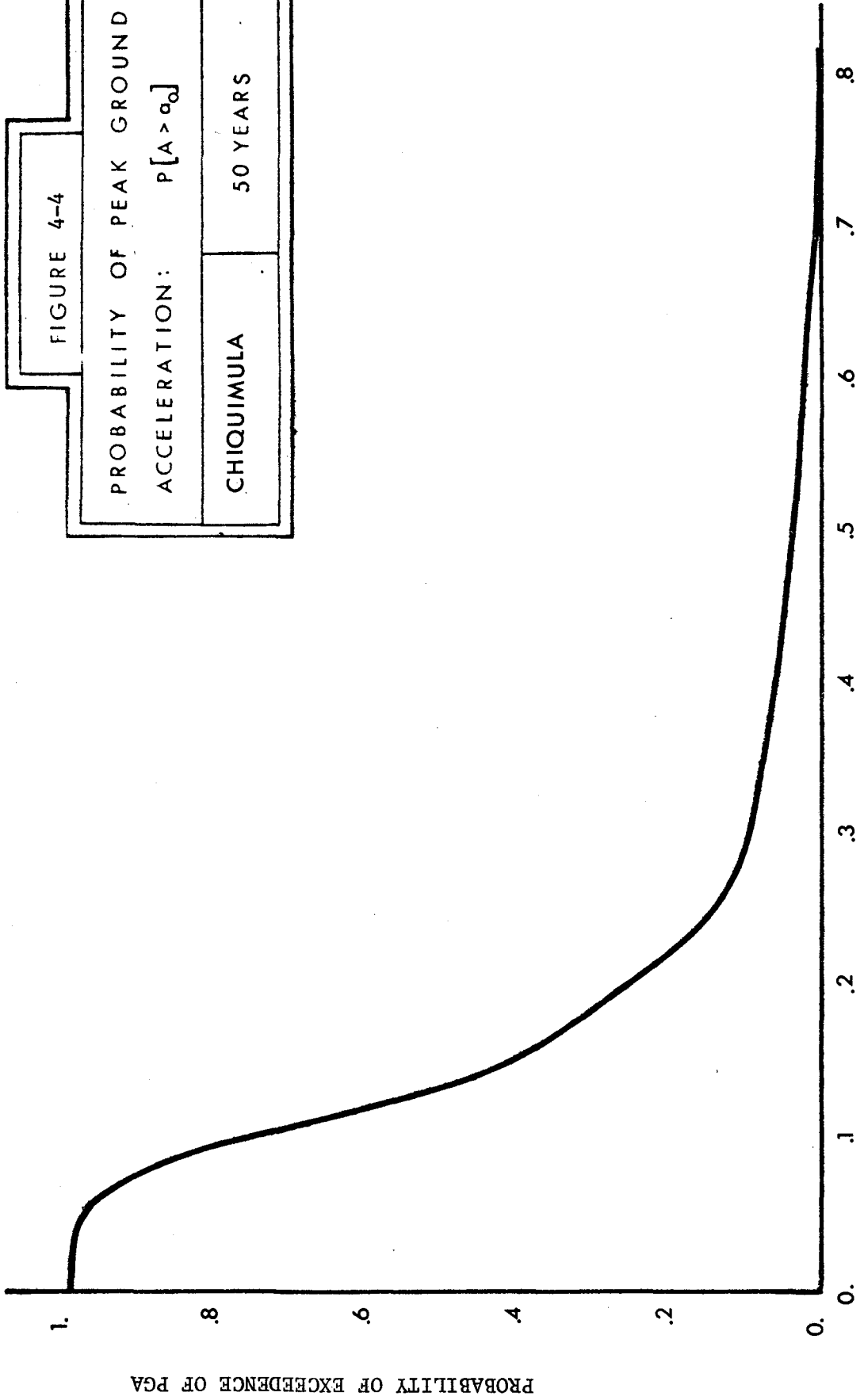


FIGURE 4-4

PROBABILITY OF PEAK GROUND  
ACCELERATION:  $P[A > a_d]$

CHIQUIMULA

50 YEARS

PROBABILITY OF EXCEEDENCE OF PGA

PEAK GROUND ACCELERATION IN G UNITS



FIGURE 4-5

|   |          |
|---|----------|
| PROBABILITY OF PEAK GROUND ACCELERATION: $P[A > a_0]$ |          |
| PUERTO BARRIOS  | 50 YEARS |

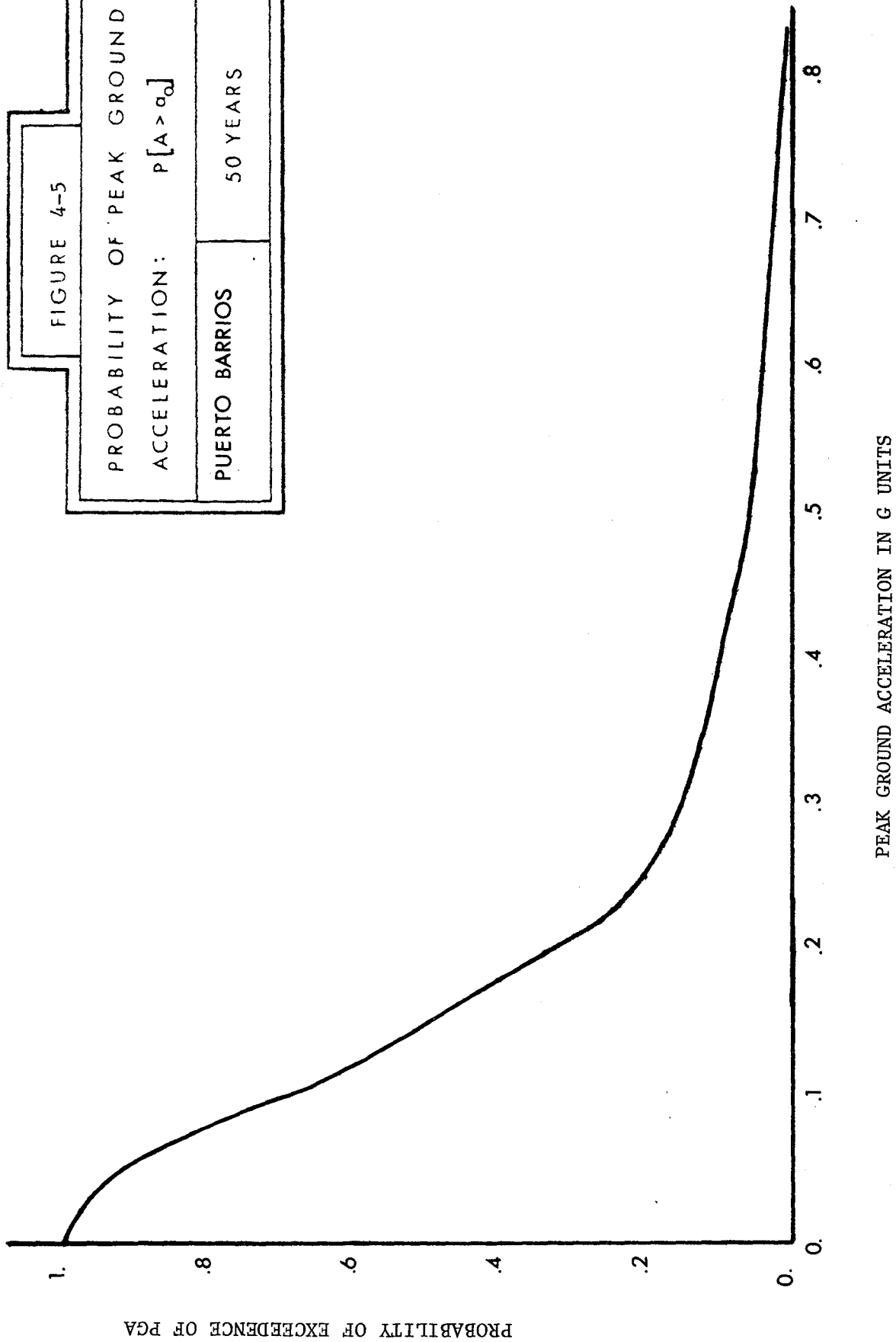
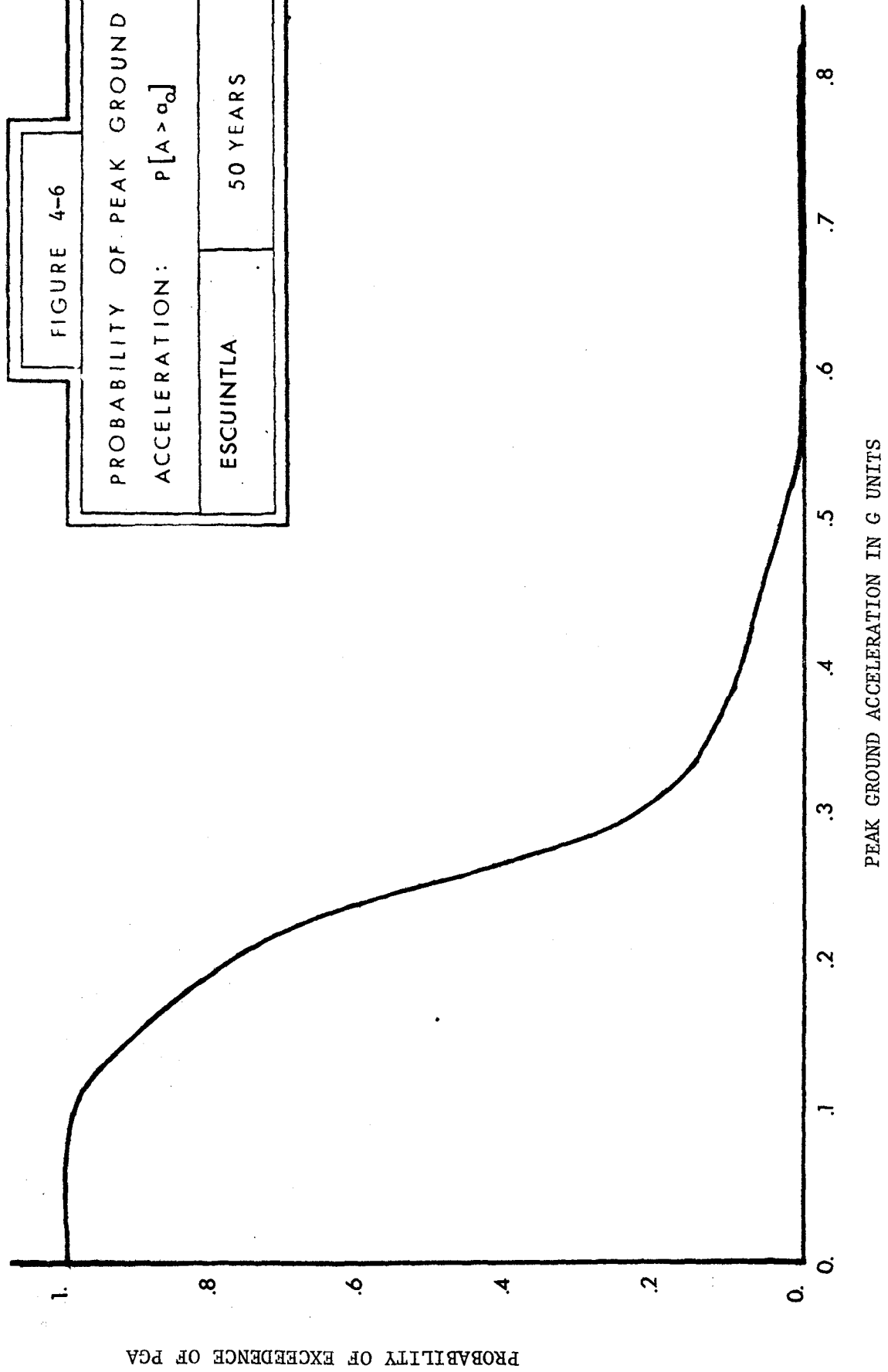


FIGURE 4-6

|   |          |
|---|----------|
| PROBABILITY OF PEAK GROUND ACCELERATION: $P[A > a_0]$ |          |
| ESCUINTLA   | 50 YEARS |



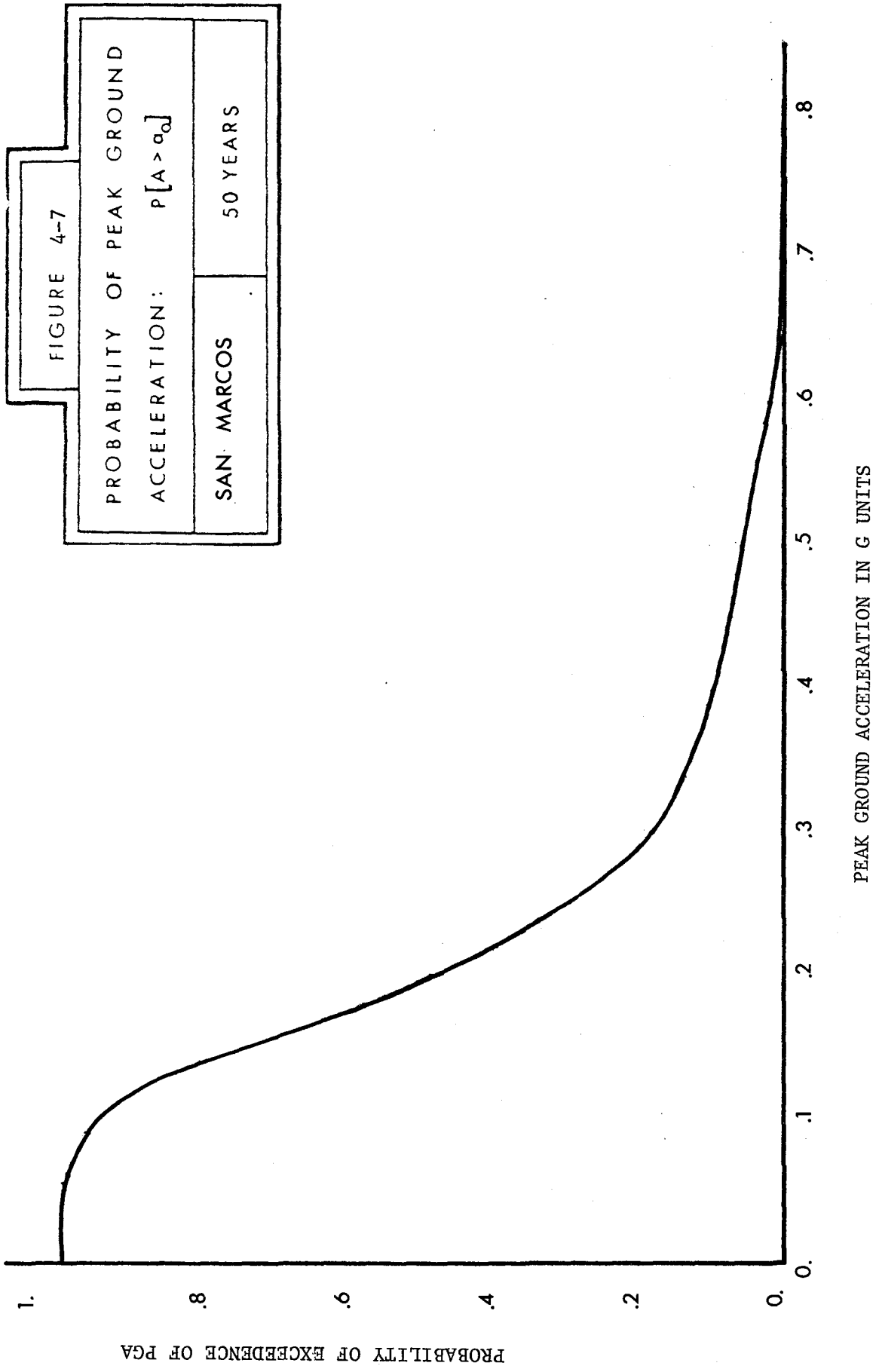


FIGURE 4-7

|   |          |
|---|----------|
| PROBABILITY OF PEAK GROUND ACCELERATION: $P[A > a_d]$ |          |
| SAN MARCOS  | 50 YEARS |

PROBABILITY OF EXCEEDENCE OF PGA

PEAK GROUND ACCELERATION IN G UNITS

PROBABILITY OF EXCEEDENCE OF PGA

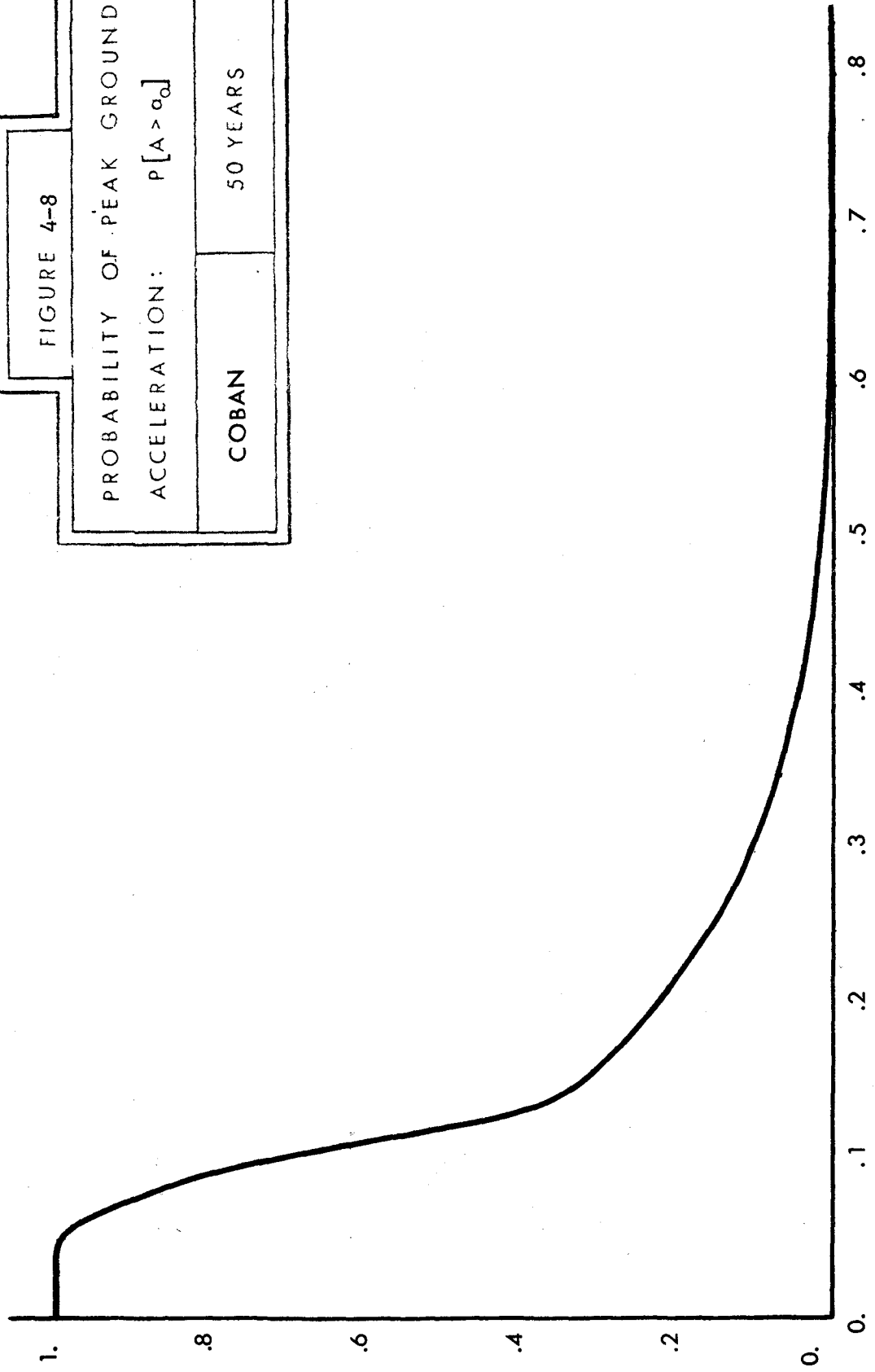


FIGURE 4-8

PROBABILITY OF PEAK GROUND  
ACCELERATION:  $P[A > a_d]$

COBAN

50 YEARS

PEAK GROUND ACCELERATION IN G UNITS

CHAPTER V

SEISMIC RISK ZONING

Scope

In this chapter, the concept of return period and economic life is discussed in detail. The iso-acceleration maps are used to develop acceleration zone graphs for seven cities in Guatemala.

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V-1 Concept of Return Period and Acceleration Zone Graphs (AZG)

In deriving the probabilistic loading at a given site as a function of time, it is assumed that the forecasting process is Poisson. This process implies that the events are independent in time and space. Using this assumption and an appropriate attenuation relationship, the iso-acceleration maps for the country were developed. As mentioned in Chapter IV, the cumulative distribution function of the peak ground acceleration  $A$ , were obtained for eight cities. Consider the cumulative distribution function of peak ground acceleration in Guatemala City for an exposure time of 50 years. (See Figure 4-1.)

Then 
$$P_{50} (A > 0.20g) = 0.76 \qquad 5-1$$

Equation 5-1 can be interpreted in the following way: "For Guatemala City, there is a 76% chance that during the next 50 years, the peak ground acceleration of 0.20g will be exceeded at least once."

Thus, there is a 24% chance that for Guatemala City, 0.20g peak ground acceleration will not be exceeded a single time.

Hence,

$$P (\text{Zero exceedence of } 0.20g \text{ in } 50 \text{ years}) = 0.24$$

From the Binomial Probability Law, it is known that for independent trials with probability of success  $p$  at each trial, the probability of  $r$  successes in  $n$  trials is given by

$$P_n(r) = \binom{n}{r} p^r (1-p)^{n-r} \quad 5-2$$

where

$$r = 0, 1, \dots, n; \quad n = r, r+1, r+2, \dots$$

and

$$\binom{n}{r} = \frac{n!}{r! (n-r)!}$$

Let each trial be a one-year duration for which the level of peak ground acceleration is under consideration. Define success as that event when the peak ground acceleration for a given trial (year) exceeds  $0.2g$ . Thus, the probability of zero exceedence of level  $0.2g$  in 50 years is the same as the probability of 0 successes in 50 trials. Hence, from Eq. 5-2:

$$P_{50}(0) = \binom{50}{0} p^0 (1-p)^{50}$$

$$P_{50}(0) = (1-p)^{50}$$

However,

$$P_{50}(0) = 0.24$$

$$(1-p)^{50} = 0.24$$

or

$$p = 0.028.$$

Thus, for Guatemala City, there is a 2.8% chance that in any given year, a peak ground acceleration of  $0.20g$  will be exceeded.

However, the return period is defined as

$$\text{Return Period} = \text{RP} = \frac{1}{p} \qquad 5-3$$

Thus, the return period RP in Guatemala City for a peak ground acceleration of 0.20g is  $\frac{1}{0.028} \approx 36$  years.

It should be pointed out that this return period of 36 years corresponding to 0.2g, obtained by using the cumulative distribution function (CDF) of PGA at Guatemala City for a 50 year exposure time does not change if we use the CDF with a different future exposure time period.

Thus, using the CDFs for all the cities in Guatemala considered in Chapter IV, a table can be developed for peak ground acceleration and return period. Table 5-1 is a general table giving this relationship for the cities considered. The following statements should be understood in using the concept of return period:

- (1) A return period is the mean (or average) waiting time for an event of interest. Thus, the average (waiting) time between 2 events producing 0.20g in Guatemala City is approximately 36 years.
- (2) The probability that an event corresponding to a return period RP will occur in any given year is given by  $p = \frac{1}{\text{RP}}$ . Thus, probability of exceeding 0.20g in Guatemala City in any given year is  $\frac{1}{36} \approx .028$ .
- (3) The probability that not a single event of the RP type will occur in RP years is given by  $\frac{1}{e}$  where  $e = 2.718$ , the Napierian base. Thus, probability that in 36 years, there will not be a single event producing a peak ground acceleration of 0.20g in Guatemala City is given by  $\frac{1}{e} \approx 0.36$ .

Hence, there is 64% chance that in RP years there will be at least one event of RP type. For Guatemala City, there is a 64% chance that in 36 years

TABLE 5-1

Return Period in Years

| PGA in g units | Guatemala City (1) | Quezaltenango (2) | Mazatenango (3) | Chiquimula (4) | Puerto Barrios (5) | Escuintla (6) | San Marcos (7) | Coban (8) |
|----------------|--------------------|-------------------|-----------------|----------------|--------------------|---------------|----------------|-----------|
| 0.05           | 1                  | 1                 | 1               | 10             | 20                 | 1             | 1              | 9         |
| 0.10           | 13                 | 8                 | 13              | 37             | 41                 | 11            | 17             | 37        |
| 0.15           | 23                 | 37                | 26              | 82             | 72                 | 21            | 38             | 124       |
| 0.20           | 36                 | 76                | 53              | 158            | 120                | 32            | 76             | 192       |
| 0.25           | 44                 | 153               | 103             | 302            | 225                | 63            | 141            | 287       |
| 0.30           | 52                 | 248               | 205             | 441            | 297                | 178           | 250            | 432       |
| 0.35           | 86                 | 337               | 325             | 571            | 385                | 381           | 361            | 696       |
| 0.40           | 393                | 448               | 540             | 731            | 494                | 571           | 468            | 1848      |
| 0.45           | 603                | 604               | 1254            | 933            | 645                | 807           | 610            |           |
| 0.50           | 1620               | 818               |                 | 1201           | 813                | 1302          | 820            |           |
| 0.55           | 2169               | 1183              |                 | 1583           | 990                |               | 1186           |           |
| 0.60           |                    | 2238              |                 | 2198           | 1213               |               | 2248           |           |
| 0.65           |                    |                   |                 | 3472           | 1504               |               |                |           |
| 0.70           |                    |                   |                 |                | 1921               |               |                |           |
| 0.75           |                    |                   |                 |                | 2607               |               |                |           |
| 0.80           |                    |                   |                 |                | 4177               |               |                |           |



0.20g peak ground acceleration will be exceeded. Consider again Table 5-1. For seismic zoning purposes, the following statements can be made:

The return period corresponding to a peak ground acceleration of 0.20g in Guatemala City is 36 years, in Escuintla is 32 years, in Quezaltenango and San Marcos is 76 years, in Mazatenango is 53 years, in Puerto Barrios is 120 years and in Chiquimula is 158 years. Thus, for each city, a graph relating the peak ground acceleration and return period can be plotted. Figures 5-1 through 5-9 show these graphs. They are referred to hereafter as Acceleration Zones Graphs. Figure 5-1 shows return period vs. peak ground acceleration for all the cities. It can be seen that for a given return period event (say, 100 years), Chiquimula has the lowest value of peak ground acceleration ( $\approx .16g$ ) and Guatemala City has the highest value of peak ground acceleration ( $.36g$ ). The values for other cities lies between these two limits. Qualitatively, it can be said that for a facility requiring a design loading corresponding to a 200 year return period, Chiquimula and Puerto Barrios have the lowest seismic zoning requirement; San Marcos, Quezaltenango, Mazatenango and Escuintla have similar zoning requirement; and finally, the highest level is for Guatemala City. This type of graph can help in macrozoning a country for a given class and use of a structure or facility.

#### V-2 Seismic Risk Zoning

In the previous section, we have seen the relationship between the peak ground acceleration and the corresponding return period for different cities of Guatemala. However, these relationships by themselves do not help in selecting a return period for a given acceptable level of risk. The next step, in any seismic zoning procedure, is to obtain a relationship between the economic (or exposure) life of a structure, the level of risk one is willing to take, and the return period consistent with the

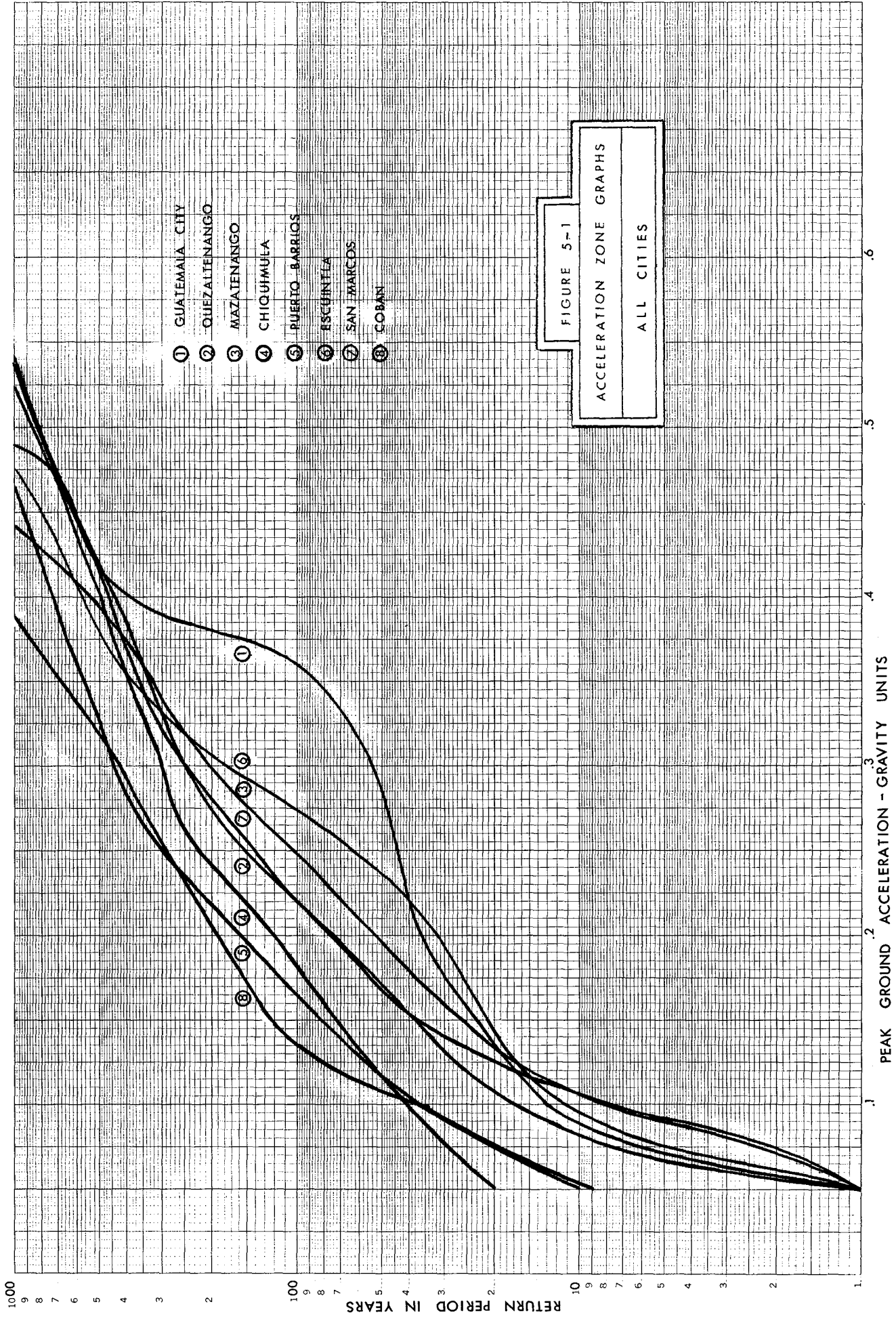


FIGURE 5-1  
ACCELERATION ZONE GRAPHS  
ALL CITIES

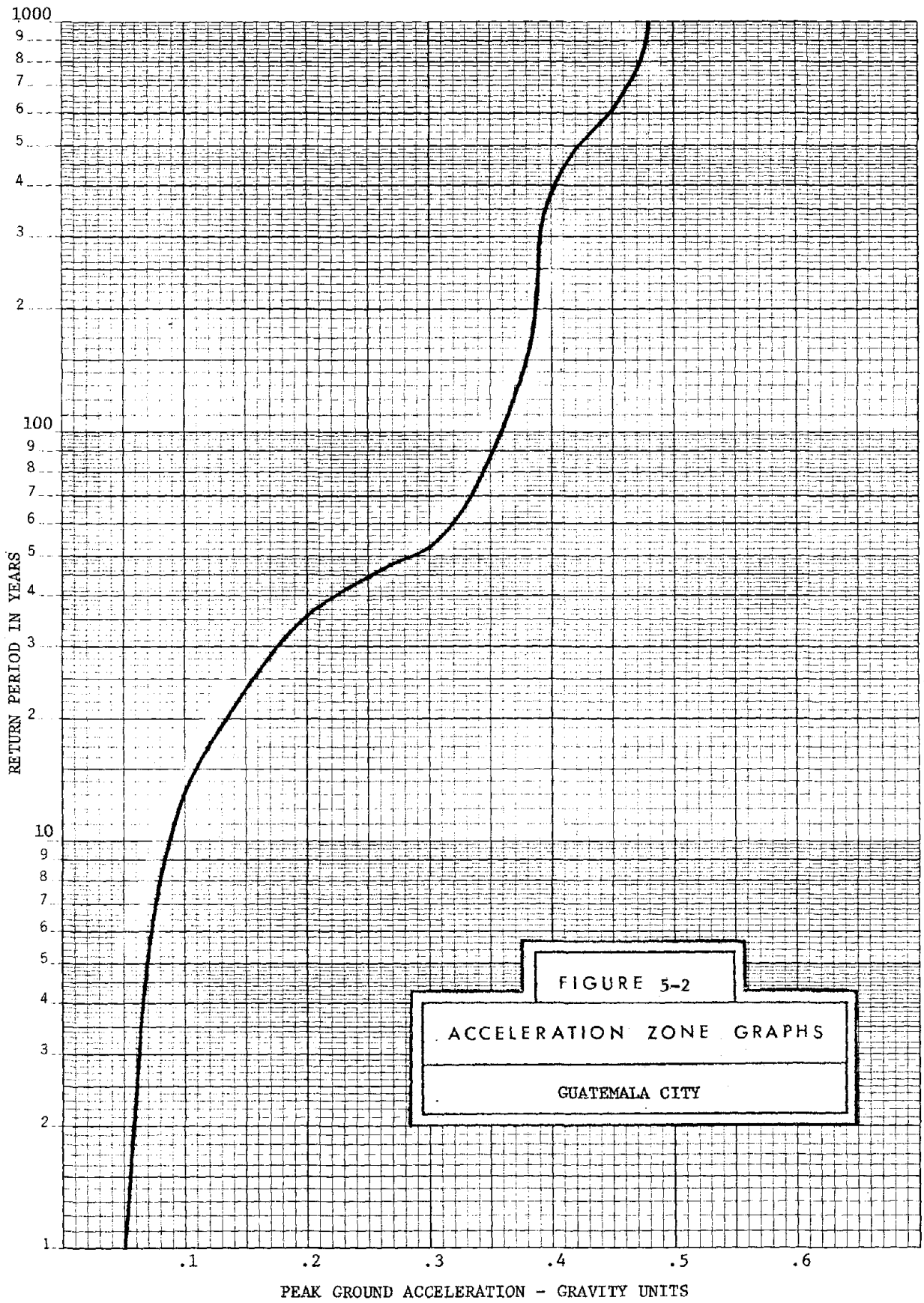


FIGURE 5-2  
ACCELERATION ZONE GRAPHS  
GUATEMALA CITY

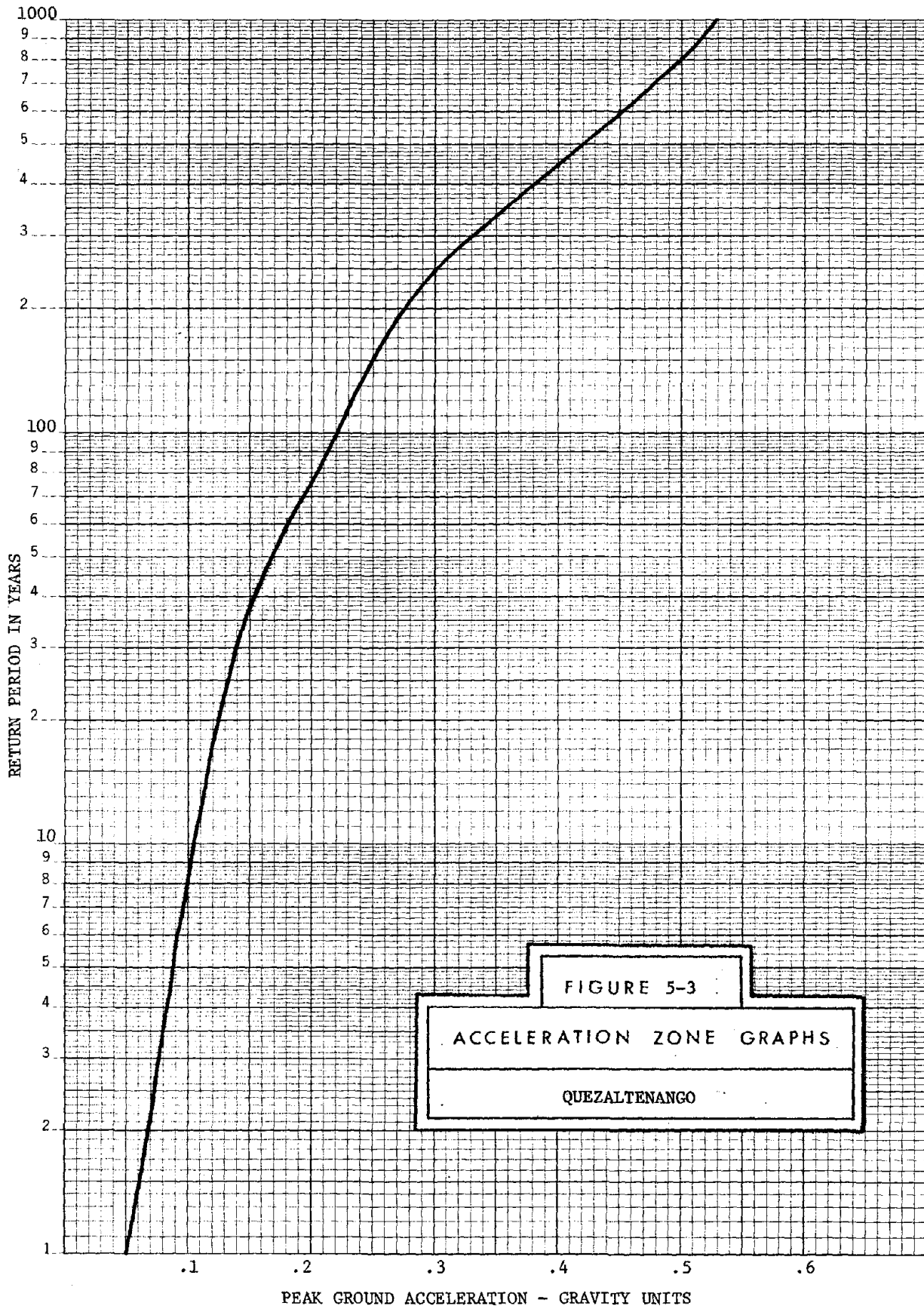


FIGURE 5-3  
ACCELERATION ZONE GRAPHS  
QUEZALTENANGO

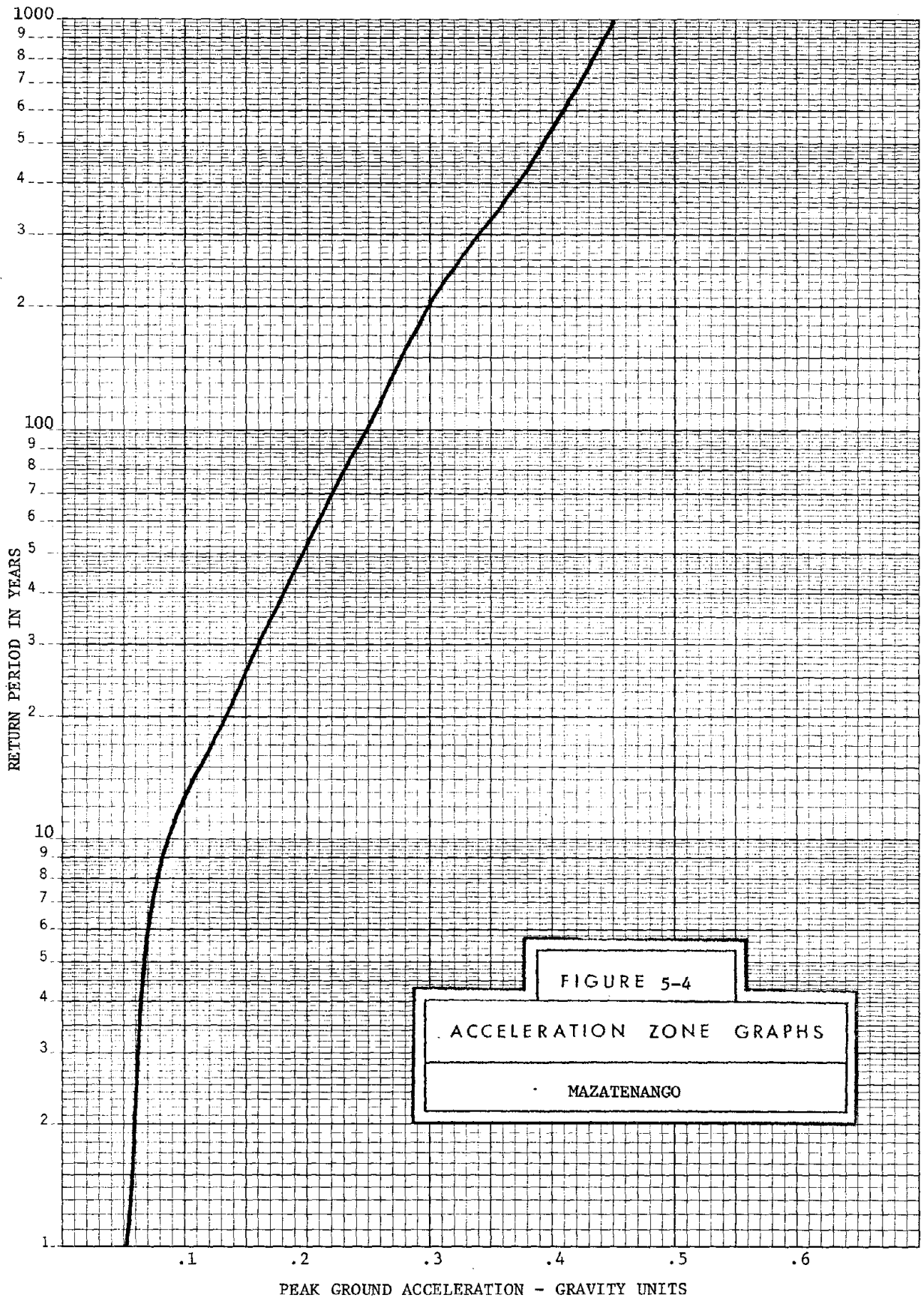


FIGURE 5-4  
ACCELERATION ZONE GRAPHS  
MAZATENANGO

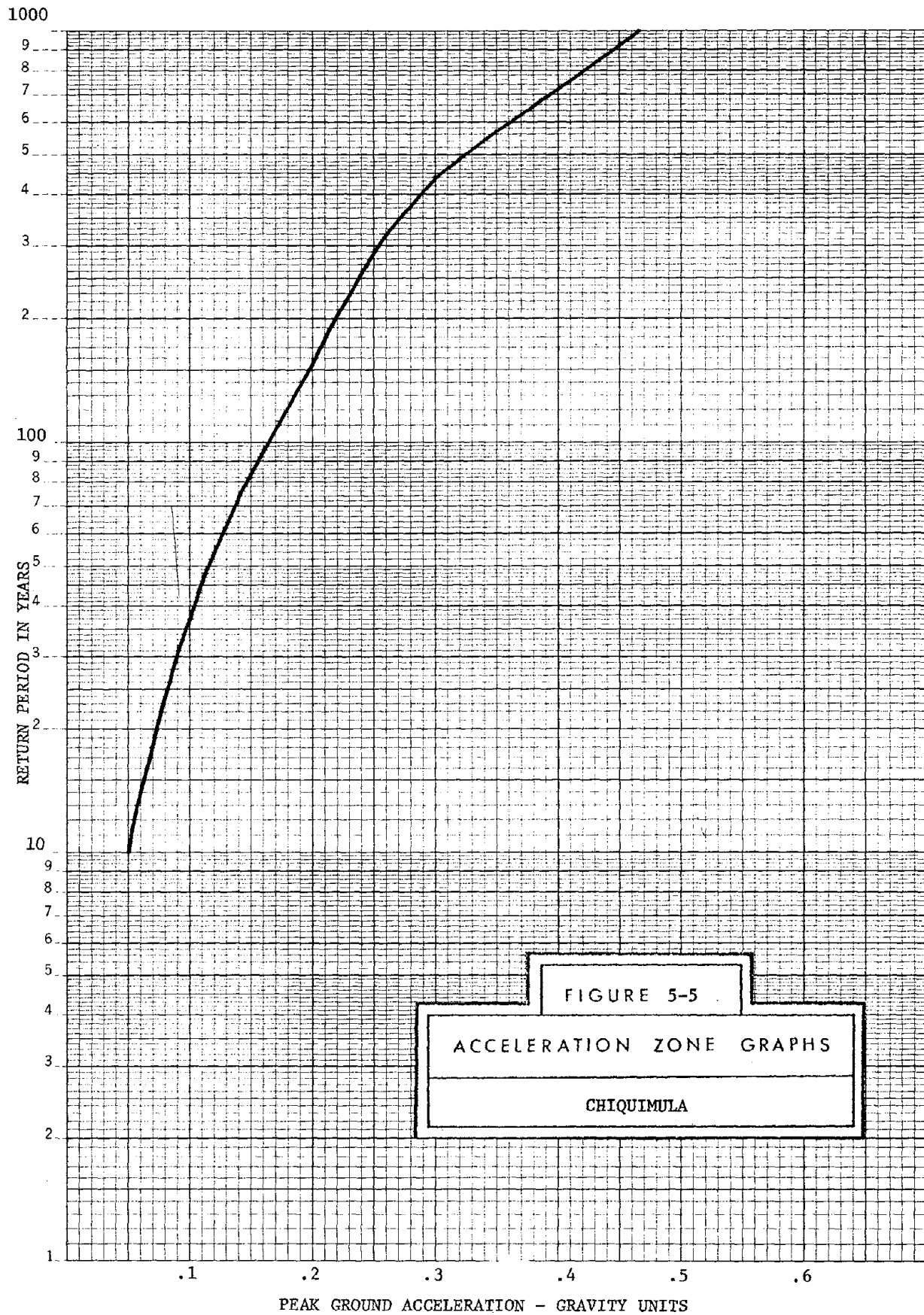


FIGURE 5-5  
 ACCELERATION ZONE GRAPHS  
 CHIQUIMULA

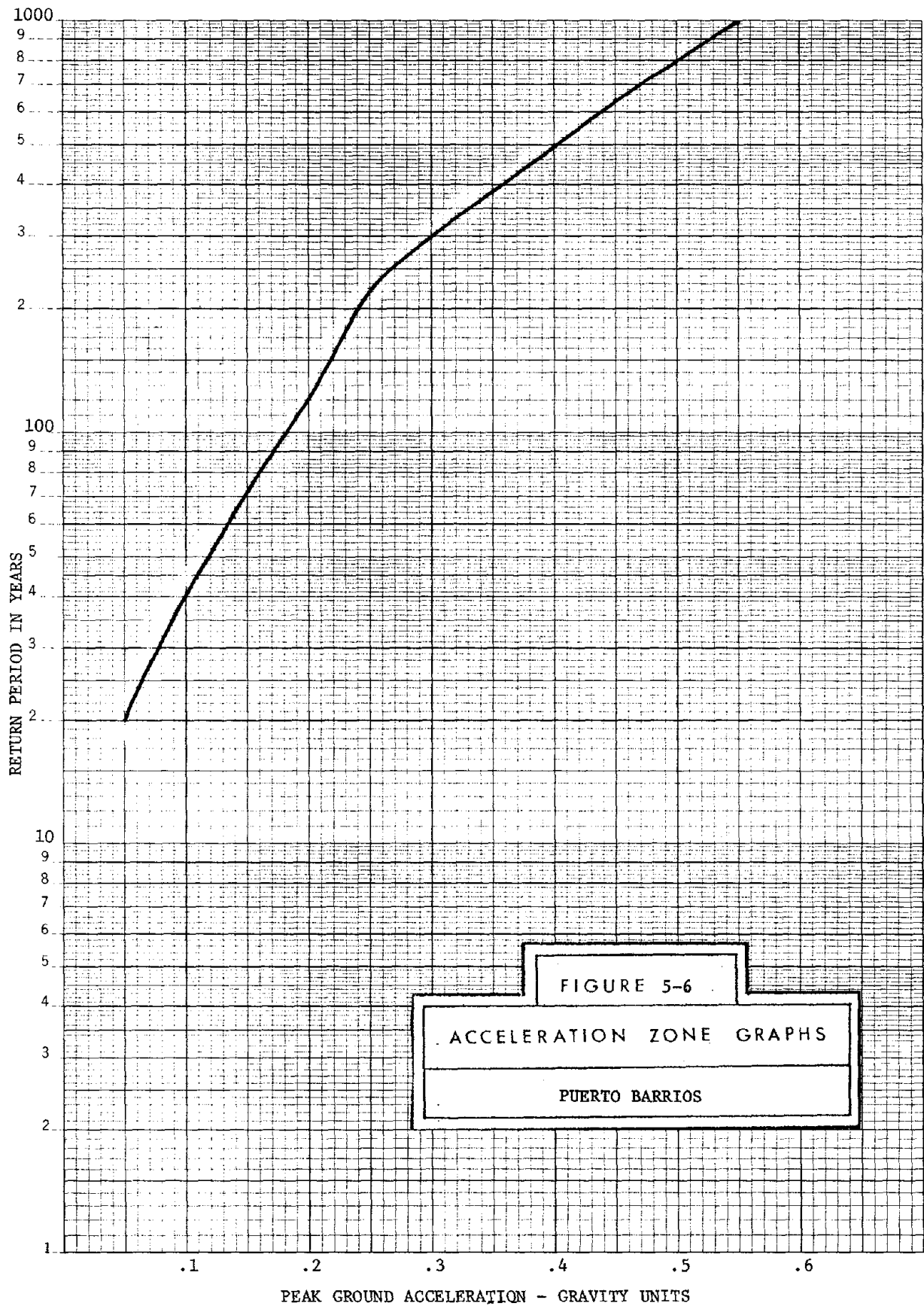


FIGURE 5-6  
ACCELERATION ZONE GRAPHS  
PUERTO BARRIOS

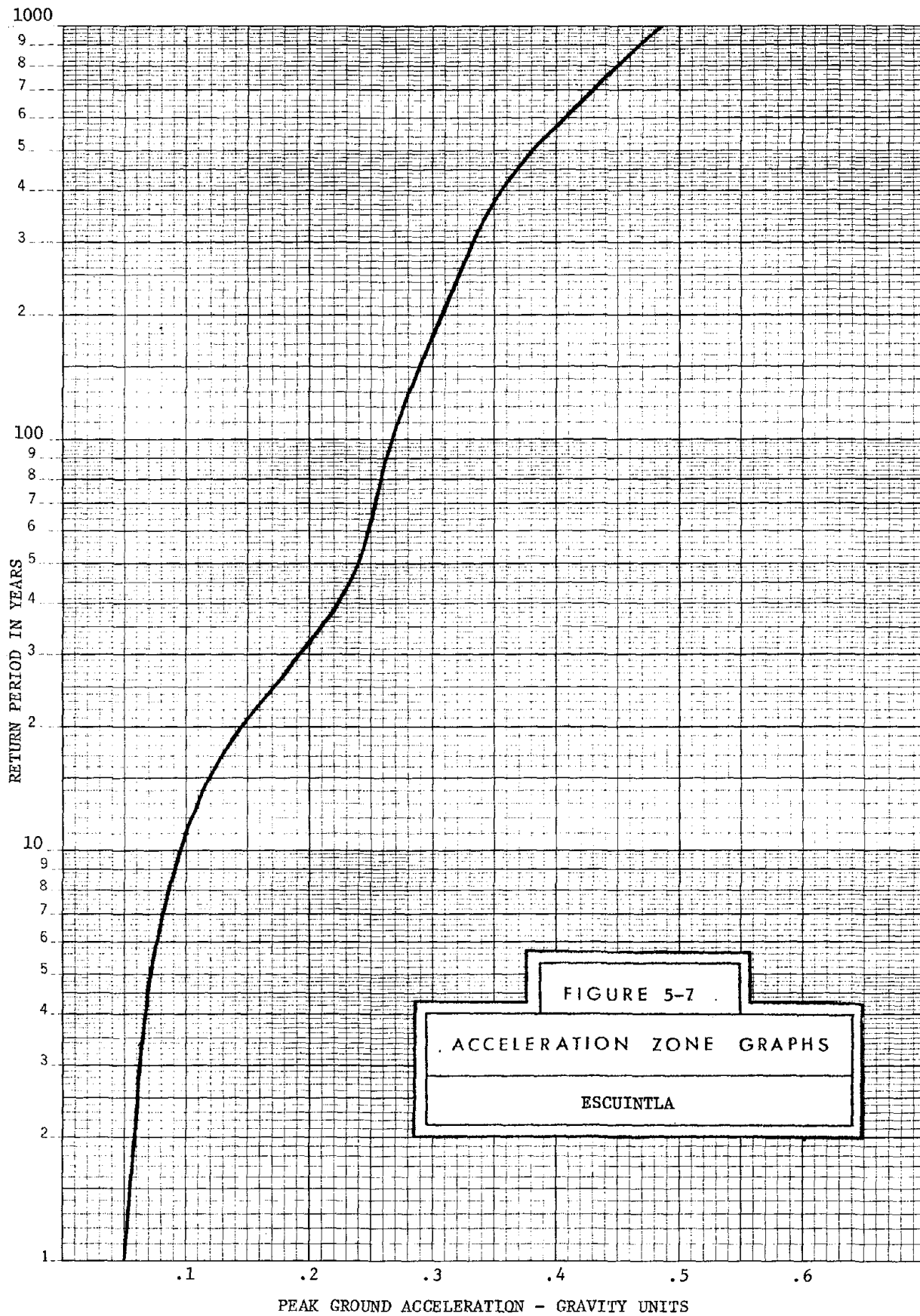


FIGURE 5-7  
ACCELERATION ZONE GRAPHS  
ESCUINTLA



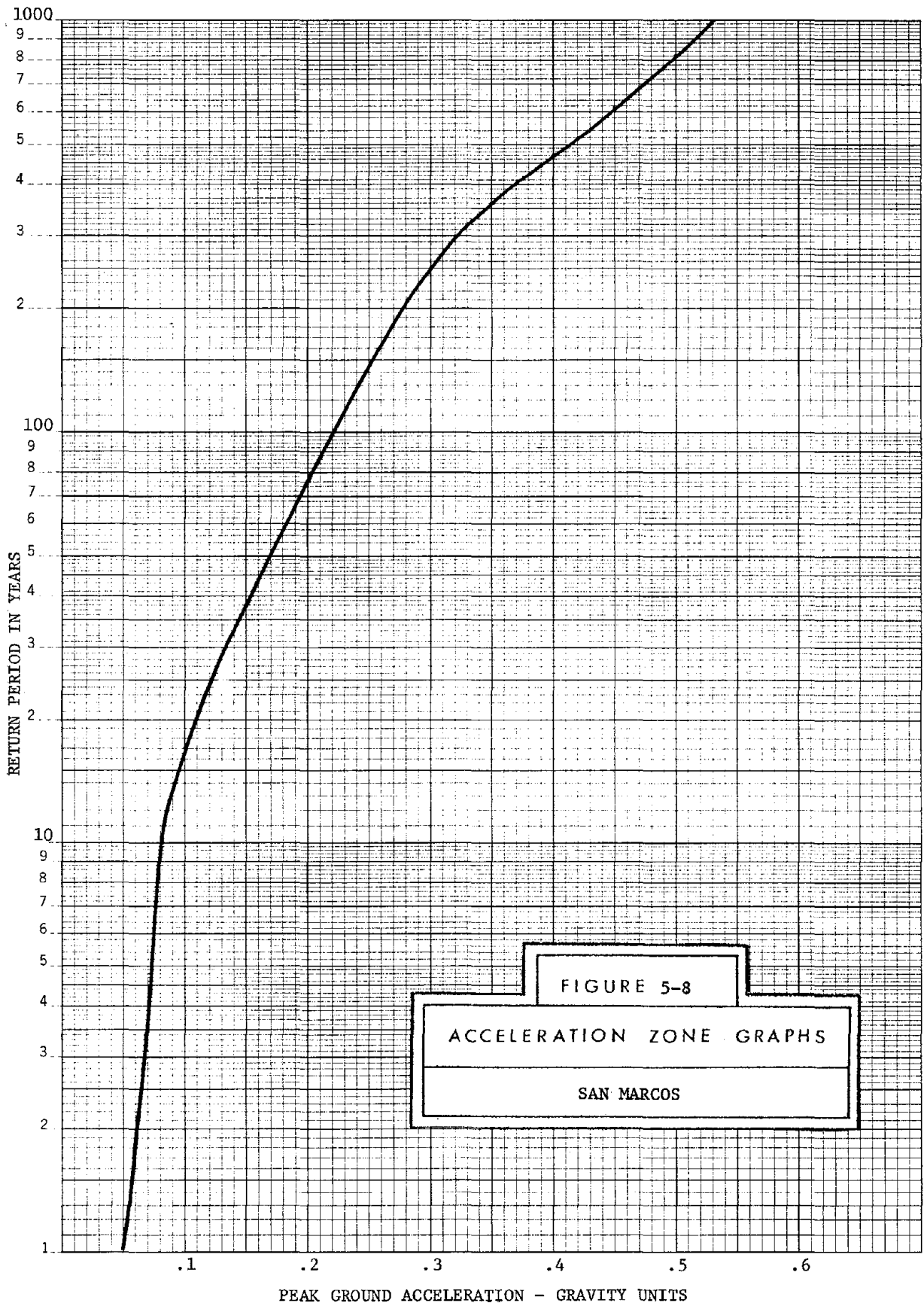


FIGURE 5-8  
 ACCELERATION ZONE GRAPHS  
 SAN MARCOS

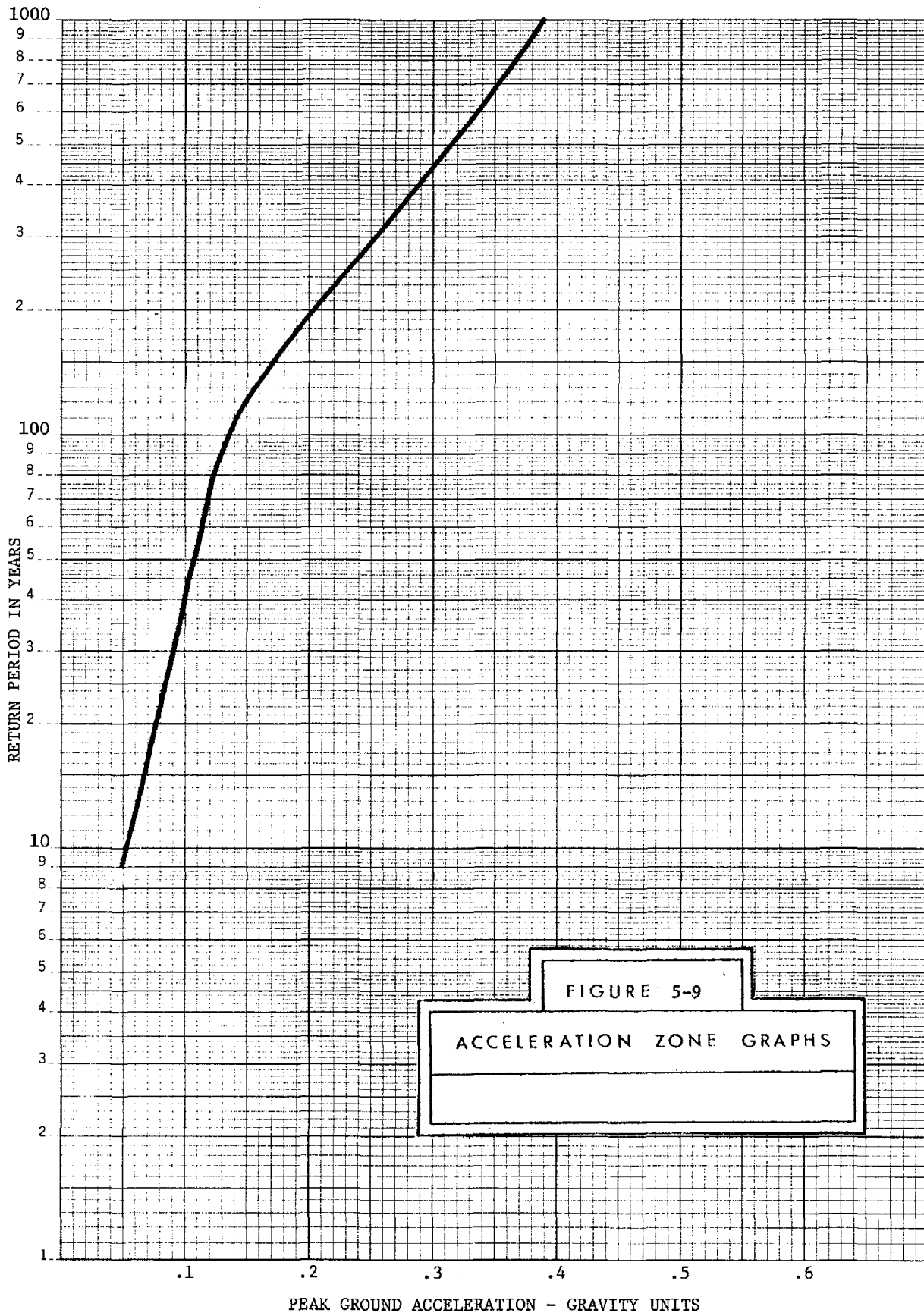


FIGURE 5-9  
ACCELERATION ZONE GRAPHS

risk and economic life. Consider again the Binomial distribution. The probability of  $r$  successes in  $n$  independent Bernoulli trials, with probability  $p$  of success at each trial, is given by

$$p_n(r) = \binom{n}{r} p^r (1-p)^{n-r} \quad \text{Eq. 5-2 repeated}$$

Thus

$$\begin{aligned} p_{10}(0) &= \binom{10}{0} (p)^0 (1-p)^{10} \\ &= (1-p)^{10} = \text{probability of zero successes} \\ &\quad \text{in ten trials (years)}. \end{aligned}$$

Let  $p(0) = (1-p)^{10}$  be equal to 0.90. Then the probability of no occurrence (or success) of a certain level of loading in ten years is given by 0.90.

or 
$$(1-p)^{10} = 0.90$$

Hence 
$$p = .01048$$

or return period  $RP = 95$  years.

Thus, for a structure whose economic life is ten years, if the acceptable risk level is 10% of exceeding the specified loading level, then the structure should be designed for a return period of 95 years. Table 5-2 gives the relationship between acceptable risk level, economic life and return period. If, for example, the acceptable risk level is 20% for a structure whose economic life is 50 years, then the loading level should correspond to a return period of 225 years. If this structure is in Guatemala City, the corresponding peak ground acceleration level is approximately 0.39g. If the same facility for the same risk level is to be built in Puerto Barrios, the corresponding peak ground acceleration level should be approximately 0.25g. Thus, for a given class and use of structure,

having the same economic life (50 years) and same acceptable risk (20%), the two consistent values of peak ground accelerations in Guatemala City and Puerto Barrios are 0.39g and 0.25g. This is the concept of consistent risk design from one seismic region to another region of different seismicity. Figure 5-10 shows the graph relating the risk level, economic life and the return period. This particular graph is independent of any region and gives return periods only as functions of risk and economic life. Such graphs can easily be codified. Once the acceptable risk level for a given economic life is selected for a given class and use of a structure, the corresponding return period is immediately obtained from Figure 5-10. Then, based on the graph of return period vs. peak ground acceleration (similar to Figures 5-1 to 5-9), the loading at a site can be determined. Let us describe this concept of risk, economic life, return period and Acceleration Zone Graphs (AZG).

As an example, consider a design of a hospital facility. Assume that the exposure time or economic life of the system is 50 years. The peak ground acceleration level for which this facility should be designed for each of three different cities is to be determined. The cities are Guatemala City, Quezaltenango and Puerto Barrios. Assume that for the hospital, which is a critical facility that must remain functional after a seismic event, the acceptable level of risk corresponding to damage is 20%. Thus, whether the planned facility is in Guatemala City, Quezaltenango, or Puerto Barrios, the engineers are willing to accept a 20% chance of damage during the 50 years economic life of the structure. Then, from Figure 5-10, the return period corresponding to the 50 year economic life and 20% risk is 225 years.

Now refer to the AZG corresponding to Guatemala City (see Figure 5-2). The peak ground acceleration for a 225 year return period in Guatemala City

TABLE 5-2

Return Period as a Function of Economic Life and  
Probability of Non-exceedence

| Probability of<br>exceeding<br>% | Economic Life<br>Years |     |     |     |     |     |
|----------------------------------|------------------------|-----|-----|-----|-----|-----|
|                                  | 10                     | 20  | 30  | 40  | 50  | 100 |
| 10                               | 95                     | 190 | 285 | 390 | 475 | 950 |
| 20                               | 45                     | 90  | 135 | 180 | 225 | 449 |
| 30                               | 29                     | 57  | 84  | 113 | 140 | 281 |
| 40                               | 20                     | 40  | 59  | 79  | 98  | 196 |
| 50                               | 15                     | 29  | 44  | 58  | 72  | 145 |
| 60                               | 11                     | 22  | 33  | 44  | 55  | 110 |
| 70                               | 9                      | 17  | 25  | 34  | 42  | 84  |
| 80                               | 7                      | 13  | 19  | 25  | 31  | 63  |
| 90                               | 5                      | 9   | 14  | 18  | 22  | 44  |
| 95                               | 4                      | 7   | 11  | 14  | 18  | 34  |
| 99                               | 3                      | 5   | 7   | 9   | 11  | 22  |
| 99.5                             | 2                      | 4   | 6   | 8   | 10  | 19  |

is 0.39g. Similarly, referring to the AZG for Quezaltenango and Puerto Barrios, the peak ground acceleration values corresponding to the 225 year return period are 0.29g and 0.25g, respectively. Thus, these three values of peak ground acceleration in the three different cities are consistent with the given acceptable risk.

As an alternate situation, consider two separate classes of structures to be built in Guatemala City. Let a school building with an economic life of 30 years have an acceptable risk level of 20%, and a warehouse with a ten year economic life have a 40% acceptable risk level. Referring to Figure 5-10, the return period for which the school should be designed is 135 years, and the return period for which the warehouse should be designed is 20 years. Again from the Guatemala City AZG (Figure 5-2), the corresponding peak ground acceleration values are 0.37g and 0.14g for the school and warehouse, respectively. If the same two facilities were to be located in Chiquimula, the corresponding peak ground acceleration values would be 0.19g and 0.07g. The major advantage of this method of zoning is that one can keep a consistent risk level from one region to another. Variations in the economic life and acceptable risk levels can be accounted for in arriving at a loading level through the return period transformation. Further application of the AZG to structural design will be presented in Part II of the total study.

It should also be pointed out that even though Puerto Barrios and Quezaltenango have smaller values of PGA corresponding to lower return periods (100 years and 200 years), the potential for major ground shaking -- above 0.5g -- does exist as can be seen from the AZG of these two cities. Thus, in terms of maximum credible ground shaking, Puerto Barrios and Quezaltenango are the two cities which should be kept in mind.

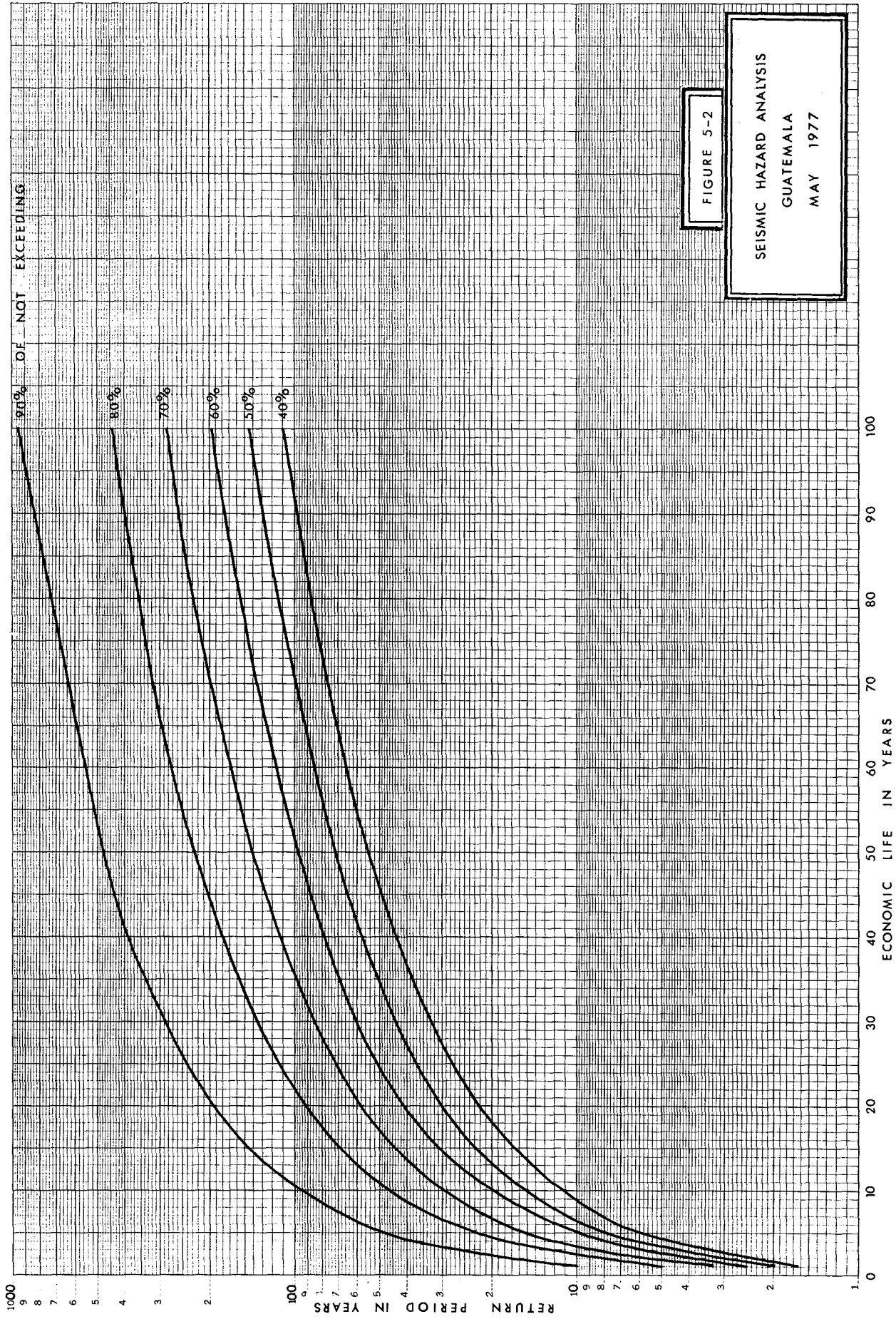


FIGURE 5-2  
SEISMIC HAZARD ANALYSIS  
GUATEMALA  
MAY 1977

Unlike the older seismic zone maps (such as the 1973 Uniform Building Code "risk" map) the recommended hazard map takes into account the frequency of seismic events, the level of "risk" one is willing to take in selecting a specific peak ground acceleration value and the future time horizons for which one wishes to consider the economic or structural life of the facility being designed.

Various questions come up regarding the reliability and long range stability of such hazard maps. Some of the questions are:

1. How reliable are the maps that are developed based on only historical data?
2. How stable are such maps? In other words, will these hazard maps change dramatically with each new future seismic event?
3. Is the formulation such that any new information available in the future can be incorporated to update the hazard maps?
4. What is the effect of local site conditions on the values obtained from these maps?

These and many such questions were discussed in Shah et al. (1975). However, in summary, the following responses can be given to the four questions posed above.

With respect to the reliability of results based on historical data, it is felt that for engineering and planning purposes and for seismic code formulation, the results presented are sufficiently reliable. The usual economic life of any engineered facility is usually less than 100 years to 200 years. In terms of geological time spans, this is a short period. Hence, we can assume that the geological processes during this short period are at a steady state. Hence, any information available from historical data can be extrapolated into similar time spans in the future. This discussion does not mean to imply that there are no errors introduced.



This possibility always exists. However, to wait for a complete geological information before developing a "seismic load" criteria for a country is unrealistic and impractical.

Concerning the stability of the hazard map, it is felt that the results presented here are quite stable. As long as the future seismic events can be assigned to any one of the sources considered in this work, the shape of the maps as well as the level of PGA's suggested should not change substantially. The only time the maps should be updated and changed is when a major seismic event occurs in a region where no previously known seismic source existed. In that case, the formulation and the computer programs are such that the suggested maps can be readily updated with the new information incorporated. Thus, in reply to the third question, such maps could be updated very easily. As a general recommendation, it is felt that such maps should be updated every five to seven years.

Effect of local site conditions (micro-characteristics) is usually felt in the amplitude of vibrations and in the frequency content of the vibration. The hazard map developed here is based on "average" soil condition. Thus, no site specific information is included in their development. In Part II of this study the effect of soft soil will be introduced. by changing the shape of the response spectrum to include higher period components. However, it should be pointed out that for important facilities such as dams, power plants, hospitals, etc., a site specific study should be conducted. Such information can then be used to modify the values suggested by the hazard map of this chapter.

In conclusion, it can be said that the seismic ground shaking hazard information developed in this study represents "a state-of-the-art" engineering solution. It is not the total information but it is one of the best that can be developed with the available knowledge and resources.

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96

## CHAPTER VI

### SUMMARY, CONCLUSION AND FURTHER RESEARCH

#### Scope

This chapter summarizes the work done in this phase of the study regarding the seismic hazard mapping of Guatemala. The conclusions drawn from this work are presented and a summary look at the phase 2 of this study is introduced.

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#### VI-1 Summary of Work on Seismic Hazard Mapping

In evaluating the future seismic hazard for Guatemala, the general geologic setting of the country, the seismological environment and all the available historical data was reviewed in this report. Two types of models were used in developing the seismic hazard maps (or iso-acceleration maps) of the country. The first model was the Poisson occurrence model where the seismic recurrence for various sources was based on past earthquake data and on a hypothesized upper Richter magnitude. Chapter IV presents these maps for four separate return periods. In Part II of this study, the relevance of obtaining iso-acceleration maps for the four return periods (50, 100, 500 and 1000 years) will be discussed. The Bayesian model, where subjective and objective information on the number of occurrences and the magnitude level is used, is presented in the Appendix B. For this model the attenuation relationship is assumed to be probabilistic with Uniform distribution and the scatter of this distribution is represented by the coefficient of variation. Iso-acceleration maps developed using this Bayesian model are also given in Appendix B. For further details on this

model and its application see Mortgat (1976) and Mortgat et al. (1977).

Using the Bayesian model, iso-duration maps are also presented in the Appendix B. At this time, explicit use of duration in seismic design decisions is not available. However, it is felt that this information is an important engineering input and will be of great interest to planners and engineers.

Based on the information from iso-acceleration maps, the acceleration zone graphs (AZG) for the following eight major urban areas was developed.

1. Guatemala City
2. Quezaltenango
3. Mazatenango
4. Chiquimula
5. Puerto Barrios
6. Esquintla
7. San Marcos
8. Copan

These acceleration zone graphs can be used to evaluate the level of peak ground acceleration as a function of return period. In a future report on phase 2 of this study, the suggested levels of peak ground accelerations for various regions of the country and for various types and uses of structures will be presented.

## VI-2 Conclusions from the Present Study

It is very interesting to note that the shape of the iso-acceleration maps obtained by using the Bayesian model (Appendix B) and the Poisson occurrence model are very similar. This is due to the following reasons:

- 1) The data base used for both models is the same.
- 2) The subjective data of the Bayesian model was developed from the regression line of the recurrence curve obtained in the Poisson occurrence model.

The value of the peak ground acceleration obtained using the Bayesian model will increase with the increase in attenuation uncertainty. The coefficient of variation used in the Appendix B maps was 0.3. Looking at this similarity and based on experience obtained in developing Nicaragua iso-acceleration maps (Shah et al., 1975) and Costa Rica iso-acceleration maps (Mortgat et al., 1977) it is reasonable to accept the iso-acceleration maps for various return periods presented in Chapter IV.

The iso-duration maps given in Appendix B can be used to evaluate damage potential and energy content of future hypothesized or forecasted seismic events.

It can be concluded from the results presented in this report that for engineering planning, sufficient data are available (together with geological information and subjective knowledge) to provide seismic "zoning" information for Guatemala. The methods presented in this report can use either objective or subjective seismological data. The method is simple and has the capabilities of being updated at regular intervals. It can be seen from the iso-acceleration maps of Chapter IV that the seismicity in the northern part of Guatemala is low. The high seismic region of the south-southwestern part of Guatemala reaches a peak near the intersection of the Motagua fault, the Benioff Zone and the Mixco fault. The probable seismic "loading" in Guatemala City, based on the available data and the resultant iso-acceleration maps, is high. In comparison to the Managua region of Nicaragua, the Guatemala City region is very similar. The seismic lateral load requirement for these various levels of peak ground accelerations will be developed in Part II of this study.

In Chapter V, the reliability and stability of the iso-acceleration maps was discussed. It should be noted once more here that the methodology available now can take into account not only the short historical data,

but also any geological or seismological information that may be available. It is felt by the authors of this report that such maps should be updated every five years so that any new information that will be available can be incorporated in the updated version.

In conclusion, it can be said that the seismic ground shaking hazard information in the form of iso-acceleration and iso-duration maps developed in this study represents the "state-of-the-art" methodology and engineering solution. It is consistent with the available information. Thus, its shortcomings are also consistent with the shortcomings of the available data. It is always possible that the results obtained by using the past knowledge can always be proven to be incorrect by the nature.

### VI-3 Further Research

In order to implement and use the seismic hazard information presented in this report, the following topics and tasks will be accomplished in part II of this study:

- A thorough study of the 1973 UBC, the 1976 UBC and the ATC-3 work currently being completed. The above codes will be evaluated with the seismicity of Guatemala in mind. Such an evaluation will permit the engineers and planners in Guatemala to appreciate and understand the current code levels and their relationships with future seismic loading demands.
- A detailed discussion on the purpose and effects of earthquake codes will be presented. This will include the history of earthquake loading criteria, the relation of design loads and quality of structures, the objectives and qualities of workable seismic codes and the role of design detailing and design forces in providing a safe economic construction.

- Introduction to the Proposed Seismic Design Provisions.  
This will take into account the current codes, their advantages, their shortcomings and the available solutions to eliminate or reduce the shortcomings. This part of the study will explain in detail the concept of acceptable risk and the associated loading levels from the hazard maps of Guatemala. The study will also develop the shape and levels of various design spectra. These spectra will include the site characteristics information.
- A detailed description of the type of structural systems, their effects on the design level and a step-by-step design procedure will be developed.
- Based on the spectral approach of seismic design, a simplified equivalent static load method, similar to the 1976 UBC method of design, will be developed. It should be emphasized here that a workable code should have the following four ingredients.
  - 1) Simplicity.
  - 2) Rationality.
  - 3) Freedom to use responsible ingenuity for special structures.
  - 4) Reward and encouragement for using dynamic analysis when merited by the complexity of a given structure.

In developing the methodology in Part II of this study, the above four ingredients will be kept in mind.

- Finally, a detailed comparison with the proposed methodology and the 1976 UBC will be made. This will be done with the seismic environment of Guatemala (as obtained in this report) in mind.

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106

APPENDIX A

POISSON MODEL

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108

## APPENDIX A

### POISSON MODEL

#### A-1 Poisson Model of Seismic Occurrences

As mentioned in Chapter II, earthquake occurrences can be modeled using the Poisson probability law. For earthquake events to follow the Poisson Model, the following assumptions must be valid:

- (1) Earthquakes are spatially independent;
- (2) Earthquakes are temporally independent;
- (3) Probability that two seismic events will take place at the same place and at the same instant of time approaches zero.

These assumptions are necessary for the formulation of the Poisson Model. The first assumption implies that occurrence or nonoccurrence of a seismic event at one site does not affect the occurrence or nonoccurrence of another seismic event at some other site. The second assumption implies that the seismic events do not have memory in time. A Markovian assumption of one-step memory in time may be a better assumption, but as mentioned previously, this assumption for large events does not introduce major errors (Gardner and Knopoff, 1974). The third assumption implies that for a small time interval,  $\Delta t$ , no more than one seismic event can occur. This assumption is considered to be realistic and to fit the physical phenomenon.

For events which satisfy the above assumptions, the Poisson law can be written as

$$P_n(t) = \frac{e^{-\lambda t} (\lambda t)^n}{n!} \quad \text{A-1}$$

where  $P_n(t)$  = Probability of having  $n$  events in time period  $t$ .

$n$  = Number of events.

$\lambda$  = Mean rate of occurrence per unit of time.

In Chapter III, it was shown how the mean number of occurrences above magnitude  $M$  for a given source can be obtained using recurrence relationships. This relationship in its general form can be stated as

$$N(M) = \phi(M, A, T) \quad A-2$$

where  $N(M)$  = Number of occurrences above Richter magnitude  $M$ .

$M$  = Richter magnitude.

$A$  = Source characteristic (area for area source, length for line source).

$T$  = Time period of data base.

As mentioned in Chapter III, a log-linear recurrence relationship is assumed for all sources. Also, for some sources, the relationship is bi-linear (two lines described by  $\alpha_1, \beta_1$ , and  $\alpha_2, \beta_2$ ). (See Table 3-3 of Chapter III.) Thus, for a given source, the two lines describing the recurrence relationship are given by:

$$\begin{aligned} \ln N'(M) &= \alpha_1' + \beta_1 M & 0 \leq M \leq M_1 \\ \ln N'(M) &= \alpha_2' + \beta_2 M & M_1 \leq M \leq M_2 \end{aligned} \quad A-3$$

where  $M_1$  is the magnitude at which the two recurrence lines intersect (see, for example, Fig. 3-4).

$M_2$  is the upper cutoff magnitude for a given source (see Table 3-3, Chapter III).

Thus, depending on the source and the value of  $M$ , the mean number of events above magnitude  $M$  for a unit area for an area source, a unit length for a



line source, and a unit time is given by:

$$N'(M) = \exp [\alpha_1' + \beta_1 M] \quad \text{A-4}$$

From equation A-1 it follows that

$$P_n(t) = \frac{\exp [- \exp (\alpha_1' + \beta_1 M)t] [\exp (\alpha_1' + \beta_1 M)t]^n}{n!} \quad \text{A-5}$$

Note that in equation A-5 above,  $\lambda$  is replaced by  $N'(M)$ . Equation A-5 gives the probability of observing  $n$  events above magnitude  $M$  in time period  $t$ , based on the seismic history of a given source.

## A-2 Source Mechanisms

Three different types of sources can be used to represent the seismicity of any region. They are point, line, and area sources. All three source mechanisms will be discussed for generality and completeness of the development, although only line sources were considered for the Guatemala region.

### a. Point Source

For this type of source, all occurrences (past and future) take place at one point. The recurrence relationship can be normalized with respect to time  $T$  as follows:

$$N'(m) = \frac{N(m)}{T} \quad \text{A-6}$$

Substituting the value of  $N'(m)$  in the Poisson law, Equation A-1 becomes:

$$P_n[M > m, t] = \frac{\exp [- N'(m)t] [N'(m)t]^n}{n!} \quad \text{A-7}$$

where the notation

$P_n[M > m, t]$  gives the probability that there will be  $n$  events

of Richter magnitude greater than  $m$  in time period  $t$ .

For engineering purposes, it is of primary interest to determine the probability of at least one event greater than  $m$  in time period  $t$ . The probability is given by

$$\begin{aligned} P(\text{at least one event of magnitude } M > m \text{ in time } t) \\ = 1 - P(\text{no earthquake of magnitude } M < m \text{ in time } t) \end{aligned}$$

Hence, from Equation A-7

$$\begin{aligned} P(\text{at least one event of magnitude } M > m \text{ in time } t) \\ = 1 - \exp [- N'(m)t] \end{aligned}$$

b. Line Source

In the case of a line source, it is assumed that epicenters lie along a linear fault pattern. For a line source of length  $L$  (fault length  $L$ ) and a data base for a time period  $T$ , the recurrence relationship can be normalized to:

$$N'(m) = \frac{N(m)}{LT} \quad (\text{Equation 3-3b repeated})$$

This  $N'(m)$  can be substituted directly in Equation A-8 to obtain the probability of at least one event of magnitude  $M > m$  in time  $t$ .

c. Area Source

When the past earthquake epicenters do not lie on a line (i.e., along a given fault line) or when there is no information on fault locations, but earthquake events are scattered over a region, the seismic source should be con-

considered as an area source. The area source can be approximated by a full circle or any section of a circle in which the epicenters are scattered. In this case, the recurrence relationship is normalized with respect to the area A and the time of the data base T.

$$N'(m) = \frac{N(m)}{AT} \quad \text{(Equation 3-3a repeated)}$$

Again, the probability of at least one event due to this area source above magnitude m in time period t is given by

$$P(\text{at least one } M > m \text{ in time } t) = 1 - \exp [- N'(m)t]$$

It should be noted that while Equation A-8 has the same form for a point source, line source and area source, in each case the normalized  $N'(m)$  has a different interpretation.

### A-3 Peak Ground Acceleration at a Site

For design purposes, it is necessary to know the loading at a site. As mentioned in Chapter III, the most commonly used parameter to describe the seismic loading at a given site is the peak ground acceleration. In order to derive the probability distribution on peak ground acceleration at a site for a future time period t, the following information is required:

1. Probabilistic formulation on Richter magnitude for a source as a function of future time t.
2. Distance from the source to the site.
3. Attenuation of peak ground acceleration from source to site.

The first parameter has already been determined in the previous section. Various attenuation formulae are available which give the relationship

between the Richter magnitude  $M$ , the epicentral distance or the hypocentral distance, and the peak ground acceleration. Donovan (1974) has summarized and compared ten attenuation relationships. These relations are repeated in Table A-1 and their graphs are shown in Figure A-1. Most of the relationships in Table A-1, can be written in the following general form:

$$a = \frac{b_1 \exp(b_2 m)}{(R_h + b_4)^{b_3}} \quad \text{A-9}$$

where  $a$  = Peak Ground Acceleration (PGA) in  $\text{cm/sec}^2$  or  $g$

$R_h$  = Hypocentral distance from source to site in km or miles

$m$  = Richter magnitude

$b_1$ ,  $b_2$ ,  $b_3$  and  $b_4$  are constants depending on the region

Using the general form of the attenuation relationship given by Equation A-9, the probability distribution on peak ground acceleration for a site due to the three types of seismic sources can be obtained.

#### Point Source

The probability of at least one event greater than  $m$  in time  $t$  from a point source such as the one shown on Figure A-2 is:

$$P[M > m, t] = 1 - \exp[-N'(m)t] \quad \text{A-10}$$

Substituting the recurrence relationships given by Equation 3-2, the probability distribution is derived as follows:

$$P[M > m, t] = 1 - \exp[-\exp(\alpha' + \beta m)t] \quad \text{A-11}$$

where  $\alpha'$  = normalized regression constant.

TABLE A-1  
Attenuation Equations

| <u>Data Source</u>                             | <u>Equation</u>   | <u>Reference</u>            |
|--|---|-----------------------------|
| 1. San Fernando Earthquake<br>February 9, 1971 | $y = 186206 R^{-1.83}$  | --                          |
| 2. California Earthquakes                      | $y = \frac{981 y_0}{1 + \frac{R'}{h}^2}$ <p style="text-align: center;">where <math>\log y_0 = -(\bar{b}+3) + 0.81m - 0.027m^2</math></p> <p style="text-align: center;"><math>\bar{b}</math> is a site factor</p>  | Blume (1965)                |
| 3. California Earthquakes                      | Graphical Presentation  | Housner (1965)              |
| 4. California & Japanese                       | $y = \frac{5}{T_G} 10^{0.61m - P \log R + Q}$ <p style="text-align: center;">where <math>P = 1.66 + \frac{3.60}{R}</math></p> <p style="text-align: center;"><math>Q = 0.167 - \frac{1.83}{R}</math></p> <p style="text-align: center;"><math>T_G</math> = fundamental period of site</p> | Kanai (1966)                |
| 5. Cloud (1963)                                | $y = \frac{6.77 e^{1.64m}}{1.1e^{1.1m} + R^2}$  | Milne & Davenport<br>(1969) |
| 6. Cloud (1963)<br>Housner (1962)              | $y = 1230 e^{0.8m} (R+25)^{-2}$   | Esteve (1970)               |
| 7. U.S.C. & G.S.                               | $\log_{10} y = 6.5 - 2 \log_{10} (R'+80)$   | Cloud & Perez<br>(1971)     |
| 8. 11 Selected Records                         | Graphical Presentation  | Schnabel & Seed<br>(1972)   |
| 9. 303 Instrumental Values                     | $y = 1300 e^{0.67m} (R+25)^{-1.6}$  | --                          |
| 10. Western U. S. Records                      | $y = 18.9 e^{0.8m} (R^2+400)^{-1}$  | --                          |

$y$  is  $\text{cm/sec}^2$   
 $R$  is kilometers (distance to causative fault)  
 $R'$  is miles (epicentral distance)  
 $h$  is miles (focal depth)  
 $m$  is magnitude

\* Taken from Donovan 1974

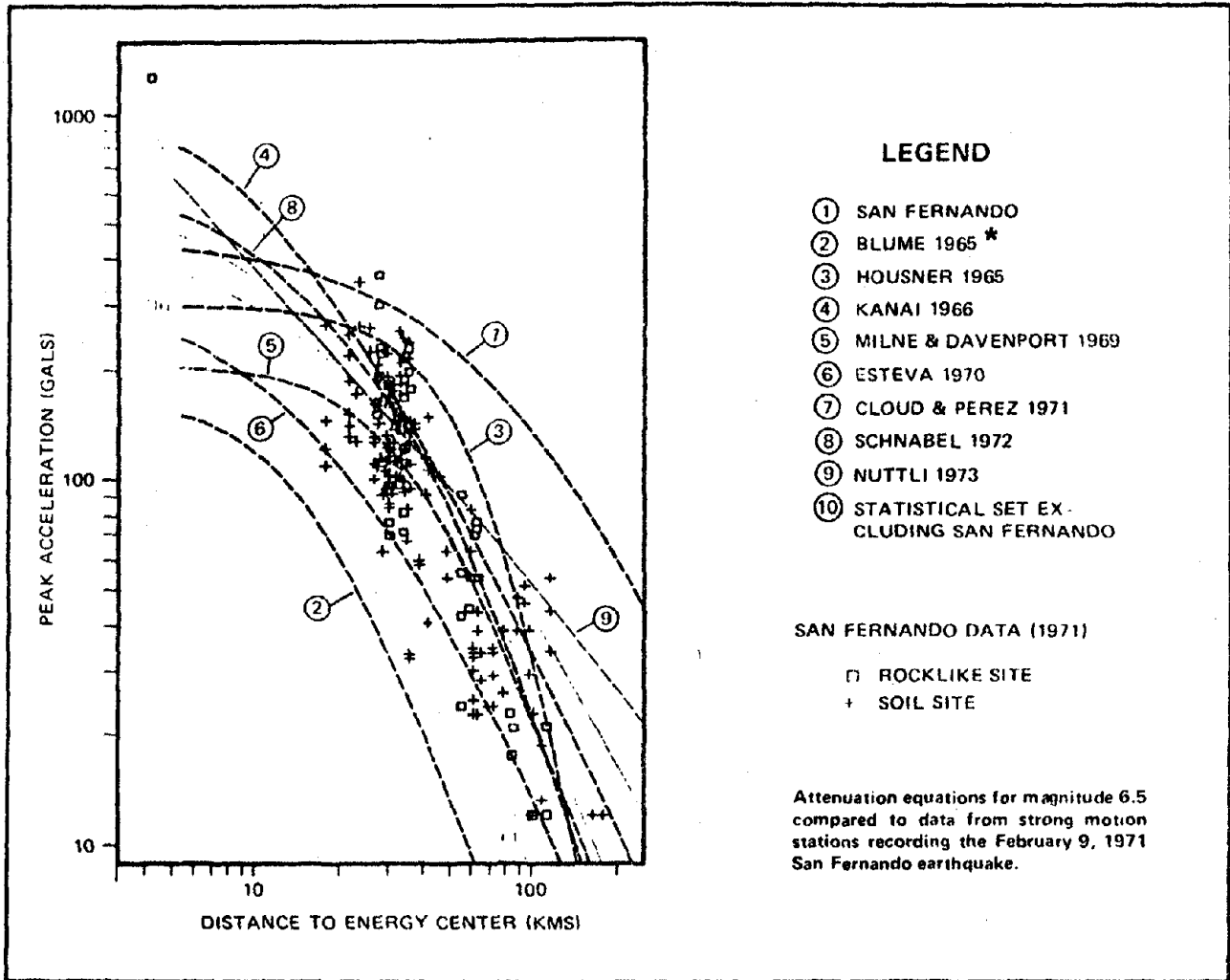
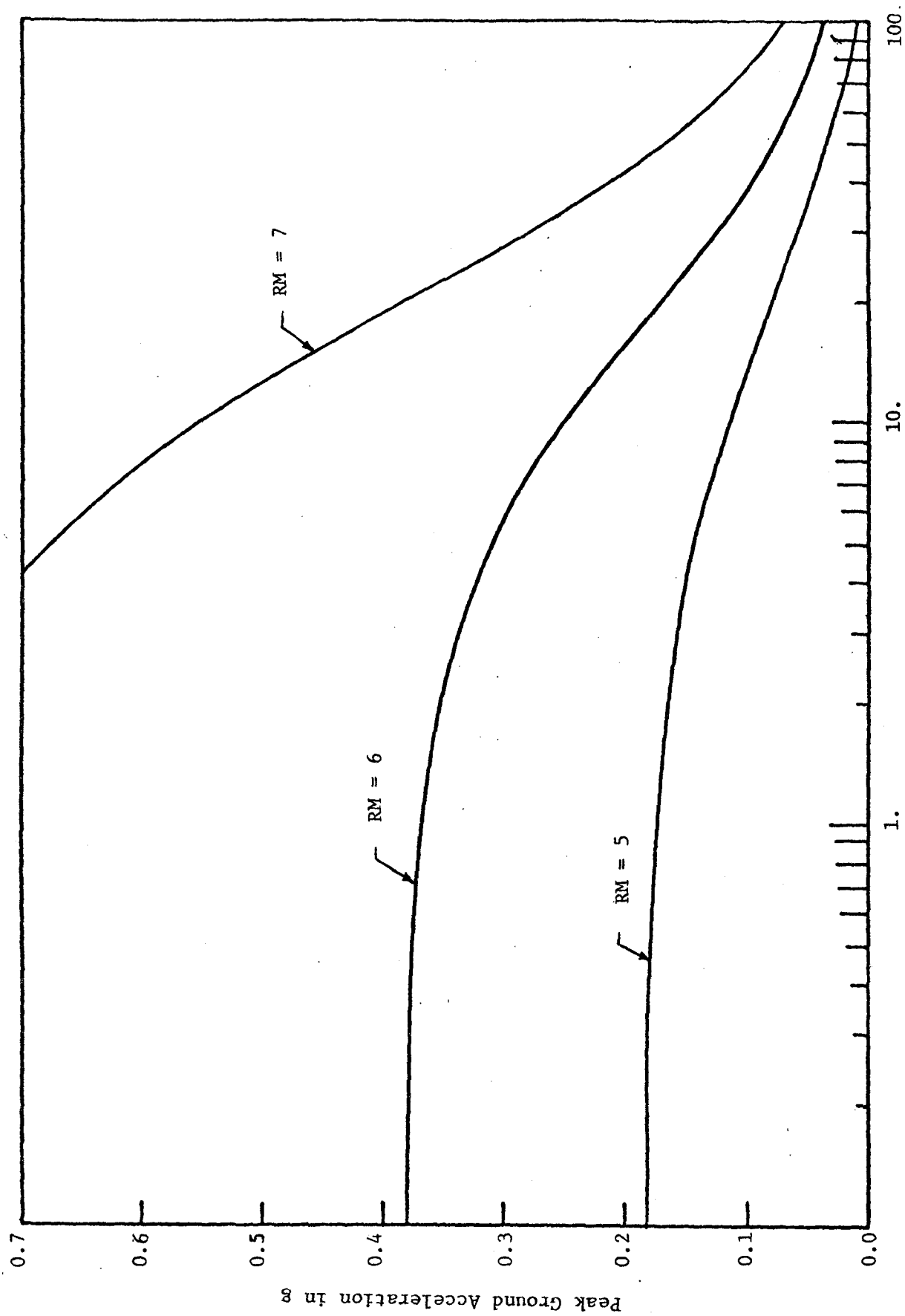


FIGURE A-1a Attenuation Relationships.  
(Taken from Shah et al., 1975).

\* This relation can provide results which are very close to mean data behavior if the soil characteristics for the region are recognized. The soil characteristics input was not used in the preparation of this figure from Shah et al., 1975.



1. Hypocentral Distance -  $R_h$  in km.

FIGURE A-1b Esteva's 1973 Attenuation Relationship

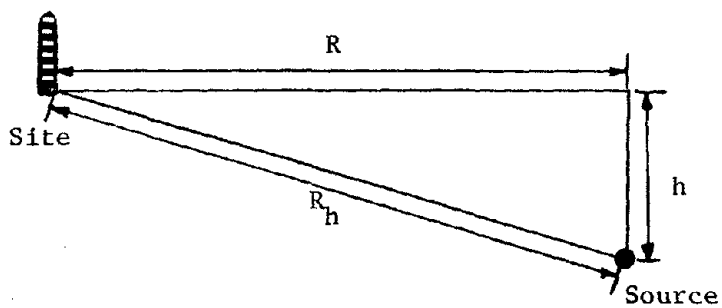
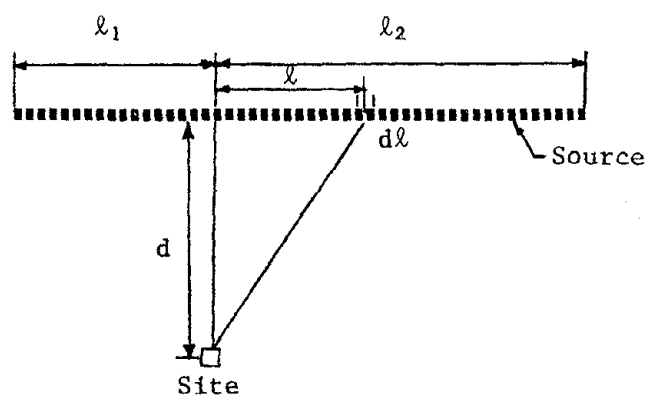
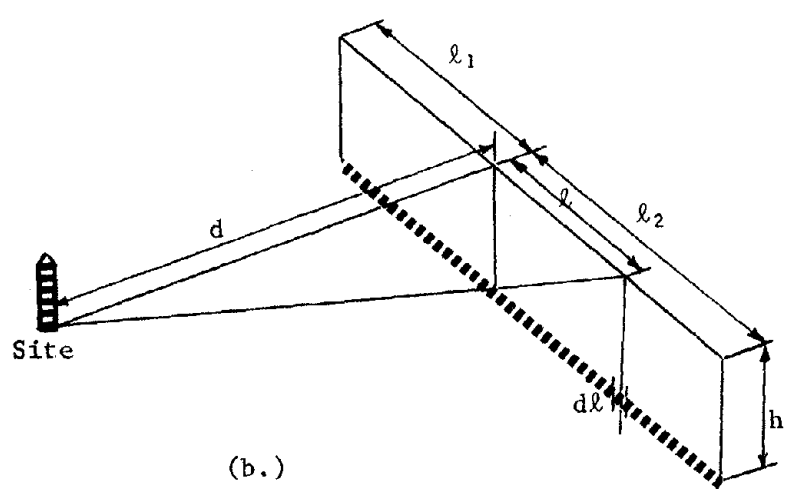


FIGURE A-2 Point Source: Top View



(a.)



(b.)

FIGURE A-3 Line Source: (a.) Top View; (b.) Isometric View



To determine the probability distribution on peak ground acceleration  $a$ , the attenuation relationship is used as shown below:

$$\begin{aligned}
 P[A > a, t] &= P \left[ \frac{b_1 \exp(b_2 M)}{(R_h + b_4)^{b_3}} > a, t \right] \\
 &= P \left[ M > \ln \left[ \frac{a}{b_1} \left[ R_h + b_4 \right]^{b_3} \right]^{\frac{1}{b_2}}, t \right] \quad \text{A-12}
 \end{aligned}$$

The last expression in Equation A-12 can be recognized as the distribution on magnitude. Thus, Equation A-12 becomes

$$P[A > a, t] = 1 - \exp \left[ -e^{\alpha'} \left[ \frac{a}{b_1} \right]^{\frac{\beta}{b_2}} \left[ R_h + b_4 \right]^{\frac{\beta b_3}{b_2}} t \right] \quad \text{A-13}$$

By denoting

$$\gamma = e^{\alpha'} \quad \text{A-14}$$

$$\delta = \frac{\beta}{b_2} \quad \text{A-15}$$

$$\rho = \frac{\beta b_3}{b_2} = \delta b_3 \quad \text{A-16}$$

one finally obtains

$$P[A > a, t] = 1 - \exp \left[ -\gamma \left[ \frac{a}{b_1} \right]^{\delta} \left[ R_h + b_4 \right]^{\rho} t \right] \quad \text{A-17}$$

### Line Source

Most of the earthquake epicenters around the world are generally located along the major fault systems. Thus, the usual case of epicenters falling along a line gives rise to the so-called line source. The line source can be divided into  $\kappa$  small segments of length  $\Delta l_i$ ,  $1 \leq i \leq \kappa$ .

Let  $E_{\ell_i}$  be the event that no earthquake with a Richter magnitude greater than  $m$  occurs on the element  $\Delta\ell_i$  in time  $t$ . Then, from the Poisson condition on spatial independence the following is true:

$$P \left[ \sum_{i=1}^K E_{\ell_i} \right] = \sum_{i=1}^K P \left[ E_{\ell_i} \right] \quad \text{A-18}$$

The probability of having an earthquake of Richter magnitude smaller than or equal to  $m$  due to the entire line source is:

$$\begin{aligned} P[M \leq m, t] &= \lim_{\Delta\ell \rightarrow 0} \sum_{i=1}^K P \left[ E_{\ell_i} \right] = \lim_{\Delta\ell \rightarrow 0} \exp \left[ - \sum_{i=1}^K N'(m) \Delta\ell_i \cdot t \right] = \\ &= \exp \left[ - \int_{\ell_1}^{\ell_2} N'(m) d\ell \cdot t \right] \end{aligned} \quad \text{A-19}$$

where  $\Delta\ell = \max_{1 \leq i \leq K} \Delta\ell_i$

The log-linear recurrence relationship is recalled to be

$$\ln N'(m) = \alpha' + \beta m$$

where  $N'(m)$  and  $\alpha'$  are normalized with respect to the length of the line source and the time period of the data. Substituting for  $N'(m)$  into Equation A-19 one gets:

$$P[M \leq m, t] = \exp \left[ - \int_{\ell_1}^{\ell_2} \exp [\alpha' + \beta m] d\ell \cdot t \right] \quad \text{A-20}$$

To obtain the probability of having a maximum peak ground acceleration  $a$ , the attenuation relationship (Equation A-9) is used as follows:

$$\begin{aligned}
P[A \leq a, t] &= P \left[ \frac{b_1 \exp(b_2 M)}{[R_h + b_4]^{b_3}} \leq a, t \right] = P \left[ M \leq \ln \left[ \frac{a}{b_1} \right]^{b_2} [R_h + b_4]^{\frac{b_3}{b_2}}, t \right] \\
&= \exp \left[ -e^{\alpha'} \left[ \frac{a}{b_1} \right]^{\frac{\beta}{b_1}} \int_{\ell_1}^{\ell_2} [R_h + b_4]^{\frac{\beta b_3}{b_2}} d\ell \cdot t \right] \quad \text{A-21}
\end{aligned}$$

Letting  $e^{\alpha'}$ ,  $\frac{\beta}{b_1}$  and  $\frac{\beta b_3}{b_2}$  be denoted by  $\gamma$ ,  $\delta$  and  $\rho$ , respectively, and noting from the geometry of the line source (Figure A-3) that

$$R_h = [d^2 + \ell^2 + h^2]^{1/2} \quad \text{A-22}$$

one can write:

$$P[A \leq a, t] = \exp \left[ -\gamma \left[ \frac{a}{b_1} \right]^{\delta} t \int_{\ell_1}^{\ell_2} \left[ [d^2 + \ell^2 + h^2]^{1/2} + b_4 \right]^{\delta} d\ell \right] \quad \text{A-24}$$

Alternatively,

$$P[A > a, t] = 1 - \exp \left[ -\gamma \left[ \frac{a}{b_1} \right]^{\delta} t \int_{\ell_1}^{\ell_2} \left[ [d^2 + \ell^2 + h^2]^{1/2} + b_4 \right]^{\delta} d\ell \right] \quad \text{A-25}$$

### Area Source

The peak ground acceleration cumulative distribution function due to an area source at a site can be obtained in a manner similar to that for the line source. Figure A-4 shows schematically the geometry of the area source. Consider an elemental area  $\Delta A_i = R_i \Delta R_i \Delta \theta_i$ . Let  $E_{A_i}$  be the event that no earthquake occurring on the element  $\Delta A_i$  will have a Richter magnitude larger than  $m$  in time  $t$ . Then, again from the Poisson condition on spatial independence it follows:

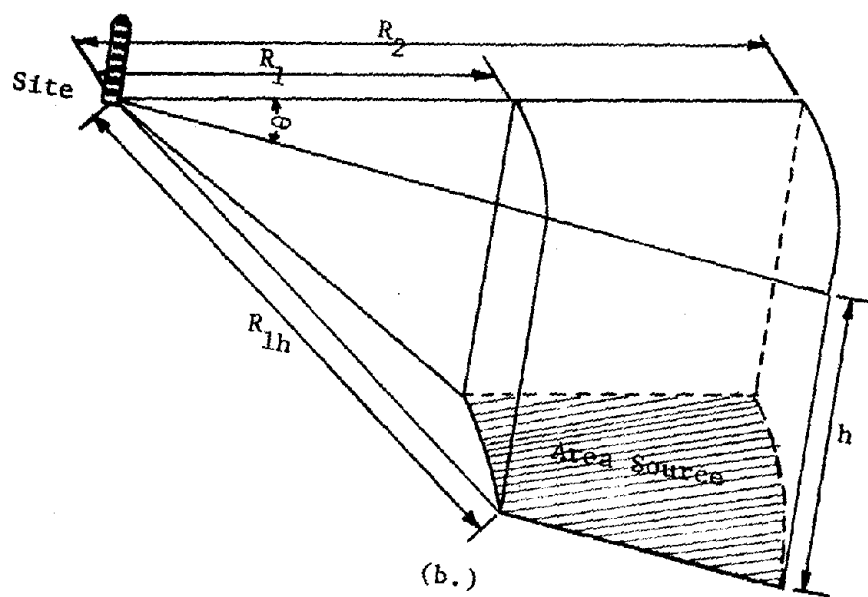
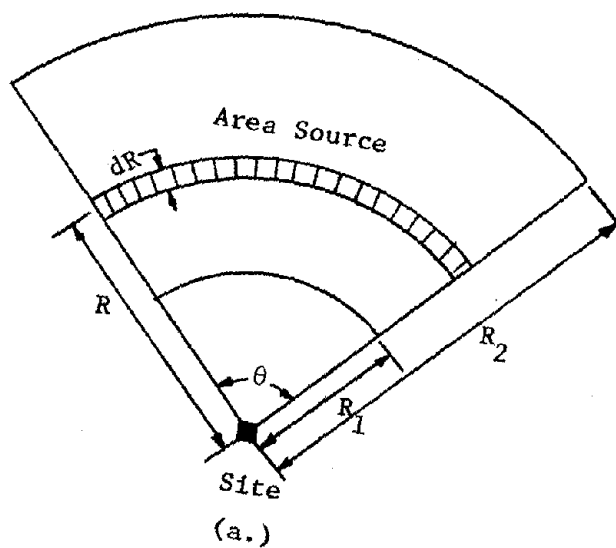


FIGURE A-4 Area Source: (a.) Top View; (b.) Isometric View

$$P \left[ \sum_{i=1}^K E_{A_i} \right] = \sum_{i=1}^K P[E_{A_i}] \quad A-26$$

As  $\Delta A \rightarrow 0$ ,  $\Delta R$  and  $\Delta \theta$  both approach 0. The probability of having an earthquake of Richter magnitude smaller than or equal to  $m$  due to the entire area source is then:

$$\begin{aligned} P[M \leq m, t] &= \lim_{\Delta A \rightarrow 0} \sum_{i=1}^K P[E_{A_i}] = \lim_{\Delta A \rightarrow 0} \exp \left[ - \sum_{i=1}^K N'(m) \Delta A_i t \right] \\ &= \exp \left[ - \int_0^\theta \int_{R_1}^{R_2} N'(m) R dR d\theta \cdot t \right] \end{aligned} \quad A-27$$

Where  $N'(m)$  is the mean rate of occurrence for the area source as defined by Equation 3-3a.

Substituting for  $n'(m)$  in Equation A-27 above

$$P[M \leq m, t] = \exp \left[ - \int_0^\theta \int_{R_1}^{R_2} \exp(\alpha' + \beta m) R dR d\theta \cdot t \right] \quad A-28$$

Using the attenuation relationship of Equation A-9

$$\begin{aligned} P[A \leq a, t] &= P \left[ \frac{b_1 \exp(b_2 M)}{(R_h + b_4)^{b_3}} \leq a, t \right] \\ &= P \left[ M \leq \ln \left[ \frac{a}{b_1} \right]^{\frac{1}{b_2}} \left[ R_h + b_4 \right]^{\frac{b_3}{b_2}}, t \right] \end{aligned} \quad A-29$$

This equation is of the same form as the equation for the distribution on Richter magnitude (see Equation A-8) and thus can be written as:

$$\begin{aligned}
P[A \leq a, t] &= \exp \left[ -e^{\alpha'} \left[ \frac{a}{b_1} \right]^{\frac{\beta}{b_2}} t \int_0^\theta d\theta \int_{R_1}^{R_2} [R^2 + h^2 + b_4]^{\frac{\beta b_3}{b_2}} R dR \right] \\
&= \exp \left[ -e^{\alpha'} \left[ \frac{a}{b_1} \right]^{\frac{\beta}{b_2}} t \theta \int_{R_1}^{R_2} [R^2 + h^2 + b_4]^{\frac{\beta b_3}{b_2}} R dR \right] \quad \text{A-30}
\end{aligned}$$

Let  $\gamma = e^{\alpha'}$

$$\delta = \beta/b_2$$

$$\rho = \beta b_3/b_2 \quad \text{as before}$$

and

$$R_h = \sqrt{R^2 + h^2} \quad \text{A-31}$$

Then

$$P[A \leq a, t] = \exp \left[ -\gamma \left[ \frac{a}{b_1} \right]^\delta t \theta \int_{R_1}^{R_2} (R_h + b_4)^\rho R dR \right] \quad \text{A-32}$$

and

$$P[A > a, t] = 1 - \exp \left[ -\gamma \left[ \frac{a}{b_1} \right]^\delta t \theta \int_{R_1}^{R_2} (R_h + b_4)^\rho R dR \right] \quad \text{A-33}$$

In general, a site may be surrounded by any or all three types of sources discussed in this section. If there are

NP point sources

NL line sources

NA area sources

the cumulative distribution function of peak ground acceleration at a site is then given by:

$$\begin{aligned}
P[\Lambda > a, t] = 1 - \exp & \left[ - \sum_{i=1}^{NP} \gamma_i \left[ \frac{a}{b_1} \right]^{\delta_i} t [R_{h_i} + b_4]^{\rho_i} \right. \\
& - \sum_{j=1}^{NL} \gamma_j \left[ \frac{a}{b_1} \right]^{\delta_j} t \int_{\ell_{1j}}^{\ell_{2j}} \left[ [d_j^2 + \ell^2 + h_j^2]^{1/2} + b_4 \right]^{\rho_j} d\ell \\
& \left. - \sum_{k=1}^{NA} \gamma_k \left[ \frac{a}{b_1} \right]^{\delta_k} t \theta_k \int_{R_{1k}}^{R_{2k}} [R_{h_k} + b_4]^{\rho_k} R dR \right] \quad A-34
\end{aligned}$$

In Equation A-34, the summation over  $i$  is for all point sources, that over  $j$  is for all line sources, and over  $k$  is for all area sources.

As it was shown in Chapter III, eighteen line sources were formulated for the Guatemalan region, based on past data. Any part of the country is affected by these sources, depending upon the proximity of the site to the source location.

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APPENDIX B

BAYESIAN HAZARD MODEL

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## B-1 Introduction

Bayesian probabilistic approach has the distinct advantage that it can include subjective information acquired through experience together with the quantitative data. As an alternative to the Poisson modeling presented in Appendix A, a Bayesian model will be used to develop hazard maps in terms of peak ground acceleration and duration of future earthquakes. The method for obtaining these maps will be only reviewed in the present appendix. Its detailed development is given in Mortgat (1976).

## B-2 Procedure

The procedure for estimating seismic risk for peak ground accelerations in essence consists of the following steps:

### STEP 1. Identification of Earthquake Sources

Based on the geology and historic seismicity of the area sources are identified as point sources, line sources (faults) or area sources. The maximum earthquakes associated with each source are established from historic seismicity and geology (in terms of magnitude or intensity). Typical examples of this approach are Cornell and Vanmarcke (1969), Kiremidjian (1975), Wiggins (1975), Shah et al. (1975), Algermissen (1969).

### STEP 2. Selection of Attenuation Relationship

Using one of the numerous empirical attenuation relationships, the peak accelerations at a given site due to earthquakes of various sizes occurring at different locations are estimated. The attenuation relationship is based on data of non-uniform

quality. Most procedures utilize only the mean curves determined from a regression analysis.

### STEP 3. Recurrence of Earthquakes

The recurrence of earthquakes of various sizes is estimated based primarily on the historic seismicity. A straight line or a set of straight lines is fitted on the data using regression analysis. This method usually results in unacceptable uncertainties for large magnitudes where the data is scarce. Some variations have been proposed. For example Wiggins (1975) uses a Bayesian procedure at the level of the results once the analysis is complete.

### STEP 4. Results

Utilizing the computations in Steps (1), (2) and (3) the probability that a certain acceleration will not be exceeded within a given time period  $t$  is determined. The results of the evaluation are presented in terms of iso-acceleration or iso-intensity curves for selected levels of probability and time periods.

## B-3 Seismic Mapping Model

The following paragraphs present the general model used for seismic mapping in the current study.

### Source Location

The location of the sources is made using the recorded hypocentral position of past earthquakes for the period of the records. The spatial distribution of hypocenters is then divided in different sources as a function of their shapes and seismicity (defined as the number of occurrences per unit area or unit length).

### Seismic Model

In the data presently available, the most commonly used measure of energy release is the Richter magnitude. In this model, the seismicity of the sources is described by the distribution on the number of occurrences for each magnitude. This is different from the current practice where the source seismicity is described by the probability of generating an event larger than a given magnitude and not by a distribution on all the events.

**Occurrences:** Assuming that earthquake occurrences form a Poisson process with mean rate of occurrence independent of magnitude, a distribution is obtained on the number of occurrences for a given period of time for a given source. The assumption of spatial and temporal independence is fairly well verified by data and is a common accepted practice in seismology. Moreover the amount of dependence due to the dual mechanism of stress accumulation and release has not been determined as yet.

**Magnitude:** Given that an event has occurred, a distribution on the magnitude is determined from past data. The magnitudes are discretized every  $1/4$  point as it is commonly done in data recording. This representation has the advantage of getting away from the data fitting. This method is specially valuable for regions where little data is available. The probability corresponding to each magnitude can be used in a Bernoulli trial where one outcome will be an event of the magnitude considered (success) and the other an event of any other magnitude (failure). The following question can then be answered: "Given that  $N$  earthquakes will occur, what is the probability that there will be  $0, 1, 2, \dots, N$  events of any given magnitude?" Combining those Binomial conditional distributions with the Poisson distribution on occurrences, the distribution on the number of occurrences for each magnitude is obtained.

### Bayesian Statistics

Bayesian statistics is applied to the Poisson and Binomial laws to eliminate some of the incompleteness of the data. For example, considering the fault length and the type of fault, geologists can determine the maximum magnitude earthquake the source can generate. This information has to be taken into consideration even if no such earthquake has been recorded in the data. This is done by assuming the mean rate of occurrences of the Poisson law to be a Gamma probability distribution (Benjamin and Cornell, 1970) and the probability of success of the Binomial law to be a Beta probability distribution (Benjamin and Cornell, 1970). This method has the advantage of including personal experience together with the data as well as updating the distribution as new data is made available.

### Mapping Parameters

Two mapping parameters are used, namely the PGA and the duration of the ground motion. PGA was used since no other attenuation relationship is readily available in the literature. The methodology allows for the use of a more stable parameter such as rms that would certainly improve the reliability of the model.

### Attenuation Relationship

The Esteva (1973) relationship is used to relate PGA to magnitude as a function of distance. Attenuation on duration is obtained using the relation developed by Bolt (1973).

Both relationships are treated probabilistically to take into consideration uncertainty in the attenuation decay.

Combining the distributions that describe the seismicity at the sources with the two transfer functions, the probability of exceedance of any PGA or duration is obtained at the site.

TABLE B-1

Duration Attenuation

| M<br>distance | 5.5 | 6.0 | 6.5 | 7.0 | 7.5 | 8.0 | 8.5 |
|---------------|-----|-----|-----|-----|-----|-----|-----|
| 10            | 8   | 12  | 19  | 26  | 31  | 34  | 35  |
| 25            | 4   | 9   | 15  | 24  | 28  | 30  | 32  |
| 50            | 2   | 3   | 10  | 22  | 26  | 28  | 29  |
| 75            | 1   | 1   | 5   | 10  | 14  | 16  | 17  |
| 100           | 0   | 0   | 1   | 4   | 5   | 6   | 7   |
| 125           | 0   | 0   | 1   | 2   | 2   | 3   | 3   |
| 150           | 0   | 0   | 0   | 1   | 2   | 2   | 3   |
| 175           | 0   | 0   | 0   | 0   | 1   | 2   | 2   |
| 200           | 0   | 0   | 0   | 0   | 0   | 1   | 2   |

The Bayesian forecast model discussed in the previous section is applied to Guatemala. The data base and the eighteen seismic sources described in Chapter III are used directly in this model. For each source the following information is gathered:

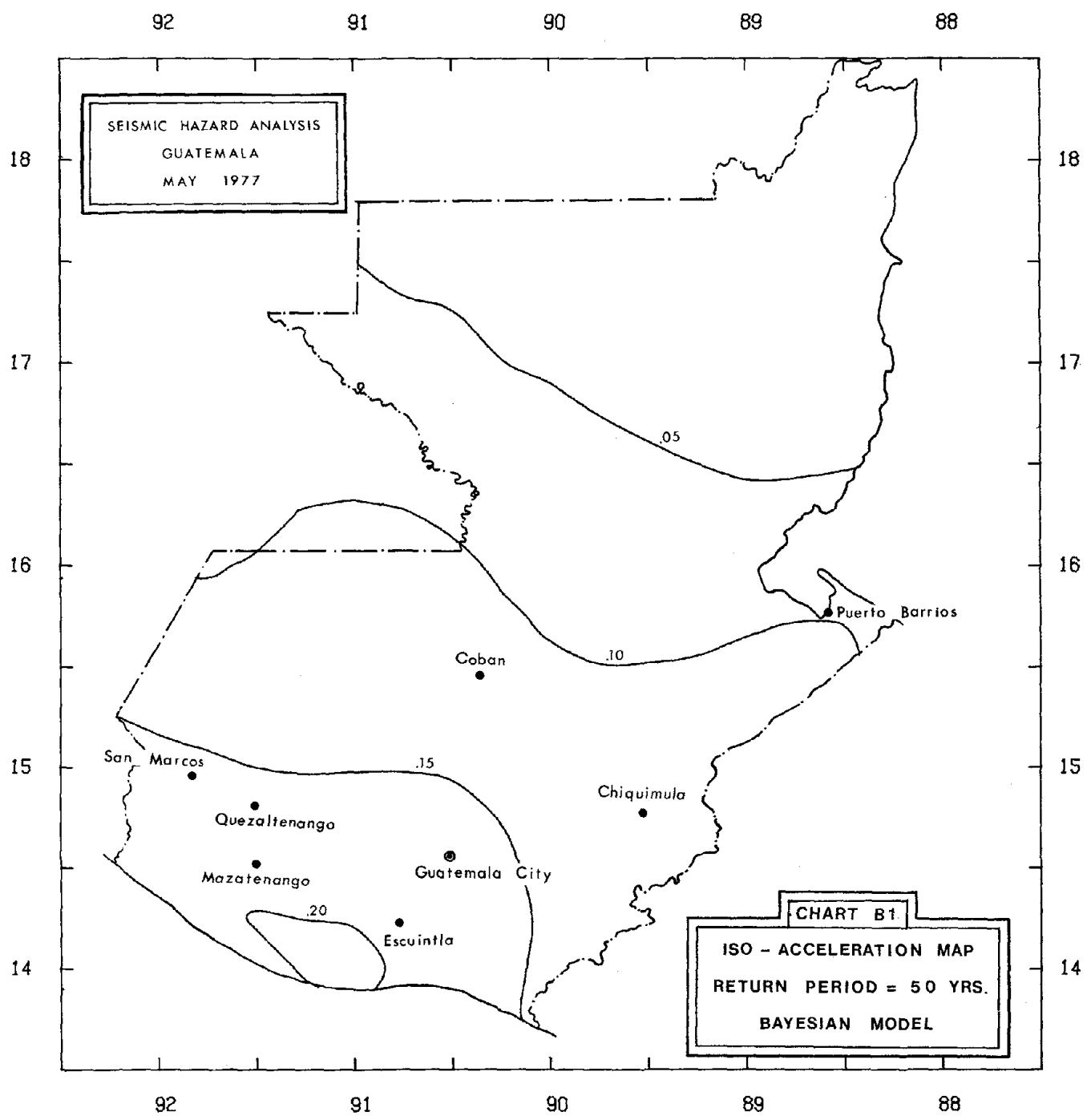
- (1) The number of recorded earthquakes  $N$  for discretized intervals of Richter magnitude  $M$ . An increment of  $0.25M$  is used for the discretization.
- (2) The time period of the data  $T$ .
- (3) The value of the parameters  $\lambda'$  and  $\nu'$  of the Gamma distribution (Benjamin and Cornell, 1970) determined from subjective input.
- (4) The values of the parameters  $\eta'_{M_i}$  and  $\xi'_{M_i}$  of the Beta distribution (Benjamin and Cornell, 1970) which also come from subjective input.

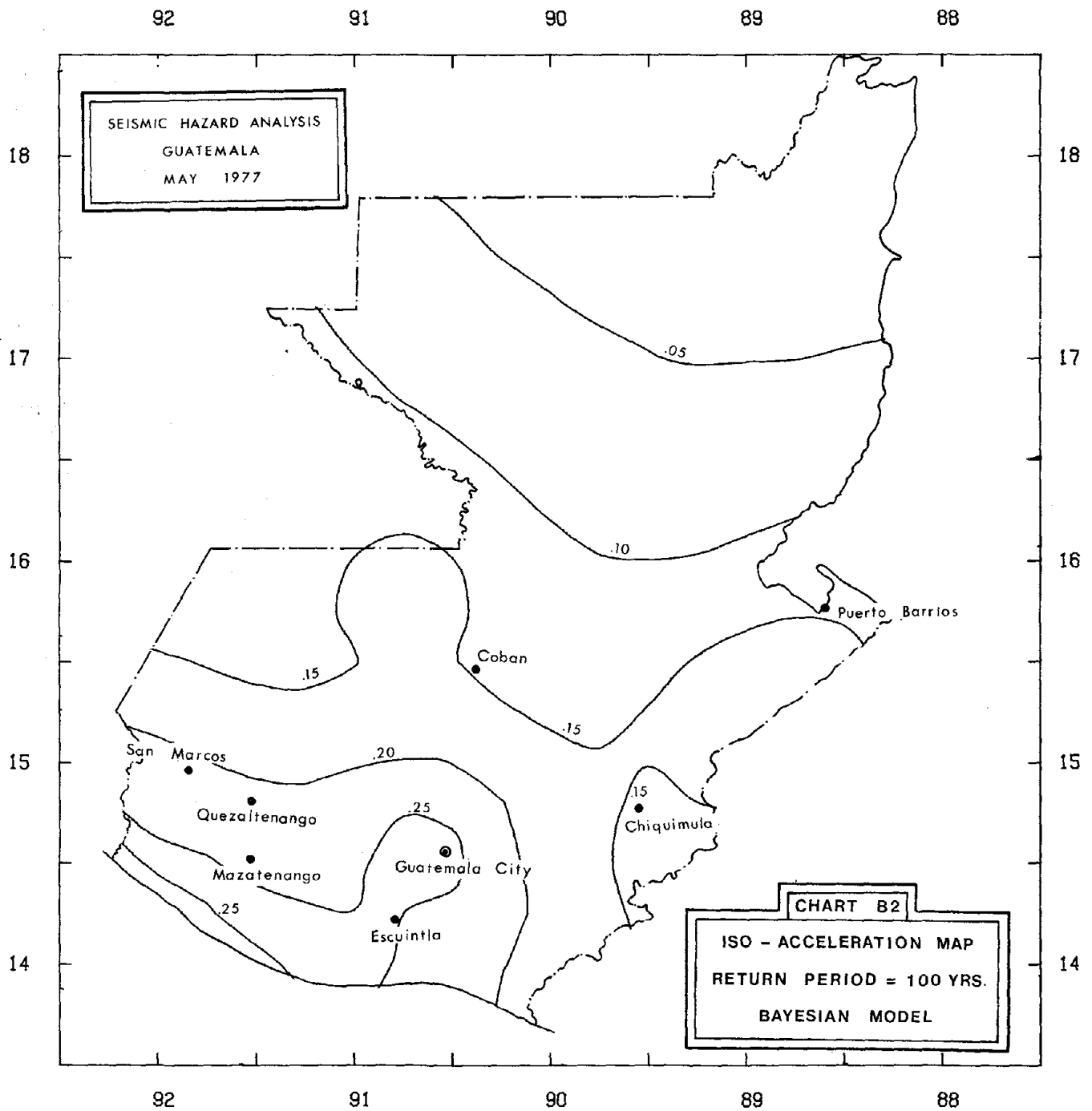
The values of all of these parameters are determined for each source and are listed in Tables B2 - B19. This procedure developed by Mortgat (1976) has a uniform distribution associated with the peak ground acceleration attenuation relationship (Equation 4-1). The coefficient of variation for this distribution is assumed to be 0.30.

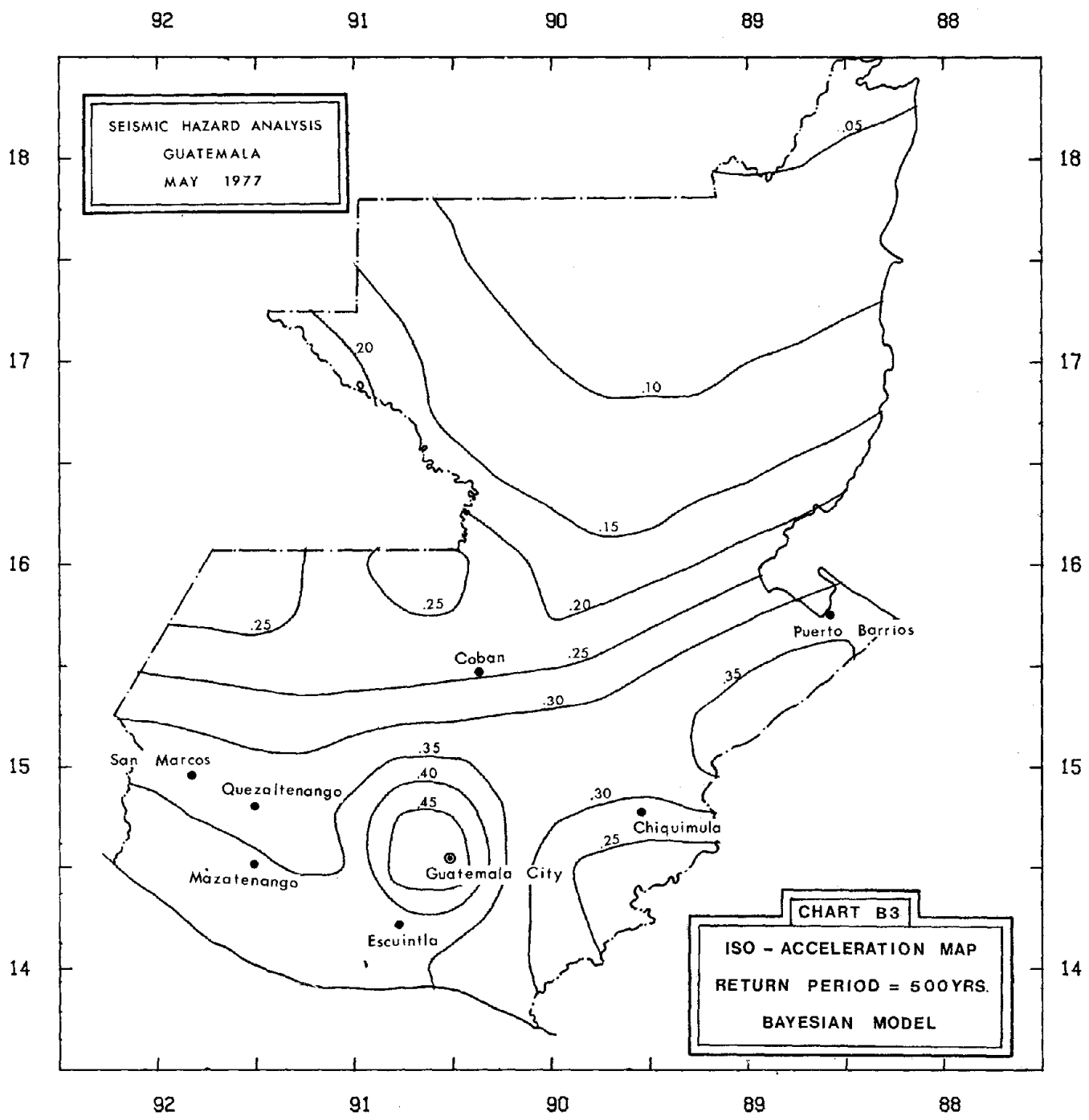
For the duration maps, Table B1 lists the relationship between the Richter magnitude, distance from the source to the site and the duration (from Bolt, 1973). A coefficient of variation of 0.30 is also applied to these duration values.

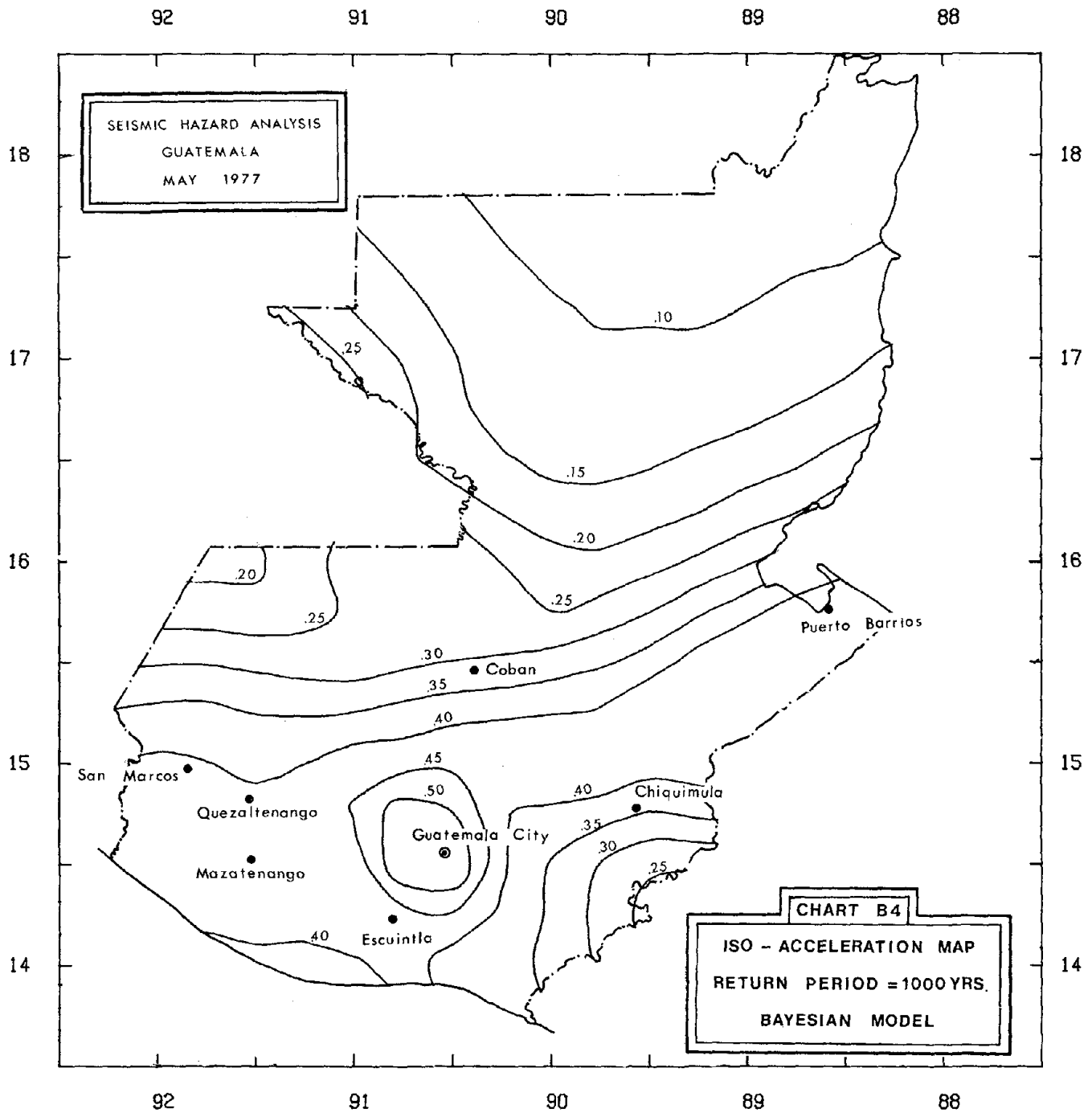
Charts B1 to B4 show the resulting iso-acceleration maps for return periods of 50, 100, 500 and 1000 years. The iso-duration maps for the same return periods are represented in Charts B5 to B8. A low seismicity region in the northern part of Guatemala can be observed in all maps both for duration and acceleration. The high regions in the south-south-west portion and along the Cayman trough is also common to all maps.

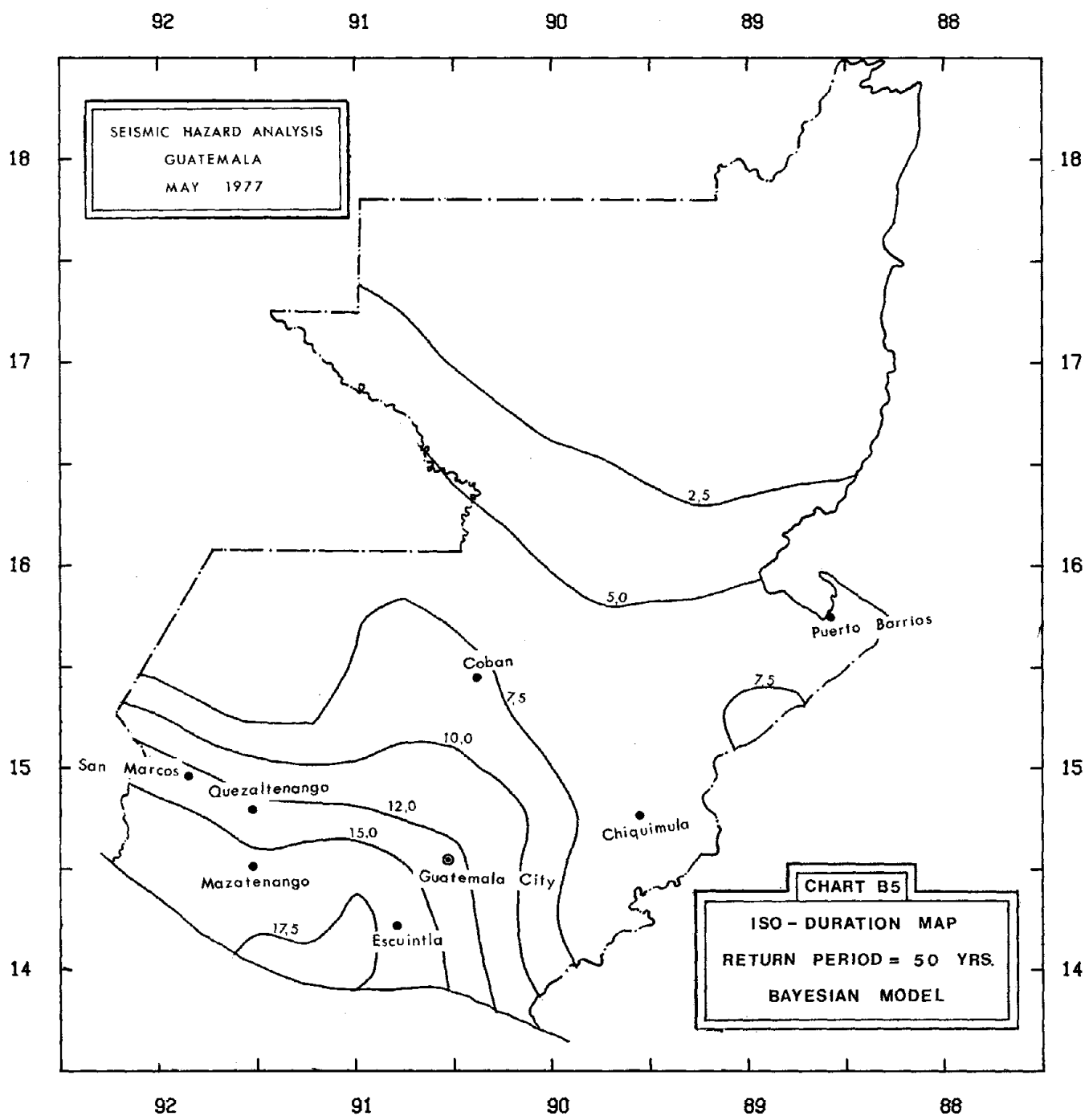


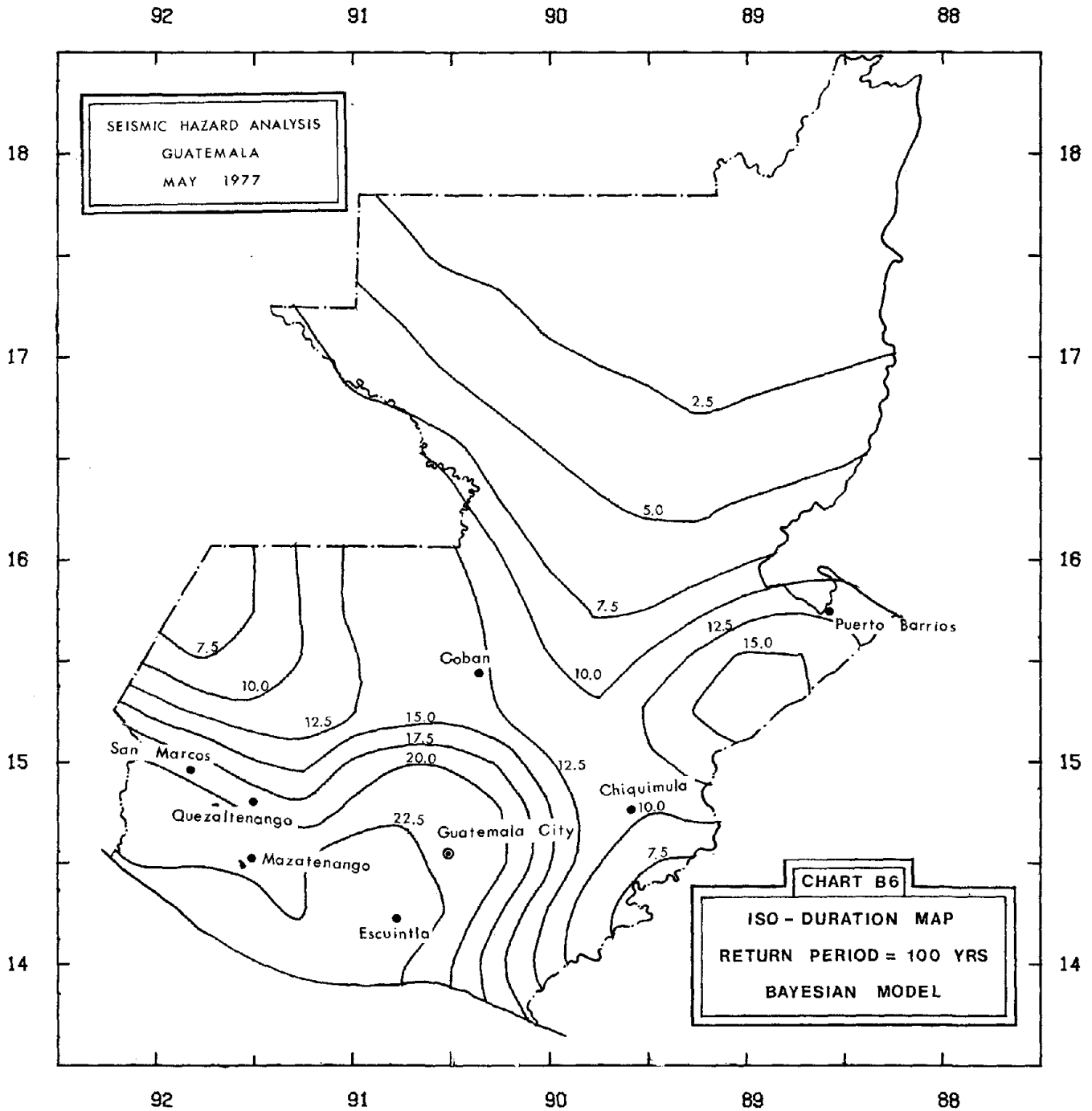


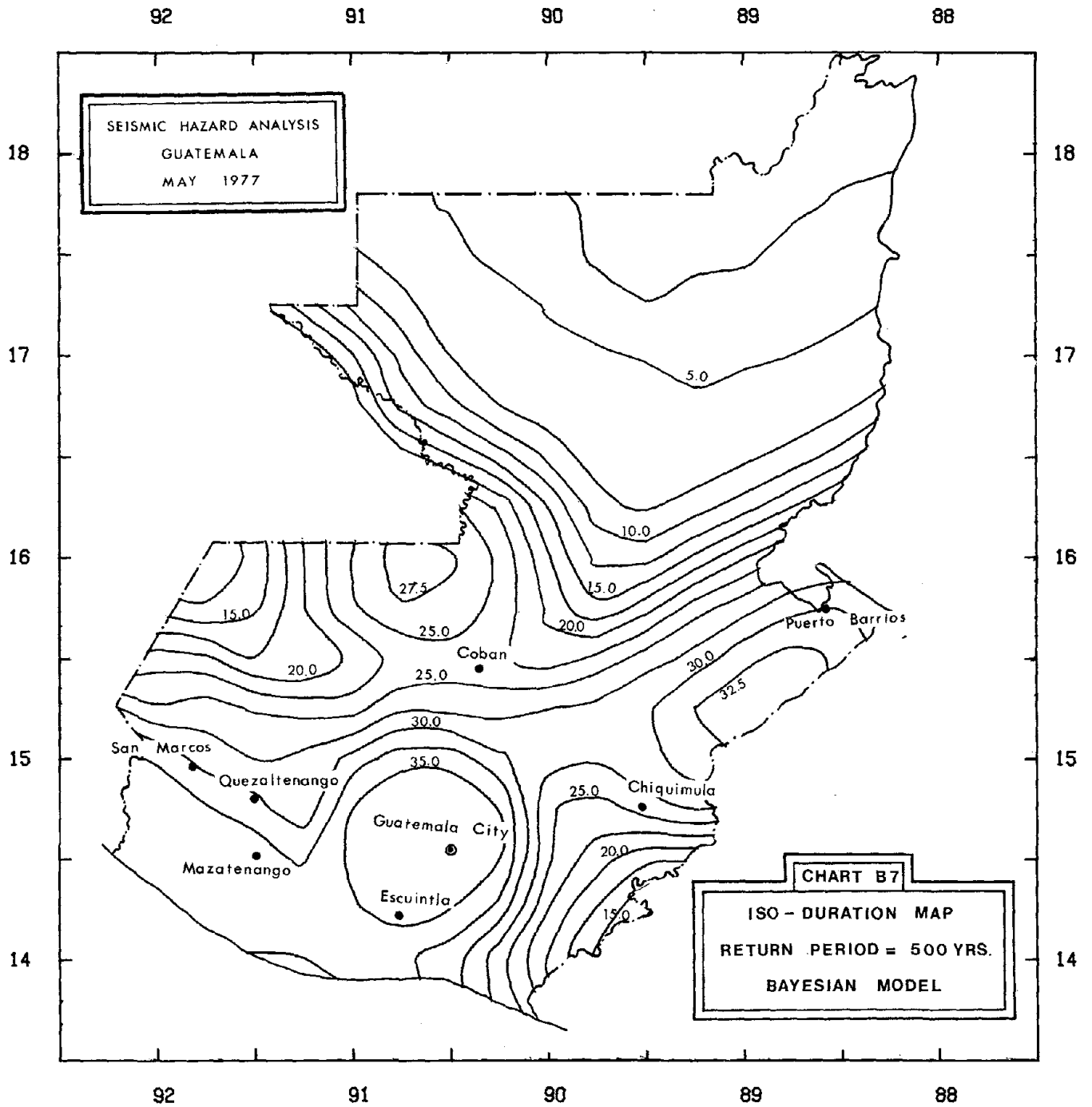


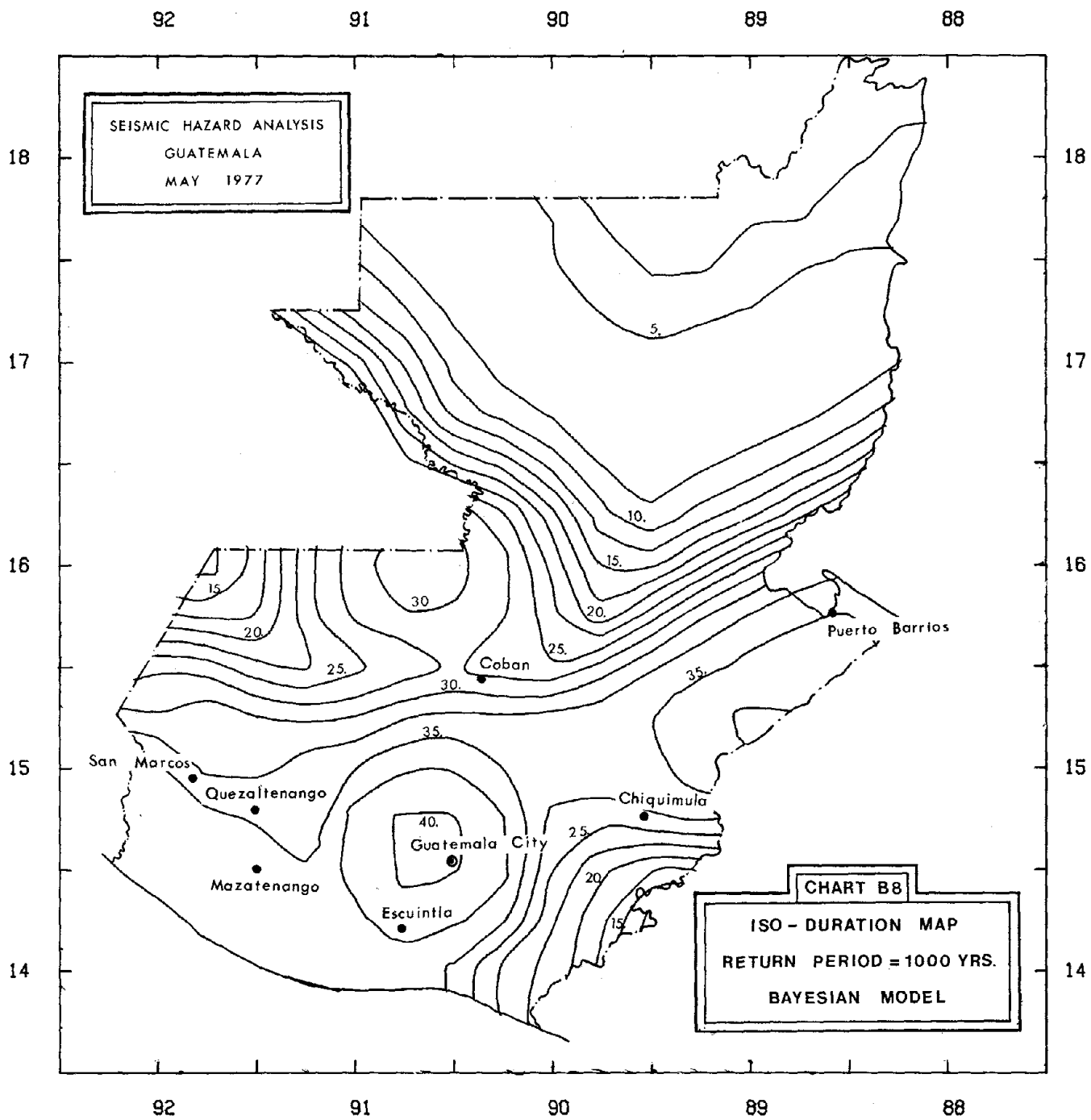














Comparing the iso-acceleration maps from the Bayesian model to the iso-acceleration maps from the original Poisson model (Charts 7 to 10), the acceleration values are higher for the maps in Charts B1 to B4. Also, in Charts 7 to 10 the effect of some of the faults (e.g., Mixco, Motagua, Polochic) has been accentuated while the effect of some others (e.g., Jocotan) has been reduced. Overall however, the maps appear quite similar. It should be pointed out that the advantages of the Bayesian approach are:

- (1) subjective information can be included
- (2) it is easier to incorporate additional earthquake data.

In a manner similar to Charts 7 to 10, the iso-acceleration and iso-duration maps resulting from the Bayesian approach can be used for zoning purposes. Again, their application for design and analysis of structures in Guatemala will be discussed at length in the Part II report of the present study.

TABLE B2  
Bayesian Data Information

| Source 1  |   |   |  |  |
|---|---|---|--|--|
| Time Data Base (T) : 75 years<br>Number of Recorded Events (N) : 12<br>$\nu'$ from log-linear fit : 14.14<br>$\lambda'' = \lambda' + T = 75 + 75 = 150$<br>$\nu'' = \nu' + N = 14.14 + 12 = 26.14$<br>$\eta''_{M_i} = \eta'_{M_i} + N = 14.14 + 12 = 26.14$ |   |   |  |  |
| Richter Magnitude ( $M_i$ )   | Nb of Recorded Occurrences in $M_i$ bands ( $R_{M_i}$ ) | Cumulative Nb of Occurrences (log-linear fit Fig. ) ( $N_c$ ) | Nb of Occurrences in $M_i$ bands (log-linear fit) ( $\xi'_{M_i}$ ) | $\xi'_{M_i} + R_{M_i}$ ( $\xi''_{M_i}$ ) |
| 4.00  | 2   | 14.14   | 2.55   | 4.55                                     |
| 4.25  | 1   | 11.59   | 2.09   | 3.09                                     |
| 4.50  | 2   | 9.50  | 1.71   | 3.71                                     |
| 4.75  | 0   | 7.79  | 1.41   | 1.41                                     |
| 5.00  | 2   | 6.38  | 1.15   | 3.15                                     |
| 5.25  | 0   | 5.23  | 0.94   | 0.94                                     |
| 5.50  | 2   | 4.29  | 0.78   | 2.78                                     |
| 5.75  | 1   | 3.51  | 0.63   | 1.63                                     |
| 6.00  | 0   | 2.88  | 0.52   | 0.52                                     |
| 6.25  | 0   | 2.36  | 0.42   | 0.42                                     |
| 6.50  | 1   | 1.94  | 0.35   | 1.35                                     |
| 6.75  | 0   | 1.59  | 0.29   | 0.29                                     |
| 7.00  | 0   | 1.30  | 0.24   | 0.24                                     |
| 7.25  | 0   | 1.06  | 0.19   | 0.19                                     |
| 7.50  | 1   | 0.87  | 0.15   | 1.15                                     |
| 7.75  | 0   | 0.72  | 0.13   | 0.13                                     |
| 8.00  | 0   | 0.59  | 0.11   | 0.11                                     |
| 8.25  | 0   | 0.48  | 0.09   | 0.09                                     |
| 8.50  | 0   | 0.39  | 0.39   | 0.39                                     |

TABLE B3  
Bayesian Data Information

| Source 2  |   |   |  |  |
|---|---|---|--|--|
| Time Data Base (T) : 75 years<br>Number of Recorded Events (N) : 6<br>$\nu'$ from log-linear fit : 7.09<br>$\lambda'' = \lambda' + T = 75 + 75 = 150$<br>$\nu'' = \nu' + N = 7.09 + 6 = 13.09$<br>$\eta''_{M_i} = \eta'_{M_i} + N = 7.09 + 6 = 13.09$ |   |   |  |  |
| Richter Magnitude ( $M_i$ )   | Nb of Recorded Occurrences in $M_i$ bands ( $R_{M_i}$ ) | Cumulative Nb of Occurrences (log-linear fit Fig. ) ( $N_c$ ) | Nb of Occurrences in $M_i$ bands (log-linear fit) ( $\xi'_{M_i}$ ) | $\xi'_{M_i} + R_{M_i}$ ( $\xi''_{M_i}$ ) |
| 4.00  | 1   | 7.09  | 1.28   | 2.28                                     |
| 4.25  | 0   | 5.81  | 1.05   | 1.05                                     |
| 4.50  | 1   | 4.76  | 0.86   | 1.86                                     |
| 4.75  | 0   | 3.90  | 0.70   | 0.70                                     |
| 5.00  | 1   | 3.20  | 0.58   | 1.58                                     |
| 5.25  | 0   | 2.62  | 0.47   | 0.47                                     |
| 5.50  | 2   | 2.15  | 0.39   | 2.39                                     |
| 5.75  | 0   | 1.76  | 0.32   | 0.32                                     |
| 6.00  | 1   | 1.44  | 0.26   | 1.26                                     |
| 6.25  | 0   | 1.18  | 0.21   | 0.21                                     |
| 6.50  | 0   | 0.97  | 0.18   | 0.18                                     |
| 6.75  | 0   | 0.79  | 0.14   | 0.14                                     |
| 7.00  | 0   | 0.65  | 0.12   | 0.12                                     |
| 7.25  | 0   | 0.53  | 0.09   | 0.09                                     |
| 7.50  | 0   | 0.44  | 0.08   | 0.08                                     |
| 7.75  | 0   | 0.36  | 0.07   | 0.07                                     |
| 8.00  | 0   | 0.29  | 0.29   | 0.29                                     |

TABLE B4  
Bayesian Data Information

| Source 3  |   |   |   |  |
|---|---|---|---|--|
| Time Data Base (T) : 75 years<br>Number of Recorded Events (N) : 9<br>$\nu'$ from log-linear fit : 8.89<br>$\lambda'' = \lambda' + T = 75 + 75 = 150$<br>$\nu'' = \nu' + N = 8.89 + 9 = 17.89$<br>$\eta''_{M_i} = \eta'_{M_i} + N = 8.89 + 9 = 17.89$ |   |   |   |  |
| Richter Magnitude ( $M_i$ )   | Nb of Recorded Occurrences in $M_i$ bands ( $R_{M_i}$ ) | Cumulative Nb of Occurrences (log-linear fit Fig. ) ( $N_c$ ) | Nb of Occurrences in $M_i$ bands (log-linear fit) ( $\epsilon'_{M_i}$ ) | $\epsilon'_{M_i} + R_{M_i}$ ( $\epsilon''_{M_i}$ ) |
| 4.00  | 1   | 8.89  | 1.47  | 2.47   |
| 4.25  | 1   | 7.42  | 1.22  | 2.22   |
| 4.50  | 2   | 6.20  | 1.02  | 3.02   |
| 4.75  | 0   | 5.18  | 0.86  | 0.86   |
| 5.00  | 2   | 4.32  | 0.71  | 2.71   |
| 5.25  | 0   | 3.61  | 0.60  | 0.60   |
| 5.50  | 0   | 3.01  | 0.55  | 0.55   |
| 5.75  | 0   | 2.52  | 0.42  | 0.42   |
| 6.00  | 1   | 2.10  | 0.34  | 1.34   |
| 6.25  | 1   | 1.76  | 0.29  | 1.29   |
| 6.50  | 0   | 1.47  | 0.22  | 0.22   |
| 6.75  | 1   | 1.22  | 0.20  | 1.20   |
| 7.00  | 0   | 1.02  | 0.17  | 0.17   |
| 7.25  | 0   | 0.85  | 0.14  | 0.14   |
| 7.50  | 0   | 0.71  | 0.11  | 0.11   |
| 7.75  | 0   | 0.60  | 0.10  | 0.10   |
| 8.00  | 0   | 0.50  | 0.50  | 0.50   |

TABLE B5  
Bayesian Data Information

| Source 4  |   |   |   |  |
|---|---|---|---|--|
| Time Data Base (T) : 75 years<br>Number of Recorded Events (N) : 17<br>$\nu'$ from log-linear fit : 21.53<br>$\lambda'' = \lambda' + T = 75 + 75 = 150$<br>$\nu'' = \nu' + N = 21.53 + 17 = 38.53$<br>$\eta''_{M_i} = \eta'_{M_i} + N = 21.53 + 17 = 38.53$ |   |   |   |  |
| Richter Magnitude ( $M_i$ )   | Nb of Recorded Occurrences in $M_i$ bands ( $R_{M_i}$ ) | Cumulative Nb of Occurrences (log-linear fit Fig. ) ( $N_c$ ) | Nb of Occurrences in $M_i$ bands (log-linear fit) ( $\epsilon'_{M_i}$ ) | $\epsilon'_{M_i} + R_{M_i}$ ( $\epsilon''_{M_i}$ ) |
| 4.00  | 2   | 21.53   | 2.80  | 4.80   |
| 4.25  | 1   | 18.73   | 2.44  | 3.44   |
| 4.50  | 2   | 16.29   | 2.12  | 5.12   |
| 4.75  | 0   | 14.17   | 1.85  | 1.85   |
| 5.00  | 2   | 12.32   | 1.50  | 3.50   |
| 5.25  | 0   | 10.82   | 1.50  | 1.50   |
| 5.50  | 1   | 9.32  | 1.21  | 2.21   |
| 5.75  | 0   | 8.11  | 1.06  | 1.06   |
| 6.00  | 2   | 7.05  | 0.91  | 2.91   |
| 6.25  | 1   | 6.14  | 0.92  | 1.92   |
| 6.50  | 3   | 5.22  | 2.18  | 5.18   |
| 6.75  | 1   | 3.04  | 1.27  | 2.27   |
| 7.00  | 1   | 1.77  | 0.74  | 1.74   |
| 7.25  | 1   | 1.03  | 0.43  | 1.43   |
| 7.50  | 0   | 0.60  | 0.25  | 0.25   |
| 7.75  | 0   | 0.35  | 0.15  | 0.15   |
| 8.00  | 0   | 0.20  | 0.20  | 0.20   |

TABLE B6  
Bayesian Data Information

| Source 5  |   |   |  |                                      |
|---|---|---|--|--------------------------------------|
| Time Data Base (T) : 75 years<br>Number of Recorded Events (N) : 5<br>$\nu'$ from log-linear fit : 5.52<br>$\lambda'' = \lambda' + T = 75 + 75 = 150$<br>$\nu'' = \nu' + N = 5.53 + 5 = 10.53$<br>$\eta''_{M_i} = \eta'_{M_i} + N = 5.53 + 5 = 10.53$ |   |   |  |                                      |
| Richter Magnitude ( $M_i$ )   | Nb of Recorded Occurrences in $M_i$ bands ( $R_{M_i}$ ) | Cumulative Nb of Occurrences (log-linear fit Fig. ) ( $N_c$ ) | Nb of Occurrences in $M_i$ bands (log-linear fit) ( $E'_{M_i}$ ) | $E'_{M_i} + R_{M_i}$ ( $E''_{M_i}$ ) |
| 4.00  | 1   | 5.53  | 1.57   | 2.57                                 |
| 4.25  | 1   | 3.96  | 1.14   | 2.14                                 |
| 4.50  | 0   | 2.82  | 0.80   | 0.80                                 |
| 4.75  | 2   | 2.02  | 0.58   | 2.58                                 |
| 5.00  | 1   | 1.44  | 0.41   | 1.41                                 |
| 5.25  | 0   | 1.03  | 0.29   | 0.29                                 |
| 5.50  | 0   | 0.74  | 0.21   | 0.21                                 |
| 5.75  | 0   | 0.53  | 0.15   | 0.15                                 |
| 6.00  | 0   | 0.38  | 0.38   | 0.38                                 |

TABLE B7  
Bayesian Data Information

| Source 6  |   |   |  |                                      |
|---|---|---|--|--------------------------------------|
| Time Data Base (T) : 75 years<br>Number of Recorded Events (N) : 3<br>$\nu'$ from log-linear fit : 2.74<br>$\lambda'' = \lambda' + T = 75 + 75 = 150$<br>$\nu'' = \nu' + N = 2.74 + 3 = 5.74$<br>$\eta''_{M_i} = \eta'_{M_i} + N = 2.74 + 3 = 5.75$ |   |   |  |                                      |
| Richter Magnitude ( $M_i$ )   | Nb of Recorded Occurrences in $M_i$ bands ( $R_{M_i}$ ) | Cumulative Nb of Occurrences (log-linear fit Fig. ) ( $N_c$ ) | Nb of Occurrences in $M_i$ bands (log-linear fit) ( $E'_{M_i}$ ) | $E'_{M_i} + R_{M_i}$ ( $E''_{M_i}$ ) |
| 4.00  | 1   | 2.74  | 0.41   | 1.41                                 |
| 4.25  | 1   | 2.33  | 0.34   | 1.34                                 |
| 4.50  | 0   | 1.99  | 0.30   | 0.30                                 |
| 4.75  | 0   | 1.69  | 0.25   | 0.25                                 |
| 5.00  | 0   | 1.44  | 0.22   | 0.22                                 |
| 5.25  | 0   | 1.22  | 0.18   | 0.18                                 |
| 5.50  | 0   | 1.04  | 0.15   | 0.15                                 |
| 5.75  | 1   | 0.89  | 0.14   | 1.14                                 |
| 6.00  | 0   | 0.75  | 0.11   | 0.11                                 |
| 6.25  | 0   | 0.64  | 0.10   | 0.10                                 |
| 6.50  | 0   | 0.54  | 0.54   | 0.54                                 |

TABLE B8  
Bayesian Data Information

| Source 7  |   |   |   |  |
|---|---|---|---|--|
| Time Data Base (T) : 75 years<br>Number of Recorded Events (N) : 5<br>$\nu'$ from log-linear fit : 4.22<br>$\lambda'' = \lambda' + T = 75 + 75 = 150$<br>$\nu'' = \nu' + N = 4.22 + 5 = 9.22$<br>$\eta''_{M_i} = \eta'_{M_i} + N = 4.22 + 5 = 9.22$ |   |   |   |  |
| Richter Magnitude ( $M_i$ )   | Nb of Recorded Occurrences in $M_i$ bands ( $R_{M_i}$ ) | Cumulative Nb of Occurrences (log-linear fit Fig. ) ( $N_c$ ) | Nb of Occurrences in $M_i$ bands (log-linear fit) ( $\epsilon'_{M_i}$ ) | $\epsilon'_{M_i} + R_{M_i}$ ( $\epsilon''_{M_i}$ ) |
| 4.00  | 1   | 4.22  | 0.62  | 1.62   |
| 4.25  | 1   | 3.60  | 0.53  | 1.53   |
| 4.50  | 1   | 3.07  | 0.45  | 1.45   |
| 4.75  | 0   | 2.62  | 0.39  | 0.39   |
| 5.00  | 1   | 2.23  | 0.32  | 1.32   |
| 5.25  | 0   | 1.91  | 0.28  | 0.28   |
| 5.50  | 0   | 1.63  | 0.24  | 0.24   |
| 5.75  | 0   | 1.39  | 0.21  | 0.21   |
| 6.00  | 0   | 1.18  | 0.17  | 0.17   |
| 6.25  | 0   | 1.01  | 0.15  | 0.15   |
| 6.50  | 1   | 0.86  | 0.13  | 0.13   |
| 6.75  | 0   | 0.73  | 0.10  | 1.10   |
| 7.00  | 0   | 0.63  | 0.63  | 0.63   |

TABLE B9  
Bayesian Data Information

| Source 8  |   |   |   |  |
|---|---|---|---|--|
| Time Data Base (T) : 75 years<br>Number of Recorded Events (N) : 6<br>$\nu'$ from log-linear fit : 7.22<br>$\lambda'' = \lambda' + T = 75 + 75 = 150$<br>$\nu'' = \nu' + N = 7.22 + 6 = 13.22$<br>$\eta''_{M_i} = \eta'_{M_i} + N = 7.22 + 6 = 13.22$ |   |   |   |  |
| Richter Magnitude ( $M_i$ )   | Nb of Recorded Occurrences in $M_i$ bands ( $R_{M_i}$ ) | Cumulative Nb of Occurrences (log-linear fit Fig. ) ( $N_c$ ) | Nb of Occurrences in $M_i$ bands (log-linear fit) ( $\epsilon'_{M_i}$ ) | $\epsilon'_{M_i} + R_{M_i}$ ( $\epsilon''_{M_i}$ ) |
| 4.00  | 1   | 7.22  | 1.00  | 2.00   |
| 4.25  | 0   | 6.22  | 0.85  | 0.85   |
| 4.50  | 1   | 5.37  | 0.75  | 1.75   |
| 4.75  | 0   | 4.62  | 0.63  | 0.63   |
| 5.00  | 1   | 3.99  | 0.55  | 1.55   |
| 5.25  | 0   | 3.44  | 0.48  | 0.48   |
| 5.50  | 0   | 2.96  | 0.41  | 0.41   |
| 5.75  | 1   | 2.55  | 0.35  | 1.35   |
| 6.00  | 2   | 2.20  | 0.30  | 2.30   |
| 6.25  | 0   | 1.90  | 0.26  | 0.26   |
| 6.50  | 0   | 1.64  | 0.23  | 0.23   |
| 6.75  | 0   | 1.41  | 0.19  | 0.19   |
| 7.00  | 0   | 1.22  | 0.17  | 0.17   |
| 7.25  | 0   | 1.05  | 0.15  | 0.15   |
| 7.50  | 0   | 0.90  | 0.90  | 0.9  |

TABLE B10  
Bayesian Data Information

| Source 9  |   |   |  |  |
|---|---|---|--|--|
| Time Data Base (T) : 75 years<br>Number of Recorded Events (N) : 9<br>$\nu'$ from log-linear fit : 9.29<br>$\lambda'' = \lambda' + T = 75 + 75 = 150$<br>$\nu'' = \nu' + N = 9.29 + 9 = 18.29$<br>$\eta''_{M_i} = \eta'_{M_i} + N = 9.29 + 9 = 18.29$ |   |   |  |  |
| Richter Magnitude ( $M_i$ )   | Nb of Recorded Occurrences in $M_i$ bands ( $R_{M_i}$ ) | Cumulative Nb of Occurrences (log-linear fit Fig. ) ( $N_c$ ) | Nb of Occurrences in $M_i$ bands (log-linear fit) ( $\epsilon''_{M_i}$ ) | $\epsilon'_{M_i} + R_{M_i}$ ( $\epsilon''_{M_i}$ ) |
| 4.00  | 2   | 9.29  | 1.04   | 3.04   |
| 4.25  | 1   | 8.25  | 0.93   | 1.93   |
| 4.50  | 0   | 7.32  | 0.82   | 0.82   |
| 4.75  | 0   | 6.50  | 0.72   | 0.72   |
| 5.00  | 0   | 5.78  | 0.65   | 0.65   |
| 5.25  | 1   | 5.13  | 0.72   | 1.72   |
| 5.50  | 2   | 4.41  | 0.82   | 2.82   |
| 5.75  | 0   | 3.59  | 0.67   | 0.67   |
| 6.00  | 1   | 2.92  | 0.55   | 1.55   |
| 6.25  | 1   | 2.38  | 0.45   | 1.45   |
| 6.50  | 0   | 1.93  | 0.35   | 0.35   |
| 6.75  | 0   | 1.57  | 0.29   | 0.29   |
| 7.00  | 0   | 1.28  | 0.24   | 0.24   |
| 7.25  | 0   | 1.04  | 0.19   | 0.19   |
| 7.50  | 1   | 0.85  | 0.16   | 1.16   |
| 7.75  | 0   | 0.69  | 0.13   | 0.13   |
| 8.00  | 0   | 0.56  | 0.56   | 0.56   |

TABLE B11  
Bayesian Data Information

| Source 10   |   |   |  |  |
|---|---|---|--|--|
| Time Data Base (T) : 75 years<br>Number of Recorded Events (N) : 7<br>$\nu'$ from log-linear fit : 7.58<br>$\lambda'' = \lambda' + T = 75 + 75 = 150$<br>$\nu'' = \nu' + N = 7.58 + 7 = 14.58$<br>$\eta''_{M_i} = \eta'_{M_i} + N = 7.58 + 7 = 14.58$ |   |   |  |  |
| Richter Magnitude ( $M_i$ )   | Nb of Recorded Occurrences in $M_i$ bands ( $R_{M_i}$ ) | Cumulative Nb of Occurrences (log-linear fit Fig. ) ( $N_c$ ) | Nb of Occurrences in $M_i$ bands (log-linear fit) ( $\epsilon''_{M_i}$ ) | $\epsilon'_{M_i} + R_{M_i}$ ( $\epsilon''_{M_i}$ ) |
| 4.00  | 0   | 7.58  | 0.65   | 0.65   |
| 4.25  | 1   | 6.93  | 0.60   | 1.60   |
| 4.50  | 0   | 6.33  | 0.54   | 0.54   |
| 4.75  | 1   | 5.79  | 0.50   | 1.50   |
| 5.00  | 1   | 5.29  | 0.45   | 1.45   |
| 5.25  | 0   | 4.84  | 0.82   | 0.42   |
| 5.50  | 0   | 4.42  | 1.13   | 1.13   |
| 5.75  | 2   | 3.29  | 1.08   | 3.08   |
| 6.00  | 0   | 2.21  | 0.72   | 0.72   |
| 6.25  | 1   | 1.49  | 0.49   | 1.49   |
| 6.50  | 1   | 1.00  | 0.33   | 1.33   |
| 6.75  | 0   | 0.67  | 0.22   | 0.22   |
| 7.00  | 0   | 0.45  | 0.15   | 0.15   |
| 7.25  | 0   | 0.30  | 0.10   | 0.10   |
| 7.50  | 0   | 0.20  | 0.20   | 0.20   |

TABLE B12  
Bayesian Data Information

| Source 11   |   |   |  |                                      |
|---|---|---|--|--------------------------------------|
| Time Data Base (T) : 75 years<br>Number of Recorded Events (N) : 5<br>$\nu'$ from log-linear fit : 4.43<br>$\lambda'' = \lambda' + T = 75 + 75$<br>$\nu'' = \nu' + N = 4.43 + 5 = 9.43$<br>$\eta''_{M_i} = \eta'_{M_i} + N = 4.43 + 5 = 9.43$ |   |   |  |                                      |
| Richter Magnitude ( $M_i$ )   | Nb of Recorded Occurrences in $M_i$ bands ( $R_{M_i}$ ) | Cumulative Nb of Occurrences (log-linear fit Fig. ) ( $N_c$ ) | Nb of Occurrences in $M_i$ bands (log-linear fit) ( $E'_{M_i}$ ) | $E'_{M_i} + R_{M_i}$ ( $E''_{M_i}$ ) |
| 4.00  | 3   | 4.43  |  |                                      |
| 4.25  | 0   | 3.87  | 0.56   | 3.56                                 |
| 4.50  | 0   | 3.38  | 0.49   | 0.49                                 |
| 4.75  | 0   | 2.96  | 0.42   | 0.42                                 |
| 5.00  | 0   | 2.58  | 0.38   | 0.38                                 |
| 5.25  | 0   | 2.26  | 0.32   | 0.32                                 |
| 5.50  | 0   | 2.26  | 0.29   | 0.29                                 |
| 5.75  | 0   | 1.97  | 0.25   | 0.25                                 |
| 6.00  | 0   | 1.72  | 0.21   | 0.21                                 |
| 6.25  | 1   | 1.51  | 0.19   | 1.19                                 |
| 6.50  | 1   | 1.32  | 0.17   | 1.17                                 |
| 6.75  | 0   | 1.15  | 0.15   | 0.15                                 |
| 7.00  | 0   | 1.00  | 0.12   | 0.12                                 |
| 7.25  | 0   | 0.88  | 0.11   | 0.11                                 |
| 7.50  | 0   | 0.77  | 0.10   | 0.10                                 |
| 7.75  | 0   | 0.67  | 0.09   | 0.09                                 |

TABLE B13  
Bayesian Data Information

| Source 12 - Benioff 1   |   |   |  |                                      |
|---|---|---|--|--------------------------------------|
| Time Data Base (T) : 75 years<br>Number of Recorded Events (N) : 122<br>$\nu'$ from log-linear fit : 111.21<br>$\lambda'' = \lambda' + T = 75 + 75 = 150$<br>$\nu'' = \nu' + N = 111.21 + 122 = 233.21$<br>$\eta''_{M_i} = \eta'_{M_i} + N = 111.21 + 122 = 233.21$ |   |   |  |                                      |
| Richter Magnitude ( $M_i$ )   | Nb of Recorded Occurrences in $M_i$ bands ( $R_{M_i}$ ) | Cumulative Nb of Occurrences (log-linear fit Fig. ) ( $N_c$ ) | Nb of Occurrences in $M_i$ bands (log-linear fit) ( $E'_{M_i}$ ) | $E'_{M_i} + R_{M_i}$ ( $E''_{M_i}$ ) |
| 4.00  | 41  | 111.21  | 26.96  | 67.96                                |
| 4.25  | 9   | 84.25   | 20.43  | 29.43                                |
| 4.50  | 19  | 63.82   | 15.47  | 34.47                                |
| 4.75  | 5   | 48.35   | 11.73  | 16.73                                |
| 5.00  | 15  | 36.62   | 8.87   | 23.87                                |
| 5.25  | 1   | 27.75   | 6.73   | 7.73                                 |
| 5.50  | 10  | 21.02   | 5.10   | 15.10                                |
| 5.75  | 5   | 15.92   | 3.86   | 8.86                                 |
| 6.00  | 6   | 12.06   | 2.92   | 8.92                                 |
| 6.25  | 2   | 9.14  | 2.22   | 4.22                                 |
| 6.50  | 4   | 6.92  | 1.68   | 5.68                                 |
| 6.75  | 1   | 5.24  | 1.27   | 2.27                                 |
| 7.00  | 1   | 3.97  | 0.96   | 1.96                                 |
| 7.25  | 3   | 3.01  | 0.73   | 3.73                                 |
| 7.50  | 0   | 2.28  | 0.55   | 0.55                                 |
| 7.75  | 0   | 1.73  | 0.42   | 0.42                                 |
| 8.00  | 0   | 1.31  | 0.32   | 0.32                                 |
| 8.25  | 0   | 0.99  | 0.24   | 0.24                                 |
| 8.50  | 0   | 0.75  | 0.18   | 0.18                                 |

TABLE B14  
Bayesian Data Information

| Source 13 - Benioff 2   |   |   |   |  |
|---|---|---|---|--|
| Time Data Base (T) : 75 years<br>Number of Recorded Events (N) : 45<br>$\nu'$ from log-linear fit : 40.17<br>$\lambda'' = \lambda' + T = 75 + 75 = 150$<br>$\nu'' = \nu' + N = 40.17 + 45 = 85.17$<br>$\eta''_{M_i} = \eta'_{M_i} + N = 40.17 + 45 = 85.17$ |   |   |   |  |
| Richter Magnitude ( $M_i$ )   | Nb of Recorded Occurrences in $M_i$ bands ( $R_{M_i}$ ) | Cumulative Nb of Occurrences (log-linear fit Fig. ) ( $N_c$ ) | Nb of Occurrences in $M_i$ bands (log-linear fit) ( $\epsilon'_{M_i}$ ) | $\epsilon'_{M_i} + R_{M_i}$ ( $\epsilon''_{M_i}$ ) |
| 4.00  | 17  | 40.17   | 9.10  | 26.10  |
| 4.25  | 8   | 31.07   | 10.71   | 18.71  |
| 4.50  | 4   | 20.36   | 7.02  | 11.02  |
| 4.75  | 5   | 13.34   | 4.60  | 9.60   |
| 5.00  | 6   | 8.74  | 3.02  | 9.02   |
| 5.25  | 3   | 5.72  | 1.97  | 4.97   |
| 5.50  | 0   | 3.75  | 1.29  | 1.29   |
| 5.75  | 0   | 2.46  | 0.85  | 0.85   |
| 6.00  | 1   | 1.61  | 0.56  | 1.56   |
| 6.25  | 1   | 1.05  | 0.36  | 1.36   |
| 6.50  | 0   | 0.69  | 0.24  | 0.24   |
| 6.75  | 0   | 0.45  | 0.15  | 0.15   |
| 7.00  | 0   | 0.30  | 0.11  | 0.11   |
| 7.25  | 0   | 0.19  | 0.06  | 0.06   |
| 7.50  | 0   | 0.13  | 0.05  | 0.05   |
| 7.75  | 0   | 0.08  | 0.03  | 0.03   |
| 8.00  | 0   | 0.05  | 0.02  | 0.02   |
| 8.25  | 0   | 0.03  | 0.01  | 0.01   |
| 8.50  | 0   | 0.02  | 0.02  | 0.02   |

TABLE B15  
Bayesian Data Information

| Source 14 - Benioff 3   |   |   |   |  |
|---|---|---|---|--|
| Time Data Base (T) : 75 years<br>Number of Recorded Events (N) : 93<br>$\nu'$ from log-linear fit : 84.08<br>$\lambda'' = \lambda' + T = 75 + 75 = 150$<br>$\nu'' = \nu' + N = 84.08 + 93 = 177.08$<br>$\eta''_{M_i} = \eta'_{M_i} + N = 84.08 + 93 = 177.08$ |   |   |   |  |
| Richter Magnitude ( $M_i$ )   | Nb of Recorded Occurrences in $M_i$ bands ( $R_{M_i}$ ) | Cumulative Nb of Occurrences (log-linear fit Fig. ) ( $N_c$ ) | Nb of Occurrences in $M_i$ bands (log-linear fit) ( $\epsilon'_{M_i}$ ) | $\epsilon'_{M_i} + R_{M_i}$ ( $\epsilon''_{M_i}$ ) |
| 4.00  | 22  | 84.08   | 17.68   | 39.68  |
| 4.25  | 9   | 66.40   | 13.97   | 22.97  |
| 4.50  | 18  | 52.43   | 11.03   | 29.03  |
| 4.75  | 6   | 41.40   | 8.71  | 14.71  |
| 5.00  | 14  | 32.69   | 6.88  | 20.88  |
| 5.25  | 4   | 25.81   | 5.43  | 9.43   |
| 5.50  | 4   | 20.38   | 4.28  | 8.28   |
| 5.75  | 3   | 16.10   | 3.39  | 6.39   |
| 6.00  | 4   | 12.71   | 6.02  | 10.02  |
| 6.25  | 4   | 6.69  | 3.51  | 7.51   |
| 6.50  | 2   | 3.18  | 1.67  | 3.67   |
| 6.75  | 3   | 1.51  | 0.79  | 3.79   |
| 7.00  | 0   | 0.72  | 0.38  | 0.38   |
| 7.25  | 0   | 0.34  | 0.18  | 0.18   |
| 7.50  | 0   | 0.16  | 0.08  | 0.08   |
| 7.75  | 0   | 0.08  | 0.04  | 0.04   |
| 8.00  | 0   | 0.04  | 0.02  | 0.02   |
| 8.25  | 0   | 0.02  | 0.01  | 0.01   |
| 8.50  | 0   | 0.01  | 0.01  | 0.01   |
|   |   | 150   |   |  |



TABLE B16  
Bayesian Data Information

| Source 15 - Benioff 4   |   |   |  |                                      |
|---|---|---|--|--------------------------------------|
| Time Data Base (T) : 75 years<br>Number of Recorded Events (N) : 130<br>$\nu'$ from log-linear fit : 110.86<br>$\lambda'' = \lambda' + T = 75 + 75 = 150$<br>$\nu'' = \nu' + N = 110.86 + 130 = 240.86$<br>$n''_{M_i} = n'_{M_i} + N = 110.86 + 130 = 240.86$ |   |   |  |                                      |
| Richter Magnitude ( $M_i$ )   | Nb of Recorded Occurrences in $M_i$ bands ( $R_{M_i}$ ) | Cumulative Nb of Occurrences (log-linear fit Fig. ) ( $N_c$ ) | Nb of Occurrences in $M_i$ bands (log-linear fit) ( $E'_{M_i}$ ) | $E'_{M_i} + R_{M_i}$ ( $E''_{M_i}$ ) |
| 4.00  | 35  | 110.86  | 17.00  | 52.00                                |
| 4.25  | 7   | 93.86   | 14.39  | 21.39                                |
| 4.50  | 19  | 79.47   | 12.19  | 31.19                                |
| 4.75  | 12  | 67.28   | 10.32  | 22.32                                |
| 5.00  | 11  | 56.96   | 8.74   | 19.74                                |
| 5.25  | 5   | 48.22   | 10.99  | 15.99                                |
| 5.50  | 9   | 37.23   | 8.19   | 17.19                                |
| 5.75  | 7   | 28.04   | 6.92   | 13.92                                |
| 6.00  | 8   | 21.12   | 5.22   | 13.22                                |
| 6.25  | 6   | 15.90   | 3.92   | 9.92                                 |
| 6.50  | 3   | 11.98   | 2.86   | 5.86                                 |
| 6.75  | 3   | 9.02  | 2.23   | 5.23                                 |
| 7.00  | 2   | 6.79  | 1.67   | 3.67                                 |
| 7.25  | 0   | 5.12  | 1.27   | 1.27                                 |
| 7.50  | 0   | 3.85  | 0.95   | 0.95                                 |
| 7.75  | 0   | 2.90  | 0.71   | 0.71                                 |
| 8.00  | 1   | 2.19  | 0.54   | 1.54                                 |
| 8.25  | 2   | 1.65  | 0.41   | 2.41                                 |
| 8.50  | 0   | 1.24  | 1.24   | 1.24                                 |

TABLE B17  
Bayesian Data Information

| Source 16 - Benioff 5   |   |   |  |                                      |
|---|---|---|--|--------------------------------------|
| Time Data Base (T) : 75 years<br>Number of Recorded Events (N) : 63<br>$\nu'$ from log-linear fit : 59.46<br>$\lambda'' = \lambda' + T = 75 + 75 = 150$<br>$\nu'' = \nu' + N = 59.46 + 63 = 122.46$<br>$n''_{M_i} = n'_{M_i} + N = 59.46 + 63 = 122.46$ |   |   |  |                                      |
| Richter Magnitude ( $M_i$ )   | Nb of Recorded Occurrences in $M_i$ bands ( $R_{M_i}$ ) | Cumulative Nb of Occurrences (log-linear fit Fig. ) ( $N_c$ ) | Nb of Occurrences in $M_i$ bands (log-linear fit) ( $E'_{M_i}$ ) | $E'_{M_i} + R_{M_i}$ ( $E''_{M_i}$ ) |
| 4.00  | 13  | 59.46   | 8.57   | 21.57                                |
| 4.25  | 6   | 50.89   | 7.33   | 13.33                                |
| 4.50  | 8   | 43.56   | 6.28   | 14.28                                |
| 4.75  | 4   | 37.28   | 7.17   | 11.17                                |
| 5.00  | 6   | 30.11   | 7.23   | 13.23                                |
| 5.25  | 3   | 22.88   | 5.49   | 8.49                                 |
| 5.50  | 8   | 17.39   | 4.17   | 12.17                                |
| 5.75  | 3   | 13.22   | 3.17   | 6.17                                 |
| 6.00  | 5   | 10.05   | 2.42   | 7.42                                 |
| 6.25  | 2   | 7.63  | 1.83   | 3.83                                 |
| 6.50  | 2   | 5.80  | 1.39   | 3.39                                 |
| 6.75  | 1   | 4.41  | 1.06   | 2.06                                 |
| 7.00  | 0   | 3.35  | 0.80   | 0.80                                 |
| 7.25  | 0   | 2.55  | 0.62   | 0.62                                 |
| 7.50  | 1   | 1.93  | 0.46   | 1.46                                 |
| 7.75  | 0   | 1.47  | 0.35   | 0.35                                 |
| 8.00  | 0   | 1.12  | 0.27   | 0.27                                 |
| 8.25  | 0   | 0.85  | 0.20   | 0.20                                 |
| 8.50  | 1   | 0.65  | 0.65   | 1.65                                 |

TABLE B18  
Bayesian Data Information

| Source 17 - Benioff 6   |   |   |  |  |
|---|---|---|--|--|
| Time Data Base (T) : 75 years<br>Number of Recorded Events (N) : 19<br>$v'$ from log-linear fit : 18.32<br>$\lambda'' = \lambda' + T = 75 + 75 = 150$<br>$v'' = v' + N = 18.32 + 19 = 37.32$<br>$n''_{M_i} = n'_{M_i} + N = 18.32 + 19 = 37.32$ |   |   |  |  |
| Richter Magnitude ( $M_i$ )   | Nb of Recorded Occurrences in $M_i$ bands ( $R_{M_i}$ ) | Cumulative Nb of Occurrences (log-linear fit Fig. ) ( $N_c$ ) | Nb of Occurrences in $M_i$ bands (log-linear fit) ( $\epsilon''_{M_i}$ ) | $\epsilon'_{M_i} + R_{M_i}$ ( $\epsilon''_{M_i}$ ) |
| 4.00  | 5   | 18.32   | 2.86   | 7.86   |
| 4.25  | 1   | 15.46   | 2.42   | 3.42   |
| 4.50  | 2   | 13.04   | 2.03   | 4.03   |
| 4.75  | 1   | 11.01   | 1.72   | 2.72   |
| 5.00  | 2   | 9.29  | 1.46   | 3.46   |
| 5.25  | 2   | 7.83  | 1.22   | 3.22   |
| 5.50  | 0   | 6.61  | 1.65   | 1.65   |
| 5.75  | 3   | 4.96  | 2.14   | 5.14   |
| 6.00  | 1   | 2.82  | 1.21   | 2.21   |
| 6.25  | 1   | 1.61  | 0.69   | 1.69   |
| 6.50  | 1   | 0.92  | 0.40   | 1.40   |
| 6.75  | 0   | 0.52  | 0.22   | 0.22   |
| 7.00  | 0   | 0.30  | 0.13   | 0.13   |
| 7.25  | 0   | 0.17  | 0.07   | 0.07   |
| 7.50  | 0   | 0.10  | 0.05   | 0.05   |
| 7.75  | 0   | 0.05  | 0.02   | 0.02   |
| 8.00  | 0   | 0.03  | 0.01   | 0.01   |
| 8.25  | 0   | 0.02  | 0.01   | 0.01   |
| 8.50  | 0   | 0.01  | 0.01   | 0.01   |

TABLE B19  
Bayesian Data Information

| Source 18 - Benioff 7   |   |   |  |  |
|---|---|---|--|--|
| Time Data Base (T) : 75 years<br>Number of Recorded Events (N) : 24<br>$v'$ from log-linear fit : 20.80<br>$\lambda'' = \lambda' + T = 75 + 75 = 150$<br>$v'' = v' + N = 20.80 + 24 = 44.80$<br>$n''_{M_i} = n'_{M_i} + N = 20.80 + 24 = 44.80$ |   |   |  |  |
| Richter Magnitude ( $M_i$ )   | Nb of Recorded Occurrences in $M_i$ bands ( $R_{M_i}$ ) | Cumulative Nb of Occurrences (log-linear fit Fig. ) ( $N_c$ ) | Nb of Occurrences in $M_i$ bands (log-linear fit) ( $\epsilon''_{M_i}$ ) | $\epsilon'_{M_i} + R_{M_i}$ ( $\epsilon''_{M_i}$ ) |
| 4.00  | 7   | 20.80   | 3.03   | 10.03  |
| 4.25  | 3   | 17.77   | 2.59   | 5.59   |
| 4.50  | 2   | 15.18   | 2.21   | 4.21   |
| 4.75  | 1   | 12.97   | 1.88   | 2.88   |
| 5.00  | 1   | 11.09   | 1.62   | 2.62   |
| 5.25  | 1   | 9.47  | 1.38   | 2.38   |
| 5.50  | 2   | 8.09  | 1.17   | 3.17   |
| 5.75  | 1   | 6.92  | 1.38   | 2.38   |
| 6.00  | 2   | 5.54  | 1.53   | 3.53   |
| 6.25  | 2   | 4.01  | 1.10   | 3.10   |
| 6.50  | 0   | 2.91  | 0.80   | 0.80   |
| 6.75  | 2   | 2.11  | 0.58   | 2.58   |
| 7.00  | 0   | 1.53  | 0.42   | 0.42   |
| 7.25  | 0   | 1.11  | 0.31   | 0.31   |
| 7.50  | 0   | 0.80  | 0.22   | 0.22   |
| 7.75  | 0   | 0.58  | 0.16   | 0.16   |
| 8.00  | 0   | 0.42  | 0.11   | 0.11   |
| 8.25  | 0   | 0.31  | 0.09   | 0.09   |
| 8.50  | 0   | 0.22  | 0.22   | 0.22   |

APPENDIX C

THE FEBRUARY 4, 1976 GUATEMALAN EARTHQUAKE

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## THE FEBRUARY 4, 1976 GUATEMALAN EARTHQUAKE

### C-1 Introduction:

On February 4, 1976 at 3:01 a.m. (local Guatemalan time), central Guatemala experienced a major earthquake. The surface-wave magnitude ( $M_s$ ) of this earthquake was 7.5, and the body-wave magnitude ( $M_b$ ), 5.8. The hypocentral location was  $15.32^\circ\text{N}$  latitude and  $89.08^\circ\text{W}$  longitude, with a focal depth of about five kilometers. Numerous aftershocks followed the main earthquake, the largest of which occurred on February 6, 1976, causing additional damage to and around Guatemala City. The  $M_b$  of the February 6 earthquake was about 5.8. Damage due to the earthquake resulted from 1) direct seismic shaking of structures, particularly those of poor construction such as adobe buildings, 2) fault rupture, 3) earthquake induced lateral spreading and cracking of unconsolidated deposits and 4) a variety of earthquake triggered landslides.

### C-2 Intensity:

The maximum intensity of shaking within Guatemala City and the Mixco area, based on the nature and amount of damage to man-made structures whose foundations did not fail, was IX on the Modified Mercalli Scale, with small pockets of higher intensity occurring in different zones throughout the city. MM Intensity IX is defined as follows: "severe damage to well built and ordinary masonry structures, collapse of unreinforced masonry (adobe) structures, some damage to earthquake resistant structures" (see Section C-3).

Outside of Guatemala City, Modified Mercalli Intensity VI was reached over an area of 33,000 square kilometers, commonly in close proximity to the rupture zone (see Figure C-1). MM Intensity VI is defined as "objects fall from walls and shelves, some damage to poorly constructed masonry (adobe) structures".

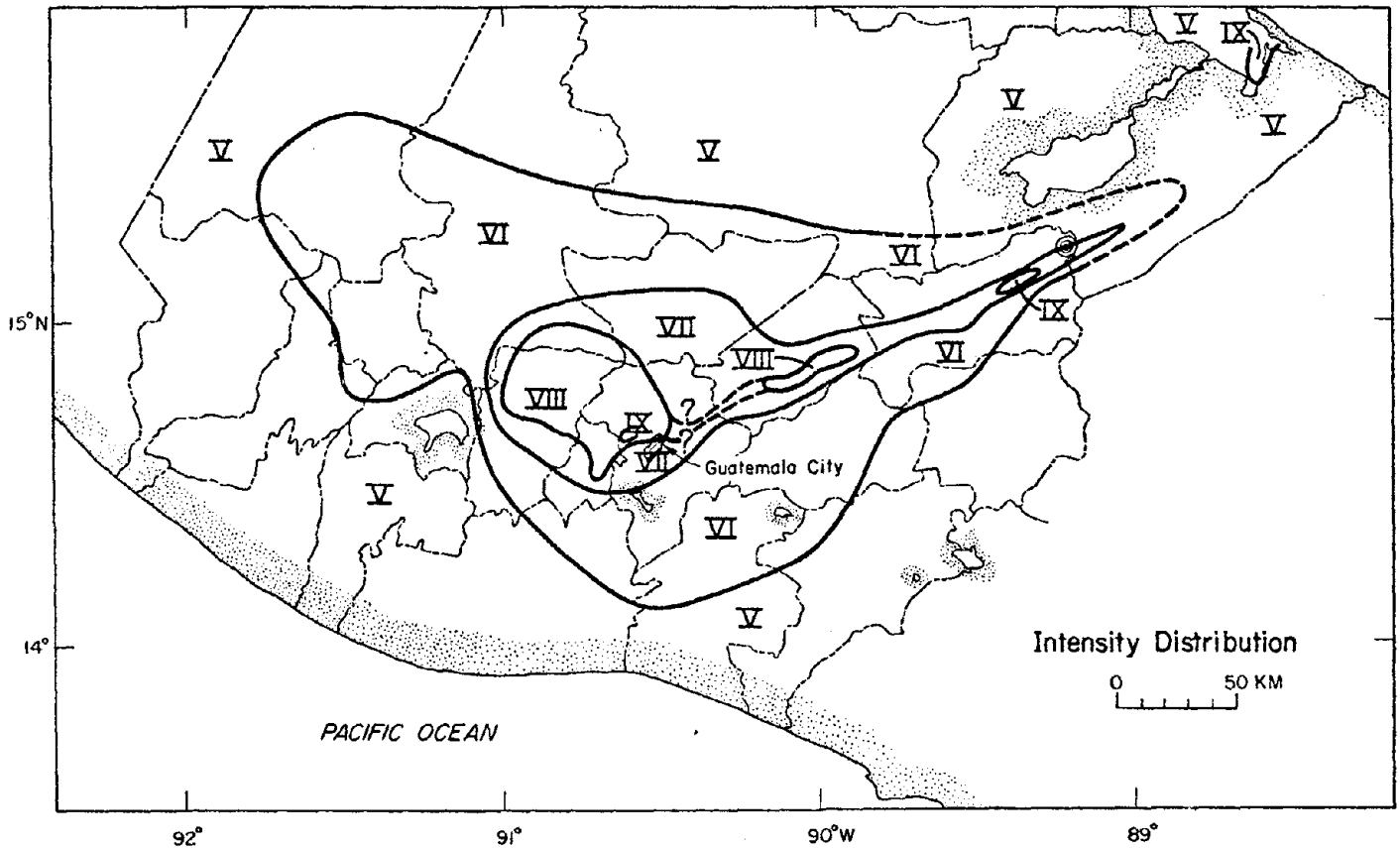


Figure C-1. Isoseismal Map showing Modified Mercalli intensity distribution in Guatemala from the February 4, 1976 earthquake.

indicates epicenter location.  
(From USGS PP1002)

C-3 Damage:

The zone of significant damage due to the earthquake and after-shocks extends about 300 kilometers in the east-west direction and 70

kilometers in the north-south direction, coinciding approximately with the region of fault rupture (see Figure C-2). The felt area for this earthquake was at least 100,000 square kilometers.

Severe damage occurred in the areas around and including Guatemala City, Joyobaj and Chimaltenango, and from El Progreso west to Gualan. A few towns up to 200 kilometers away from the instrumental epicenter were completely destroyed.

Foundation failures involving downslope slumping and/or sliding, particularly of loosely consolidated pumiceous pyroclastic rocks of the Guatemalan Highlands, lateral spreading, liquefaction and differential compaction of unconsolidated or poorly consolidated deposits damaged numerous buildings and roads throughout central Guatemala. (See Figure C-3).

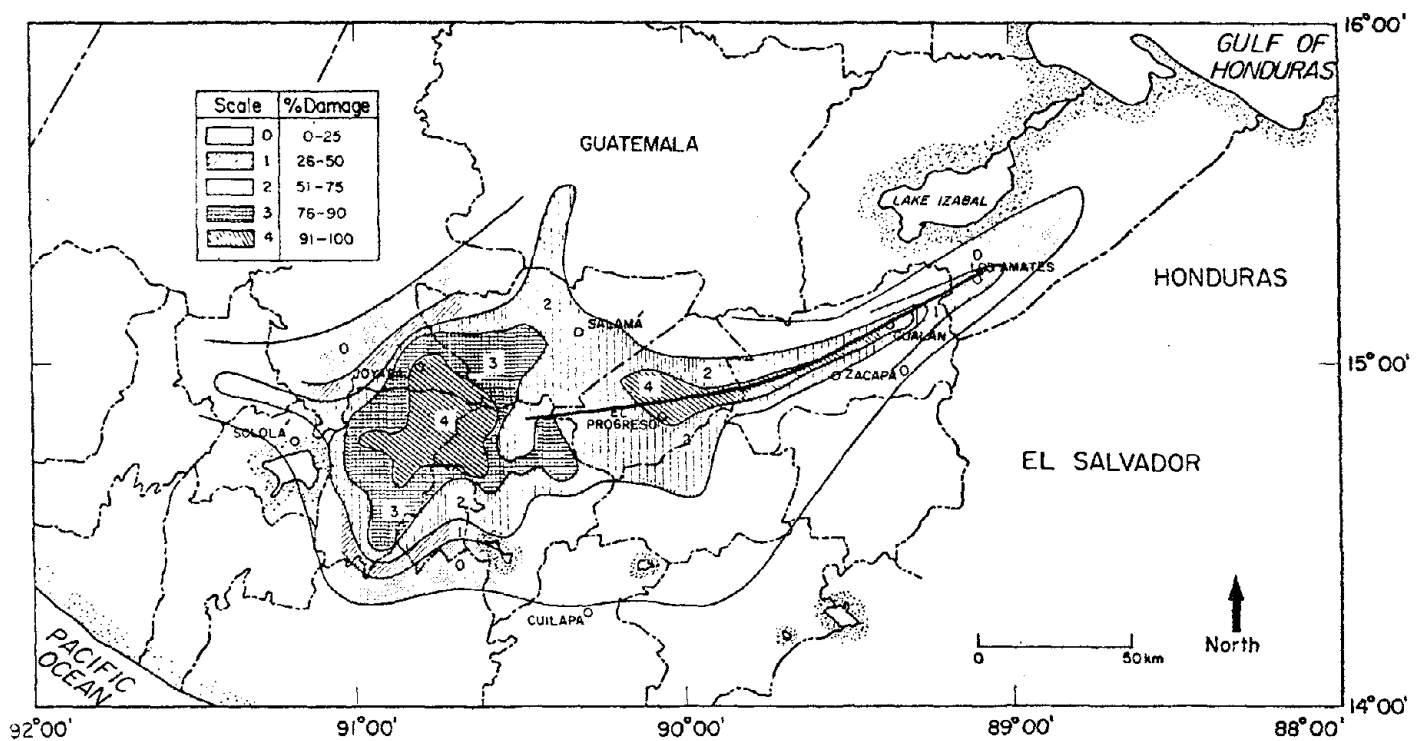


Figure C-2. Contour map showing damage to adobe-type structures in Guatemala due to the February 4, 1976 earthquake. (From USGS PP1002)

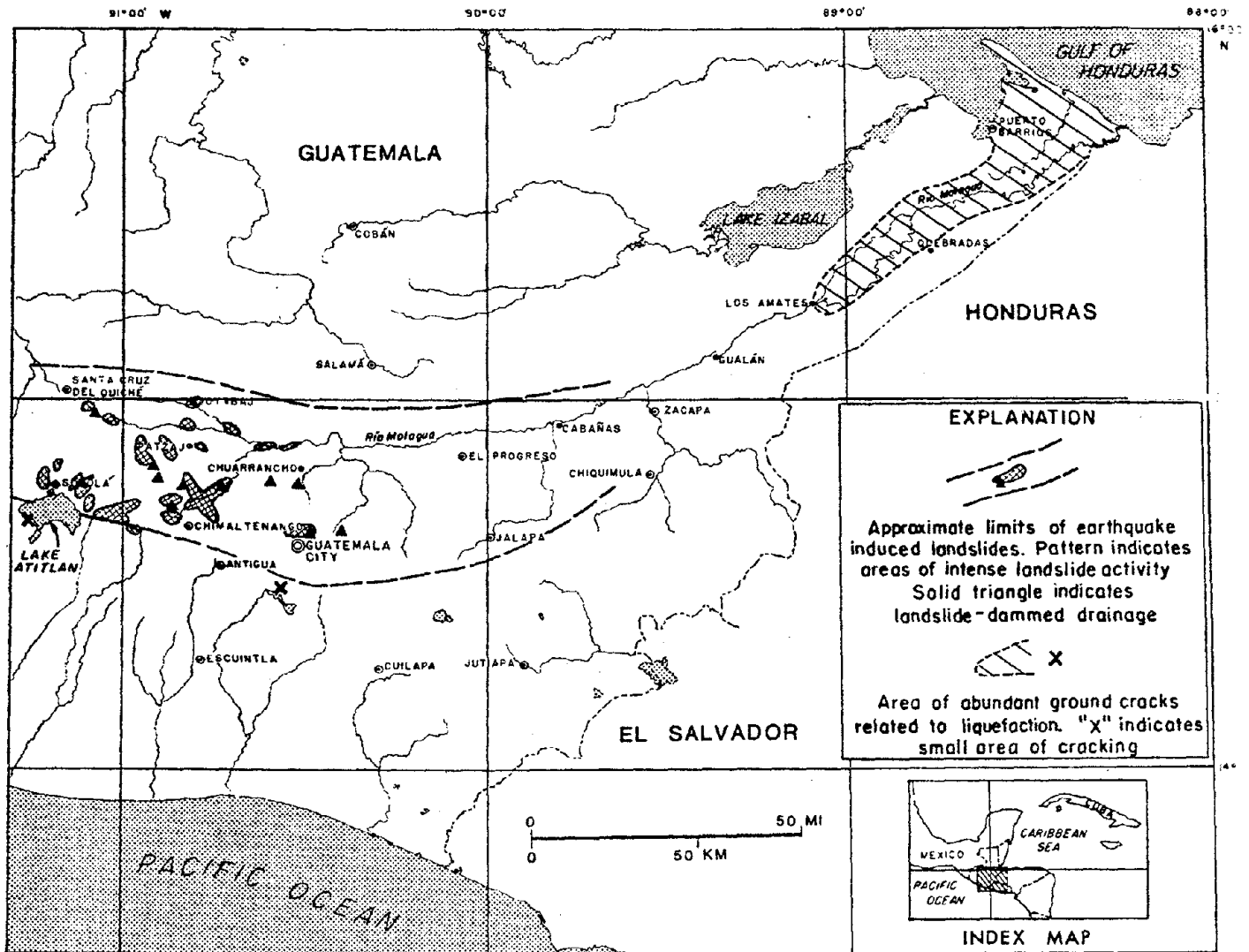


Figure C-3. Areas of earthquake-induced landsliding and of ground cracks probably related to liquefaction of unconsolidated deposits. (From USGS PP1002)

C-4 Hypocenter Location:

The hypocenter location of the main shock was determined by the National Earthquake Information Service (NEIS). The epicenter of this event was located near Los Amates, within the Motagua Valley, about 157 kilometers northeast of Guatemala City. The wide distribution of the 90 recording stations used in locating the epicenter gives reasonable confidence in the epicenter solution. However, a location bias of tens



of kilometers is possible, due to undetected seismic wave travel-time anomalies.

The depth of the hypocenter could not be reliably determined instrumentally because of the location of the seismic recording stations and the nature of the seismic waves recorded. A shallow depth of about five kilometers was assigned because of the surface faulting observed accompanying the earthquake and because of the shallow (0-12 km) depth of the aftershocks.

C-5 Nature and Amount of Fault Movement:

Movement along the Motagua fault was predominantly left lateral strike-slip, although minor vertical displacement was also observed. The length of surface faulting observed during field investigations by Plafker and others was 230 kilometers in the Motagua Valley and mountainous area west of the valley (see Figure C-4). The length of the main break, based on the occurrence of aftershocks is about 270 kilometers.

The main trace of the surface fault consists of right stepping en echelon fractures, forming a zone one to three meters wide. These fractures have a more northerly orientation than the overall trend of the fault zone.

The maximum horizontal displacement observed along the main trace of the fault was 325 centimeters, in the area between El Progreso and Chuarrancho (Espinosa, 1976). The average horizontal displacement was about 100 centimeters.

Subsidiary faults and splays of the Motagua fault appeared to be scarce. Secondary faulting (faults which underwent surface displacement approximately concurrent with that on the main fault but which at the surface do not join the main fault), however, occurred associated with

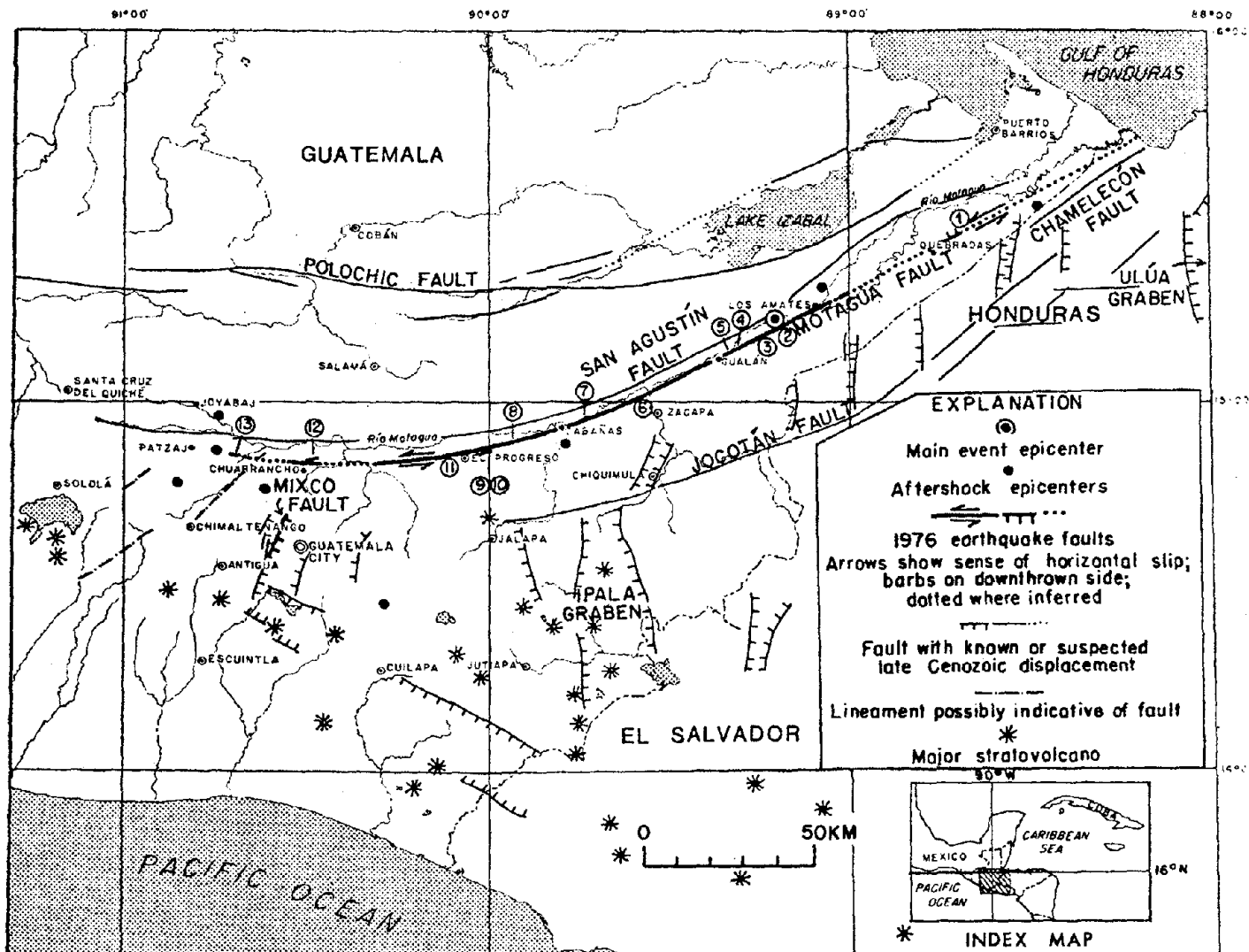


Figure C-4. Fault map of central Guatemala showing faults active during February 1976 earthquakes (heavy line) and other faults in the area (light lines). Epicenter of main event and major after shocks also shown (circles). (From USGS PP1002)

the earthquake. This faulting ruptured the ground surface in the Mixco area, west of Guatemala City, as much as 30 kilometers from the main fault trace. Ground rupture occurred along three generally north to north-north-east trending zones with predominantly vertical displacement on the order of 5 to 12 centimeters.

C-6 Cause of the Earthquake:

The February 4, 1976 earthquake occurred due to sinistral (left lateral) motions along the active plate boundary between the North American and Caribbean Plates. The North American Plate is moving westward relative to the Caribbean Plate, producing a strain accumulation along the margin of the two plates. When this situation reaches a critical point, the crust yields, producing an earthquake, as occurred on February 4, 1976. The faulting and earthquakes associated with the Mixco fault are likely due to extension produced by the plate motions and complexities related to the Cocos-Caribbean-North American triple junction.

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APPENDIX D

EARTHQUAKE DATA

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164

## Data File Indices

CGS: Coast and Geodetic Survey

CGS-B: Coast and Geodetic Survey  
Seismological Bulletin

CGSPDE: Coast and Geodetic Survey  
Preliminary Determination of Epicenters

ERL: Environmental Research Laboratories (NOAA)

NOS: National Ocean Survey (NOAA)

ISS: World Tape of Epicenters

GUTE: Gutenberg

























\*\*\*\*\*EARTHQUAKE DATA SORTED BY JAWGUES\*\*\*\*\*

EARTHQUAKE DATA SORTED BY JAWGUES

Table with columns: GUTE, S2, POLOCHIC, S3, VOLCAND, S1, MOTAGUA, SA, SOURCE, DAY, MONTH, YEAR, HOUR, MINUTE, SECOND, LONGITUDE, DEPTH, MAG, M A G N I T U D E

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Table with columns: S3, VOLCAND, S1, MOTAGUA, SA, SOURCE, DAY, MONTH, YEAR, HOUR, MINUTE, SECOND, LONGITUDE, DEPTH, MAG, M A G N I T U D E

Table with multiple columns: CGS-#, N, 37, 38, 39, 40, 41, 42, 43, 44, 45, 46, 47, 48, 49, 50, 51, 52, 53, 54, 55, 56, 57, 58, 59, 60, 61, 62, 63, 64, 65, 66, 67, 68, 69, 70, 71, 72, 73, 74, 75, 76, 77, 78, 79, 80, 81, 82, 83, 84, 85, 86, 87, 88, 89, 90, 91, 92, 93, 94, 95. Includes station identifiers like S4 - COASTLINE, S5 - COASTLINE, S6 - GUATEMALA-MEXICO, S7 - GUATEMALA - MEXICO, S8 - MIXCO.









Table with columns: Station, Frequency, Power, Time, etc. Rows include station identifiers (e.g., ERL, CGS, GTE), frequencies, power levels, and time values.





|        |   |      |     |       |       |    |    |    |    |            |     |       |     |       |     |     |
|--------|---|------|-----|-------|-------|----|----|----|----|------------|-----|-------|-----|-------|-----|-----|
| CGS 45 | 0 | 3.12 | 112 | 400MB | CGS-B | 44 | 13 | 03 | 54 | 26.8       | 14  | 300W  | 104 | 430MB | CGS | 425 |
| CGSDE  | 0 | 3.11 | 111 | 520MB | CGS-B | 66 | 15 | 08 | 54 | 44.9       | 102 | 470MB | CGS | 428   |     |     |
| GUTE   | 0 | 6.25 | 110 | 0     | ERL   | 12 | 02 | 17 | 53 | 01.1       | 105 | 440MB | CGS | 430   |     |     |
| GUTE   | 0 | 6.75 | 110 | 0     | CGSDE | 12 | 08 | 17 | 43 | 15.2       | 103 | 480MB | CGS | 432   |     |     |
| CGS 95 | 0 | 3.45 | 115 | 420MB | CGSDE | 12 | 08 | 17 | 43 | 15.2       | 106 | 440MB | CGS | 435   |     |     |
| CGS-B  | 0 | 3.13 | 114 | 440MB | CGS-B | 23 | 08 | 06 | 17 | 24.4       | 109 | 440MB | CGS | 439   |     |     |
| CGS-B  | 0 | 3.78 | 119 | 560MB | ERL   | 5  | 28 | 05 | 07 | 59.2       | 103 | 310MB | CGS | 450   |     |     |
| CGS-B  | 0 | 3.00 | 119 | 400MB | CGS   | 70 | 13 | 07 | 22 | 52.1       | 108 | 450MB | CGS | 455   |     |     |
| CGS 5  | 0 | 3.12 | 111 | 400MB | CGS   | 70 | 13 | 07 | 22 | 52.1       | 104 | 480MB | CGS | 455   |     |     |
| CGS    | 0 | 3.23 | 115 | 410MB | ERL   | 7  | 15 | 07 | 17 | 12.3       | 109 | 490MB | CGS | 455   |     |     |
| CGS    | 0 | 3.45 | 115 | 410MB | ERL   | 7  | 15 | 07 | 17 | 12.3       | 109 | 490MB | CGS | 455   |     |     |
| CGS    | 0 | 3.79 | 119 | 540MB | CGS-B | 23 | 08 | 06 | 17 | 24.4       | 109 | 480MB | CGS | 450   |     |     |
| CGS    | 0 | 3.92 | 118 | 470MB | CGS   | 31 | 04 | 04 | 09 | 07.1       | 100 | 470MB | CGS | 450   |     |     |
| CGS    | 0 | 3.29 | 118 | 470MB | GUTE  | 1  | 07 | 02 | 01 | 30.9       | 105 | 430MB | CGS | 452   |     |     |
| CGS    | 0 | 4.12 | 112 | 490MB | CGS-B | 17 | 02 | 05 | 03 | 34.1       | 104 | 390MB | CGS | 450   |     |     |
| CGS    | 0 | 3.70 | 115 | 490MB | CGSDE | 17 | 02 | 05 | 03 | 34.1       | 104 | 400MB | CGS | 450   |     |     |
| CGS    | 0 | 4.12 | 115 | 460MB | CGSDE | 17 | 02 | 05 | 03 | 34.1       | 104 | 400MB | CGS | 450   |     |     |
| CGS    | 0 | 3.92 | 118 | 450MB | CGSDE | 10 | 04 | 04 | 07 | 19.6       | 109 | 480MB | CGS | 444   |     |     |
| CGS    | 0 | 3.45 | 115 | 420MB | CGSDE | 12 | 10 | 07 | 23 | 01.0       | 108 | 410MB | CGS | 444   |     |     |
| CGS    | 0 | 3.92 | 118 | 420MB | CGSDE | 11 | 01 | 07 | 23 | 01.0       | 108 | 410MB | CGS | 444   |     |     |
| CGS    | 0 | 3.95 | 115 | 420MB | CGSDE | 21 | 01 | 07 | 23 | 01.0       | 108 | 410MB | CGS | 444   |     |     |
| CGS    | 0 | 4.61 | 111 | 470MB | CGSDE | 14 | 04 | 07 | 23 | 01.0       | 108 | 410MB | CGS | 444   |     |     |
| CGS    | 0 | 3.95 | 115 | 470MB | CGSDE | 21 | 01 | 07 | 23 | 01.0       | 108 | 410MB | CGS | 444   |     |     |
| CGS    | 0 | 4.28 | 115 | 450MB | CGSDE | 14 | 04 | 07 | 23 | 01.0       | 108 | 410MB | CGS | 444   |     |     |
| CGS    | 0 | 3.95 | 115 | 450MB | CGSDE | 30 | 01 | 09 | 15 | 35.24.4    | 100 | 480MB | CGS | 444   |     |     |
| CGS    | 0 | 4.61 | 111 | 470MB | GUTE  | 28 | 01 | 09 | 15 | 35.24.4    | 100 | 480MB | CGS | 444   |     |     |
| CGS    | 0 | 4.95 | 120 | 470MB | CGS   | 28 | 01 | 09 | 15 | 35.24.4    | 100 | 480MB | CGS | 444   |     |     |
| CGS    | 0 | 3.29 | 124 | 510MB | GUTE  | 17 | 04 | 19 | 42 | 20.0       | 109 | 500MB | CGS | 425   |     |     |
| CGS    | 0 | 4.78 | 123 | 410MB | GUTE  | 17 | 04 | 19 | 42 | 20.0       | 109 | 500MB | CGS | 425   |     |     |
| CGS    | 0 | 7.50 | 123 | 500MB | GUTE  | 06 | 08 | 42 | 22 | 33.6       | 0   | 0     | CGS | 625   |     |     |
| CGS    | 0 | 3.12 | 122 | 400MB | GUTE  | 08 | 08 | 42 | 22 | 33.6       | 0   | 0     | CGS | 625   |     |     |
| CGS    | 0 | 5.88 | 128 | 440MB | CGS   | 19 | 04 | 02 | 23 | 00.0       | 25  | 380MB | CGS | 650   |     |     |
| CGS    | 0 | 3.79 | 127 | 430MB | CGS-B | 13 | 12 | 06 | 05 | 17.29.4    | 48  | 390MB | CGS | 650   |     |     |
| CGS    | 0 | 3.62 | 128 | 400MB | CGS-B | 14 | 10 | 06 | 04 | 00.17.1    | 48  | 400MB | CGS | 650   |     |     |
| CGS    | 0 | 6.25 | 120 | 390MB | GUTE  | 14 | 10 | 06 | 04 | 00.17.1    | 48  | 400MB | CGS | 650   |     |     |
| CGS    | 0 | 3.00 | 126 | 400MB | CGS   | 24 | 04 | 35 | 18 | 51.40.0    | 50  | 0     | CGS | 625   |     |     |
| CGS    | 0 | 3.45 | 126 | 390MB | CGS   | 24 | 04 | 35 | 18 | 51.40.0    | 50  | 0     | CGS | 625   |     |     |
| CGS    | 0 | 4.95 | 124 | 430MB | CGS-B | 16 | 02 | 05 | 00 | 08.41.1    | 33  | 410MB | CGS | 625   |     |     |
| CGS    | 0 | 4.45 | 123 | 510MB | CGS-B | 23 | 09 | 02 | 20 | 18.00.0    | 25  | 0     | CGS | 625   |     |     |
| CGS    | 0 | 3.00 | 124 | 480MB | CGS-B | 03 | 05 | 06 | 12 | 45.30.3    | 53  | 430MB | CGS | 625   |     |     |
| CGS    | 0 | 3.79 | 128 | 440MB | CGS-B | 11 | 01 | 05 | 04 | 10.04.4    | 14  | 500MB | CGS | 625   |     |     |
| CGS    | 0 | 4.28 | 123 | 440MB | CGSDE | 14 | 07 | 07 | 18 | 02.08.7    | 17  | 460MB | CGS | 412   |     |     |
| CGS    | 0 | 3.45 | 123 | 420MB | CGS-B | 17 | 09 | 03 | 03 | 48.09.7    | 14  | 390MB | CGS | 300   |     |     |
| CGS    | 0 | 5.88 | 122 | 490MB | NDS   | 41 | 11 | 06 | 71 | 11.51.40.1 | 18  | 450MB | CGS | 300   |     |     |
| CGS    | 0 | 4.62 | 128 | 490MB | CGSDE | 09 | 11 | 03 | 02 | 36.55.2    | 19  | 490MB | CGS | 461   |     |     |
| CGS    | 0 | 3.45 | 128 | 430MB | CGSDE | 20 | 12 | 07 | 35 | 03.23.54.0 | 19  | 460MB | CGS | 461   |     |     |
| CGS    | 0 | 3.62 | 125 | 450MB | NDS   | 3  | 24 | 12 | 07 | 37.4       | 13  | 420MB | CGS | 441   |     |     |
| CGS    | 0 | 3.95 | 121 | 450MB | GUTE  | 12 | 01 | 05 | 29 | 29.3       | 11  | 560MB | CGS | 445   |     |     |
| CGS    | 0 | 5.60 | 130 | 470MB | ERL   | 32 | 04 | 08 | 26 | 49.5       | 10  | 0     | CGS | 525   |     |     |
| CGS    | 0 | 4.28 | 131 | 500MB | CGS   | 2  | 26 | 08 | 08 | 42.37.0    | 14  | 460MB | CGS | 495   |     |     |
| CGS    | 0 | 4.78 | 130 | 500MB | CGS   | 2  | 26 | 08 | 08 | 42.37.0    | 14  | 460MB | CGS | 495   |     |     |
| CGS    | 0 | 5.94 | 138 | 500MB | CGSDE | 21 | 01 | 06 | 08 | 13.12.4    | 15  | 480MB | CGS | 412   |     |     |
| CGS    | 0 | 3.12 | 137 | 638W  | CGSDE | 18 | 01 | 08 | 08 | 08.48.9    | 11  | 380MB | CGS | 300   |     |     |
| CGS    | 0 | 3.12 | 138 | 638W  | CGSDE | 18 | 01 | 08 | 08 | 08.48.9    | 11  | 380MB | CGS | 300   |     |     |
| CGS    | 0 | 3.12 | 137 | 638W  | CGSDE | 18 | 01 | 08 | 08 | 08.48.9    | 11  | 380MB | CGS | 300   |     |     |
| CGS    | 0 | 3.12 | 138 | 638W  | CGSDE | 18 | 01 | 08 | 08 | 08.48.9    | 11  | 380MB | CGS | 300   |     |     |



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APPENDIX E  
COMPUTER PROGRAMS

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188



C 200 CONTINUE  
 RETURN  
 END  
 \$DATA

```

***** STANFORD UNIVERSITY *****
PROGRAM SORT.DEPTH
***** THIS PROGRAM SORTS EARTHQUAKES BY DEPTH. THE BAND THICKNESS IS SET
SET TO 10 KM. *****
INPUT FORMAT *****
THE DATA IS READ FROM DISK, ONE CARD PER EVENT. (8A4,I3,I1A4)

COL VARIABLE NAME VARIABLE DESCRIPTION
1-32 XIN1 COL 1/32 READ IN A FORMAT
33-35 NDP DEPTH OF HYPOCENTER
36-79 XIN3 COL 36/79 READ IN A FORMAT

OUTPUT
*****
THE OUTPUT IS SAVED ON DISK AND CONTAINS THE DATA SORTED BY
BANDWIDTH OF 10 KM STARTING AT DEPTH 0
*****
DIMENSION XINF(21,180,8),NDEP(21,180)
*,XINF3(21,180,11),IND(21),XIN1(8),XIN3(11)
*****

DC 130 I=1,21
130 IND(I)=0
C
100 READ(5,1000,END=99)XIN1,NDP,XIN3
1000 FORMAT(8A4,I3,I1A4)
N=(NDP/10)+1
IF(N.GT.21)N=21
IND(N)=IND(N)+1
NDEP(N,NDP)=NDP
DJ 400 J=1,8
400 XI*FL(N,NDP,J)=XIN1(J)
DC 500 J=1,11
500 XINF(N,NDP,J)=XIN3(J)
GO TO 100
C
99 CONTINUE
DC 200 N=1,21
NDP=(N-1)+10
NDEP(N)
IF(NEQ,0) GO TO 210
WRITE(11,2010) NDEP,NDP
2010 FORMAT(I5,'EQ AT DEPTH',I5,' PLUS',
PRINT,NDP,' EQ AT DEPTH',NDP,' PLUS'
DC 220 I=1,NDP
220 WRITE(11,2000)(XINF(N,I,J),J=1,8),NDEP(N,I),
(XINF3(N,I,J),J=1,11)
2000 FORMAT('8A4,I3,I1A4)
220 CONTINUE
PRINT,
GO TO 200
210 PRINT,NO EQ AT DEPTH',NDP,' PLUS'

```



```

XR=XB+X(IR)
YB=YB+Y(IR)
DB=DB+D(IR)
IF(D(IR).EQ.0.0)NBZE=NBZE+1
Y2=Y2+Y(IR)*Y(IR)
X2=X2+X(IR)*X(IR)
XY=XY+X(IR)*Y(IR)
CONTINUE
XNBRC1=NBRC-NBZE
XNBRC=NBRC
YB=YB/XNBRC
DB=DB/XNBRC1
ALPHA=(XY-XNBRC*XB*YB)/(X2-XNBRC*XB*XB)
PRINT,'X,Y,SLOPE',XB,YB,ALPHA,ALPH1,DB
PRINT,'NUMBER OF NO DEPTH',NRZE
DO 130 IR=1,NBRC
X(IR)=X(IR)-XB
Y(IR)=Y(IR)-YB
XT=XT+X(IR)*X(IR)
YT=YT+Y(IR)*Y(IR)
XYT=XYT+X(IR)*Y(IR)
CONTINUE
ALPH=XYT/XT
ALPH1=XYT/YT
PRINT,ALPH,ALPH1
100 CONTINUE
RETURN
END
$DATA

```

```

***** STANFORD UNIVERSITY *****
PROGRAM LOCATE LINE *****
***** THIS PROGRAM FITS A LINE THROUGH A SET OF POINTS IN A PLANE USING *****
***** REGRESSION ANALYSIS. IT IS USED TO LOCATE A SEISMIC LINE SOURCE *****
***** GIVEN A NUMBER OF EPICENTERS. *****
***** INPUT FORMAT *****
*****
C COL VARIABLE NAME VARIABLE DESCRIPTION
C 1- 5 NBST 1 CARD (115) NB OF SOURCES TO BE LOCATED
C 2- SOURCE IDENTIFICATION 1 CARD (15,18A4)
C 1- 5 NBRC NB OF RECORDS IN THE SOURCE
C 6-77 HEDI IDENTIFICATION OF THE SOURCE
C 3- RECORD CARDS NBRC CARDS (4F10.0)
C 1-10 Y( ) LATITUDE OF THE EPICENTER
C 11-20 X( ) LONGITUDE OF THE EPICENTER
C 21-30 D( ) DEPTH OF THE HYPOCENTER
C 31-40 XM( ) RICHTER MAGNITUDE
C REPEAT CARDS 2/3 NRST TIMES FOR THE NBST SOURCES
C OUTPUT
C *****
C THE OUTPUT DISPLAYS THE SOURCE IDENTIFICATION, THE NUMBER OF
C RECORDS FOR IN THE SOURCE, THE COORDINATES OF THE CENTROID
C OF THE DATA AS WELL AS THE SLOPE AND THE DEPTH OF THE LINE.
C *****
C IMPLICIT REAL*8 (A-H,C-Z)
C DIMENSION HEDI(18),X(200),Y(200),D(200),XM(200)
C REAL*4 HEDI
C *****
C READ(5,1000)NBST
C DO 100 IS=1,NBST
C XR=0.00
C YB=0.00
C DB=0.00
C Y2=0.00
C X2=0.00
C XY=0.00
C XT=0.00
C YT=0.00
C XYT=0.00
C READ(5,1000)NBRC,HEDI
C FORMAT(15,18A4)
C WRITE(6,2000)NBRC,HEDI
C FORMAT('NUMBER OF RECORDS',16,3X,18A4)
C NBZE=0
C DO 110 IR=1,NBRC
C READ(5,1010)Y(IP),X(IP),D(IP),XM(IP)
C 1010 FORMAT(19X,F5.3,3X,F5.3,14X,F3.0,25X,F3.2)
C X(IP)=-X(IP)

```

\$DATA

```
C *****
C PROGRAM LINE,INTEP
C *****
C STANFORD UNIVERSITY
C *****
C THIS PROGRAM COMPUTES THE INTERSECTION OF TWO LINES IN THE SAME
C PLANE. IT IS USED TO OBTAIN THE INTERSECTION OF TWO LINE SOURCES
C AT THE SAME DEPTH.
C *****
C INPUT FORMAT
C *****
C
C COL VARIABLE NAME VARIABLE DESCRIPTION
C
C 1- IDENTIFICATION I CARD (20A4) IDENTIFICATION OF THE FIRST LINE
C 1-80 HED1
C
C 2- LINE PARAMETERS I CARD (3F10.0)
C 1-10 X1 X COORDINATE OF A POINT OF THE LINE
C 11-20 Y1 Y COORDINATE OF THE SAME POINT
C 21-30 AL1 SLOPE OF THE LINE
C
C 3- IDENTIFICATION I CARD (20A4) IDENTIFICATION OF THE SECOND LINE
C 1-80 HED2
C
C 2- LINE PARAMETERS I CARD (3F10.0)
C 1-10 X2 X COORDINATE OF A POINT OF THE LINE
C 11-20 Y2 Y COORDINATE OF THE SAME POINT
C 21-30 AL2 SLOPE OF THE LINE
C
C OUTPUT
C *****
C THE OUTPUT DISPLAYS THE IDENTIFICATION OF THE LINES AND
C THEIR THE COORDINATES OF THEIR INTERSECTION POINT.
C *****
C DIMENSION HED1(20),HED2(20)
C *****
C *****
C 100 READ(5,1000,END=99)HED1
C READ(5,1001)X1,Y1,AL1
C DO 200 I=1,2
C READ(5,1000)HED2
C READ(5,1001)X2,Y2,AL2
C 1000 FORMAT(20A4)
C 1001 FORMAT(8F10.0)
C
C X=(AL1*X1-AL2*X2-Y1+Y2)/(AL1-AL2)
C IF(Y2.EQ.0.) X=X2
C Y=AL1*(X-X1)+Y1
C
C WRITE(6,2000)HED1,HED2
C 2000 FORMAT('20A4)
C PRINT,'LINE1',X1,Y1,AL1
C PRINT,'LINE2',X2,Y2,AL2
C PRINT,'INTERSECTION',X,Y
C
C 200 CONTINUE
C GO TO 100
C 99 RETURN
C END
```

```

C ***** STANFORD UNIVERSITY *****
C PROGRAM LINE LENGTH
C *****
C THIS PROGRAM COMPUTES THE LENGTH OF A SEGMENT OF LINE BETWEEN
C TWO HORIZONTAL COORDINATES.
C *****
C INPUT FORMAT
C *****
C CCL VARIABLE NAME VARIABLE DESCRIPTION
C 1-- IDENTIFICATION 1 CARD (20A4)
C 1-80 HED1 LINE IDENTIFICATION
C 2-- LINE PARAMETERS 1 CARD (5F10.0)
C 1-10 X1 X COORDINATE OF A POINT OF THE LINE
C 11-20 Y1 Y COORDINATE OF THE SAME POINT
C 21-30 ALPH SLOPE OF THE LINE
C 31-40 XS X COORD OF ORIGIN OF SEGMENT
C 41-50 YX X COORD OF END OF SEGMENT
C OUTPUT
C *****
C THE OUTPUT DISPLAYS THE LINE IDENTIFICATION AND PARAMETERS AS
C AS WELL AS THE LENGTH OF THE SEGMENT.
C *****
C DIMENSION HED1(20)
C *****
100 READ(5,1001,END=99)HED1
1001 FORMAT(20A4)
1000 READ(5,1000)X1,Y1,ALPH,XS,XE
1000 FORMAT(8F10.0)
YS=ALPH*(XS-X1)+Y1
YE=ALPH*(XE-X1)+Y1
XLEN=SQRT((XS-XE)**2+(YS-YE)**2)
2000 WRITE(6,2000)HED1
FCRMT(10,20A4)
PRINT,'DATA',X1,Y1,ALPH
GC TO 100
99 RETURN
END
$DATA

```

```

*****
REGRESSION ANALYSIS
STANFORD UNIVERSITY
*****
THIS PROGRAM IS TO BE USED FOR REGRESSION ANALYSIS BETWEEN
ONE DEPENDENT VARIABLE AND ONE INDEPENDENT VARIABLE.
THE LINAP SCALE IS USED FOR THE INDEPENDENT VARIABLE (X) AND THE
LN SCALE FOR THE DEPENDENT VARIABLE (Y).
BOTH VARIABLES ARE INPUT ON THE LINEAR SCALE
IF THE DEPENDENT VARIABLE IS NOT DETERMINED, THE PROGRAM WILL
COMPUTE IT AS (1-CDF) USING A INCREMENT RMIC OF THE DEPENDENT VARIABLE
IT CAN EITHER FIT ONE OR TWO LINES ON THE DATA WITH A BREAKING
POINT INPUT AS RMBK
INPUT FORMAT
*****
COL VARIABLE VARIABLE DESCRIPTION
1--IDENTIFICATION CARD 1 CARD (5F10.0,2I5,5A4)
1-10 RMMN MINIMUM RM VALUE
11-20 RMIC RM INCREMENT TO BE USED TO COMPUTE CDF
21-30 RMBK BREAK OFF MAGNITUDE
IF (RMBK.EQ.0.) ONLY ONE LINE WILL BE FIT
31-40 A AREA
41-50 T TIME
51-55 SKIPCD A AND T ARE USED TO NORMALIZE ALPHA
RMMN,RMIC ARE INPUT AS ZEROS
= 0 ONLY X IS INPUT (AS RM), Y IS COMPUTED FROM CDF
RMMN,RMIC ARE TO BE INPUT
= 0 INTERVALS WITH NO EVENT WILL NOT BE
CONSIDERED IN REGRESSION ANALYSIS
= 1 ALL INTERVALS WILL BE CONSIDERED
61-80 HED TITLE
2-- X AND Y INPUT CARDS NBRC CARDS (ZF10.0)
INPUT IF SKIPCD=1
1-10 X INDEPENDENT VARIABLE
11-20 Y DEPENDENT VARIABLE
3-- RM INPUT CARDS NBRC CARDS (50X,F3.1)
SET 1002 FORMAT AS A FUNCTION OF DATA
INPUT IF SKIPCD=0
1-10 RM INDEPENDENT VARIABLE
4--END OF FILE CARD (ONLY IF RM IS INPUT) ONE CARD (50X,F3.1)
51-53 RM VALUE GREATER THAN 9.
NOTE
IF RM IS INPUT, SEVERAL FILES CAN BE RUN AT ONE TIME BY REPEATING
THE INPUT SEQUENCE DESCRIBED ABOVE.
OUTPUT
*****
THE OUTPUT DISPLAYS THE INPUT, THE INTERVAL FREQUENCY, THE
FREQUENCY OF OCCURRENCE ABOVE A GIVEN MAGNITUDE, THE PERCENTAGE
COEFFICIENTS ALPHA AND BETA AND ALPHA NORMALIZED WITH RESPECT
TO TIME AND AREA OR LENGTH.
*****
DIMENSION RM(2500),X(50),Y(50),YY(50),IHIS(50),HED(15),
1 ALPHA(2),BETA(2)
INTEGER SKIPCD,SKIPDZ
*****
111 READ(5,1000,END=99)RMMN,RMIC,RMBK,A,T,SKIPCD,SKIPDZ,HED
1000 FORMAT(5F10.0,2I5,5A4)
2000 FORMAT(1,IPREGRESSION ANALYSIS',/OLINEAR-LN SCALE',/O*7A4//)
IF(SKIPCD.EQ.0) GO TO 200
C NBRC=1
C 100 READ(5,1001,END=104)X(NBRC),YY(NBRC)
NBRC=NBRC+1
GO TO 100
C 104 NBRC=NBRC-1
1001 FORMAT(2F10.0)
2002 FORMAT('NUMBER OF RECORDS',IS/'OUTPUT VARIABLES',/
*'INDEPENDENT DEPENDENT')
DO 150 IX=1,NBRC
150 WRITE(6,2011)X(IX),YY(IX)
2011 FORMAT(' ',ZF11.2)
GO TO 201
200 CONTINUE
C COMPUTE HISTOGRAM
C CAUTION FORMAT SET TO READ DATA FROM TAPE U599
C NBRC=0
C RMLW=10.
C 112 NBRC=NBRC+1
109 READ(5,1002)RM(NBRC)
1002 FORMAT(74X,F3.2)
RMIN(NBRC)=FLOAT((FIX(RMIN(NBRC)*10.+5011))/10.
IF(RMIN(NBRC).GE.9.) GO TO 105
IF(RMIN(NBRC).LT.RMMN) GO TO 110
IF(RMIN(NBRC).LT.RMLW) RMLW=RM(NBRC)
GO TO 112
110 WRITE(6,2009) NBRC,RM(NBRC)
2009 FORMAT(' RECORD',I5,' HAS MAGNITUDE',F10.2,' IT IS DISREGARDED')
GO TO 109
C
C 105 NBRC=NBRC-1
RCK1=RMMN
WRITE(6,2001)NBRC,A,T
2001 FORMAT('NUMBER OF RECORDS ',I5,'0AREA',I0X,F10.2,/
*'OTIME (YEARS)',F11.2)
IF(NBRC.LT.2) GO TO 111

```

```

*****
REGRESSION ANALYSIS
STANFORD UNIVERSITY
*****
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IT CAN EITHER FIT ONE OR TWO LINES ON THE DATA WITH A BREAKING
POINT INPUT AS RMBK
INPUT FORMAT
*****
COL VARIABLE VARIABLE DESCRIPTION
1--IDENTIFICATION CARD 1 CARD (5F10.0,2I5,5A4)
1-10 RMMN MINIMUM RM VALUE
11-20 RMIC RM INCREMENT TO BE USED TO COMPUTE CDF
21-30 RMBK BREAK OFF MAGNITUDE
IF (RMBK.EQ.0.) ONLY ONE LINE WILL BE FIT
31-40 A AREA
41-50 T TIME
51-55 SKIPCD A AND T ARE USED TO NORMALIZE ALPHA
RMMN,RMIC ARE INPUT AS ZEROS
= 0 ONLY X IS INPUT (AS RM), Y IS COMPUTED FROM CDF
RMMN,RMIC ARE TO BE INPUT
= 0 INTERVALS WITH NO EVENT WILL NOT BE
CONSIDERED IN REGRESSION ANALYSIS
= 1 ALL INTERVALS WILL BE CONSIDERED
61-80 HED TITLE
2-- X AND Y INPUT CARDS NBRC CARDS (ZF10.0)
INPUT IF SKIPCD=1
1-10 X INDEPENDENT VARIABLE
11-20 Y DEPENDENT VARIABLE
3-- RM INPUT CARDS NBRC CARDS (50X,F3.1)
SET 1002 FORMAT AS A FUNCTION OF DATA
INPUT IF SKIPCD=0
1-10 RM INDEPENDENT VARIABLE
4--END OF FILE CARD (ONLY IF RM IS INPUT) ONE CARD (50X,F3.1)
51-53 RM VALUE GREATER THAN 9.
NOTE
IF RM IS INPUT, SEVERAL FILES CAN BE RUN AT ONE TIME BY REPEATING
THE INPUT SEQUENCE DESCRIBED ABOVE.
OUTPUT

```

```

WRITE(6,2112)RMNI,RMIC
FORMAT('MINIMUM MAGNITUDE',I2X,F10.2 /
* MAGNITUDE INCREMENT FOR CDF ',F10.2)
DO 120 I=1,50
RMCK=RMCK1+RMIC
IF(RMLV.LT.(RMCK-.05)) GO TO 121
RMCK1=RMCK
120 WRITE(6,2220) RMNI, RMCK
2220 FORMAT(' THERE IS NO RECORD BETWEEN MAGNITUDE',F8.2,' AND',F8.2)
C
121 CONTINUE
WRITE(6,2004)
FORMAT('DEARTHQUAKE MAGNITUDES')
2004 WRITE(6,2003)(RM(I),IX=1,NBRC)
2003 FORMAT(' ',I3F10.2)
RMIC2=RMIC+.5
X(I)=RMCK1+RMIC2
NBIC=0
DO 102 I=1,49
X(I+1)=X(I)+RMIC
102 THIS(I)=0
C
RMNI=RMCK1-.1
DO 103 IX=1,NBRC
NBKC=(IPM(IX)-RMNI)/RMIC)+.999
THIS(NBKC)=THIS(NBKC)+1
IF(NBKC.GT.NBIC) NBIC=NBKC
103 CONTINUE
C
IF(SKIPD.EQ.1.OR.SKIPC.D.EQ.1) GO TO 133
C
I=0
131 I=I+1
134 IF(I.GE.NBIC) GO TO 133
IF(THIS(I).NE.0) GO TO 131
NBIC=NBIC-I
DO 132 K=1,NBIC
X(K)=X(K+1)
132 THIS(K)=THIS(K+1)
NBIC=NBIC-I
GO TO 134
C
133 CONTINUE
IF(NBIC.LT.2) GO TO 111
C
COMPUTE CDF
YY(NBIC)=THIS(NBIC)
NBIC1=NBIC+1
DO 130 IX=2,NBIC
KX=NBIC1-IX
YY(KX)=YY(KX+1)+FLOAT(THIS(KX))
130 YY(KX)=YY(KX+1)+FLOAT(THIS(KX))
2005 FORMAT('O',IX,FM,10X,INTERVAL',10X,CUMULATIVE FREQUENCY', /
* INTERVAL',5X,FREQUENCY',10X,OCCURRENCES ABOVE RM', /)
DO 160 IX=1,NBIC
RM1=X(IX)+RMIC2
RM2=X(IX)+RMIC2-.01
160 WRITE(6,2006)RM1,RM2,THIS(IX),YY(IX),PHI
2006 FORMAT('F7.2,' - ',F4.2,I8,I2X,F10.0,8X,F6.2)
C
201 DO 202 IX=1,NBIC
202 Y(IX)=ALOG(YY(IX))
CHECK WHETHER THERE WILL BE ONE OR TWO REGRESSION LINES
C
IR=1
NBPT=NBIC
IF(RMBK.EQ.0.) GO TO 300
C
NRPT=0.
DO 320 IX=1,NBIC
IF(X(IX).GT.RMBK) GO TO 340
320 NRPT=NRPT+1
C
340 NRPT1=NBIC-NBPT+1
WRITE(6,2110)RMBK,NBPT,NBPT1
2110 FORMAT('TWO STRAIGHT LINES WILL BE USED TO FIT THE DATA', /
1. BREAK POINT MAGNITUDE',F8.2,' ',15, ' POINTS IN THE FIRST LINE', /
2. ', ',15, ' POINTS IN THE SECOND LINE', )
GO TO 300
C
330 NBPT1=NBPT-I
NBPT=NBIC-NBPT1
RMBK=0.
IR=IR+1
DO 350 IX=1,NBPT
X(IX)=X(IX+NBPT1)
350 Y(IX)=Y(IX+NBPT1)
C
300 WRITE(6,2130)IR
2130 FORMAT('OINTERCEPT AND SLOPE OF LINE',I5)
C
COMPUTE ALPHA INTERCEPT
C
COMPUTE BETA SLOPE
C
COMPUTE MEAN
SMX=0.
SMY=0.
SMX=0.
SMY=0.
PRINT,'NBPT',NBPT
WRITE(6,3333)(Y(I),I=1,NBPT)
WRITE(6,3333)(X(I),I=1,NBPT)
3333 FORMAT(' ',I3F10.5)
DO 140 IX=1,NBPT
SMX=SMX+X(IX)*X(IX)
SMY=SMY+X(IX)*Y(IX)
SMX=SMX+X(IX)
SMY=SMY+Y(IX)
140 SMY=SMY+Y(IX)
XNRPT=NRPT
YN=SMY/XNRPT
XM=SMX/XNRPT
BETA(IP)=(SMY-XNRPT*YM*XM)/(SMX-XNRPT*XM*XM)
ALPHA(IP)=(YM-BETA(IP)*XM)
WRITE(6,2007)ALPHA(IP),BETA(IP)
2007 FORMAT('OALPHA',F10.6/'OBETA ',F10.6)
C
IF(A.FO.0.OR.T.EQ.0.) GO TO 360
ALPHA=ALPHA(IP)-ALOG(A*Y)
WRITE(6,2008)ALPHA
2008 FORMAT('ONORMALIZED ALPHA',F10.6 //)

```

```

C 360 CONTINUE
C IF(RMK.NF.0.) GO TO 330
C IF(IR.E0.1) GO TO 111
C COMPUTE INTERSECTION POINT IF THERE ARE TWO LINES
C XM=ABS((ALPHA(1)-ALPHA(2))/(BETA(1)-BETA(2)))
C XLN=ALPHA(1)+BETA(1)*XM
C WRITE(6,2120)XM,XLN
C 2120 FORMAT('INTERSECTION POINT:', MAGNITUDE',F8.2,/, LN OF N ',F8.2)
C GO TO 111
C 99 RETURN
C END
C $DATA

```

```

C *****
C PROGRAM ACC.-LINE AREA ***** STANFORD UNIVERSITY *****
C *****
C THIS PROGRAM COMPUTES PROBABILITY OF EXCEEDANCE OF GIVEN PEAK
C SEISMIC ACCELERATIONS AT SPECIFIED SITES DUE TO LINE OR AREA
C SEISMIC SOURCES.
C INPUT FORMAT *****
C *****
C 1.- IDENTIFICATION CARD 1 CARD 20A4
C
C 1-80 HED1 TITLE
C
C 2.- ATTENUATION CONSTANTS 1 CARD 3F10.0
C
C 1-10 R1
C 11-20 B2
C 21-30 B3
C 31-40 B4
C 41-50 DELTA STEP SIZE FOR LINE INTEGRATION
C 51-60 DELTAC STEP SIZE FOR CIRCLE INTEGRATION
C
C ATTENUATION FORMULA OF THE TYPE
C ACC=B1*EXP(B2*MAG)/(R+B4)**B3
C
C 3.- PROBLEM DESCRIPTION 4 CARDS
C
C CARD 1 415 NUMBER OF LINE SOURCES
C 1-5 NL NUMBER OF AREA SOURCES
C 6-10 NA NUMBER OF TIME PERIODS
C 11-15 NT NUMBER OF ACC
C 16-20 NY
C
C CARD 2 TIME PERIODS, NT VALUES, 8 VALUES PER CARD 8F10.0
C 1-10 T(1) PERIOD 1
C 11-20 T(2) PERIOD 2
C
C ..... T(NT) PERIODS NT
C
C CARD 3 PEAK GROUND ACC., NY VALUES, 8 VALUES PER CARD 8F10.0
C 1-10 Y(1) PGA 1
C 11-20 Y(2) PGA 2
C
C ..... Y(NY) PGA NY
C
C CARD 4 SEARCHING GRID DESCRIPTION 8F10.0
C 1-10 XBEGIN ORIGIN OF GRID X COORD
C 11-20 YBEGIN ORIGIN OF GRID Y COORD
C 21-30 XEND END OF GRID X COORD
C 31-40 YEND END OF GRID Y COORD
C
C ORIGIN AND END SHOULD COINCIDE IF ONLY
C ONE LOCATION IS REQUIRED
C
C 41-50 DX X INCREMENT
C 51-60 DY Y INCREMENT
C
C 4.- LINE SOURCES PROPERTIES 3 CARDS PER SOURCE

```

```

C ***** IDENTIFICATION OF THE LINE SOURCE *****
C
C CARD 1 HED2
C 1-80
C
C CARD 2 7F10.0
C
C PROPERTIES ON THE LEFT OF BREAKING POINT
C ALPHA( ) NORMALIZED INTERCEPT
C RETAI( ) SLOPE
C XL1( ) X COORD FOR ORIGIN
C XL2( ) X COORD OF END
C YL1( ) Y COORD OF ORIGIN
C YL2( ) Y COORD OF END
C ML( ) DEPTH OF LINE
C
C CARD 3 3F10.0
C
C PROPERTIES ON THE RIGHT OF BREAKING POINT
C INPUT THE SAME VALUES FOR ALPHA AND BETA
C AND DUMMY BREAKING POINT IF THERE IS NO BREAKING POINT
C ALPHA2( ) NORMALIZED INTERCEPT
C BETA2( ) SLOPE
C RML( ) BREAKING POINT MAGNITUDE
C
C 5.- AREA SOURCES PROPERTIES 3 CARDS PER SOURCE
C
C CARD 1 20A4
C 1-80 HED AREA SOURCE IDENTIFICATION
C
C CARD 2 6F10.0
C
C PROPERTIES ON THE LEFT OF BREAKING POINT
C ALPHA( ) NORMALIZED INTERCEPT
C RETAI( ) SLOPE
C XL1( ) X COORD OF CIRCLE CENTER
C YL1( ) Y COORD OF CIRCLE CENTER
C XL2( ) RADIUS OF CIRCLE
C ML( ) DEPTH OF AREA SOURCE
C
C CARD 3 3F10.0
C
C PROPERTIES ON THE RIGHT OF BREAKING POINT
C INPUT SAME VALUES FOR ALPHA AND BETA AND
C DUMMY BREAKING POINT IF THERE IS NO BREAKING POINT
C ALPHA2( ) NORMALIZED INTERCEPT
C BETA2( ) SLOPE
C RML( ) BREAKING POINT MAGNITUDE
C
C OUTPUT *****
C
C DEPEND UPON THE INPUT, THE OUTPUT WILL DISPLAY FROM ONE
C PROBABILITY OF EXCEEDANCE AT ONE POINT TO A WHOLE CDF FOR A
C WHOLE GRID OF POINTS
C
C *****
C IMPLICIT REAL*8 (A-H,O-Z) *****
C DIMENSION HED1(20),HED2(20) *****
C COMMON/SOURCE/ALPHA(20),ALPHA2(20),RETAI(20),RETA2(20),RML(20),
C ML(20) *****
C COMMON/ATTEN/R1,R2,B3,B4,DELTA,DELTAC *****
C DIMENSION XL1(20),YL1(20),XL2(20),YL2(20),T(10),Y(30),PDF(30),
C FY1(30),FY2(30) *****
C *****
C READ AND WRITE ATTENUATION CONSTANTS *****

```

```

C
  READ(5,1040)HED1
1043 FORMAT(20A4)
  WRITE(6,1041)HED1
1041 FORMAT('0',20A4)
  READ(5,1000)R1,R2,R3,R4,DELTA,DELTA2
  WRITE(6,1020)R1,R2,R3,R4,DELTA,DELTA2
1020 FORMAT('0',ATTENUATION CONSTANTS,/,/,81=,E15.7,3X,82=,E15.7,
  3X,83=,E15.7,3X,84=,E15.7 / DELTA =,E15.7,3X,DELTA2 =,
  1 E15.7 /)
1000 FORMAT(REFIO-0)
C
  READ AND WRITE GEOMETRIC CONSTANTS
  READ(5,1001)NL,NA,NT,NY
1001 FORMAT(10I5)
  READ(5,1000)(T(I),I=1,NT)
  WRITE(6,2210)
  WRITE(6,2220)(T(I),I=1,NT)
  NY=NY+1
  READ(5,1000)(Y(I),I=2,NY)
  WRITE(6,2230)
  WRITE(6,2220)(Y(I),I=2,NY)
2220 FORMAT(' ',10F10.2)
2230 FORMAT('0ACCELERATIONS')
  YG(1)=0.000
2222 FORMAT('0',SITE LOCATION,/,/,X=,F10.3,5X,Y=,F10.3)
350 READ(5,1000) XBEGIN,YBEGIN,XEND,YEND,DX,DY
  NYMAX=DABS((XBEGIN-XEND)/DX)*1.900
  NYMAX=DABS((YBEGIN-YEND)/DY)*1.900
  WRITE(6,1021)NL,NA,NYMAX,NT
1021 FORMAT('0',GEOMETRIC CONSTANTS,/,/,NL=,15,5X,NA = ,15,5X,
&NYMAX= ,15,5X,NYMAX=,15,5X,NT=,15 /)
  IF(NL.LE.0) GO TO 610
C
  READ AND WRITE LINE SOURCE PROPERTIES
  WRITE(6,1024)
1024 FORMAT('0LINE SOURCES,/, ***** /)
  DO 616 I=1,NL
  READ(5,1040)HED2
  WRITE(6,1041)HED2
  WRITE(6,1020)T(I)
  WRITE(6,1020)Y(I)
  WRITE(6,1020)X(I)
  WRITE(6,2222)X,Y
  PDF(I)=0.000
  WRITE(6,1022)
  WRITE(6,1023)ALPHA(I),BETA(I),XLI(I),YL2(I),HL(I)
  EX,XLI,11X,XL2,11X,YLI,11X,YL2,12X,HL)
1023 FORMAT(' ',E13.5,6F14.5)
  WRITE(6,1031)
  WRITE(6,1023)ALPHA2(I),BETA2(I),RML(I)
1031 FORMAT('0 SECOND REGRESSION CONSTANTS,/,/,2X,ALPHA2,9X,
&L2,10X,MR)
  616 CONTINUE
  DO 700 IL=1,NL
  IF(XLI(IL).LT.XL2(IL)) GO TO 700
  XD=XLI(IL)
  YD=YLI(IL)
  XLI(IL)=XL2(IL)
  YLI(IL)=YL2(IL)
  CALL CONSTL(X,Y,XLI(IL),YLI(IL),XL2(IL),YL2(IL),
  YL2(IL)=XD
  YL2(IL)=YD
700 CONTINUE
610 IF(NA.LE.0) GO TO 551
C
  READ AND WRITE AREA SOURCE PROPERTIES
  IL=NL-1
  NN=NL-NA
  WRITE(6,2124)
2124 FORMAT('0 AREA SOURCES,/, ***** /)
  DO 615 I=1,NN
  READ(5,1040)HED2
  READ(5,1000) ALPHA(I),BETA(I),XLI(I),YLI(I),XL2(I),HL(I)
  WRITE(6,1041)HED2
  WRITE(6,2120)
  WRITE(6,1023) ALPHA(I),BETA(I),XLI(I),YLI(I),XL2(I),HL(I)
  WRITE(6,2123) ALPHA2(I),BETA2(I),RML(I)
  615 CONTINUE
2120 FORMAT('0,2X,ALPHA,10X,BETA,11X,XO,
&L2,XO,12X,R,13X,HA)
2123 FORMAT('0 SECOND REGRESSION CONSTANTS,/,/,2X,ALPHA2,10X,
&BETA2,10X,MR)
  551 CONTINUE
C
  ITERATION ON TIME PERIODS
  DO 370 IT=1,NT
  WRITE(6,225)
  WRITE(6,3000)T(IT)
  WRITE(6,2020)T(IT)
5020 FORMAT('0 TIME PERIOD,3X,F6.2)
3000 FORMAT('0, TIME PERIOD=,F10.2, VRS,/)
C
  ITERATION ON GRID
  ITERATION IN THE Y DIRECTION
  Y=YBEGIN
  DO 385 IY=1,NYMAX
  X=XBEGIN
  DO 380 IX=1,NYMAX
  WRITE(6,2222)X,Y
  PDF(I)=0.000
  ITERATION ON PGA
  DO 360 II=2,NY
  YGII=YG(II)
  YGIBI=YGII/BI
  SUM1=0.000
  SUM1=0.000
  ITERATION ON LINE SOURCES
  IF(NL.EQ.0) GO TO 441
  DO 351 IL=1,NL
  CALL CONSTL(X,Y,XLI(IL),YLI(IL),XL2(IL),YL2(IL),

```



```

1 HLIL=HL(IL)
  GAM22=DEXPIALPHA2(IL)
  DEL22=BETA2(IL)/B2
  RHO22=DEL22*B3
  FFY=(DLIL*DLIL)+(HLIL*HLIL)
  RH=(DSORT(FFY)+B4)**RHO22
  EXCK=RH**GAM22*(YGTI1)*DEL22*(AL2*AL1)*T(IT)
  IF(EXCK.LE.1.D-5) GO TO 351
  CALL INTGNTIL,XL1(IL),YL1(IL),XL2(IL),YL2(IL),AL1,AL2,DL1,
1 HL(IL),AIL,BIL,YG11,X1,Y,FFY,GAM22,DEL22,RHO22,FF,FFI)
  SUM=SUM+FF
  SUMI=SUMI+FFI
351 CONTINUE
  C
  C
  C AREA SOURCE COMPUTATION
  IEND=NL+NA
  DO 850 IA=1,IEND
  I1=NL+1
  I2=NL+2
  DO 850 JA=1,I2END
  C CHECK WHETHER THE CIRCLE CONTRIBUTES ANY TO THE PROB
  R1A=XL2(JA)
  H1L=HL(JA)
  GAM22=DEXPIALPHA2(JA)
  DEL22=BETA2(JA)/B2
  RHO22=DEL22*B3
  RH=(DSORT((X-XL1(JA))*(X-XL1(JA))+(Y-YL1(JA))*(Y-YL1(JA))))
  RH=(DIST+B4)**RHO22
  EXCK=RH**GAM22*(YGTI1)*DEL22*(6.283100*R1A*IA)*T(IT)
  IF(EXCK.LE.1.D-5) GO TO 850
  SUMP=0.D0
  SUMPI=0.D0
  NR=R1A/DELTAC*0.900
  DR=R1A/NR
  DRI=DPA*.500
  C
  C ITERATION ON THE NB OF SEGMENTS IN THE CIRCLE
  DO 202 IR=1,NR
  DL1=DABS(DIST-DP1)
  DL2=DIST+DRI
  AL=DSORT(R1A*PIA-DRI*DR1)
  FFY=DL1*DL1+HLIL*HLIL
  CALL INTGN(IA,0.D0,DL1,AL,DL1,AL,DL1,0.D0,DL1,HLIL,0.D0,DL1,
1 YGTI,0.D0,0.D0,FFY,GAM22,DEL22,RHO22,FF,FFI)
  SUMP=SUMP+FF*2.D0
  SUMPI=SUMPI+FFI*2.D0
  FFY=DL2*DL2+HLIL*HLIL
  CALL INTGN(IA,0.D0,DL2,AL,DL2,AL,DL2,0.D0,DL2,HLIL,0.D0,DL2,
1 YGTI,0.D0,0.D0,FFY,GAM22,DEL22,RHO22,FF,FFI)
  SUMP=SUMP+FF*2.D0
  SUMPI=SUMPI+FFI*2.D0
  DRI=DPI+DR
202 CONTINUE
  C
  C SUMP=SUMP*DP
  SUMPI=SUMPI*DR
  SUM=SUM+SUMP
  SUMI=SUMI+SUMPI
1 AL1,AL2,DL1L,AIL,BIL)
  C
  C 442 CONTINUE
  FV1(I1)=DEXP(-SUM*F1(I1))
  FV2(I1)=1.000-FV1(I1)
  PDF(I1)=-SUM*F1(I1)*FV1(I1)
360 CONTINUE
  C
  C WRITE(10,50001)(FV2(J),J=2,NV)
  NNY=(NV-2)/10+1
  K=2
  NVI=11
  DO 375 JJ=1,NNY
  IF(JJ.EQ.NNY) NVI=NY
  WRITE(6,2261)(YGI(J),J=K,NVI)
  WRITE(6,2271)(FY2(J),J=K,NVI)
  WRITE(6,2281)(FYI(J),J=K,NVI)
  WRITE(6,2291)(PDF(J),J=K,NVI)
  NVI=NYI+10
  K=K+10
375 CONTINUE
5010 FORMAT('SITE ',ZF10.2)
5000 FORMAT(10F8.5)
226 FORMAT('0',PGA=' ',10F10.4)
227 FORMAT(' ',P(Y>Y0)',10F10.4)
228 FORMAT(' ',P(Y<Y0)',10F10.4)
229 FORMAT(' ',PDF(Y)',10F10.4)
  C
  C MOVE TO THE NEXT POINT ALONG THE GRID
380 X=X+DX
385 Y=Y+DY
370 CONTINUE
365 CONTINUE
225 FORMAT('I',*****P O R B A B I L I T Y D I S T R I B U T
*****//)
  STOP
  END
1 SURPOUTINE CONSTL(X,Y,XL1,YL1,XL2,YL2,
  AL1,AL2,DL1L,AM,B)
  C
  C *****
  C IMPLICIT REAL*8 (A-H,O-Z)
  C *****
  C AM = SLOPE OF LINE DEFINED BY (X1,Y1),(X2,Y2)
  C B = Y-INTERSEPT
  C
  C AM=(YL2-YL1)/(XL2-XL1)
  C B=YL1-AM*XL1
  C XBAP=(X+AM*Y-B*AM)/(AM*AM+1.000)
  C YBAR=(AM*X+Y*AM+AM*B)/(AM*AM+1.000)
  C DL1=DSORT((X-XBAR)*(X-XBAR)+(Y-YBAR)*(Y-YBAR))
  C
  C CHECK POSITION OF (XBAR,YBAR) ALONG LINE W.R.T. END
  C OF POINTS OF LINE SOURCE
  C AL1=DSORT((XL1-XBAR)*(XL1-XBAR)+(YL1-YBAR)*(YL1-YBAR))
  C AL2=DSORT((XL2-XBAR)*(XL2-XBAR)+(YL2-YBAR)*(YL2-YBAR))
  C
  C IF(XBAR.LT.XL1) AL1=-AL1

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```

IF(XPAP.LE.XL2) RETURN
DUM = AL1
AL1 = -AL2
AL2 = DUM
RETURN
END
SUBROUTINE INTGNIIL,X1,Y1,XL2,YL2,AL1,AL2,DLIL,HLIL,ATL,BIL,
1 YG11,X,Y,FEY,GAM22,DEL22,RHO22,FF,FF1)
C THIS SUBROUTINE COMPUTES THE INTEGRAL ALONG A LINE
C IT IS CALLED FOR LINE SOURCES AS WELL AS FOR CIRCLES SOURCES
C THE PARAMETERS ALPHA AND BETA TO BE USED ARE FIRST CHECKED
C *****
C IMPLICIT REAL*8 (A-H,O-Z)
C INTEGER CASE
C DIMENSION XX(4),YY(4),DEL(2),GAM(2),RHO(2)
C COMMON/SOURCE/ALPHA1(20),ALPHA2(20),BETA1(20),BETA2(20),RML(20),
C HL(20)
C COMMON/ATTEN/R1,B2,B3,B4,DELTA,DELTAC
C *****
C DETERMINE LINE PARAMETERS
C GAM11=DEXP(ALPHA1/IL)
C DEL11=RETA1(IL)/B2
C RHO11=DEL11*B3
C NBSG=1
C XX(1)=XL1
C YY(1)=YL1
C *****
C COMPUTE EPICENTRAL DISTANCE RECK AT WHICH AN EARTHQUAKE
C OF BREAKPOINT MAGNITUDE RML(1) WILL GENERATE AN ACCELERATION
C YG(1)
C IF RECK IS IMAGINARY OR SMALLER THAN DLIL, ALPHA2 AND BETA2
C ARE USED. OTHERWISE POINTS WHERE ALPHA AND BETA CHANGE VALUE
C ARE DETERMINED ON THE LINE
C TERM=(B1*DEXP(B2*RML(1)/YG11)**(1.00/83)
C IF((TERM-B4).LE.HLIL) GO TO 830
C TERM=(TERM-B4)*(TERM-B4)-(HLIL*HLIL)
C *****
C THE EPICENTRAL DISTANCE IS NOT IMAGINARY
C RECK=DSORT(TERM)
C IF(RECK.LE.DLIL) GO TO 830
C *****
C THE EPICENTRAL DISTANCE IS NOT SMALLER THAN DLIL, POINTS OF
C INTERSECTION WITH THE LINE ARE COMPUTED
C TERM1=AL*(BIL-Y)-X
C TERM2=X*X+BIL*BIL+Y*(Y-2.00*BIL)-RECK*RECK
C TERM3=AL*AL+1.00
C TERM2=DSORT((TERM1*TERM1-TERM2*TERM3)
C *****
C COORDINATES OF THE POINTS OF INTERSECTION
C X11=1-TERM1/TERM2/TERM3
C X12=1-TERM1-TERM2/TERM3
C Y11=AL*X11+BIL
C Y12=AL*X12+BIL
C IF(X11.LT.X12) GO TO 150
C STOP=X11
C X11=X12
C X12=STOP

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```

STOP=Y11
Y11=Y12
Y12=STOP
150 CONTINUE
C *****
C CHECK WHETHER THE POINTS OF INTERSECTION ARE ON THE SEGMENT
C CONSIDERED AND DETERMINE THE USE OF ALPHA AND BETA
C *****
C IF(XL1.LT.X11) GO TO 800
C IF(XL1.GT.X12) GO TO 830
C IF(XL2.GT.X12) GO TO 820
C *****
C CASE=1
C GC TO 850
C *****
C 800 IF(XL2.LT.X11) GO TO 830
C IF(XL2.LT.X12) GO TO 840
C *****
C CASE=5
C NBSG=3
C XX(2)=X11
C YY(2)=Y11
C XX(3)=X12
C YY(3)=Y12
C XX(4)=XL2
C YY(4)=YL2
C GO TO 870
C *****
C 830 CASE=3
C 850 XX(2)=XL2
C YY(2)=YL2
C GO TO 870
C *****
C 820 CASE=2
C X11=X12
C Y11=Y12
C GO TO 860
C *****
C 840 CASE=4
C 860 NBSG=2
C XX(2)=X11
C YY(2)=Y11
C XX(3)=XL2
C YY(3)=YL2
C *****
C 870 CONTINUE
C I=1
C IF(CASE.GT.2) I=2
C DEL(I)=DEL11
C GAM(I)=GAM11
C RHO(I)=RHO11
C I=I+1
C IF(I.EQ.3) I=1
C DEL(I)=DEL22
C GAM(I)=GAM22
C RHO(I)=RHO22
C *****
C FF=0.00
C FF1=0.00
C AL=-AL1

```

```

C      DO 900 IXSG=1,NBSG
      TERM=0.DO
      IX=IXSG
      IF(IX.EQ.3) IX=1
C      COMPUTE LENGTH OF SEGMENT CONSIDERED AND DETERMINE
C      NUMBER OF INCREMENTS
      XLEN=DSORT((XX(IXSG)-XX(IXSG+1))*(XX(IXSG)-XX(IXSG+1))+
      1      (YY(IXSG)-YY(IXSG+1))*(YY(IXSG)-YY(IXSG+1)))
      NBIC=(XLEN/DELTA)*.9DO
      IF(NBIC.EQ.0) NBIC=1
      XNBIC=NBIC
      DELTA=XLEN/XXNBIC
C      START DO LOOP AT 1/2 DELTA
      DELDO=DEL(IX)
      RHDDO=RH(IX)
      AL=AL+.5DO*DELTA
C      DO 100 IXIC=1,NBIC
      RH=AL*AL+FFY
      TERM=TERM+DSORT(RH)*B4)**RHDDO
      AL=AL+DELTA
      100 CONTINUE
C      AL=AL-.5DO*DELTA
      TERM=TERM+GAM(IX)*((YGII/RI)**DELDO)*DELTA
      FF=FF+TERM
      FF1=FF1+TERM*DELDO/YGII
      900 CONTINUE
      RETURN
      END
C      $DATA

```

```

PR(I)=1.00
AC(1)=0.00
AC(2)=ACST
DO 120 I=2,NBAC
120 AC(I)=AC(I)+ACIC
C
C PRINT,'AC',AC
C
DO 100 IP=1,NBPD
READ(5,1000)HEDI
1000 FORMAT(20A4)
2000 WRITE(6,2000)HEDI
2000 FORMAT(' ',20A4)
WRITE(10,1000)HEDI
DO 100 IY=1,NBYX
DO 110 IX=1,NBXX
READ(5,1020)HEDI
C1020 FORMAT(/,20A4)
C
1030 READ(5,1030)(PB(I),I=2,NBAC1)
1030 FORMAT(10F8.0)
C
DO 150 KP=2,NBAC1
IF(PBCK.GE.PB(KP)) GO TO 160
150 CONTINUE
C
WRITE(6,2020)HEDI
2020 FORMAT('0*** ATTENTION ***// EXCEDANCE IS LARGER THAN THE INPUT'
1 /, '20A4)
ACPB(IX)=AC(NBAC1)
GO TO 110
C
160 ACPB(IX)=AC(KP-1)+((AC(KP)-AC(KP-1))*(PB(KP-1)-PBCK))/
1 (PB(KP-1)-PB(KP))
110 CONTINUE
C
WRITE(6,2040)(ACPB(I),I=1,NBXX)
2040 FORMAT(' ',12F10.6)
WRITE(10,3000)(ACPB(I),I=1,NBXX)
3000 FORMAT(8F10.6)
100 CONTINUE
RETURN
END
$DATA

```

```

C *****
C PROGRAM COMST,PROB
C STANFORD UNIVERSITY
C *****
C
C GIVEN THE CDF ON ACCELERATION P(A,GT,A0) FOR EACH NODE OF A GRID
C THIS PROGRAM FINDS THE ACCELERATION (A0) CORRESPONDING TO A GIVEN
C PROBABILITY USING LINEAR INTERPOLATION. THE GRID IS COVERED
C STARTING FROM BOTTOM LEFT NODE, ITERATIONS ARE MADE FIRST IN
C THE HORIZONTAL DIRECTION
C
C *****
C INPUT FORMAT
C *****
C
C COL VARIABLE NAME VARIABLE DESCRIPTION
C
C 1.- GENERAL INFORMATION 1 CARD (4I5)
C 1-5 NBPD NB OF RUNS REQUIRED
C 6-10 NBAC NB OF ACC GIVEN FOR EACH NODE
C 11-15 NBXX NB OF COLUMNS
C 16-20 NBYX NB OF ROWS
C
C 2.- CDF PARAMETERS 1 CARD (3F10.0)
C 1-10 ACST SMALLEST ACC IN CDF
C 11-20 ACIC ACC INCREMENT IN CDF
C 21-30 PRCK PROBABILITY A0
C
C 3.- IDENTIFICATION 1 CARD (20A4)
C 1-40 HEDI RUN IDENTIFICATION
C
C 4.- CDF CARDS (NBXX*NBYX) CARDS OR SET OF CARDS (10FB.0)
C 1-10 PR(I) PROBABILITY OF EXCEDANCE CORRESPONDING
C TO THE SMALLEST ACCELERATION
C
C ... PB(I) PROBABILITY OF EXCEDANCE CORRESPONDING
C TO THE LARGEST ACCELERATION
C
C OUTPUT
C *****
C THE OUTPUT IS SAVED ON DISK AND CONTAINS THE RUN IDENTIFICATION
C AND THE ACCELERATION AT EACH NODE CORRESPONDING TO THE GIVEN
C PROBABILITY OF EXCEDANCE. FORMAT (8F10.6)
C *****
C IMPLICIT REAL*8 (A-H,O-Z)
C DIMENSION AC(30),PB(30),ACPB(40),HEDI(20),HEDI2(20)
C *****
C
C READ(5,1010)NBPD,NBAC,NBXX,NBYX
1010 FORMAT(8I5)
WRITE(6,1010)NBPD,NBAC,NBXX,NBYX
2010 FORMAT('0NBPD',I6,' NBAC',I6,' NBXX',I6,' NBYX',I6)
C
READ(5,1040)ACST,ACIC,PRCK
1040 FORMAT(8F10.0)
WRITE(6,2030)ACST,ACIC,PRCK
2030 FORMAT('0ACST',F10.2,' ACIC',F10.2,' PRCK',F10.2)
C
NBAC1=NBAC+1

```

```

C ***** STANFORD UNIVERSITY *****
C PROGRAM ACC TO INT
C *****
C THIS PROGRAM FINDS THE MODIFIED MERCALLI INTENSITIES CORRESPONDING
C TO A CDF ON ACCELERATION USING THE GUTENBERG-RICHTER RELATIONSHIP
C A MONTF-CAPLO PROCESS IS USED. CDF P(A.GT.AOI) IS INPUT.
C *****
C INPUT FORMAT
C *****
C COL VARIABLE NAME VARIABLE DESCRIPTION
C 1-- GENERAL 1 CARD (315)
C 1-5 NR (NUMBER OF POINTS IN CDF AS READ IN)*2
C 6-10 IV RANDOM GENERATOR INIATOR
C 11-15 NRPT NB OF CDF TO BE READ IN
C 2-- ACCELERATIONS 8 VALUES PER CARD (8F10.0)
C 1-10 X(4) SMALLEST ACC IN CDF
C ... X(1)
C ... X(NB*2)
C ... X(NB*4)
C ... X(NB*8)
C THE REMAINING DATA IS READ FROM DISK
C 3-- MODE IDENTIFICATION 1 CARD IMAGE (2044)
C 1-90 HEDI IDENTIFICATION
C 4-- CDF P(A.GT.AOI) 8 VALUES PER CARD IMAGE (8F10.0)
C 1-10 X(3) P(A.GT.X(4))
C ... X(1)
C ... X(NB*1)
C ... X(NB*2)
C OUTPUT
C *****
C THE OUTPUT DISPLAYS THE INPUT, THE NUMBER OF OCCURRENCES OF
C THE MODIFIED MERCALLI INTENSITY IN THE SIMULATION AND THE
C CORRESPONDING PERCENTAGE.
C *****
C DIMENSION X1(44),INTEN(15),HEDI(20)
C COMMON/POI/IV,DELPOI
C *****
C READ(5,1000)NB,IY,NBPT
C NR2=NB*4
C NR1=NR*1
C X1(1)=0.
C X1(2)=0.
C 1000 FORMAT(3I5)
C READ(5,1100)(X1(I),I=4,NR2,2)
C 1100 FORMAT(F10.0)
C X1(NB*3)=100.
C CALL RANDKITY,YFL,0)
C YFL=-.01
C READ(10,1020)HEDI
C WRITE(6,2010) HEDI
C D= 400 IP=1,NRPT

```

```

1020 FORMAT(20A4)
READ(10,1200)(X1(I),I=3,NR1,2)
1200 FORMAT(10F8.0)
DO 60 I=3,NR1,2
60 X1(I)=1.-X1(I)
DO 220 I=1,NB1,2
220 X1(I)=X1(I)*100.
C
C WRITE(6,2010)HEDI
2010 FORMAT(' ',//',',20A4)
WRITE(6,2050)
2050 FORMAT('0',7X,'CDF',7X,'ACC')
C
C WRITE(6,2020)X1
2020 FORMAT(' ',2F10.3)
DO 90 IT=1,12
90 INTEN(IT)=0
C
DO 100 IT=1,1000
CALL CDF(X1,NB2,1,ACC)
INT=3.*(ALOG10(ACC*1000.))+.5)+.5
IF(INT.LT.1) GO TO 100
PRINT *,ACC,INT,ACC,INT
INTEN(INT)=INTEN(INT)+1
100 CONTINUE
C
TOT=0.
DO 110 INT=1,12
PRINT, INT,NB OF OCC,INT,INTEN(INT)
110 TOT=TOT+FLOAT(INTEN(INT))
C
C WRITE(6,2060)
2060 FORMAT('0',INTENSITY PERCENTAGE//',',19,F13.3)
SUB=FLOAT(INTEN(INT))/TOT
DO 120 INT=1,12
WRITE(6,2030)INT,SUB
120 WRITE(6,2030)INT,SUB
400 CONTINUE
2030 FORMAT(' ',19,F13.3)
RETURN
END
SUBROUTINE CDF(X1,KEND,ICK,AMS)
C THIS SUBROUTINE FINDS THE VALUE CORRESPONDING TO THE RANDOM
C VARIABLE ON THE CDF, LINEAR FITTING IS APPLIED IF INDEX.EO.1
C *****
C COMMON/POI/IV,DELPOI
C DIMENSION X1(KEND)
C *****
C KEN=KEND/2
C CALL RANDKITY,YFL,0)
C YFL=YFL+.01
C DO 100 I=1,KEN
C YFL=DELPOI*(X1(2*I-1)/100.)) GO TO 110
100 CONTINUE
WRITE(6,1000)
1000 FORMAT(27H0** ERROR IN CDF INPUT ** //

```

```

1 264 LAST PERCENTAGE INPUT IS NOT 100.01 //)
I=I-1
110 IF(IICK-1) 200,300,120
200 ANS = XI(2*I)
RETURN
C
C LINEAR FITTING HALF DISTANCE BETWEEN INPUT POINTS
12) IF(I.EQ.KENJGO TO 140
IF(I.LT.3) GO TO 130
XT=(XI(2*I)+XI(2*I-2))*5
YT=XI(2*I-1)
XB=(XI(2*I-4)+XI(2*I-2))*5
YB=XI(2*I-3)
GO TO 150
C
130 XT=(XI(2*I)+XI(2*I-2))*5
YT=XI(2*I-1)
XB=XI(2*I-2)
YB=0.
GO TO 150
140 XT=XI(2*I)
YT=100.
160 XB=(XI(2*I-4)+XI(2*I-2))*5
YB=XI(2*I-3)
GO TO 150
C
C LINEAR FIT THROUGH INPUT POINTS
300 XT=XI(2*I)
YT=XI(2*I-1)
XB=XI(2*I-2)
YB=XI(2*I-3)
C
C CHECK FOR VERTICAL INTERPOLATION
150 ANS=XT
IF(XT-XB).LT..00001) RETURN
C
A=(YT-YB)/(XT-XB)
ANS=(YEL*100.-(YB-A*XT))/A
RETURN
END
SUBROUTINE RANDK (IY,YEL,INDEX)
C
C THIS SUBROUTINE GENERATES RANDOM NUMBERS
IY=IY*314159269+453806245
4 IF(IY.GE.0) GO TO 6
5 IY=IY+2147483647+1
6 CONTINUE
IF(INDEX.GT.0) GO TO 8
YEL=FLNAT(IY)*.4656613E-9
9 RETURN
END
$DATA

```

```

*****
PROGRAM PLOT.EPI
*****
STANFORD UNIVERSITY
*****
THIS PROGRAM PLOTS THE EARTHQUAKE EPICENTERS ON A GRID (MAP)
GIVEN THE LATITUDE AND LONGITUDE. DIFFERENT SYMBOLS ARE
USED FOR THE DIFFERENT MAGNITUDES.
THE PLOT WILL BE ROTATED 90 DEGREES COUNTERCLOCKWISE IF IT IS TOO
HIGH TO FIT ON THE PLOTTER. IF BOTH DIMENSIONS ARE TOO BIG, AN
MESSAGE WILL BE GENERATED AND THE PLOT ABORTED.
*****
INPUT FORMAT
*****
COL VARIABLE NAME VARIABLE DESCRIPTION
1-- IDENTIFICATION CARD 1 CARD (80A1) RUN IDENTIFICATION
1-80 HED
2-- PLOT FLAGS 1 CARD (715) NO OF PLOTS
1-5 NRPL -0, WILL PLOT EO OF MAG 3
6-10 SKIP3 -1, WILL NOT PLOT EO OF MAG 3
11-15 SKIP4 EO OF MAG 4
16-20 SKIP5 EO OF MAG 5
21-25 SKIP6 EO OF MAG 6
26-30 SKIP7 EO OF MAG 7
31-35 SKIP8 EO OF MAG 8
3-- GRID PARAMETERS 1 CARD (8F10.0)
THE ORIGIN OF THE GRID IS DEFINED AS
THE BOTTOM LEFT CORNER; THE END AS THE
TOP RIGHT. THE VERTICAL HAS TO POINT
NORTH FOR THE DATA TO BE READ IN
CONSISTENTLY.
1-10 XMIN1 X COORD OF ORIGIN
11-20 YMIN1 Y COORD OF ORIGIN
21-30 XMAX1 X COORD OF END
31-40 YMAX1 Y COORD OF END
41-50 INDGX LENGTH OF ONE DEGREE IN INCHES FOR
THE PLOT (X DIRECTION).
51-60 INDGY LENGTH OF ONE DEGREE IN INCHES FOR
THE PLOT (Y DIRECTION)
61-70 CAL PLOTTER TO BE USED (10. OR 29.)
71-80 DIVR IF ZERO, WILL BE SET TO 10.
THE FRAME COORDINATES ARE ROUNDED OFF
TO THE CLOSEST (DEGREE/DIVR)
IF INPUT AS 0, WILL BE SET TO 1
3-- EPICENTER COORDINATES 1 EPICENTER PER CARD IMAGE (3F10.0)
INFORMATION READ FROM DISK
X( ) X COORDINATE (LONGITUDE) OF EPICENTER
Y( ) Y COORDINATE (LATITUDE) OF EPICENTER
RA( ) MAGNITUDE
OUTPUT
*****
A FILE IS CREATED ON DISK, IT HAS TO BE TRANSFERRED ON TAPE AND
PLOTTED ON A 10 OR 29 INCH CALCOMP PLOTTER

```

```

C *****
C DIMENSION AMODES(200),END1(2),END2(2),XX4(500),YY4(500),XX5(500)
C ,YY5(500),XX6(300),YY6(300),XX7(100),YY7(100),XX8(100),YY8(100),
C &X(2050),Y(2050),PM(2050),XX(200),YY(200),XX3(500),YY3(500)
C INTEGER,SKIP3,SKIP4,SKIP5,SKIP6,SKIP7,SKIP8
C REAL INRAS,INDG,INDGX,INDGY
C LOGICAL*1 HED(70)
C *****
C READ(5,1001)HED
C FORMAT(80A1)
C READ(5,1002)NRPL,SKIP3,SKIP4,SKIP5,SKIP6,SKIP7,SKIP8
C FORMAT(10I5)
C READ(5,1000)XMIN1,YMIN1,XMAX1,YMAX1,INDGX,INDGY,CAL,DIVR
C IF(DIVR.EQ.0.) DIVR=1.
C FORMAT(8F10.0)
C INRAS=100.
C IF(CAL.EQ.29.) INRAS=200.
C XRAS=(ABS(XMAX1-XMIN1)*INDGX+1.5)*INRAS
C YRAS=(ABS(YMAX1-YMIN1)*INDGY+1.5)*INRAS
C XMIN1=XMIN1*DIVR
C XMAX1=XMAX1*DIVR
C YMIN1=YMIN1*DIVR
C YMAX1=YMAX1*DIVR
C INDGX=INDGX/DIVR
C INDGY=INDGY/DIVR
C ORIENTATE FRAME
C DELTAX=-.75/INDGX
C DELTAY=-.75/INDGY
C IF(YRAS.LE.(CAL*INRAS)) GO TO 112
C IF(XRAS.LE.(CAL*INRAS)) GO TO 110
C WRITE(6,2100)
C 2100 FORMAT('OTHER SCALE IS TOO LARGE FOR THE PLOTTER, REDUCE INDG PER D
C EGREE')
C STOP
C INTERCHANGE X AND Y SCALES
C 110 WRITE(6,2101)
C 2101 FORMAT('OTHER PLOT WILL BE ROTATED BY 90 DEGREES')
C STORE=XMIN1
C XMIN1=YMAX1
C YMIN1=XMAX1
C YMAX1=YMIN1
C YMIN1=STOR
C DELTAX=DELTAY
C DELTAY=STOR
C STORE=XRAS
C XRAS=YRAS+3.5*INRAS
C YRAS=STOR
C GO TO 111
C 112 XRAS=XRAS+3.5*INRAS
C 111 CALL MODESG(AMODES,'MORTGAT,BIN 502',15)
C DO 500 IP=1,NRPL
C AMODES(45)=1.
C XMAX=XMAX1

```

```

YMAX=YMAXI
XMIN=XMINI
YMIN=YMINI
C
1030 READ(10,1030)NBRC,HED
C
IF(YRAS.LE.(CAL*(R*AS))) GO TO 100
DO 120 IX=1,NBRC
120 READ(10,1010)X(I),Y(I),RM(I)
1010 FORMAT(19X,F5.3,3X,F5.3,42X,F3.2)
GO TO 130
C
100 CONTINUE
DO 140 IX=1,NBRC
140 READ(10,1010)Y(I),X(I),RM(I)
C
CHECK SIGNS
130 IF(XMAXI.GT.XMINI) GO TO 150
XMAX=-XMAXI
XMIN=-XMINI
DO 160 IX=1,NBRC
160 X(IX)=-X(IX)
C
150 XMAX=XMAX*DELTA*4.25/.75
XMIN=XMIN*DELTA
IF(YMAXI.GT.YMINI) GO TO 170
YMAX=-YMAXI
YMIN=-YMINI
DO 180 IX=1,NBRC
180 Y(IX)=-Y(IX)
170 YMAX=YMAX*DELTA
YMIN=YMIN*DELTA
552 WRITE(6,551)(X(IX),Y(IX),RM(IX)),IX=1,NBRC
551 FORMAT(' ',12F10.3)
C
AMODES(97)=XRAS
AMODES(179)=YRAS
CALL SURJEG(AMODES,XFMIN,YFMIN,XFMAX,YFMAX)
CALL OBJEGG(AMODES,0.,0.,XRAS,YRAS)
C
IF(AMODES(12).NE.AMODES(13))WRITE(6,2102)AMODES(12),AMODES(13)
2102 FORMAT('OTHE SCALES ARE NOT THE SAME',2F10.6)
FND1(1)=IFIX(YMIN)
FND2(1)=IFIX(XMAX)
FND1(2)=IFIX(YMAX)
FND2(2)=IFIX(XMIN)
C
DELTA2=1./(.6.*INDGX)
XMUL=4.
DELTA1=4.*DELTA2
DELTA3=3.*DELTA2
SIGN=-1.
C
I=1
VAP=FND1(2)
FIX=FND2(2)
C
DO 400 K=1,2
400 CONTINUE
C
DO 500 K=1,I
500 WRITE(6,550)XX(K),YY(K)
550 FORMAT(' ',2F10.3,16)
I=I-3
CALL LINESG(AMODES,I,XX,YY)
C
I3=0
I4=0
I5=0
I6=0
I7=0
I8=0
DO 300 IX=1,NBRC
IRM=RM(IX)-1.99

```



```

IF (IRM.LT.1) GO TO 300
GO TO (230,240,250,260,270,280),IRM
C
XX3(I3)=X(IX)
YY3(I3)=Y(IX)
GO TO 300
C
I4=I4+1
XX4(I4)=X(IX)
YY4(I4)=Y(IX)
GO TO 300
C
I5=I5+1
XX5(I5)=X(IX)
YY5(I5)=Y(IX)
GO TO 300
C
I6=I6+1
XX6(I6)=X(IX)
YY6(I6)=Y(IX)
GO TO 300
C
I7=I7+1
XX7(I7)=X(IX)
YY7(I7)=Y(IX)
GO TO 300
C
I8=I8+1
XX8(I8)=X(IX)
YY8(I8)=Y(IX)
300 CONTINUE
AMODES(40)=8.*INRAS/100.
AMODES(41)=14.*INPAS/100.
C
AMODES(45)=.5
AMODES(84)=12.
IF (SKIP3.EQ.0.AND.I3.GT.0) CALL POINTG(AMODES,I3,XX3,YY3)
IF (SKIP3.EQ.0.AND.I3.GT.0) PRINT,'I3 HAS BEEN CALLED'
IF (SKIP4.EQ.0.AND.I4.GT.0) CALL POINTG(AMODES,I4,XX4,YY4)
IF (SKIP4.EQ.0.AND.I4.GT.0) PRINT,'I4 HAS BEEN CALLED'
C
AMODES(84)=15.
IF (SKIP5.EQ.0.AND.I5.GT.0) CALL POINTG(AMODES,I5,XX5,YY5)
IF (SKIP5.EQ.0.AND.I5.GT.0) PRINT,'I5 HAS BEEN CALLED'
C
AMODES(45)=.35
AMODES(84)=40.
IF (SKIP6.EQ.0.AND.I6.GT.0) CALL POINTG(AMODES,I6,XX6,YY6)
IF (SKIP6.EQ.0.AND.I6.GT.0) PRINT,'I6 HAS BEEN CALLED'
C
AMODES(45)=2.0
AMODES(84)=12.
IF (SKIP7.EQ.0.AND.I7.GT.0) CALL POINTG(AMODES,I7,XX7,YY7)
IF (SKIP7.EQ.0.AND.I7.GT.0) PRINT,'I7 HAS BEEN CALLED'
C
AMODES(45)=.65
AMODES(84)=40.
IF (SKIP8.EQ.0.AND.I8.GT.0) CALL POINTG(AMODES,I8,XX8,YY8)
IF (SKIP8.EQ.0.AND.I8.GT.0) PRINT,'I8 HAS BEEN CALLED'
C
XT=XMAX-.5/INDGX
AMODES(50)=90.
AMODES(46)=90.
YT=YMIN+DELTA1
AMODES(45)=.5
XT=XFMAX-.3./INDGX
CALL VEC SG(AMODES,XT,YT,I6,' . MAGNITUDE 4+')
C
XT=XT+DELTA1
CALL VEC SG(AMODES,XT,YT,I6,' + MAGNITUDE 5+')
C
XT=XT+DELTA1
CALL VEC SG(AMODES,XT,YT,I6,' X MAGNITUDE 6+')
C
XT=XT+DELTA1
AMODES(45)=1.0
CALL VEC SG(AMODES,XT,YT,I6,' . MAGNITUDE 7+')
C
XT=XT+DELTA1
AMODES(45)=.55
CALL VEC SG(AMODES,XT,YT,I6,' X MAGNITUDE 8+')
C
XT=XT+DELTA1
AMODES(45)=1.5
CALL VEC SG(AMODES,XT,YT,48,HEO)
CALL PICTRG(AMODES)
PPRINT,'MAGNITUDE 3'
PRINT,'MAGNITUDE 4 +'
IF (I3.NE.0)WRITE(6,555)(XX3(I),YY3(I),I=1,I3)
IF (I4.NE.0)WRITE(6,555)(XX4(I),YY4(I),I=1,I4)
PRINT,'MAGNITUDE 5 +'
IF (I5.NE.0)WRITE(6,555)(XX5(I),YY5(I),I=1,I5)
PRINT,'MAGNITUDE 6 +'
IF (I6.NE.0)WRITE(6,555)(XX6(I),YY6(I),I=1,I6)
PRINT,'MAGNITUDE 7 +'
IF (I7.NE.0)WRITE(6,555)(XX7(I),YY7(I),I=1,I7)
PRINT,'MAGNITUDE 8 +'
IF (I8.NE.0)WRITE(6,555)(XX8(I),YY8(I),I=1,I8)
500 CONTINUE
555 FORMAT(' ',I2F10.3)
CALL EXITG(AMODES)
RETURN
END
$DATA

```

```

IF (IRM.LT.1) GO TO 300
GO TO (230,240,250,260,270,280),IRM
C
XX3(I3)=X(IX)
YY3(I3)=Y(IX)
GO TO 300
C
I4=I4+1
XX4(I4)=X(IX)
YY4(I4)=Y(IX)
GO TO 300
C
I5=I5+1
XX5(I5)=X(IX)
YY5(I5)=Y(IX)
GO TO 300
C
I6=I6+1
XX6(I6)=X(IX)
YY6(I6)=Y(IX)
GO TO 300
C
I7=I7+1
XX7(I7)=X(IX)
YY7(I7)=Y(IX)
GO TO 300
C
I8=I8+1
XX8(I8)=X(IX)
YY8(I8)=Y(IX)
300 CONTINUE
AMODES(40)=8.*INRAS/100.
AMODES(41)=14.*INPAS/100.
C
AMODES(45)=.5
AMODES(84)=12.
IF (SKIP3.EQ.0.AND.I3.GT.0) CALL POINTG(AMODES,I3,XX3,YY3)
IF (SKIP3.EQ.0.AND.I3.GT.0) PRINT,'I3 HAS BEEN CALLED'
IF (SKIP4.EQ.0.AND.I4.GT.0) CALL POINTG(AMODES,I4,XX4,YY4)
IF (SKIP4.EQ.0.AND.I4.GT.0) PRINT,'I4 HAS BEEN CALLED'
C
AMODES(84)=15.
IF (SKIP5.EQ.0.AND.I5.GT.0) CALL POINTG(AMODES,I5,XX5,YY5)
IF (SKIP5.EQ.0.AND.I5.GT.0) PRINT,'I5 HAS BEEN CALLED'
C
AMODES(45)=.35
AMODES(84)=40.
IF (SKIP6.EQ.0.AND.I6.GT.0) CALL POINTG(AMODES,I6,XX6,YY6)
IF (SKIP6.EQ.0.AND.I6.GT.0) PRINT,'I6 HAS BEEN CALLED'
C
AMODES(45)=2.0
AMODES(84)=12.
IF (SKIP7.EQ.0.AND.I7.GT.0) CALL POINTG(AMODES,I7,XX7,YY7)
IF (SKIP7.EQ.0.AND.I7.GT.0) PRINT,'I7 HAS BEEN CALLED'
C
AMODES(45)=.65
AMODES(84)=40.
IF (SKIP8.EQ.0.AND.I8.GT.0) CALL POINTG(AMODES,I8,XX8,YY8)
IF (SKIP8.EQ.0.AND.I8.GT.0) PRINT,'I8 HAS BEEN CALLED'
C
XT=XMAX-.5/INDGX
AMODES(50)=90.
AMODES(46)=90.
YT=YMIN+DELTA1
AMODES(45)=.5

```

```

C C ***** STANFORD UNIVERSITY *****
C C PROGRAM BYS.ACC.DUR *****
C C THIS PROGRAM COMPUTES THE PROBABILITY OF EXCEEDANCE OF GIVEN PEAK *****
C C GROUND ACCELERATIONS AND STRONG MOTION DURATION AT SPECIFIED *****
C C SITES DUE TO LINE OR AREA SEISMIC SOURCES. *****
C C CDF OF PGA IS COMPUTED WITH INCREMENTS OF .02 G *****
C C CDF OF DURATION IS COMPUTED WITH INCREMENTS OF 2 SEC *****
C C INPUT FORMAT *****
C C *****
C C 1.- IDENTIFICATION CARD 1 CARD 2044
C C
C C 1-80 HED1 TITLE
C C
C C 2.- ATTENUATION CONSTANTS 1 CARD 3F10.0
C C
C C 1-10 81
C C 11-20 82
C C 21-30 83
C C 31-40 84
C C 41-50 DELTA
C C 51-60 DELTAC
C C
C C STEP SIZE FOR LINE INTEGRATION
C C STEP SIZE FOR CIRCLE INTEGRATION
C C UNITS OF LENGTH USED IN THE PROGRAM
C C ATTENUATION FORMULA OF THE TYPE
C C ACC=81*EXP(0.2*MAG)/(R+8)*J**83
C C UPAC
C C
C C RATIO OF UPPER BOUND TO MEAN OF UNIFORM
C C DISTRIBUTION PLACED ON PGA ATTENUATION
C C RELATIONSHIP
C C RATIO OF UPPER BOUND TO MEAN OF UNIFORM
C C DISTRIBUTION PLACED ON DURATION ATTENUATION
C C RELATIONSHIP
C C
C C 3.- PROBLEM DESCRIPTION 3 CARDS
C C
C C CARD 1 615
C C 1- 5 NL NUMBER OF LINE SOURCES
C C 6-10 NA NUMBER OF AREA SOURCES
C C 11-15 NT NUMBER OF TIME PERIODS
C C 16-20 NY NUMBER OF ACC
C C 21-25 NBGO NUMBER OF GRIDS
C C 26-30 SKSAVE IF 0 WILL SAVE RESULTS ON DEVICE SPECIFIED
C C BY JCL FOR DEVICES 11,12 ETC. FOR FIRST,
C C SECOND, ETC. TIME PERIODS.
C C IF 1 WILL NOT SAVE RESULTS
C C
C C CARD 2 TIME PERIODS, NT VALUES, 8 VALUES PER CARD 8F10.0
C C T(1) PERIOD 1
C C 11-20 T(2) PERIOD 2
C C ***** T(NT) PERIODS NT
C C
C C 4.- LINE SOURCES PROPERTIES AT LEAST 3 CARDS PER SOURCE
C C
C C CARD 1
C C 1-80 HED2 IDENTIFICATION OF THE LINE SOURCE
C C
C C CARD 2 7F10.0,110
C C GEOMETRIC DESCRIPTION AND PARAMETERS OF POISSON MODEL

```

```

C C X XORI J X COORD OF ORIGIN
C C 11-20 XXEDA J X COORD OF END
C C 21-30 YYORI J Y COORD OF ORIGIN
C C 31-40 YYEDI J Y COORD OF END
C C 41-50 DPT J DEPTH OF LINE
C C 51-60 TMDAL J TIME DATA BASE
C C 61-70 XNBDA( J NO OF EVENTS RECORDED ABOVE M=4 ON THIS SOURCE
C C 71-80 NB NO OF DIFFERENT MAG GREATER OR EQUAL TO 4.0
C C ON THIS SOURCE (.25 INCREMENT)
C C
C C CARD 3 8 VALUES PER CARD 8F10.0
C C PARAMETERS OF BERNOULLI MODEL
C C XNBEV( J NO OF EVENTS BERNOULLI TRIAL IS BASED ON
C C 11-20 XNBMG(1) NO OF SUCCESSSES FOR M=4.0
C C 21-30 XNBMG(2) NO OF SUCCESSSES FOR M=4.25
C C
C C 5.- AREA SOURCES PROPERTIES AT LEAST 3 CARDS PER SOURCE
C C
C C CARD 1
C C 1-80 HED2 AREA SOURCE IDENTIFICATION
C C
C C CARD 2 7F10.0,110
C C GEOMETRIC DESCRIPTION AND PARAMETERS OF POISSON MODEL
C C X XORI J X COORD OF CIRCLE CENTER
C C 11-20 XXEDI J Y COORD OF CIRCLE CENTER
C C 21-30 YYORI J RADIUS OF CIRCLE
C C 31-40 DPT( J LEAVE BLANK
C C 41-50 TMDAL J DEPTH OF CIRCLE
C C 51-60 TMDAL J TIME DATA BASE
C C 61-70 XNBDA( J NO OF EVENTS RECORDED ABOVE M=4 ON THIS SOURCE
C C 71-80 NB NO OF DIFFERENT MAG GREATER OR EQUAL TO 4.0
C C ON THIS SOURCE (.25 INCREMENT)
C C
C C CARD 3 8 VALUES PER CARD 8F10.0
C C PARAMETERS OF BERNOULLI MODEL
C C XNBEV( J NO OF EVENTS BERNOULLI TRIAL IS BASED ON
C C 11-20 XNBMG(1) NO OF SUCCESSSES FOR M=4.0
C C 21-30 XNBMG(2) NO OF SUCCESSSES FOR M=4.25
C C
C C 6.- GRID DESCRIPTION 2 CARDS PER GRID
C C
C C CARD 1 IDENTIFICATION CARD 2044
C C 1-80 HED2 NAME OF GRID
C C
C C CARD 2 SEARCHING GRID DESCRIPTION 8F10.0
C C 11-20 XBEGIN ORIGIN OF GRID X COORD
C C 11-20 YBEGIN ORIGIN OF GRID Y COORD
C C 21-30 XEND END OF GRID X COORD
C C 31-40 YEND END OF GRID Y COORD
C C
C C X INCREMENT
C C 41-50 DX X INCREMENT
C C 51-60 DY Y INCREMENT
C C
C C IF ONLY ONE LOCATION IS REQUIRED, INPUT THE
C C SITE COORDINATES AS XBEGIN AND XEND, AND
C C INPUT XEND,YEND,DX,DY AS ZEROS

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C      OUTPUT
C      *****
C      DEPENDING UPON THE INPUT, THE OUTPJT WILL DISPLAY FROM ONE
C      PROBABILITY OF EXCEEDANCE AT ONE POINT TO A WHOLE CDF FOR A
C      WHOLE GRID OF POINTS
C      *****
C      NOTE: THE PRESENT DIMENSIONS ALLOW FOR 30 SEISMIC SOURCES,
C      AND 5 TIME PERIODS
C      *****
C      IMPLICIT REAL*8 (A-H,O-Z)
C      *****
C      LOGICAL*1 USE
C      DIMENSION HED1(20),HED2(20)
C      COMMON/BRNUI/XNBV(30),XNBTP(30),XNBMG(30,20)
C      COMMON/ATTEN/BI,B2,B3,B4,DELTA,DELTA1,DELTA2,DELTA3,DELTA4,DELTA5,DELTA6,DELTA7,DELTA8,DELTA9,DELTA10,DELTA11,DELTA12,DELTA13,DELTA14,DELTA15,DELTA16,DELTA17,DELTA18,DELTA19,DELTA20
C      COMMON/GEOM/ XCOR(30),YYOR(30),XXED(30),YYED(30),DP(30),
C      TMDA(30),XNBDA(30)
C      COMMON/ACV/PBACAV(10),B1EB2M(20),PBDC(1000),PBZROC(30,20)
C      COMMON/ACIN/ACIN(10),PBORAV(10),DR50(20)
C      COMMON/ACIN/IKACIN,NBACAV(10),NBDRAV(51,3)
C      DIMENSION TML(0),ACUT(10),XXMGST(20),DR50(120)
C      INTEGER SKSAVE
C      DATA DR50/1.500,1.500,1.500,1.500,2.000,2.000,2.000,2.000,2.500,
C      1 3.000,5.500,10.00,17.00,
C      1 22.00,24.2500,26.00,27.2500,28.00,28.500,29.00,29.500/
C      *****
C      READ AND WRITE ATTENUATION CONSTANTS
C      *****
C      DO 97 I=1,20
C      97 DR50(I)=DR50(I)
C      DO 98 IX=1,20
C      XIX=IX
C      98 XXMGST(IX)=3.7500+XIX*.2500
C      YGMN=-.0500
C      READ(5,1040)HED1
C      1040 FORMAT(20A4)
C      WRITE(6,1041)HED1
C      1041 FORMAT('0',20A4)
C      READ(5,1000)BI,B2,B3,B4,DELTA,DELTA1,DELTA2,DELTA3,DELTA4,DELTA5,DELTA6,DELTA7,DELTA8,DELTA9,DELTA10,DELTA11,DELTA12,DELTA13,DELTA14,DELTA15,DELTA16,DELTA17,DELTA18,DELTA19,DELTA20
C      IF(IUPDR.EQ.0.00) UPAC=1.500
C      WRITE(6,1020)BI,B2,B3,B4,DELTA,DELTA1,DELTA2,DELTA3,DELTA4,DELTA5,DELTA6,DELTA7,DELTA8,DELTA9,DELTA10,DELTA11,DELTA12,DELTA13,DELTA14,DELTA15,DELTA16,DELTA17,DELTA18,DELTA19,DELTA20
C      1020 FORMAT('0','ATTENUATION CONSTANTS',//,' 81',E15.7,3X,'82',E15.7,
C      1 3X,'83',E15.7,3X,'84',E15.7 / DELTA1 =,E15.7 / DELTA2 =,
C      1 1,E15.7 /)
C      1000 FORMAT(8F10.0)
C      YGMNCK=YGMN/UPAC
C      READ AND WRITE GEOMETRIC CONSTANTS
C      READ(5,1001)NBLN,NBAR,NBTR,NBGO,SKSAVE
C      WRITE(6,2030)NBLN,NBAR,NBTR,NBGO
C      2030 FORMAT('0',NUMBER OF LINE SOURCES',15 /
C      1 15 /
C      2 15 /
C      3 15 /
C      4 15 /)
C      IF (SKSAVE.EQ.0) WRITE(6,2031)
C      2031 FORMAT('USAVE OUTPUT ON CARDS',/)
C      NMSC=NBLN+NBAR
C      1001 FORMAT(10I5)

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READ(5,1000)TMDA(1),I=1,NBTRM)
WRITE(6,2210)
WRITE(6,2220)TMDA(1),I=1,NBTRM)
2220 FORMAT(' ',10F10.2)
2210 FORMAT('0',TIME PERIODS')
2222 FORMAT('0','SITE LOCATION',//,' ',X=',F10.3,5X,'Y=',F10.3)
IF(NBLN.LE.0) GO TO 610
C
C      READ AND WRITE LINE SOURCE PROPERTIES
C      *****
C      WRITE(6,1024)
C      1024 FORMAT('0',LINE SOURCES',//, ***** /)
C      DO 616 IX=1,NBLN
C      READ(5,1040)HED2
C      READ(5,1010)XXCOR(IX),XXED(IX),YYOR(IX),YYED(IX),DP(IX),TMDA(IX),
C      1 XNBDA(IX),NB
C      1010 FORMAT(7F10.0,1I0)
C      READ(5,1000)XNBV(IX),(XNBMG(IX,I),I=1,NB)
C      XNBTP(IX)=NB
C      WRITE(6,1041)HED2
C      WRITE(6,1022)
C      1022 FORMAT(' ',7X,'XCOR',7X,'YOR',7X,'XED',7X,'YED',5X,'DEPTH',
C      1 5X,'OCCURRENCE TIME',5X,'TIME BASE',5X,'NB OF OCC')
C      WRITE(6,1023)XXCOR(IX),YYOR(IX),XXED(IX),YYED(IX),DP(IX),TMDA(IX),
C      1 XNBDA(IX)
C      1023 FORMAT(' ',5F10.3,25X,2F10.3)
C      WRITE(6,1031)XNBTP(IX),XNBV(IX)
C      1031 FORMAT(' NB OF DIFF MAG.',F10.2 /
C      1 1, TOTAL NB OF MAG.',F10.2 /
C      2 1, DISTRIBUTION OF MAG. ')
C      IXED=0
C      DO 616 IXST=1,NB,10
C      IXED=IXED+10
C      IF(NB-IXST.LT.9)IXED=NB
C      WRITE(6,1025)(XNBMG(IX,I),I=IXST,IXED)
C      616 WRITE(6,1025)(XNBMG(IX,I),I=IXST,IXED)
C      1025 FORMAT(' ',10F12.2)
C
C      DO 700 IL=1,NBLN
C      IF(XXCOR(IL).LT.XXED(IL)) GO TO 700
C      XE=XXCOR(IL)
C      YD=YYOR(IL)
C      XXOR(IL)=XXED(IL)
C      YYOR(IL)=YYED(IL)
C      XXED(IL)=XE
C      YYED(IL)=YD
C      700 CONTINUE
C      610 IF(NBAR.LE.0) GO TO 551
C
C      READ AND WRITE AREA SOURCE PROPERTIES
C      IXI=NBLN+1
C      WRITE(6,2124)
C      2124 FORMAT(' '//, AREA SOURCES',//, ***** /)
C
C      DO 615 IX=IXI,NBSC
C      READ(5,1040)HED2
C      READ(5,1010)XXCOR(IX),YYOR(IX),XXED(IX),YYED(IX),DP(IX),TMDA(IX),
C      1 XNBDA(IX),NB
C      READ(5,1000)XNBV(IX),(XNBMG(IX,I),I=1,NB)
C      XNBTP(IX)=NB

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5020 FORMAT('TIME PERIOD',F6.2,' XB=',F7.3,' XE=',F7.3,' YB=',F7.3,'
6 ' YE=',F7.3,' DX=',F5.2,' DY=',F5.2)
650 CONTINUE
651 CONTINUE
C
C ITERATION UN GRID
C ITERATION IN THE Y DIRECTION
Y=YBEGIN
DO 385 IY=1,NYMAX
C
C ITERATION IN THE X DIRECTION
X=XBEGIN
DO 380 IX=1,NKMAX
WRITE(6,222) X,Y
C
C INITIALIZE ACCELERATION DISTRIBUTION
ACPDF(I)=0.00
DRPDF(I)=0.00
DO 390 I=2,200
390 ACPDF(I)=1.00
DO 391 I=2,100
391 DRPDF(I)=1.00
C
C ITERATION ON LINE SOURCES
IF(NBLN.EQ.0) GO TO 441
DO 351 IXSC=1,NBSC
CALL CONSTL(IXSC,X,Y)
351 CONTINUE
C
C *** AREA SOURCE COMPUTATION
441 IF(NBAR.EQ.0) GO TO 851
SIGN=-1.00
II=NBLN+1
IEND=NBLN+NBAR
DO 850 IXSC=11,IEVD
C
C CHECK WHETHER THE CIRCLE CONTRIBUTES TO SITE ACCELERATION
RIA=XXED(IXSC)
HLIL=DP(IXSC)
DIST=DSJRT((X-XXOR(IXSC))*(X-XXOR(IXSC))*(Y-YYOR(IXSC))*
1 (Y-YYOR(IXSC)))
C
C *** DETERMINE WHETHER CIRCLE WILL BE CONSIDERED
RPMX=3.7500+XNBTP(IXSC)*.2500
RH=DIST-RIA
IF(RH.LT.0.00) RH=0.00
RH=DSORT(RH+RH+HLIL*HLIL)
YGCK=(BL*DEXP(B2*RPMX))/(RH+84)**83
IF(YGCK.LT.YGMNCK) GO TO 850
C
C NR=RIA/DELTAC*0.900
DRI=DELTAC*.500
C
C ITERATION ON THE NB OF SEGMENTS IN THE CIRCLE
DO 202 IR=1,NR
DO 203 IS=1,2
DRI=SIGN*DRI
DLIL2=(DIST+DRI)*(DIST+DRI)

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WRITE(6,1041)HED2
WRITE(6,2120)
2120 FORMAT(' ',7X,'XOR',7X,'YOR',4X,'RADIUS',5X,'DEPTH',
1 5X,'OCCURRENCE TIME',5X,'TIME BASE',5X,'NB OF OCC')
WRITE(6,1027)XOR(IX),YOR(IX),XED(IX),DP(IX),TMDA(IX),
1 XNBDA(IX)
1027 FORMAT (' ',4F10.3,25X,2F10.3)
IXED=0
WRITE(6,1031)XNBTP(IX),XNBVE(IX)
DO 615 IAST=1,NB,10
IXED=IXED+10
IF(NB-IXST.LT.9)IXED=NB
WRITE(6,1025)(XMGST(I),I=IXST,IXED)
615 WRITE(6,1025)(XNBMG(IX,I),I=IAST,IXED)
C
C 551 CONTINUE
C
C INITIALIZE DISTRIBUTION ON ACCELERATIONS AND OTHER VARIABLES
CALL INITIA
DO 123 I=1,30
DO 123 J=1,20
123 PBZROC(I,J)=0.00
C
C COMPUTE DISTRIBUTION ON THE NUMBER OF EVENTS FOR EACH SOURCE
DO 100 IXSC=1,NBSC
CALL BERNUI(IXSC,NBLN,TH(I),TMDA(IXSC),XNBDA(IXSC),XNBVE(IXSC))
100 CONTINUE
C
C ITERATION ON THE NUMBER OF GRIDS
DO 365 IAGD=1,NBGD
WRITE(6,1041)HED2
READ(5,1000) XBEGIN,YBEGIN,XEND,YEND,DX,DY
IF(XEND.EQ.0.00) XEND=XBEGIN
IF(YEND.EQ.0.00) YEND=YBEGIN
IF(DX.EQ.0.00) DX=1.00
IF(DY.EQ.0.00) DY=1.00
NMAX=DABS((XBEGIN-YEND)/DX)+1.900
NYMAX=DABS((YBEGIN-XEND)/DY)+1.900
WRITE(6,1021)NBLN,NBAR,NKMAX,NYMAX,NBTH
1021 FORMAT(' ',5X,5X,'NYMAX=',15,5X,'NA = ',15,5X,
6'NXMAX = ',15,5X,'NYMAX=',15,5X,'NT=',15 / )
C
IF(SKSAVE.NE.0) GO TO 651
DO 650 IT=1,NBTH
IMRT=10+2*(IT-1)
IMRT1=IMRT+1
WRITE(IMRT,1040)HED1
WRITE(IMRT1,1040)HED1
WRITE(IMRT,5021)
5021 FORMAT('ACCELERATION DISTRIBUTION')
WRITE(IMRT1,5022)
5022 FORMAT('DURATION DISTRIBUTION')
WRITE(IMRT,5020)TH(IT),XBEGIN,XEND,YBEGIN,XEND,YEND,DX,DY
WRITE(IMRT1,5020)TH(IT),XBEGIN,XEND,YBEGIN,XEND,YEND,DX,DY

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AL=(R(A*KIA-DR)*DR1)
IF(AL-LE-0.00) GO TO 203
AL=DSQRTIAL)
RHITZ=DLIL2+HLIL*HLIL
COMPUTE EPICENTRAL DISTANCE RECK AT WHICH AN EARTHQUAKE
OF LARGEST MAGNITUDE RMXH WILL GENERATE AN ACCELERATION
YG(II)
C IF RECK IS IMAGINARY OR SMALLER THAN DLIL, THE SOURCE IS
C DISREGARDED. OTHERWISE POINTS WHERE THE CONTRIBUTION BECOMES
C ZERO ARE DETERMINED ON THE LINE. SEGMENTS BEYOND THESE POINTS
C ARE DISREGARDED.
C TERM=(B1*DEXP(LB2*RMXH)/YGMNCK)**(1.00/B3)
IF((TERM-B4).LE.-HLIL) GO TO 203
TERM=(TERM-B4)**(TERM-B4)-((HLIL*HLIL)
C
C THE EPICENTRAL DISTANCE IS NOT IMAGINARY
C IF(TERM.LE.-DLIL2) GO TO 203
C THE SEGMENT WILL BE PARTIALLY OR TOTALLY CONSIDERED
C
C DETERMINE INTERSECTION BETWEEN RMXH AND THE SEGMENT
C AL1=DSUKT(TERM-DLIL2)
IF(AL1.LT.ALJAL=AL1)
C
C DUMMY=0.00
CALL INTGCL(IJXG,AL,DUMMY,RHITZ)
C
C 203 CONTINUE
DR1=DABS(DR1)
DR1=DR1*DELTAC
C
C 202 CONTINUE
C
C 850 CONTINUE
851 CONTINUE
C
C *** COMPUTE CDF OF EXCEEDING AT LEAST ONE TIME
N1=100
N2=101
DO 442 I=1,99
N2=N2-1
ACPDF(N1)=ACPDF(N1)*ACPDF(N2)
DRPDF(N1)=DRPDF(N1)*DRPDF(N2)
ACPDF(N2)=1.00-ACPDF(N1)
DRPDF(N2)=1.00-DRPDF(N1)
DRPDF(I)=1.00-ACPDF(I)
WRITE(6,2301)
FORMAT('ACCELERATIONS')
2301 FORMAT('OACCELERATIONS')
IST=1
IED=10
XIC=.0200
ACST=-.0200
DO 400 IS=1,3
DO 410 IM=1,10
XIX=IM
ACUT(IM)=ACST+XIX*XIC
410 CONTINUE
WRITE(6,2300) ACUT
WRITE(6,2300)(ACPDF(I),I=1,IED)
FORMAT(' ',10F12.5)
IST=1ST*10

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IED=IED+10
ACST=ACUT(10)
400 CONTINUE
C
2302 FORMAT('DURATIONATIONS')
WRITE(6,2302)
FORMAT('OACCELERATIONS')
IST=1
IED=10
XIC=2.00
ACST=-.200
DO 420 IS=1,3
DO 430 IM=1,10
XIX=IM
ACUT(IM)=ACST+XIX*XIC
430 CONTINUE
WRITE(6,2300) ACUT
WRITE(6,2300)(DRPDF(I),I=1,IED)
IST=1ST*10
IED=IED+10
ACST=ACUT(10)
420 CONTINUE
IF(LKSAVE.EQ.1) GO TO 380
WRITE(1,MT,2303)(ACPDF(I),I=1,30)
WRITE(1,MT,2303)(DRPDF(I),I=1,30)
2303 FORMAT(10F8.5)
C
380 X=X+DX
385 Y=Y+DY
365 CONTINUE
C
RETURN
END
SUBROUTINE INITIA
C
C *****
IMPLICIT REAL*8 (A-H,O-Z)
COMMON/ATTEN/B1,B2,B3,B4,DELTA,DELTA,YGMN, YGMNCK,JPAC,UPDR
COMMON/ACV/PBACAV(10),BLEB2M(20),PBUC(1000),PBZROC(30,20)
I ,ACPDF(200),DRPDF(100),PBORAV(51),OR50(20)
COMMON/ACTN/IXACN,PBACAV(10),3),NBORAV(51,3)
*****
2000 WRITE(6,2000) UPAC,UPDR
FORMAT('UPPER LIMIT OF UNIFORM DISTRIBUTION',2F10.2)
C
C *** COMPUTE BAND & PROB OF JNIF. DISTRIBUTION FOR ACC
XIC=.0200
DO 100 IX=1,101
XIX=IX-1
ACAV=XIX*XIC
HIGH=ACAV*UPAC
IXACED=(50.*HIGH)*1.5
IXACST=IX+IX-IXACED
NBACAV(IX,1)=IX
NBACAV(IX,2)=IXACST
NBACAV(IX,3)=IXACED
XIXX=IXACED-IXACST*1
PBACAV(IX)=1.00/XIXX
WRITE(6,2010)X,ACAV,PBACAV(IX),(NBACAV(IX,I),I=1,3)

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C2010 FORMAT( ' ,15,2F14.5,3I10)
100 CONTINUE
C *** COMPUTE HAND & PROB OF UNIF DISTRIBUTION FOR DURATION
C XIC=2.00
DO 110 IX=1,51
XIX=IX-1
ACAV=XIX**IX
HIGH=ACAV*UPDR
IXACED=(1.50*HIGH)+1.51
IXACST=IX*IX-IXACED
NBORAV(IX,1)=IX
NBORAV(IX,2)=IXACST
NBORAV(IX,3)=IXACED
XIXX=IXACED-IXACST+1
C
C PBDRAW(IX)=1.00/AXIX
C WRITE(6,2010)IX,ACAV,PBDRAW(IX),(NBORAV(IX,1),1,1,3)
C2010 FORMAT( ' ,15,2F14.5,3I10)
110 CONTINUE
XXST=3.7500
XXIC=-2.500
DO 200 IX=1,20
XIX=IX
XIXM=XIX*(XIX**XIX)
BLEBZM(IX)=BL*DEXP(0.2*XIXM)
C
C WRITE(6,2020)IX,XIXM,BLEBZM(IX)
C2020 FORMAT( ' ,F10.2,E15.7)
200 CONTINUE
C
C IXACST=(50.00*YGMN)+1.500
DO 300 IX=2,30
IXACMN=IX
IF(NBACAV(IX,3).GT.IXACST) GO TO 310
300 CONTINUE
310 IXACMN=NBACAV(IXACMN-1,1)
C
C RETURN
C END
C SUBROUTINE BERNUI(IXSC,NBLN,THRK,TPR,XNPR,EVNB)
C
C *****
C IMPLICIT REAL*8 (A-H,I,P-Z)
C COMMON/ATTEN/BL,82,83,84,DELTA,DEL TAC,YGMN,YGMNCK,JPAC,UPDR
C COMMON/BERNUJ/ANBEV(30),ANBP(30),XNBMG(30,20)
C COMMON/ACVR/PBACAV(10),BLEBZM(20),PBOC(1000),PBZROC(30,20)
C ACPDF(200),DRPDF(100),PBDRAW(51),DR50(20)
C COMMON/GEOM/ XOK(30),YYOR(30),XXEU(30),YED(30),DP(30),
C TMDAL(30),XNBDA(30)
C DIMENSION PBMG(3,200),PBMGTT(3)
C LOGICAL*1 NONE
C *****
C
C1001 FORMAT(3F10.0,2I5)
C INITIALIZE VARIABLES
DO 101 IX=1,1000
101 PBOC(IX)=0.00
DO 102 IX=1,200
DO 102 I=1,3

```

```

C WRITE(6,2003)IXMG,PBZROC(IXSC,IXMG)
C PBZROC(IXSC,IXMG)=1.00-(PBZROC(IXSC,IXMG)+PBZROC(1))
C
C 200 CONTINUE
C 2003 FORMAT(' ',I5,7G15.7)
C WRITE(6,2004)PBZROC(IXSC,IXMG),IXMG=1,NBMG)
C 2004 FORMAT(10F12.6)
C
C RETURN
C END
C SUBROUTINE CONSTL(IL,X,Y)
C *****
C IMPLICIT REAL*8 (A-H,O-Z)
C INTEGER CASE
C COMMON/GEUM/ XXOR(30),YYOR(30),XXED(30),YYED(30),DP(30),
C 1 TMDA(30),XNDA(30)
C COMMON/ATTEN/BI,B2,B3,B4,DELTA,DELTA2,DELTA3,DELTA4,DELTA5,DELTA6,DELTA7,DELTA8,DELTA9,DELTA10,DELTA11,DELTA12,DELTA13,DELTA14,DELTA15,DELTA16,DELTA17,DELTA18,DELTA19,DELTA20,DELTA21,DELTA22,DELTA23,DELTA24,DELTA25,DELTA26,DELTA27,DELTA28,DELTA29,DELTA30,DELTA31,DELTA32,DELTA33,DELTA34,DELTA35,DELTA36,DELTA37,DELTA38,DELTA39,DELTA40,DELTA41,DELTA42,DELTA43,DELTA44,DELTA45,DELTA46,DELTA47,DELTA48,DELTA49,DELTA50,DELTA51,DELTA52,DELTA53,DELTA54,DELTA55,DELTA56,DELTA57,DELTA58,DELTA59,DELTA60,DELTA61,DELTA62,DELTA63,DELTA64,DELTA65,DELTA66,DELTA67,DELTA68,DELTA69,DELTA70,DELTA71,DELTA72,DELTA73,DELTA74,DELTA75,DELTA76,DELTA77,DELTA78,DELTA79,DELTA80,DELTA81,DELTA82,DELTA83,DELTA84,DELTA85,DELTA86,DELTA87,DELTA88,DELTA89,DELTA90,DELTA91,DELTA92,DELTA93,DELTA94,DELTA95,DELTA96,DELTA97,DELTA98,DELTA99,DELTA100
C *****
C AIL = SLOPE OF LINE DEFINED BY (XL,Y1),(X2,Y2)
C BIL = Y-INTERSEPT
C
C *** SET GEOMETRY LINE PARAMETERS
C XLI=XXOR(IL)
C YLI=YYOR(IL)
C XL2=XXED(IL)
C YL2=YYED(IL)
C DPIL=DP(IL)
C AIL=(YL2-YLI)/(XL2-XLI)
C BIL=YLI-AIL*XLI
C XBAR=(X+AIL*Y-BIL*AIL)/(AIL*AIL+1.000)
C YBAR=(AIL*X+Y+AIL*BIL)/(AIL*AIL+1.000)
C DLIL=DSQRT(1-X-ABAR)*(X-ABAR)+(Y-YBAR)*(Y-YBAR)
C RHITZ=DLIL*DLIL*DPIL*DPIL
C
C RMMX=3.7500*XNBTP(IL)*.2500
C *** THERE IS NO CUTOFF MAGNITUDE FOR THE SOURCE, SKIP FOLLOWING ANALYSIS
C
C COMPUTE EPICENTRAL DISTANCE RECK AT WHICH AN EARTHQUAKE
C OF CUTOFF MAGNITUDE RMMX WILL GENERATE AN ACCELERATION
C YG(II)
C IF RECK IS IMAGINARY OR SMALLER THAN DLIL, THE SOURCE IS
C DISREGARDED. OTHERWISE POINTS WHERE THE CONTRIBUTION BECOMES
C ZERO ARE DETERMINED ON THE LINE. SEGMENTS BEYOND THESE POINTS
C ARE DISREGARDED.
C TERM=(B1+DEXP(B2*RMMX)/YGMNCK)*L.00/B3)
C IF(TERM-B4)-LE-DPIL) GO TO 830
C TERM=(TERM-B4)*(TERM-B4)-(DPIL*DPIL)
C
C THE EPICENTRAL DISTANCE IS NOT IMAGINARY
C RECK=DSQRT(TERM)
C IF(RECK.LE.DLIL) GO TO 830
C
C THE EPICENTRAL DISTANCE IS NOT SMALLER THAN DLIL, POINTS OF
C INTERSECTION WITH THE LINE ARE COMPUTED
C TERM=AIL*(BIL-Y)-X
C TERM2=X*X+BIL*BIL+Y*(Y-Z.00*BIL)-RECK*RECK

```

```

DO 200 IXMG=1,NBMG
XXXI=XNBMG(IXMG)
C
C PBCI=G(ETA)/G(XI)*G(ETA-XI)
C PBCI=DLGAMA(ETA)-DLGAMA(XXXI)-DLGAMA(ETA-XXXI)
C
C *** PROB OF ZERO SUCCESS GIVEN N OCCUR.
C P(O/N)=EXP(PBCT)*G(XI)*G(N+ETA-XI)/G(N+ETA)
C *** PROB OF K SUCCESSSES GIVEN N OCCUR.
C P(K/N)=CNR*EXP(PBCT)*G(XI)*G(N+ETA-R-XI)/G(N+ETA)
C
C NBOCI=NBOC-I
C DO 300 N=1,XST,NBOCI
C XTETT=0.00
C XIN=N
C
C *** PROB OF ZERO SUCCESS
C XNPR=XN+ETA-XXXI
C DEN=DLGAMA(XN+ETA)
C PROB=PBCT+DLGAMA(XXXI)+DLGAMA(XNPR)-DEN
C IF(PROB.LT.10.00) GO TO 300
C PBZROC(IXSC,IXMG)=PBZROC(IXSC,IXMG)+DEXP(PROB)*PBOC(N+1)
C PBMG(IXMG,1)=PBMG(IXMG,1)+DEXP(PROB)*PBOC(N+1)
C XTETT=DEXP(PROB)
C XTETT=XTETT+XTETT
C PBMGTT(IXMG)=PBMGTT(IXMG)+DEXP(PROB)*PBOC(N+1)
C GO TO 300
C
C *** DO LOOP ON THE NUMBER OF SUCCESSSES
C 300 TERM=CNR
C *** PROB OF I SUCCESSSES GIVEN N OCCUR.
C 305 TERM=XN
C NONE=.TRUE.
C PROB=DLGAMA(TERM)+PBCT+DLGAMA(1+XXXI)+DLGAMA(XNPR-I)-DEN
C IF(PROB.LT.10.00) GO TO 312
C PBMG(IXMG,2)=PBMG(IXMG,2)+DEXP(PROB)*PBOC(N+1)
C PBMGTT(IXMG)=PBMGTT(IXMG)+DEXP(PROB)*PBOC(N+1)
C XTETT=DEXP(PROB)
C XTETT=XTETT+XTETT
C 312 IF(N.EQ.1)GO TO 300
C *** PROB OF 2,3,...N SUCCESSSES GIVEN N OCCUR.
C DO 310 IR=2,N
C XI=N-IR+1
C XIR=IR
C TERM=TERM*XI/XIR
C DLTER=DLG(TERM)
C PROB=DLTER+PBCT+DLGAMA(1+XXXI)+DLGAMA(XNPR-XIR)-DEN
C IF(PROB.LT.10.00) GO TO 311
C PBMG(IXMG,3)=PBMG(IXMG,3)+DEXP(PROB)*PBOC(N+1)
C PBMGTT(IXMG)=PBMGTT(IXMG)+DEXP(PROB)*PBOC(N+1)
C XTETT=DEXP(PROB)
C XTETT=XTETT+XTETT
C NONE=.FALSE.
C WRITE(6,2003)IK,PBMG(IXMG,IR+1)
C GO TO 310
C 311 IF(.NOT.NONE) GO TO 300
C 310 CONTINUE
C 300 CONTINUE
C
C PBT=0.00
C DO 313 IX=1,NBOC
C PBT=PBT+PBMG(IXMG,IX)
C 313 WRITE(6,2003)IX,PBMG(IXMG,IX)

```

```

C C TERM3=AL*AL*1.00
C C TERM2=DSQRT(TERM1*TERM1-TERM2*TERM2)
C C COORDINATES OF THE POINTS OF INTERSECTION
C C X11=(1-TERM1+TERM2)/TERM3
C C Y11=AL*X11+B1L
C C IF(X11.LT.X12) GO TO 150
C C STOR=X11
C C X11=X12
C C X12=STOR
C C STOR=Y11
C C Y11=Y12
C C Y12=STOR
C C 150 CONTINUE
C C CHECK WHETHER THE POINTS OF INTERSECTION ARE ON THE SEGMENT
C C CONSIDERED AND DETERMINE THE USE OF ALPHA AND BETA
C C IF(X11.LT.X11) GO TO 800
C C IF(X11.GT.X12) GO TO 830
C C IF(X12.GT.X12) GO TO 820
C C CASE=1
C C *** SOURCE TOTALLY INCLUDED
C C GO TO 850
C C 800 IF(X12.LT.X11) GO TO 830
C C IF(X12.LT.X12) GO TO 840
C C CASE=5
C C *** SOURCE CUTOFF ON LEFT AND RIGHT ENDS
C C XL2=X12
C C YL2=Y12
C C GO TO 860
C C 830 CASE=3
C C *** SOURCE TOTALLY DISREGARDED
C C RETURN
C C 820 CASE=2
C C *** SOURCE CUTOFF ON THE RIGHT END
C C XL2=X12
C C YL2=Y12
C C GOTO 850
C C 840 CASE=4
C C *** SOURCE CUTOFF ON THE LEFT END
C C X11=X11
C C Y11=Y11
C C 850 CONTINUE
C C CHECK POSITION OF (XBAR,YBAR) ALONG LINE W.R.T. END
C C OF POINTS OF LINE SOURCE
C C AL=DSQRT((X11-XBAR)*(X11-XBAR)+(Y11-YBAR)*(Y11-YBAR))
C C AL2=DSQRT((X12-XBAR)*(X12-XBAR)+(Y12-YBAR)*(Y12-YBAR))
C C XLEN=DSQRT((AL2-AL)*(AL2-AL)+(YL2-YL1)*(YL2-YL1))
C C IF(XBAR.GT.X11) AL1=-AL1

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C C IF(XBAR.GT.X12) AL2=-AL2
C C CALL INTGLN(IL,XLEN,AL1,RHITZ)
C C RETURN
C C SUBROUTINE INTGLN(IASC,XLEN,AL,RHITZ)
C C *****
C C IMPLICIT REAL*8(A-H,O-7)
C C COMMON/ALVR/PBACAV(10),BLEZM(20),PB0C(100),PBZROC(30,20)
C C 1,ACPDF(20),DRPDF(10),PBDKAV(5),DR50(20)
C C COMMON/ACLN/IXACHN,NBACAV(10),J,NBDRAV(5),3
C C COMMON/ATTEN/B1,B2,B3,B*,DELTA,DELTA*,YGMN,YGMNCK,JPAC,JPDR
C C COMMON/BRNUJ/XNBEV(30),XNBTP(30),XNBMG(30,20)
C C DIMENSION ACUT(10)
C C *****
C C NBMG=XNBTP(IASC)
C C NB1C=(XLEN/DELTA)+.500
C C IF(NB1C.EQ.0) GO TO 110
C C START DO LOOP AT 1/2 DELTA
C C AL=AL+.500*DELTA
C C DO 100 IXIC=1,NB1C
C C RH=RHITZ+AL*AL
C C RHRT=DSQRT(RH)
C C RH=(RHRT+B*)*(RHRT+B*)
C C RHRT=RHRT*108.8800
C C DO LOOP ON ALL POSSIBLE MAGNITUDES
C C IXMG=NBMG+1
C C DO 200 IX=1,NBMG
C C IXMG=IXMG-1
C C AC=B1E2MH(IXMG)/RH
C C IXAC=(50.00*AC)+1.500
C C PRINT,'IXAC,IXMG',AC,IXAC,IXMG
C C IF(IXAC.LI.IXACMN) GO TO 100
C C IXACPB=NBALAV(IXAC,1)
C C IXACST=NBACAV(IXAC,2)
C C IXACED=NBACAV(IXAC,3)
C C PBZR=PBZROC(IXAC,IXMG)
C C PROB=1.00-(PBACAV(IXACPB)*PBZR)
C C PRINT,'300',IXACPB,IXACST,PROB,PBACAV(IXACPB),PBZROC(IXAC,IXMG)
C C DO 300 IXACVR=IXACST,IXACED
C C ACPDF(IXACVR)=ACPDF(IXACVR)*PROB
C C 300 CONTINUE
C C *** DURATIONS USING THE SAME VARIABLES AS ACC
C C AC=DR50(IXMG)*.1500*(50.00-RHRT)
C C IF(RHRT.GT.50.00) AC=DR50(IXMG)/(1.00+12.00*(RHRT-50.00))*
C C 1 (RHRT-50.00)+1.00)
C C IXAC=AC*.500+.15100
C C IXACPB=NBDRAV(IXAC,1)
C C IXACST=NBDRAV(IXAC,2)
C C IXACED=NBDRAV(IXAC,3)
C C PROB=1.00-(PBDRAV(IXACPB)*PBZR)
C C DO 310 IXACVR=IXACST,IXACED

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DRPDF(IACVR)=DRPDF(IACVK)*PRUB
310 CONTINUE
C
200 CONTINUE
100 AL=AL+DELTA
DO 500 IX=1,100
IF(ACPDF(IX).LT.1.0-5) ACPDF(IX)=0.00
500 CONTINUE
C
110 CONTINUE
IST=1
IED=10
XIC=2.00
ACST=-2.00
DO 400 IS=1,3
DO 410 IX=1,10
XIX=IX
ACUT(IX)=ACST+XIX*XIC
410 CONTINUE
WRITE(6,1000) DRPDF(IX),I=IST,IED
WRITE(6,1000) DRPDF(IX),I=IST,IED
C1000 FORMAT(10F12.5)
IST=IST+10
IED=IED+10
ACST=ACUT(10)
C 400 CONTINUE
RETURN
END
SUBROUTINE INTGCL(IASC,XLEN,AL,RHITZ)
C *****
IMPLICIT REAL*8(A-H,O-Z)
COMMON/ACVR/PBACAV(101),PBEBZM(20),PBRC(1000),PBZRC(10,20)
1 ,ACPDF(200),DRPDF(100),PBDRAV(51),DR50(20)
COMMON/ACIN/IAACMN,NBACAV(101,3),NBORAV(51,3)
COMMON/ATTEN/81,82,83,84,DELTA,DELTA2,DELTA3,DELTA4,DELTA5,DELTA6,DELTA7,DELTA8,DELTA9,DELTA10
COMMON/BKNUJ/XNBEV(30),XNBT(30),XNBMG(30,20)
DIMENSION ACUT(10)
C *****
NBMG=XNBT(IASC)
NBIC=(XLEN/DELTA)+500
IF(NBIC.EQ.0) GO TO 110
C
C START DO LOOP AT 1/2 DELTA
AL=AL+.500*DELTA
C
DO 100 IXIC=1,NBIC
RH=RHITZ+AL*AL
RHPT=DSURT(RH)
RH=(RHRT+84)*(RHRT+84)
RHRT=RHRT+108.8800
C
DO LOOP ON ALL POSSIBLE MAGNITUDES
IXMG=NBMG+1
DO 200 IX=1,NBMG
IXMG=IXMG-1
C *** ACCELERATIONS
AC=81EBZM(IASC)/RH

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IXAC=(50.00*AC)+1.500
PRINT,IXAC,IXMG,AC,IXAC,IXMG
IF(IXAC.LT.IXACHN) GO TO 100
IXACPB=NBACAV(IXAC,1)
IXACST=NBACAV(IXAC,2)
IXACED=NBACAV(IXAC,3)
PBZR=PBZRC(IASC,IXMG)
PROB=1.00-(PBACAV(IXACPB)*PBZR)
PRINT,300,IXACPB,IXACST,PROB,PBACAV(IXACPB),PBZRC(IASC,IXMG)
DO 300 IXACVR=IXACST,IXACED
ACPDF(IXACVR)=ACPDF(IACVR)*PROB*PROB
300 CONTINUE
C
C *** DURATIONS USING THE SAME VARIABLES AS ACC
AC=DR50(IXMG)+.1500*(50.00-RHRT)
IF(RHRT.GT.50.00) AC=DR50(IXMG)/(1.0011200*(RHRT-50.00))*
1 (RHRT-50.00)+1.00)
IXAC=AC*.500 + 1.5100
IXACPB=NBORAV(IXAC,1)
IXACST=NBORAV(IXAC,2)
IXACED=NBORAV(IXAC,3)
PROB=1.00-(PBORAV(IXACPB)*PBZR)
C
DO 310 IXACVR=IXACST,IXACED
DRPDF(IXACVR)=DRPDF(IACVR)*PROB*PROB
310 CONTINUE
C
200 CONTINUE
100 AL=AL+DELTA
C
DO 500 IX=1,100
IF(ACPDF(IX).LT.1.0-5) ACPDF(IX)=0.00
IF(DRPDF(IX).LT.1.0-5) DRPDF(IX)=0.00
500 CONTINUE
C
110 CONTINUE
IST=1
IED=10
XIC=2.00
ACST=-2.00
DO 400 IS=1,3
DO 410 IX=1,10
XIX=IX
ACUT(IX)=ACST+XIX*XIC
C 410 CONTINUE
WRITE(6,1000) DRPDF(IX),I=IST,IED)
C1000 FORMAT(10F12.5)
IST=IST+10
IED=IED+10
ACST=ACUT(10)
C 400 CONTINUE
RETURN
END
//GO,FT10F001 DD UNIT=DISK,DCB=CARU,DSN=WL.EI.G35.AC300,
// DISP=(NEW,KEEP),SPACE=(3120,(20,5),RLSE),VOL=SER=PUB005
//GO,FT11F001 DD UNIT=DISK,DCB=CARU,DSN=WL.EI.G35.DG300,
// DISP=(NEW,KEEP),SPACE=(3120,(20,5),RLSE),VOL=SER=PUB005
//GO.SYSIN DD *

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*****STANFORD UNIVERSITY*****
PROGRAM PLOT-ISO
THIS PROGRAM PLOTS THE ISO ACCELERATION OR ISO INTENSITY
LINES OVER A GRID (MAP) GIVEN THE ACCELERATIONS AT EACH NODE OF THE
GRID. THE ORIGIN OF THE GRID IS THE BOTTOM LEFT CORNER, THE END
THE TOP RIGHT CORNER. THE VALUES HAVE TO BE INPUT FILING COLUMNS
FROM BOTTOM TO TOP AS ONE MOVES FROM LEFT TO RIGHT: FIRST VALUE
INPUT IS THUS THE ORIGIN AND LAST THE END.
THE PLOT WILL BE ROTATED 90 DEGREES COUNTERCLOCKWISE IF IT IS TOO
HIGH TO FIT ON THE PLOTTER. IF BOTH DIMENSIONS ARE TOO BIG, AN
MESSAGE WILL BE GENERATED AND THE PLOT ABORTED.
*****INPUT FORMAT*****
*****
COL VARIABLE NAME VARIABLE DESCRIPTION
1- IDENTIFICATION CARD 1 CARD (80AL)
1-80 HED RJN IDENTIFICATION
2- PLOT FLAGS 1 CARD (515,6F10.0)
1- 5 SKIPAC =0, WILL PLOT ISO ACCELERATION LINES
=1, WILL PLOT ISO INTENSITY LINES
5-10 NF NB OF FALUTS TO BE PLOTTED
11-15 NBPL NB OF ROWS IN INPUT
16-20 NR NB OF COLUMNS IN INPUT
21-25 NC DISTANCE (DEGREE) BETWEEN TWO COLUMNS
26-35 DX DISTANCE (DEGREE) BETWEEN TWO ROWS
36-45 DY INTERVAL BETWEEN TWO ISO LINES
46-55 DC MULTIPLE OF CONTOUR TO BE LABELED IIZ
56-65 XMDC FORMAT) FOR EXAMPLE IF XMDC=2,
EVERY OTHER DC WILL BE LABELED
3- GRID PARAMETES 1 CARD ((8F10.0)
THE ORIGIN OF THE GRID IS DEFINED AS
THE BOTTOM LEFT CORNER; THE END AS THE
TOP RIGHT. THE VERTICAL HAS TO POINT
NORTH FOR THE DATA TO BE READ IN
CONSISTENTLY.
1-10 XMINI X COORD OF ORIGIN
11-20 YMINI Y COORD OF ORIGIN
21-30 XMAXI X COORD OF END
31-40 YMAXI Y COORD OF END
41-50 INDGX LENGTH OF ONE DEGREE IN INCHES FOR
THE PLOT (X DIRECTION).
51-60 INDGY LENGTH OF ONE DEGREE IN INCHES FOR
THE PLOT (Y DIRECTION)
61-70 CAL PLOTTER TO BE USED (11, OR 33.)
IF ZERO, WILL BE SET TO 11.
71-80 DIVR THE FRAME COORDINATES ARE ROUNDED OFF
TO THE CLOSEST (DEGREE/DIVR)
IF INPUT AS 0, WILL BE SET TO 1
3- FAULT COORDINATES NF CORDS (4F10.0) NO INPUT IF NF=0
1-10 YF11( ) Y COORD OF ORIGIN OF FAULT 1
11-20 XF11( ) X COORD OF ORIGIN OF FAULT 1
21-30 YF21( ) Y COORD OF END OF FAULT 1
*****
C 31-40 XF21( ) X COORD OF END OF FAULT 2
REPEAT NF TIMES
C C
C 4- ACCELERATION AT THE NJDES 8 VALUES PER CARD (8F10.0)
1-10 ACC(1,IC) ACC AT ORIGIN (IC=1)
... ACC(1R,IC)
... ACC(NR,IC) ACC AT TOP OF FIRST COLUMN (IC=1)
REPEAT NC TIMES
C C
C OUTPUT
*****
C A FILE IS CREATED ON DISK, IT HAS TO BE TRANSFERRED ON TAPE AND
PLOTTED ON A 11 OR 33 INCH CALCOMP PLOTTER
*****COMMON /PLT7/ CONTX(100),CONTY(100),DX,DY,ERROR,DFPCLY
DIMENSION AMODES(200),END1(2),END2(2),XX(1000),YY(1000),A(1000)
1,AA(40,40),XF1(20),XF2(20),YF1(20),YF2(20),MED1(20)
INTEGER*4 NR,NC,ROTA,SKIPAC
REAL RASIN,INDG,INDGX,INDGY
LOGICAL*1 HED2(80)
EQUIVALENCE(A(1),AA(1,1))
*****
C C
C 1001 READ(5,1001)HED1
FORMAT(20A4)
WRITE(6,2202)HED1
2202 FORMAT(' ',20A4)
READ(5,1003)SKIPAC,NF,NBPL,NR,NC,DX,DY,DC,XMDC
READ(5,1000)XMIN,YMIN,XMAX,YMAX,INDGX,INDGY,CAL,DIVR
IF(DIVR.EQ.0.0) DIVR=1.
IF(CAL.EQ.0.0) CAL =11.
PRINT,XMIN,YMIN,XMAX,YMAX,INDGX,INDGY,CAL,DIVR
PRINT,NBPL,NR,NC,DX,DY,DC
C C
C 1000 FORMAT(8F10.0)
1003 FORMAT(5I5,6F10.0)
RASIN=100.
ICAL=1
IF(CAL.EQ.33.) RASIN=200.
IF(CAL.EQ.33.) ICAL=2
C C
C DIVR IS USED TO TO BE ABLE TO PLOT PARTS OF DEGREES, USALLY DEGREE/10
XMIN=IFIX(XMIN*DIVR+.5)
YMIN=IFIX(YMIN*DIVR+.5)
XMAX=IFIX(XMAX*DIVR+.5)
YMAX=IFIX(YMAX*DIVR+.5)
INDGX=INDGX/DIVR
INDGY=INDGY/DIVR
DX=DX*DIVR
DY=DY*DIVR
XNAS=(ABS(XMAX-XMIN)*INDGX+1.5)*RASIN
YNAS=(ABS(YMAX-YMIN)*INDGY+1.5)*RASIN
ROTA=0
C C
C ORIENTATE FRAME
DELTAX=-.75/INDGX
DELTAY=-.75/INDGY
IF(YNAS.LE.-(ICAL*RASIN)) GO TO 100
IF(XNAS.LE.-(ICAL*RASIN)) GO TO 110
WRITE(6,2100)

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2100 FORMAT('THE SCALE IS TOO LARGE FOR THE PLOTTER, REDUCE INDO PER U
         DEGREE')
STOP
C
C INTERCHANGE X AND Y SCALES
110 WRITE(6,2101)
2101 FORMAT('THE PLOT WILL BE ROTATED BY 90 DEGREES COUNTERCLOCKWISE')
ROTA=1
STOR=DX
DX=DY
DY=STOR
STOR=DELTA X
DELTA X=DELTA Y
DELTA Y=STOR
STOR=XMIN
XMIN=YMAX
YMAX=XMAX
XMAX=YMIN
YMIN=STOR
STOR=XRAS
XRAS=YRAS+3.5*RASIN
YRAS=STOR
IF(INF.EQ.0) GO TO 144
C
C READ FAULTS
READ(5,1007)(YF1(I),XF1(I),YF2(I),XF2(I),I=1,NF)
1007 FORMAT(4F10.0)
C
C GO TO 130
100 XRAS=XRAS+3.5*RASIN
C
C IF(INF.EQ.0) GO TO 144
READ(5,1007)(XF1(I),YF1(I),XF2(I),YF2(I),I=1,NF)
C
C 130 DO 143 I=1,NF
XF1(I)=XF1(I)*DIVR
YF1(I)=YF1(I)*DIVR
XF2(I)=XF2(I)*DIVR
YF2(I)=YF2(I)*DIVR
143 YF2(I)=YF2(I)*DIVR
144 CONTINUE
C
C CHECK SIGNS
IF(XMAX.GT.XMIN) GO TO 150
XMAX=-XMAX
XMIN=-XMIN
IF(INF.EQ.0) GO TO 150
DO 145 I=1,NF
XF1(I)=-XF1(I)
XF2(I)=-XF2(I)
145 XF2(I)=-XF2(I)
C
150 XFMAX=XMAX+DELTA X*.25/.75
XFMIN=XMIN-DELTA X
IF(YMAX.GT.YMIN) GO TO 170
YMAX=-YMAX
YMIN=-YMIN
IF(INF.EQ.0) GO TO 170
DO 146 I=1,NF
YF1(I)=-YF1(I)
YF2(I)=-YF2(I)
146 YF2(I)=-YF2(I)
C
170 YFMAX=YMAX+DELTA Y
YFMIN=YFMIN-DELTA Y
C
C FIND INTERSECTION BETWEEN FAULTS AND FRAME
IF(INF.NE.0)CALL FRAME(NF,XMIN,YMIN,XMAX,YMAX,XF1,YF1,XF2,YF2)
C
C CALL STARTG(AMODES,ICAL,J,'LJ5BLK','LJ5BLK','LJ5BLK','LJ5BLK')
AMODES(97)=XRAS
AMODES(179)=YRAS
IF(ICAL.EQ.33.) AMODES(45)=AMODES(45)*1.5
CALL OBJEGG(AMODES,0.,0.,XRAS,YRAS)
IF(ROTA.EQ.0) GO TO 610
NSTOR=NC
NC=NR
NR=NSTOR
610 CONTINUE
C
DO 600 IP=L,NBPL
CALL READ(AA,HED2,NR,NC,ROTA)
C
C FIND LOWER AND UPPER CONTOUR VALUES
NBVL=NR*NC
CALL LIMITS(A,NBVL,SKIPAC,CMIN,DC,CMAX)
C
AMODES(50)=0.
AMODES(46)=0.
CALL SUBJEG(AMODES,XFMIN,YFMIN,XFMAX,YFMAX)
C
WRITE(6,2102)AMODES(12),AMODES(13)
2102 FORMAT('THE SCALES ARE ',2E16.5)
C
C PLOT GRID
END1(1)=IFIX(YMIN)
END2(1)=IFIX(XMAX)
END1(2)=IFIX(YMAX)
END2(2)=IFIX(XMIN)
C
DELTA2=1./18.*INDGX
XMUL=4.
DELTA1=4.*DELTA2
DELTA3=3.*DELTA2
DELTA4=1.5*AMODES(40)*AMODES(45)/(RASIN*INDGX)
DELTA5=2.0*DELTA4
SIGN=-1.
C
I=1
VAR=END1(2)
FIX=END2(2)
DO 400 K=1,2
C
401 XX(I)=FIX
YY(I)=VAR
XX(I+1)=FIX+DELTA2
YY(I+1)=VAR
XX(I+2)=FIX
YY(I+2)=VAR
XVAR=ABS(VAR/U|VR)
INTG=XVAR*.01

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```

XC=FIX-DELTA2*XMUL
YC=VAR
C
IF(IFIX(XVAR*.91).EQ.INTG)CALL INUMBG(AHODES,XC,YC,2,INTG)
I=I+3
IF(VAR.EQ.END1(K)) GO TO 402
VAR=VAR+SIGN
GO TO 401
C
402 STOR=VAR
VAR=FIX
FIX=STOR
XX(I)=VAR
YY(I)=FIX
XX(I+1)=VAR
YY(I+1)=FIX+DELTA2
XX(I+2)=VAR
YY(I+2)=FIX
XVAR=ABS(VAR/DIVR)
INTG=XVAR*.01
XC=VAR-DELTA4
YC=FIX-DELTA3
C
IF(IFIX(XVAR*.91).EQ.INTG)CALL INUMBG(AHODES,XC,YC,3,INTG)
I=I+3
IF(VAR.EQ.END2(K)) GO TO 403
VAR=VAR-SIGN
GO TO 404
C
403 SIGN=-SIGN
STOR=VAR
VAR=FIX
FIX=STOR
DELTA2=-DELTA2
XMUL=2.
DELTA3=-DELTA3
C
400 CONTINUE
DO 500 K=1,I
500 WRITE(6,550)XX(K),YY(K)
550 FORMAT(' ',2F10.3,16,F10.3)
I=I-3
CALL LINESG(AHODES,I,XX,YY)
C
C PLOT TITLE
XT=XFMAX-3.*DELTA2
AHODES(50)=90.
AHODES(46)=90.
YT=YMIN+DELTA1
CALL VEC SG(AHODES,XT,YT,48,MED2)
C
IF(INF.EQ.0) GO TO 142
PLOT FAULTS
C
PRINT,'FAULTS AS THEY ARE PLOTTED'
WRITE(6,2011)I,AF1(I),YF1(I),AF2(I),YF2(I),I=1,NF)
2011 FORMAT(' ',15,4F10.3)
CALL SEGMTG(AHODES,NF,AF1,YF1,AF2,YF2)
142 CONTINUE
C
C CHANGE SUBJECT SPACE TO PLOT COUNTERS, THE BOTTOM LEFT CORNER

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C HAS TO HAVE (0.,0.,0.) COORDINATES
XFMAX1=ABS(XFMAX-XFMIN)-DELTA2
YFMAX1=ABS(YFMAX-YFMIN)-DELTA2
PRINT,'DELTA, DELTAY, XFMAX, YFMAX, DELTAX, DELTAY, XFMAX, YFMAX'
XFMIN1=-DELTA2
YFMIN1=-DELTA2
C
PRINT,'RASOR,XRAS,YRAS',RASOR,XRAS,YRAS
CALL SUBJEG(AHODES,XFMIN1,YFMIN1,XFMAX1,YFMAX1)
PRINT,'SCALE',AHODES(12),AHODES(13)
C
CREATE COUNTERS AND PLOT THEM
AHODES(45)=AHODES(45)*.6
CALL CONTUR(AHODES,AA,NR,NC,XMDC,DELTA5)
AHODES(45)=AHODES(45)/.6
C
CALL PICTRG(AHODES)
600 CONTINUE
C
CALL EXITG(AHODES)
RETURN
END
SUBROUTINE READ1AA,MED2,NROW,NCOL,NROTT)
*****
DIMENSION AA(NROW,NCOL)
LOGICAL *1 MED2(80)
*****
THIS SUBROUTINE READS THE DATA TO BE INTERPOLATED
READ(5,1002)MED2
1002 FORMAT(80A1)
C
IF(NROTT.EQ.1) GO TO 110
C
THE DATA IS STORED IN A TWO DIMENSIONAL ARRAY, ALL THE DATA BEING AT
THE ORIGIN OF THE PLOT (BOTTOM LEFT CORNER)
C
DO 100 IR=1,NROW
100 READ(5,1000)(AA(IR,IC),IC=1,NCOL)
GO TO 99
C
THE PLOT HAS BEEN ROTATED BY 90 DEGREES COUNTERCLOCKWISE,
THE ROWS AND COLUMNS ARE INVERTED IN THE CALL STATEMENT
110 CONTINUE
DO 200 I=1,NCOL
IC=NCOL+1-I
200 READ(5,1000)(AA(IR,IC),IR=1,NROW)
59 CONTINUE
C
DO 300 IR=1,NROW
300 WRITE(6,1001)(AA(IR,IC),IC=1,NCOL)
1001 FORMAT(' ',13F10.3)
1000 FORMAT(8F10.5)
C
RETURN
END
SUBROUTINE LIMITS(A,NBVL,SKIPAC,CMIN,JC,CMAX)
*****
DIMENSION A(1)

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C *****:
C INTEGER SKIPAC
C *****:
C SEARCH FOR LOW BOUND
C IF(SKIPAC.NE.1) GO TO 110
C
C FIND INTENSITIES USING RICHTER GUTENBERG RELATIONSHIP
C DO 99 IV =1,NBVL
C A(IIV)=3.*(ALOG10(A(IIV)*1000.))+.5)
C 99 CONTINUE
C
C 110 CONTINUE
C XLOW=A(I)
C DO 100 IY=2,NBVL
C IF(A(IY).LT.XLOW) XLOW=A(IY)
C 100 CONTINUE
C
C SEARCH FOR UPPER BOUND
C XHIGH=A(I)
C DO 200 IY=2,NBVL
C IF(A(IY).GT.XHIGH) XHIGH=A(IY)
C 200 CONTINUE
C
C FIND LOW AND HIGH CONTOUR LINES TO BE PLOTTED
C CMIN=-DC
C CMIN=CMIN+DC
C IF(CMIN.LT.XLOW) GO TO 300
C
C CMAX=CMIN-DC
C CMAX=CMAX+DC
C IF(CMAX.LT.XHIGH) GO TO 400
C CMAX=CMAX-DC
C
C PRINT,XLOW,XHIGH,CMIN,CMAX,XLOW,XHIGH,CMIN,CMAX
C
C RETURN
C END
C SUBROUTINE FRAME(NF,XMIN,YMIN,XMAX,YMAX,XF1,YF1,XF2,YF2)
C *****:
C DIMENSION XF1(1),XF2(1),YF1(1),YF2(1)
C *****:
C *****:
C LIMIT THE FAULTS TO THE AREA UNDER CONSIDERATION
C DO 500 IF=1,NF
C
C CHECK THE X COORDINATES OF THE FAULT
C SIGN=1.
C CHECK=XMIN
C
C CHECK FOR LOWER AND UPPER BORN
C DO 100 K=1,2
C XF=XF1(IF)*SIGN
C
C CHECK FOR ORIGIN AND END OF SEGMENT
C DO 200 I=1,2
C IF(XF-GE-CHECK) GO TO 200
C ASL=(YF1(IF)-YF2(IF))/(XF1(IF)-XF2(IF))
C BET=YF1(IF)-ASL*XF1(IF)
C IF(I.EQ.2) GO TO 310
C
C CHANGE COORDINATES
C YF1(IF)=SIGN*CHECK
C XF1(IF)=(YF1(IF)-BET)/ASL
C GO TO 210
C 310 YF2(IF)=SIGN*CHECK
C XF2(IF)=(YF2(IF)-BET)/ASL
C 210 YF=YF2(IF)*SIGN
C SIGN=-1.
C 110 CHECK=-YMAX
C
C 500 CONTINUE
C RETURN
C END
C SUBROUTINE CONTOUR(AMODES,FAY,IRDM,ICDM,CTIC,XMDC,DELTA5)
C *****:
C PASS A SET OF MESH POINTS FXY(IRDM,ICDM) WITH INCREMENTS
C OF DX AND DY FOR AN AREA OF SEARCH FROM, ICDM WITH
C ORIGIN AT THE BOTTOM LEFT CORNER. FIND A SPECIFIC
C CONTOUR LEVEL CTLY
C
C DIMENSION FXY(IRDM,ICDM),AMODES(200)
C COMMON /PLTT/ CONTX(100),CONTY(100),DX,DY,ERROR,DEPOLY
C COMMON NBPT,NBPTSQ,ISIDE(4),IORDER,NBPTIT,NBSQFR,
C 1 ISQRW(100),ISCOL(100),IXCYST(20),IXSQBD(20)
C LOGICAL*1 OPEN
C
C SEARCH GRID FOR SMALLEST AND LARGEST VALUE
C SET VALUES OF CONTOURS TO BE PLOTTED
C PTLVAD=XMDC*CTIC
C PTLV=PTLVAD
C DSMN=(DX*DX+DY*DY)
C OFPOLY=SQRT(DSMN)*.0001
C XINC=SQRT(DSMN)*.05
C DSMN=DSMN*.02
C OFFSET=CTIC*.001
C ERROR=CTIC*.01

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ORDER=3
NBPTI=4
CTLW=FX(1,1)
CTHG=FX(1,1)
DO 21 IR=1,IRDM
DO 21 IC=1,ICDM
CTCK=FX(IR,IC)
IF(CTCK.LT.CTLW)CTLW=CTCK
IF(CTCK.GT.CTHG)CTHG=CTCK
21 CONTINUE
C *** DETERMINE FIRST CONTOUR TO BE LABELLED
IF(XMDC.EQ.0.) GO TO 300
DO 310 IA=1,100
IF(PTLV.GE.CTLW) GO TO 301
310 PTLV=PTLV+PTLVAD
301 CONTINUE
C
IF(CTIC.GT.1.) GO TO 300
PTLV=PTLV+100.
PTLVAD=PTLVAD*100.
300 CONTINUE
C *** DETERMINE NUMBER OF CONTOURS
NBCT=CTLW/CTIC
IF(NBCT.GT.-1) NBCT=NBCT+1
CTLV=FLOAT(NBCT)*CTIC
NBHG=CTHG/CTIC
IF(NBHG.LT.0) NBHG=NBHG-1
NBCT=NBHG-NBCT+1
NBCT=1
PRINT,'LM,HG,IC,NBCT*,CTL*,CTHG,CTIC,CTLV,NBCT
DO 9 IR=1,IRDM
WRITE(6,111)IFXY(IR,IC),IC-1,ICDM)
9 CONTINUE
111 FORMAT(' ',10F10.2)
C *** FIND EACH CONTOUR, STARTING WITH SMALLEST CTLV
DO 26 IXCT=1,NBCT
PRINT,'#####'*,CTLV
C *** CHECK IF ANY NODE HAS A VALUE EQUAL TO CTLV, IF SO CHANGE BY OFFSET
DO 22 IR=1,IRDM
DO 22 IC=1,ICDM
IF(IFY(IR,IC).EQ.CTLV) FXY(IR,IC)=FXY(IR,IC)+OFFSET
22 CONTINUE
C *** SEARCH ALONG BOUNDARIES FOR SQUARES CONTAINING CONTOUR EXTREMITIES
NBCT=0
NBPT=0
NBSQ=0
NBSQFR=0
IRMX=IRDM-1
ICMX=ICDM-1
C *** SEARCH ALONG VERTICAL BOUNDARIES
IC=1
IKBD=0
DO 100 IP=1,2
DO 110 IR=1,IRMX
IF((FXY(IR,IC).LT.CTLV.AND.FXY(IR+1,IC).LT.CTLV).OR.
(FXY(IR,IC).GT.CTLV.AND.FXY(IR+1,IC).GT.CTLV)) GO TO 110
NBSQ=NBSQ+1
IXSQ=(NBSQ)=IKBD*IRMX+IR
110 CONTINUE
IC=ICDM
100 IKBD=ICMX-1
C *** SEARCH ALONG HORIZONTAL BOUNDARIES
IR=1
IKBD=0
DO 120 IP=1,2
DO 130 IC=1,ICMX
IF((FXY(IR,IC).LT.CTLV.AND.FXY(IR,IC+1).LT.CTLV).OR.
(FXY(IR,IC).GT.CTLV.AND.FXY(IR,IC+1).GT.CTLV)) GO TO 130
NBSQ=NBSQ+1
IXSQ=(NBSQ)=(IC-1)*IRMX+IR-1
130 CONTINUE
IR=IRDM
120 IKBD=1
IF(NBSQ.EQ.0) GO TO 151
C *** START SEARCH WITH BOUNDARY SQUARES CONTAINING CONTOUR EXTREMITIES
DO 150 IXSQ=1,NBSQ
IC=(IXSQ-(IXSQ)-1)/IRMX
IR=(IXSQ-(IXSQ)-1)/IRMX
IC=IC+1
C *** CHECK IF SQUARE HAS ALREADY BEEN SEARCHED. IN THIS PART OF THE
C *** PROBLEM ALL THE SEARCHES ARE SUCCESSFUL AND SEARCH GOES FROM
C *** SQUARE TO SQUARE. ONLY FURTHER IS SEARCH EXHAUSTIVE
IF(NBPT.EQ.0) GO TO 161
DO 160 IXPT=1,NBPT
IF(IXSQRM(IXPT).EQ.IR.AND.ISQCUL(IXPT).EQ.IC) GO TO 150
160 CONTINUE
161 CONTINUE
C *** THIS SQUARE IS THE FIRST OF THE CURVE. TWO INTERSECTIONS SHOULD
C *** BE FOUND : ONE WITH BOUNDARY AND THE OTHER LEADING TO THE SEARCH
C *** WITHIN THE GRID
1 IR,IC,IRMX,ICMX
CALL SQUAR(FXY,IRDM,ICDM,IR,IC,0,CTLV)
IP=NBPTSQ+1
GO TO (150,11,12,11,14),IP
C *** ONE OR THREE POINTS: END CURVE
11 NBCV=NBCV+1
IXCVST(NBCV)=-(NBPT-IP+1)
GO TO 150
C *** FOUR POINTS, SKIP THE SQUARE AT THIS TIME NOT TO START
C *** SEARCH WITH FOUR POINT SQUARE
14 NBPT=NBPT-4
GO TO 150
C *** TWO POINTS : CHECK WHICH ONE IS ON THE BOUNDARY. SET IT FIRST
C *** IN ARRAY. START SEARCH WITHIN THE GRID
12 IPTBND=0
DO 15 IP=1,2
IF((IR.EQ.1.AND.ISIDE(IP).EQ.1).OR.

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1  (IR,EW,IRMX,AND,ISIDE(IP),EQ,3),OK,
2  (IC,EW,1,AND,ISIDE(IP),EQ,4),OR,
3  (IC,EW,ICM,AND,ISIDE(IP),EQ,2)) (PT=NU=IPTBN)+IP
15 CONTINUE
C
IP=IPTBND+1
GO TO (30,31,32,31),IP
GO TO 150
C
C *** THE SECOND POINT IS ON THE BOUNDARY
32 STORX=CONTX(NBPT)
STORY=CUNTY(NBPT)
CO,IX(NBPT)=CONTX(NBPT-1)
CONY(NBPT)=CONY(NBPT-1)
CONX(NBPT-1)=STORX
CONY(NBPT-1)=STORY
ISIDE(2)=ISIDE(1)
C
C *** THE FIRST POINT IS ON THE BOUNDARY (OR BOTH, IP=3)
C *** SET COUNTER TO NEGATIVE VALUE SINCE CURVE IS OPEN
31 IXCVST(NBCV)=- (NBPT-1)
IS=ISIDE(2)
C
CALL SEARCH(FXY,IRDM,ICDM,IR,IC,IS,NBCV,CTLV)
C
C *** SEARCH FOR PRESENT CURVE COMPLETE. GO TO NEXT CURVE
C *** AT SAME CTLV
C
ICNT=IABS(IXCVST(NBCV))
GO TO 150
C
150 CONTINUE
151 CONTINUE
C
C *** SEARCH EXHAUSTIVELY THE REMAINING SQUARES
DO 170 ICDD=1,ICM
IC=ICDD
IR=IRDD
C
C *** TEST IF THE SQUARE HAS BEEN SEARCHED BEFORE
IF(NBPT,EQ,0) GO TO 172
DO 171 IXPT=1,NBPT
IF(ISURJ(IXPT),EQ,IR,AND,(ISJCOL(IXPT),EQ,IC) GOTO 170
171 CONTINUE
172 CONTINUE
C
CALL SQUAR(FXY,IRDM,ICDM,IR,IC,0,CTLV)
C
IF(NBPTS,NE,2) GO TO 170
NBCV=NBCV+1
IXCVST(NBCV)=NBPT-1
IS=ISIDE(2)
C
CALL SEARCH(FXY,IRDM,ICDM,IR,IC,IS,NBCV,CTLV)
ICNT=IABS(IXCVST(NBCV))
C
170 CONTINUE
C
C *** KEEP LAST POINT + 1 IN MEMORY

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IXCVST(NBCV+1)=NBPT+1
C
C *** ALL THE CURVES CORRESPONDING TO THIS CTLV ARE DETERMINED
C *** START PLOTTING, IF NBCV,EW,0 SKIP PLOTTING
IF(NBCV,EW,0) GO TO 10
C
C *** REPLACE POINTS THAT TOO CLOSE TOGETHER BY THEIR AVERAGES
IXST=IXCVST(1)
IXSK=0
DO 200 IXCV=1,NBCV
OPEN=.FALSE.
IAB=IABS(IXCVST(IXCV+1))-1
IF(IXST,GT,0) GO TO 250
OPEN=.TRUE.
IXST=-IXST
STXX=CONTX(IXST)
STYY=CONY(IXST)
IXKP=IXST-IXSK
IXST=IXST+1
250 STXX=CONTX(IXST)
STYY=CONY(IXST)
IXKP=IXST-IXSK
IXST=IXST+1
C
C *** CHECK DISTANCE BETWEEN POINTS OF THE SAME CURVE
DO 210 IX=IXST,IXED
IXNW=IX-IXSK-1
DXX=CONTX(IX)-STXX
DYY=CONY(IX)-STYY
IF((DXX*DXX)+(DYY*DYY).GT,DSMN) GO TO 220
CONTX(IXNW)=STXX+.5*DXX
CONY(IXNW)=STYY+.5*DYY
STXX=CONTX(IXNW)
STYY=CONY(IXNW)
IXSK=IXSK+1
GO TO 210
C
220 CONTX(IXNW)=STXX
CONY(IXNW)=STYY
STXX=CONTX(IX)
STYY=CONY(IX)
210 CONTINUE
C
IF(OPEN) GO TO 225
DXX=CONTX(IXKP)-CONTX(IXED)
DYY=CONY(IXKP)-CONY(IXED)
IF((DXX*DXX)+(DYY*DYY).GT,DSMN) GO TO 225
CONTX(IXKP)=CONTX(IXKP)+.5*DXX
CONY(IXKP)=CONY(IXKP)+.5*DYY
IXSK=IXSK+1
C
C *** STORE LAST POINT OF CURVE AND START NEW ONE
225 CONTX(IXNW+1)=CONTX(IXED)
CONY(IXNW+1)=CONY(IXED)
IXCVCK=IXCVST(IXCV+1)
IXST=IXCVCK
IF(IXCVCK,GT,0) IXCVCK=IXCVCK-IXSK
IF(IXCVCK,LT,0) IXCVCK=IXCVCK+IXSK
IXCVST(IXCV+1)=IXCVCK
IXPP=IABS(IXCVCK)-1
200 CONTINUE
C
C *** PLOT THE CURVES, CHECK WHETHER LEVEL VALUES SHOULD BE PLOTTED
INLV=0
IF(IPTLV,EW,0) GO TO 201

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CTCK=CTLV
IF(CTIC.LT.1.) CTCK=CTCK*100.
IF(CTCK.LT.PTLV) GO TO 201
INLV=CTCK*.1
PTLV=PTLV+PTLVAD
201 CALL PLOT(AMODES,NBCV,INLV,XINC,DELTA)
10 CONTINUE
C
CTLV=CTLV*CTIC
26 CONTINUE
RETURN
END
SUBROUTINE SEARCH(FXY,IRDM,ICDM,IR,IC,IS,NBCV,CTLV)
C *** THE SUBROUTINE STARTS WITH A RECTANGLE HAVING ONE INTERSECTION WITH
C *** WITH CONTOUR AND SEARCHES THE WHOLE CONTOUR
C
COMMON /PLTT/ CUNT(100),CONTY(100),DX,DY,ERROR,DFP*2LY
COMMON NBPT,NBPTSQ,ISIDE(4),LORDER,NBPTIT,NBSQFR,
1 ISQRW(100),ISCOL(100),IXCVST(20),IXSUBD(20)
DIMENSION FXY(IRDM,ICDM)
C *** DETERMINE NEXT SQUARE TO SEARCH
20 CONTINUE
GO TO (31,32,33,34),IS
31 IR=IR-1
GO TO 35
32 IC=IC+1
GO TO 35
33 IR=IR+1
GO TO 35
34 IC=IC-1
C *** TEST IF THE SQUARE IS WITHIN THE MESH. IF NOT END CUREVE
C *** START NEXT CURVE FOR SAME CTLV
35 CONTINUE
IF(IC.LT.1.OR.IC.GT.(ICDM-1).OR(IR.LT.1.OR.IR.GT.(IRDM-1)))GO TO 9
C *** DETERMINE IF CURVE CLOSES ON ITSELF. IF SO DELETE LAST POINT
(SAME AS FIRST), END CURVE AND GO TO NEXT ONE FOR SAME CTLV
ISCK=IABS(IXCVST(NBCV))
IF(IR.EQ.ISQRW(ISCK).AND.IC.EQ.ISCOL(ISCK))
1 GO TO 8
C *** DETERMINE SIDE NOT TO BE SEARCHED IN NEXT SQUARE
ISKIP=IS*2
IF(ISKIP.GT.4)ISKIP=ISKIP-4
C *** SEARCH NEXT SQUARE
CALL SQUAR(FXY,IRDM,ICDM,IR,IC,ISKIP,CTLV)
IP=NBPTSQ+1
GO TO (7,11,7,6),IP
C *** NO POINTS ARE FOUND, THERE IS A PROBLEM
C *** END CURVE AND GO TO THE NEXT ONE FOR THE SAME CTLV
C *** ONE POINT FOUND, FLAG THE SIDE AND CONTINUE SEARCH
11 IS=ISIDE(1)
GO TO 20
GO TO 10

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C *** CHANGE INDEX OF FIRST POINT ON THE CURVE TO LAST SQUARE
C *** SUCH THAT EACH SQUARE IS FLAGGED(FOR EXHAUSTIVE SEARCH)
ISQRW(ISCK)=ISQRW(NBPT)
ISCOL(ISCK)=ISCOL(NBPT)
NBPT=NBPT-1
GO TO 10
CALL FOURPT(IR,IC,ISKIP)
IS=ISIDE(1)
GO TO 2J
C *** TWO POINTS ARE FOUND THERE IS A PROBLEM
C *** THREE POINTS ARE FOUND : TWO CURVES IN THE SAME SQUARE
C *** CALL ROUTINE TO HANDLE THAT PROBLEM
C
10 RETURN
END
SUBROUTINE FOURPT (IR,IC,ISKIP)
COMMON /PLTT/ CONTX(100),CONTY(100),DX,DY,ERROR,DFPOLY
COMMON NBPT,NBPTSQ,ISIDE(4),LORDER,NBPTIT,NBSQFR,
1 ISQRW(100),ISCOL(100),IXCVST(20),IXSUBD(20)
DIMENSION XX(3),YY(3),DIST(2),IRFR(20),ICFR(20),ISFRI(20),
1 ISFR(20)
DG=90./1.57095
NBPT=NBPT-3
C *** TRANSFER LAST 3 POINTS IN SMALL ARRAY
DO 10 IX=1,3
XX(IX)=CONTX(NBPT+IX)
10 YY(IX)=CONTY(NBPT+IX)
C *** CHECK IF THIS RECTANGLE HAS BEEN SEARCHED
IF(NBSQFR.EQ.0) GO TO 11
DO 20 IX=1,NBSQFR
IXSQ=IX
IF(IR.EQ.IRFR(IXSQ).AND.IC.EQ.ICFR(IXSQ)) GO TO 30
20 CONTINUE
C *** THIS SQUARE HAS NOT BEEN SEARCHED, REJECT THE POINT
C *** OPPOSITE TO ISKIP AND CHOOSE THE ONE CLOSEST TO THE
C *** STRAIGHT LINE (NBPT-1)-(NBPT)
11 ISOP=ISKIP*2
IF(ISOP.GT.4) ISOP=ISOP-4
C *** MOVE THE 2 POINTS NOT OPPOSITE IN XX,YY 1 AND 2
DO 40 IX=1,2
IF(ISIDE(IX).NE.ISOP) GO TO 40
XX(IX)=XX(3)
YY(IX)=YY(3)
ISIDE(IX)=ISIDE(3)
GO TO 50
40 CONTINUE
C *** COMPUTE ANGLE, CHOOSE SMALLEST
50 XZ=CONTX(NBPT)
YZ=CONTY(NBPT)
X1=CONTX(NBPT-1)
Y1=CONTY(NBPT-1)

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NBPTSQ=J
DO 20 IS=1,4
IF (IS.EQ.ISKIP) GO TO 20
GO TO (11,12,13,14), IS
IF ((FXY(IR,IC+1))-LT.CTLV.AND.FXY(IR,IC)-LT.CTLV).OR.
1 IF (FXY(IR,IC+1))-GT.CTLV.AND.FXY(IR,IC)-GT.CTLV) GO TO 20
NBPT = NBPT+1
NBPTSQ=NBPTSQ+1
ISIDE(NBPTSQ) = 1
CALL INTRP(FXY,IRDM,ICDM,IR,IC,1,CTLV,XINTER)
CONTX(NBPT)=XINTER
CONTY (NBPT) = (IR-1)*DX
GO TO 10
12 IF ((FXY(IR+1,IC+1))-LT.CTLV.AND.FXY(IR,IC+1)-LT.CTLV).OR.
1 (FXY(IR+1,IC+1))-GT.CTLV.AND.FXY(IR,IC+1)-GT.CTLV) GO TO 20
NBPT = NBPT+1
NBPTSQ=NBPTSQ+1
ISIDE(NBPTSQ) = 2
CALL INTRP(FXY,IRDM,ICDM,IR,IC,2,CTLV,XINTER)
CONTX(NBPT)=XINTER
CONTY (NBPT) = IC*DX
GO TO 10
13 IF ((FXY(IR+1,IC+1))-LT.CTLV.AND.FXY(IR+1,IC)-LT.CTLV).OR.
1 (FXY(IR+1,IC+1))-GT.CTLV.AND.FXY(IR+1,IC)-GT.CTLV) GO TO 20
NBPT = NBPT+1
NBPTSQ=NBPTSQ+1
ISIDE(NBPTSQ) = 3
CALL INTRP(FXY,IRDM,ICDM,IR,IC,3,CTLV,XINTER)
CONTX(NBPT)=XINTER
CONTY (NBPT) = IR*DX
GO TO 10
14 IF ((FXY(IR+1,IC)-LT.CTLV.AND.FXY(IR,IC)-LT.CTLV).OR.
1 (FXY(IR+1,IC))-GT.CTLV.AND.FXY(IR,IC)-GT.CTLV) GO TO 20
NBPT = NBPT+1
NBPTSQ=NBPTSQ+1
ISIDE(NBPTSQ) = 4
CALL INTRP(FXY,IRDM,ICDM,IR,IC,4,CTLV,XINTER)
CONTX(NBPT)=XINTER
CONTY (NBPT) = (IC-1)*DX
10 ISQRW(NBPT)=IR
ISQCOL(NBPT)=IC
C 20 CONTINUE
RETURN
END
SUBROUTINE INTRP(FXY,IRDM,ICDM,IR,IC,IS,CTLV,XINTER)
COMMON /PLTT/ CONTX(100),CONTY(100),DX,DY,ERROR,OFFPOLY
COMMON NBPT,NBPTSQ,ISIDE(4),IORDER,NBPTIT,NBSQFR,
1 ISQRW(100),ISQCOL(100),IXCVST(20),IXSQBD(20)
REAL FXY(IRDM,ICDM),MESH,X(1:5),OR(10,5),AL(5),DUM(10),F(5)
C *** DETERMINE REFERENCE CONER, DIRECTION OF INTERPOLATION
C *** AND MESH BOUNDARIES
GO TO (1,2,12),IS
C *** HORIZONTAL SIDE
1 MESH=DX
IREF=IC
IFCT=IR
IBND=ICDM
GO TO 5

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DXX=X2-X1
DYY=Y2-Y1
IF(DXX.EQ.0.) GO TO 100
A1=ATAN(DYY/DXX)
A1DG=A1*DG
GO TO 110
C *** THE SEGMENT IS VERTICAL A1=90.
C 100 A1=1.57095
C *** COMPUTE DIFFERENCE OF SLOPES BETWEEN PREVIOUS 2 POINTS (SEGMENT)
C *** AND SGMENT (NBPT, NBPT+1)
110 DO 130 IX=1,2
XC=XX(IX)
YC=YY(IX)
A2=ATAN((YC-Y2)/(XC-X2))
A2DG=A2*DG
DIST(IX)=ABS(A2-A1)
DSDG=DIST(IX)*DG
130 CONTINUE
C *** CHOOSE SMALLEST ONE, STORE REJECTED POINTS FOR LATER
IXCK=1
IF(DIST(1)-GT.DIST(2)) IXCK=2
NBPT=NBPT+1
CONTX(NBPT)=XX(IXCK)
CONTY(NBPT)=YY(IXCK)
C *** STORE INFORMATION FOR LATER
NBSQFR=NBSQFR+1
IRFR(NBSQFR)=IR
ICFR(NBSQFR)=IC
IX=1
IF(IXCK.EQ.1) IX=2
ISFR1(NBSQFR)=ISOP
ISFR2(NBSQFR)=ISIDE(IX)
ISIDE(1)=ISIDE(IXCK)
GO TO 99
C *** THIS RECTANGLE HAS ALREADY BEEN SEARCHED. TWO POINTS HAVE
C *** BEEN CHOSEN. CHOOSE THE ONE DIFFERENT FROM ISKIP
30 ISCK=ISFR1(IXSQ)
IF(ISKIP.EQ.ISCK) ISCK=ISFR2(IXSQ)
DO 60 IX=1,3
IXPT=IX
IF(ISCK.EQ.ISIDE(IXPT)) GO TO 70
60 CONTINUE
C *** STORE THE POINT IN VECTOR
70 NBPT=NBPT+1
CONTX(NBPT)=XX(IXPT)
CONTY(NBPT)=YY(IXPT)
ISIDE(1)=ISIDE(IXPT)
C 99 RETURN
END
SUBROUTINE SQAR(FXY,IRDM,ICDM,IR,IC,ISKIP,CTLV)
DIMENSION FXY(IRDM,ICDM)
COMMON /PLTT/ CONTX(100),CONTY(100),DX,DY,ERROR,OFFPOLY
COMMON NBPT,NBPTSQ,ISIDE(4),IORDER,NBPTIT,NBSQFR,
1 ISQRW(100),ISQCOL(100),IXCVST(20),IXSQBD(20)

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C *** VERTICAL SIDE
2 MESH=DY
  IREF=IR
  IFCT=IC
  I8ND=IRDM
C *** CHECK THAT ALL INTERPOLATION POINTS ARE WITHIN THE MESH
C *** IF NOT END INTERPOLATION AT THE BOUNDARY AND DECREASE
C *** CORRESPONDINGLY THE ORDER OF THE POLYNOMIAL
5 IPA0=(NBPTIT-2)/2
  ILOW=IREF-IPAD
  IHIGH=IREF+IPAD+1
  IF(ILOW.LT.1) ILOW=1
  IF(IHIGH.GT.I8ND) IHIGH=I8ND
  NPTS=IHIGH-ILOW+1
  IORDMX=NPTS+1-(NBPTIT-IORDER)
  IF(IORDMX.EQ.0) IORDMX=1
C *** DETERMINE SIDE INDEX AND EVALUATE POLYNOMIAL AT THE POINTS
  IF(I5.EQ.2.OR.I5.EQ.3) IFCT=IFCT+1
  ICNT=0
  DO 10 IXPT=ILOW, IHIGH
    ICNT=ICNT+1
    XI(ICNT, I)=1.
    IF(IORDMX.LT.2) GO TO 7
    DO 6 IORU=2, IORDMX
      6 XI(ICNT, IORD)= (FLOAT(ICNT-I)*MESH)**(IORD-1)
    7 CONTINUE
    GO TO (11, 12, 11, 12), I5
    11 F(ICNT)=FXI(IFCT, IXPT)
    GO TO 10
    12 F(ICNT)=FXI(IXPT, IFCT)
    10 CONTINUE
C *** CALL SOLVER
  CALL LSJ (X, QR, A, F, DUM, NPTS, I0, IORDMX, 5, 1, 630)
  XINTER = (FLOAT(IREF-ILOW)+5)*MESH
C *** EVALUATE POLYNOMIAL AT A FIRST GUESSED POINT AND
C *** OBTAIN ROOT BY ITERATION (CHECK AGAINST ERROR)
  IORDMX=IORDMX-1
  DO 20 ICNT=1, 20
    DERVFP = A(I2)
    FPOLY = A(1)
    DO 21 I = 1, IORDMX
      FPOLY = FPOLY + A(I+1)*XINTER**(I)
    21 CONTINUE
    DO 22 I=2, IORDMX
      DERVFP = DERVFP + FLOAT(I)*A(I+1)*XINTER**(I-1)
    22 CONTINUE
    DIFF=CTLV-FPOLY
    IF (ABS(DIFF).LT.ERROR) GO TO 25
    XINTER = XINTER + DIFF /DERVFP
  20 CONTINUE
  25 XINTER=XINTER+FLOAT(ILOW-I)*MESH
  GO TO 31
  31 RETURN
  END
SUBROUTINE LSQA(JR, X, B, AX8NRM, M, MDIM, N, MDIM, N, MDIM, P, *)
  INTEGER N, Np, PIVOT(100), MDIM, MDIM, I, J, K
  REAL A(MDIM, N), X(MDIM, P), B(MDIM, P), QR(MDIM, N), ALPHA(100), Y(100),
  IE(100), R(100), DBL*8, DBLSUM*8, YNORM, ENORM, AX8NRM(P), RNRSQ
  DO 100 I=1, M
    DO 1 J=1, N
      DO 1 I=1, M
        DO 1 I=1, M
          QR(I, J)=A(I, J)
          DO 10 K=1, P
            AX8NRM(K)=1.
            R(I)=B(I, K)
          CALL SLV(M, MDIM, N, QR, ALPHA, PIVOT, R, Y)
          DO 3 I=1, M
            DBLSUM=-8(I, K)
            DO 33 J=1, N
              DBL=A(I, J)
              DBLSUM=DBLSUM + DBL*Y(J)
            33 R(I)=-DBLSUM
          CALL SLV(M, MDIM, N, QR, ALPHA, PIVOT, R, E)
          YNORM=0.
          ENORM=0.
          DO 4 I=1, N
            YNORM=YNORM + Y(I)**2
            ENORM=ENORM + E(I)**2
          IF(ENORM.GT.YNORM*0.0625) GO TO 10
          DO 6 I=1, N
            Y(I) = Y(I) + E(I)
          DO 7 I=1, M
            DBLSUM = -8(I, K)
            DO 77 J=1, N
              DBL=A(I, J)
              DBLSUM=DBLSUM+DBL*Y(J)
            77 R(I)=-DBLSUM
          CALL SLV(M, MDIM, N, QR, ALPHA, PIVOT, R, E)
          ENORM=0.
          DO 8 I=1, N
            ENORM=ENORM + E(I)**2
          IF(ENORM.GT.1.58E-8*YNORM) GO TO 5
          DO 9 I=1, N
            XI(I, K) = Y(I)
          9 CONTINUE
          RNRSQ=0.
          DO 11 I=1, M
            RNRSQ=RNRSQ+R(I)**2
          AX8NRM(K)=SQRT(RNRSQ)
          10 CONTINUE
          RETURN
          30 RETURN
        END
        SUBROUTINE SLV(M, MDIM, N, QR, ALPHA, PIVOT, R, Y)
          INTEGER M, N, PIVOT(100), PIVOTI, MDIM, I, J, IJ, II, I1, NMINI
          REAL QR(MDIM, N), ALPHA(N), R(N), Y(100), Z(100), DBL*8, DBLSUM*8, GAMMA
          DO 1 J=1, N
            DBLSUM=0.
            DO 11 I=J, M
              DBL=QR(I, J)
              DBLSUM=DBLSUM+DBL*R(I)
            GAMMA=UBLSUM/ALPHA(IJ)*QR(I, J)
            DO 12 I=J, M
              R(IJ)=R(IJ)+GAMMA*QR(I, J, J)
            12 CONTINUE
          1 CONTINUE

```

```

Z(N)=R(N)/ALPHA(N)
NMINI=N-1
DO 2 I=1,NMINI
  I=N-I
  I1=I+1
  DBLSUM=-R(I)
  DO 22 J=I1,N
    DBL=QR(I,K)
    DBLSUM=DBLSUM+DBL*(I,J)
  DO 3 I=1,N
    PIVOTI=PIVOT(I)
    Y(PIVOTI)=Z(I)
  RETURN
END
SUBROUTINE DCP (M,MDIM,N,QR,ALPHA,PIVOT,*)
  INTEGER M,N,PIVOT(100),MDIM,I,J,K,JBAR,K1
  REAL QR(MDIM,N),ALPHA(N),SUM(100),Y(100),DBLSUM*8,DBL*8,BETA,
  1 SIGMA,ALPHA,QRKK
  DO 1 J=1,N
    PIVOT(J)=J
    DBLSUM = 0.
    DO 11 I=1,M
      DBL = QR(I,J)
      DBLSUM = DBLSUM + DBL**2
    DO 8 K=1,N
      SIGMA = SUM(K)
      JBAR = K
      IF(K.EQ.N) GO TO 5
      K1=K+1
      DO 2 J=K1,N
        IF(SIGMA.GE.SUM(J))GO TO 2
        SIGMA = SUM(J)
        JBAR = J
      CONTINUE
    IF(JBAR.EQ.K)GO TO 5
    I=PIVOT(K)
    PIVOT(K)=PIVOT(JBAR)
    PIVOT(JBAR)=I
    SUM(JBAR)=SUM(K)
    SUM(K)=SIGMA
    DO 4 I=1,M
      SIGMA= QR(I,K)
      QR(I,K)= QR(I,JBAR)
      QR(I,JBAR)=SIGMA
      DBLSUM=0.
      DO 55 I=K,M
        DBL=QR(I,K)
        DBLSUM=DBLSUM+DBL**2
      SIGMA=DBLSUM
      IF(SIGMA.EQ.0.) RETURN1
      QRKK=QR(K,K)
      ALPHA=SQRT(SIGMA)
      IF(QRKK.GE.0.)ALPHA=-ALPHA
      ALPHA(N)=ALPHA
      QR(K,K)=QRKK-ALPHA
      BETA=1./ISIGMA-QRKK*ALPHA
      IF(K.EQ.N) GO TO 8
      DO 6 J=K1,N
        DBLSUM=0.

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```

DO 66 I=K,M
  DBL=QR(I,K)
  DBLSUM=DBLSUM+DBL*(I,J)
66 6 Y(J)=BETA*DBLSUM
  DO 7 J=K1,N
    DO 72 I=K,M
      QR(I,J) = QR(I,J)-QR(I,K)*Y(J)
      SUM(J)=SUM(J)-QR(K,J)**2
    8 CONTINUE
  RETURN
END
SUBROUTINE PLOT(AMODES,NBCV,INLV,XINC,DELTA5)
  COMMON /PLIT/ CONTX(100),CDATY(100),DX,DY,ERROR,OPPLY
  COMMON NBPT,NBPTSQ,ISIDEI(4),IORDER,NBPTIT,NBSQFR,
  1 ISQRDM(100),ISQCOL(100),IXCVST(20),IXSQBD(20)
  DIMENSION AMODES(200)
  DIMENSION CNTX(100),CNTY(100),XX(3),YY(3)
  DIMENSION TPNX(100),TPNY(100),TPXD(100),TPYD(100)
  1,TPXNH(100),TPXNH(100),TPXDH(100),TPYDH(100)
  LOGICAL*1 OPEN,SKIP,NOAVG,VERTI,LEVEL
  DEG=180./3.14159
  PI=3.14159
  P12=3.14159**5
C *** DO LOOP ON THE NUMBER OF CURVES AT THIS CTLY
C DO 100 IXCV=1,NBCV
C
C *** DETERMINE CURVE PARAMETERS (OPEN,CLOSE,FIRST AND LAST POINT)
  LEVEL=.FALSE.
  IF(INLV.NE.0) LEVEL=.TRUE.
  OPEN=.FALSE.
  SKIP=.TRUE.
  IXST=IXCVST(IXCV)
  IXLW=IXST
  IXAD=1
  IF(IXST.LT.0) OPEN=.TRUE.
  IF(.NOT.OPEN) GO TO 109
  IXAD=-1
  SKIP=.FALSE.
  NOAVG=.TRUE.
  109 IXST=IABS(IXST)
  IXHG=IABS(IXCVST(IXCV+1))-1
C
C *** IF CURVE HAS TWO POINTS OR LESS, DO NOT PLOT
  NBPTCV=IXHG-IXST
  ILPT=0
  IF(NBPTCV.LT.2) GO TO 100
  IF(NBPTCV.LT.5)LEVEL=.FALSE.
  IF(NBPTCV.GT.5) ILPT=NBPTCV/2
  IF(NBPTCV.GT.20)ILPT=NBPTCV/3
  IXLT=IXHG+IXAD
  XOROD=0.
  YOROD=0.
  WM=0.
  WMNH=1.
  WMOD=0.
  XYOD=0.
  YYOD=0.
  SLOD=0.
  SNOD=0.

```

```

C *** TO THE MAIN SYSTEM. OBTAIN WEIHTED AVERAGES CNTX(I),CNTY(I)
C *** XXOD,YYOD AND XXNW,YYNW
C 400 XXBG=0.
      IPT=1
      NBIC=ABS(XXZ/XINC)+1
      XNBIC=NBIC
C *** DETERMINE HEIGHTING FACTOR AND ANGLE OF NEW WITH RESPECT
C *** TO OLD SYSTEM
C *** IF(NDAVG) GO TO 430
      IF(NBIC.EQ.0) GO TO 430
      WM=1./XNBIC
      WMOD=1.-WM
      WWIN=WM
      SLNO=SLNW-SLOD
      SLDNO=SLNO*DEG
      CSNO=COS(SLNO)
      SNNO=SIN(SLNO)
C *** DETERMINE IN OLD SYSTEM THE SLOPE OF THE VERTICAL IN NEW
      VERTI=.TRUE.
      IF(SLNO.EQ.0..OR.ABS(SLNO).EQ.PI) GO TO 430
      VERTI=.FALSE.
      AVT=TAN(SLNO+PI2)
      SLVT=SLNO+PI2
      SLOVT=SLVT*DEG
C *** DETERMINE INTERSECTIONS IN NEW AXIS
C 430 CNTX(I)=XX(I*PT)
      CNTY(I)=YY(I*PT)
      NBIC(T)=NBIC+1
      IF(NBIC.LT.2) GO TO 410
      XXIC=XX2/XNBIC
      DO 420 IXIC=2,NBIC
      XXNW=XXBG+FLOAT(IXIC-1)*XXIC
      YYNW=A3NW+XNW*XXNW+XNW*A2NW*XXNW
      TPXN(IXIC)=XXNW
      TPN(IXIC)=YYNW
C
      TPXN(IXIC)=XORNW+XXNW*CSNW-YYNW*SNNW
      TPN(IXIC)=YORNW+XXNW*SNNW+YYNW*CSNW
C *** DO NOT AVERAGE IF IT IS THE FIRST OR LAST LEG OF AN OPEN CURVE
      IF(NDAVG) GO TO 480
C
C *** DETERMINE THE COORD OF XXNW,U. IN OLD SYSTEM TO OBTAIN
C *** THE B COEFF. (BVT) IN OLD OF VERTICAL IN NEW
      XXOD=DXND+XXNW*CSNO
      IF(VERTI) GO TO 425
      YYOD=DYND+XXNW*SNNO
      BVT=YYOD-AVT*XXOD
      AA=AVT-A2OD
      RTAD=(SQRT(AA*AA+.A3OD*8VT))/12.*A3OD
      AA=AA+.5/A3OD
      XXOD=AA*RTAD
      XXX=AA-RTAD
      IF(XXOD.LT.0.) GO TO 426
      IF(1XOD.LT.DXND).OR.
      1 (XXOD.GT.(DXN3))XXOD=AA-RTAD
      GO TO 425
C
C 426 IF(1XO).GT.DXND).OR.

```

```

      CSOD=0.
      IPIC=0
C *** TRANSFER THE POINTS OF INTEREST INTO A SMALL ARRAY
C 110 IXED=IXST
      IPIC=IPIC+1
      IF(IXED.GT.IXMG)IXED=IXLM
      DO 300 IX=1,3
      XX(IX)=CONTX(IXED)
      YY(IX)=CNTY(IXED)
      IXED=IXED+1
      IF(IXED.GT.IXHG) IXED=IXLM
      300 CONTINUE
C
C *** OPERATE TRANSFER OF AXIS
      XMD=(XX(3)+XX(1))*5
      YMD=(YY(3)+YY(1))*5
      YDU=YMD-YY(2)
      XDU=XX(2)-XXMD
C *** CHECK IF AXIS SHOULD BE VERTICAL
      IF(ABS(YYDU).GE.OFFPOLY) GO TO 310
      SLNW=PI2
C *** IF MD AND Z COINCIDE, THE SLOPE IS TO BE DETERMINED
      BETWEEN MD AND 1
      IF(ABS(XXDU).GE.OFFPOLY) GO TO 311
      YDU=YMD-YY(1)
      XDU=XX(1)-YMD
      IF(ABS(YYDU).GE.OFFPOLY)SLNW=ATAN(XXDU/YYDU)
      GO TO 311
C
C 310 SLNW=ATAN(XXDU/YYDU)
C
C *** DETERMINE COORDINATES OF THE THREE POINTS IN NEW SYTEM
C 311 XORNW=XX(1)
      YORNW=YY(1)
      CSNW=COS(SLNW)
      SNNW=SIN(SLNW)
      XX2=(XX(2)-XX(1))*CSNW+(YY(2)-YY(1))*SNNW
      IF(ABS(XX2).LT.OFFPOLY) XX2=SIGN(OFFPOLY,XX2)
      YY2=-(XX(2)-XX(1))*SNNW+(YY(2)-YY(1))*CSNW
      YY3=-(XX(3)-XX(1))*SNNW+(YY(3)-YY(1))*CSNW
      IF(YY2-.5*YY3).EQ.0.) YY2=YY2+OFFPOLY
C
      SLDNW=SLNW*DEG
C
C *** OBTAIN THE COEFFICIENTS OF THE 2 ORDER POLYNOMIAL
C *** PASSING THROUGH THE 3 POINTS, A1NW=0.
      A2NW=(2.*YY2-.5*YY3)/XX2
      A3NW=(.5*YY3-YY2)/(XX2*XX2)
C
C *** IF THIS IS THE FIRST LEG OF A CLOSED CURVE, GO TO NEXT POINT
      IF(SKIP) GO TO 120
C
C *** IN NEW AXIS, OBTAIN NUMBER OF INCREMENTS (NBIC), INCREMENT
C *** SIZE (XXIC), COORD OF POINT ON CURVE (XXNW,YYNW), TRANSFER
C *** TO MAIN AXIS (CNTX(I),CNTY(I)).
C *** IN OLD (PREVIOUS) SYSTEM, FIND THE CORRESPONDING OF X COORDINATE
C *** (XXOD) OF XXNW,YYNW. OBTAIN THE INTERSECTION OF THE VERTICAL AT
C *** AT XXOD WITH THE OLD CURVE XXOD,YYOD. TRANSFER THE INTERSECTION

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1 (XX00.LT.(DXX3))XX00=AA-RTAU
425 YY00=A300*XX00+XX00+A200*XX00
C
C *** TRANSFER TO MAIN AXIS
480 TPX0(I,XIC)=XX00
TPY0(I,XIC)=YY00
TPX0(I,XIC)=XOR00+XX00+CS00-YD00*SN00
TPY0(I,XIC)=YOR00+XX00+SN00+YY00*CS00
C
C *** COMPUTE NEW WEIGHTS
419 WNN=WN*W
W00=W00*W
420 CONTINUE
410 CNTX(NBIC)=XX(IPT+1)
CNTY(NBIC)=YY(IPT+1)
C
C *** PLOT ARRAY OF NBIC POINTS GOING FROM IXST TO IXST+1
C* WRITE(6,1000)(TPXN(I),TPYN(I),I=2,NBIC)
C* WRITE(6,1000)(TPXN(I),TPYN(I),I=2,NBIC)
C* WRITE(6,1000)(TPXN(I),TPYN(I),I=2,NBIC)
C* WRITE(6,1000)(TPXN(I),TPYN(I),I=2,NBIC)
1000 FORMAT(' ',10F10.3)
IF(.NOT.LEVEL) GO TO 119
C
C *** WRITE CONTOUR LEVEL
IF(XX(2).EQ.0.-OR.IPIC.LT.ILPT) GO TO 119
XLAB=DELTA*DELTA5
IPIC=0
XXCK=XX(IPT+1)
YYCK=YY(IPT+1)
DO 125 IX=2,NBIC
IXCK=NBIC-IX+1
DELTA=(XXCK-CNTX(IXCK))*IXCK-CNTX(IXCK)+
1 (YYCK-CNTY(IXCK))*YYCK-CNTY(IXCK)
IF(DELTA.GT.XLAB) GO TO 126
125 CONTINUE
C
126 ANGL=ATAN(CNTY(IXCK)-YYCK)/(CNTX(IXCK)-XXCK)
AMODES150=ANGL*180./3.14159
AMODES160=AMODES150
XLDY=DELTA5/1.5
XLDX=XLDY
C
IXLB=IXCK
IF(XX2.LT.0.)IXLB=NBIC-IXCK
XLLB=CNTX(IXLB)+ALDX*COS(ANGL)
YLLB=CNTY(IXLB)+XLDY*SIN(ANGL)
CALL INJMBG(AMODES,XLLB,YLLB,2,INLV)
C
NBIC=IXCK
IF(NBIC.LT.2) GO TO 120
119 CALL LINESG(AMODES,NBIC,CNTX,CNTY)
C
C *** MOVE TO NEXT POINT
C *** STORE PARAMETERS FROM N# TO JD
120 IXST=IXST+1

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```

A200=A2NW
A300=A3NW
XOR00=XORNW
YOR00=YORNW
SL00=SLNW
CS00=CSNW
SN00=SNNW
DXND=XX2
DYND=YY2
DXX3=2.*XX2
SKIP=.FALSE.
NOAVG=.FALSE.

```

```

C *** CHECK WHERE ONE IS ON THE CURVE
IF(IXST-IXLT) 110,150,100
C
C *** 110 : MIDDLE OF CURVE, CONTINUE INTERPOLATION
C
C *** LAST LEG
C *** IF (OPEN) DO NOT INTERPOLATE, USE PRESENT A'S AND PLOT
C *** IF (.NOT.OPEN) AVERAGE PRESENT A'S WITH A*ST
150 CONTINUE
IF(.NOT.OPEN) GO TO 110
XXBG=XX2
IPT=2
NOAVG=.TRUE.
WNN=1.
W00=0.
WM=0.
GO TO 430
C
C *** END OF CURVE, START NEXT ONE
100 CONTINUE
RETURN
END
$DATA

```



